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# 贵州省碧云湖岩溶渗漏水文地质条件与 渗漏管道位置识别

曾洁<sup>1,2</sup>,潘晓东<sup>1,2</sup>,任坤<sup>1,2</sup>,刘伟<sup>1,2</sup>,彭聪<sup>1,2</sup>,郑智杰<sup>1,2</sup>

(1. 中国地质科学院岩溶地质研究所/自然资源部、广西岩溶动力学重点实验室/联合国教科文组织国际岩溶研究中心, 广西桂林 541004; 2. 广西平果喀斯特生态系统国家野外科学观测研究站, 广西平果 531406)

**摘要:**查找岩溶区湖泊水库渗漏通道是岩溶区湖库建设和渗漏灾害治理的“卡脖子”问题。文章以碧云湖为例,分析湖区周边水文地质条件,结合示踪试验和地球物理探测方法,确定了碧云湖渗漏水文地质条件和渗漏主管道位置,并通过钻探验证。结果表明,碧云湖渗漏主要为岩溶管道型,渗漏点和地下河管道主要发育在靠近八步复式向斜轴的部位,位于物探测线测点 540 m 处,管道深约 40 m,渗漏段相对较窄,灌浆堵漏可行。根据水文地质调查、地形地貌条件和地球物理探测结果,建议在水库下游方向的物探测线开展帷幕灌浆,以测点 540 m 为中心向两边灌浆。

**关键词:**碧云湖;岩溶渗漏;水文地质;渗漏管道;地球物理探测技术

**创新点:**采用“岩溶地质查条件—水文地质查结构—探测试验定位置”的技术方法体系查明湖库渗漏位置,为处理水库病害提供了可行措施。

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## 0 引言

我国是世界上岩溶分布面积最广、发育类型最齐全的国家,岩溶分布面积约占国土面积的三分之一<sup>[1]</sup>,以贵州、云南、广西等西南地区最为典型且连片分布。岩溶作用形成的地下溶缝、溶洞、管道等规模巨大,为储存和调蓄水资源提供了有利空间,使岩溶水成为最丰富的地下水资源<sup>[2]</sup>。同时地下溶洞管道又成为岩溶区湖泊水库的渗漏通道,引起或诱发湖泊水库干涸、透水溃坝等重大灾害<sup>[3]</sup>,给水利工程建设带来了极大挑战,针对渗漏病害探查、除险加固等方法研究已成为病险水库治理的重要课题<sup>[4]</sup>。

岩溶区湖泊水库渗漏通道主要由地下溶缝、溶洞和管道构成,空间分布不均且探测难度大,一直是岩溶区湖库建设和渗漏灾害治理“卡脖子”问题<sup>[5]</sup>。随着社会经济发展,岩溶水文地质等综合研究在工程建设中应用越来越重要<sup>[6]</sup>。针对复杂的岩溶地质特征,众多学者从岩溶水文地质条件分析<sup>[7-9]</sup>、数值模拟<sup>[10-12]</sup>、地球物理探测<sup>[13-14]</sup>、环境同位素<sup>[15-16]</sup>及综合水文地质调查<sup>[17-19]</sup>等方面着手对水库渗漏做了大量研究并提出相应处理措施<sup>[20]</sup>,为病险水库治理提供了依据。

本文以贵州省织金县碧云湖为例,通过开展湖区及周边地区岩溶水文地质调查,并辅以示踪试验,

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第一作者简介:曾洁(1986—),女,副研究员,主要从事岩溶地区水工环地质方面的研究。E-mail: [cengjie@mail.cgs.gov.cn](mailto:cengjie@mail.cgs.gov.cn)。

通信作者:潘晓东(1984—),男,教授级高级工程师,主要从事岩溶地区水工环地质方面的研究。E-mail: [pxiaodong@mail.cgs.gov.cn](mailto:pxiaodong@mail.cgs.gov.cn)。

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查明湖区及周边水文地质结构,综合地球物理探测成果,准确定位了岩溶地下通道,对渗漏区水文地质等条件进行了分析,以期为类似岩溶湖库渗漏灾害调查治理提供参考。

## 1 研究区地质背景

### 1.1 研究区概况

碧云湖位于贵州省织金县北部的八步街道境内,地理坐标:东经 $105^{\circ}45' \sim 105^{\circ}50'$ ,北纬 $26^{\circ}48' \sim 26^{\circ}51'$ 。碧云湖为天然高原岩溶湖泊,起始形成于1943年夏天,仅存7个月后就消失了。1975年5月,因排水不畅,又形成了一个南北长2500 m,东西宽1500 m,总蓄水量7000万m<sup>3</sup>的天然湖泊;从1988年以来,湖库库容虽有变化,但都有一定的规模。近年来,碧云湖水域面积逐渐萎缩,至2019年已不足20万m<sup>2</sup>,原本连在一起的碧云湖萎缩成两个分离的小湖泊。经实地踏勘,推测碧云湖水域萎缩的原因为其下伏的岩溶地下河管道贯通,岩溶水渗漏量增加所致。

针对碧云湖水域萎缩问题,通过水文地质条件分析,示踪试验、地球物理探测及钻探验证等手段,对岩溶地区湖泊渗漏问题进行了探讨,并提出了对应的防治建议。

### 1.2 水文地质背景

碧云湖位于扬子准地台、黔北台隆、遵义断拱、毕节北东向构造变形区上,湖区内未见有断裂带展布,周边亦未见有影响湖址稳定的全新活动断裂发育。湖区及周边出露碳酸盐岩地层有二叠系茅口组( $P_2m$ ),三叠系夜郎组二段( $T_{j^2}$ ),嘉陵江组一段( $T_{j^1}$ ),嘉陵江组三段( $T_{j^3}$ ),嘉陵江组四段( $T_{j^4}$ )和关岭组( $T_{2g}$ ),地层岩性以灰岩、白云岩为主,地层透水性相对较强。湖区出露非碳酸盐岩地层有二叠系龙潭组( $P_3l$ ),三叠系下统夜郎组第一段( $T_{j^1}$ ),第三段( $T_{j^3}$ )地层,嘉陵江组二段( $T_{j^2}$ ),地层岩性主要为粉砂岩、炭质页岩、泥岩等为主,透水性相对较差(图1)。

据现场勘查资料,湖区主要分布嘉陵江组三段

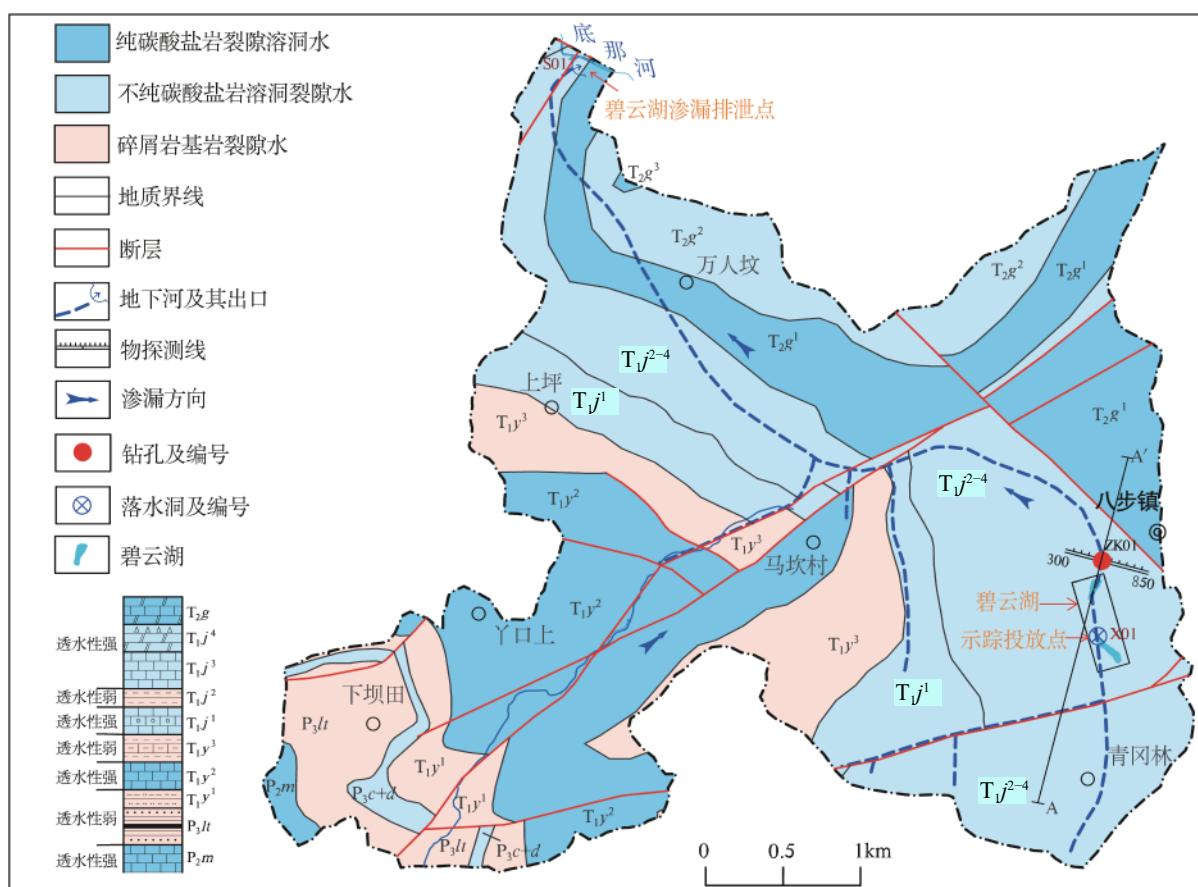


图1 研究区水文地质图

Fig. 1 Hydrogeological map of the study area

( $T_{J^3}$ )灰岩地层,下游分布嘉陵江组二段( $T_{J^2}$ )泥岩、白云质泥岩地层,厚度约20 m,相对于上覆碳酸盐岩地层,嘉陵江组二段( $T_{J^2}$ )地层为相对隔水层。宏观上来说,受地层岩性、地质构造、地貌条件及水文网络控制,碧云湖区内地下水主要赋存于地下河管道中,构成地下河系统,即八步地下河系统。此系统上游和下游地下水以集中管道形式赋存循环,中游受向斜蓄水构造控制形成一个岩溶盆地,岩溶盆地地下水呈层状展布,自上而下可划分为3个含水层段,八步岩溶地下河管道和向斜蓄水构造盆地含水层均位于三叠系嘉陵江组第三段( $T_{J^3}$ )灰岩含水层组中(图2)。湖区地下水主要为岩溶水,含水层为泥灰岩和灰岩,少部分分布于基岩裂隙和岩溶管道中。

此外,覆盖层中还分布有上层滞水,但水量小。

碧云湖补给水量大,但历史上并未形成严重内涝。洼地在暴雨后可形成一定规模的水域面积,在5~7 d内洼地积水会恢复到降雨前。主要原因在于洼地底部嘉陵江组三段( $T_{J^3}$ )灰岩地层发育有地下河管道,湖水会随着岩溶裂隙、孔隙、落水洞等补给地下水,随岩溶管道排走。但在近几年的干旱季节碧云湖能够保持在一定的水位,可以确定渗漏点及渗漏段不在碧云湖水域最低的位置。

水文地质分析结果表明,碧云湖为降雨和地表水补给,下伏嘉陵江组二段( $T_{J^2}$ )泥岩、白云质泥岩具有较好的隔水作用,可通过帷幕灌浆治理湖区的渗漏问题。

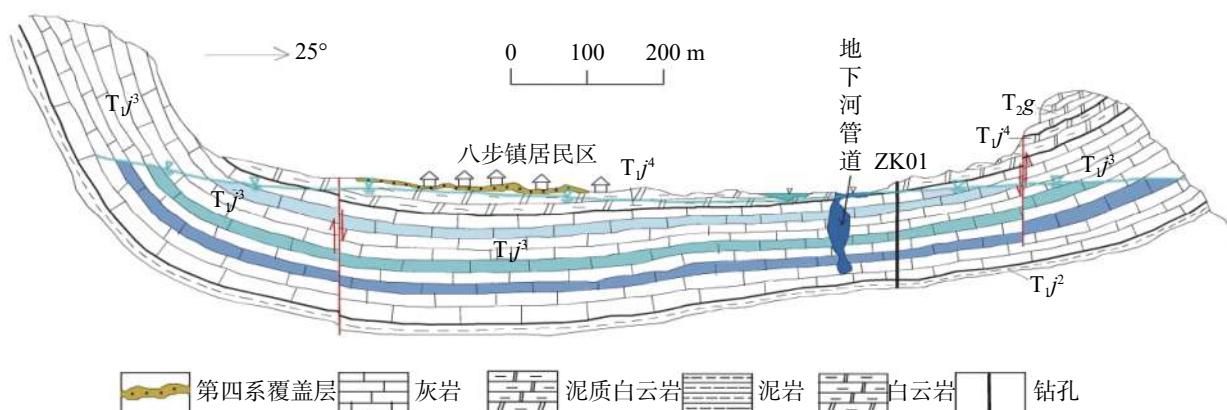


图2 碧云湖周边水文地质剖面图(A-A')

Fig. 2 Hydrogeological profile of Biyun lake (A-A')

## 2 研究方法

### 2.1 示踪试验

示踪技术是确定岩溶地下水系统连通性最有效的方法,可以为分析地下水径流、排泄及岩溶管道的分布、形态、水力特征等提供依据。本次接收点为村民饮用水源地,示踪试验选择无毒、无色的钼酸铵为示踪剂,以验证碧云湖湖区内落水洞与八步地下河出口之间的连通性。在湖区渗漏点X01进行投放,并在八步地下河出口S01进行取样检测(图1),投放点与接收点间距离约4.91 km。

本次示踪剂投放量为50 kg,将钼酸铵全部倒入桶中,加水并充分搅拌均匀溶解后,缓慢注入落水洞X01中,最后再用清水把塑料桶冲洗干净,冲洗塑料桶的水也注入落水洞X01中,以尽量保证示踪剂全

部进入地下水。示踪剂投放后开始在八步地下河出口S01取样,取样间隔为4 h,总计57组样品。样品采用30 mL聚乙烯瓶采集,为尽量减小误差,采样前先用待采水样洗涤采样瓶3~5次,采样后密封保存,并尽快送往实验室。钼离子浓度在自然资源部、广西岩溶动力学重点实验室使用美国赛默飞生产的电感耦合等离子体质谱仪(ICP-MS)进行分析测试,精度为0.01  $\mu\text{g}\cdot\text{L}^{-1}$ 。

### 2.2 地球物理探测

根据水文地质调查分析结果,结合湖区渗漏特征及地质情况,查明湖库渗漏主要为岩溶管道型渗漏,这类地质体与完整围岩相比为相对低阻和相对高介电常数。从以往物探工作成果发现,大功率充電法在较大范围内识别水库渗漏通道或地下河管道的平面位置及走向具有较高的准确性,但缺点是无法确定渗漏带的发育深度;而地质雷达法具有分辨

率高的特点,可高精度识别地下渗漏带的发育深度<sup>[13-14]</sup>。结合上述两种方法的优点,本次采用大功率充电法以及地质雷达法综合探测地下渗漏通道发育特征。

大功率充电法采用加拿大 V8 多功能电法仪,电位差分辨率  $0.1 \mu\text{V}$ , 电流测量精度  $0.1 \text{ mA}$ ; 野外工作参数: 无穷远点垂直测线, 距离测线大于  $1000 \text{ m}$ ,  $MN=10 \text{ m}$ (梯度法), 点距  $10 \text{ m}$ , 异常附近加密到  $5 \text{ m}$ 。地质雷达法采用加拿大 Ultra PulseEKKO Pro 地质雷达仪, 采用  $100 \text{ MHz}$  天线, 测点距  $1 \text{ m}$ , 计时误差小于  $1 \text{ ns}$ , 最小采样间隔达到  $0.5 \text{ ns}$ , A/D 转换不低于  $16\text{bit}$ 。

根据水文地质条件,于水库北侧下游处布设 1 条物探测线(图 1),该测线相交于经水文地质调查确定的地下水流向,目的是查明水库往下游方向渗漏通道的平面位置及发育深度。测线点号为  $290\sim850 \text{ m}$ , 其中充电法测点编号为  $290\sim850 \text{ m}$ , 测线长度为  $560 \text{ m}$ , 点距  $10 \text{ m}$ , 地质雷达法测点编号为  $450\sim650 \text{ m}$ , 测线长  $200 \text{ m}$ , 点距  $1 \text{ m}$ 。

### 3 结果与讨论

#### 3.1 岩溶渗漏方向

示踪剂投放前在八步地下河取一个样品作为背景值,检测  $\text{Mo}$  离子浓度为  $84.9 \mu\text{g}\cdot\text{L}^{-1}$ 。示踪试验检测结果如图 3 所示,在示踪剂投放  $32 \text{ h}$  后  $\text{Mo}$  浓度出现了明显的上升,达到  $93.2 \mu\text{g}\cdot\text{L}^{-1}$ ,说明地下河出口对渗漏点响应迅速。 $68 \text{ h}$  后  $\text{Mo}$  离子浓度达到了峰值  $146.0 \mu\text{g}\cdot\text{L}^{-1}$ ,随后降低。在第  $80 \text{ h}$   $\text{Mo}$  离子浓度从  $138 \mu\text{g}\cdot\text{L}^{-1}$  急聚减低至  $43.2 \mu\text{g}\cdot\text{L}^{-1}$ ,这是暴雨后地表水快速补给岩溶含水层造成的稀释作用导致的。经示踪试验验证,碧云湖湖水经 X01 落水洞渗漏后,

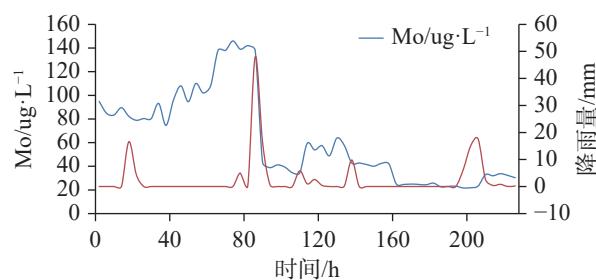


图 3 示踪接收点  $\text{Mo}$  浓度变化曲线图

Fig. 3 Varavition curve of Mo concentration at the tracer receiving point

沿着岩溶渗漏带向北东方向径流,最终排泄于八步地下河出口 S01。

#### 3.2 岩溶渗漏带位置与深度

##### 3.2.1 岩溶渗漏带位置

大功率充电法电位梯度值相对极小值的区域共 6 处(图 4),分别位于测点  $410 \text{ m}$ 、 $460 \text{ m}$ 、 $540 \text{ m}$ 、 $630 \text{ m}$ 、 $680 \text{ m}$  和  $720 \text{ m}$  处。电位曲线在测点  $540 \text{ m}$ 、 $680 \text{ m}$  和  $720 \text{ m}$  附近出现电位相对极大值,且同位置处电位梯度曲线接近零值,推断为强岩溶发育带或地下河管道,其中测点  $540 \text{ m}$  附近为电位最大值点。

大功率充电法结果中,地下河管道异常特征为:电位曲线图上出现电位相对极大值、电位梯度图上出现电位梯度相对极小值。综合电位梯度曲线和电法曲线结果,推断测点  $540 \text{ m}$  附近为测区近南北向主管道,为碧云湖水库的主要渗漏通道,可能有大型溶洞发育,其余异常段推测为强岩溶发育带,可为地下水的存储和运移提供条件。

##### 3.2.2 岩溶渗漏带深度

地质雷达探测地下河管道异常特征为地质雷达

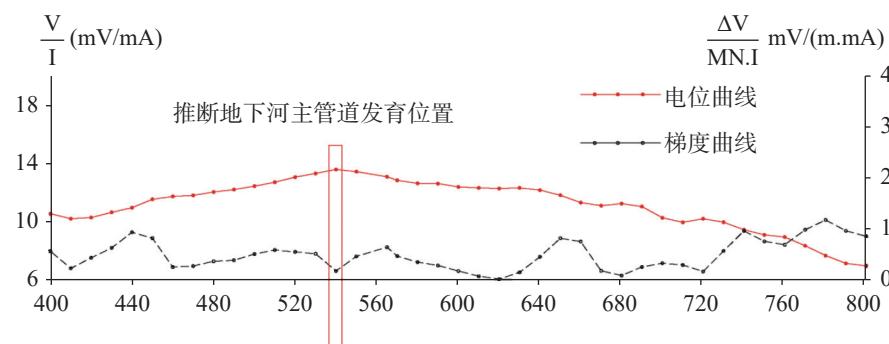


图 4 大功率充电法结果图

Fig. 4 Results by high-power charging method

时间(深度)剖面出现电磁波强反射带。根据钻孔资料校正,地质雷达参数  $V$  取值  $0.060 \text{ m}\cdot\text{ns}^{-1}$ ,由测线地质雷达时间(深度)剖面(图5)可知:在 20~28 m 深度区间呈现相对电磁波强反射面,推断为泥岩与灰岩的接触面;与测线上其它测点相比,测点 540~550 m 段,在 30~40 m 深度区间呈现相对电磁波强反射带,推断为强岩溶发育带,与大功率充电法探测强岩溶渗漏带位置基本一致。在测点 460 m 段附近,在 20~90 m 深度区间也出现电磁波强反射带,结合该测点段靠近织金—八步公路主干道,车辆较多,推断为电磁波强干扰带。

### 3.3 岩溶渗漏带验证

结合水文地质调查和地球物理探测结果分析,在测点 540 m 处实施 ZK01 号钻孔验证(图6)。ZK01 号钻孔钻探进尺 66 m, 0~8.5 m 为第四系覆盖层, 8.5~22.5 m 为三叠系下统嘉陵江组第四段白云岩层, 22.5~44.3 m 为三叠系下统嘉陵江组第三段灰岩层, 44.3~63.2 m 为三叠系下统嘉陵江组第二段泥岩, 63.2~66.0 m 为三叠系下统嘉陵江组第一段灰岩。在 8.5~11.0 m、38.0~43.2 m 处发生掉钻现象, 遇水呈流塑状, 38.0~43.2 m 可判断为溶洞, 且 38.0~43.2 m 处洞穴充填泥, 验证了物探推测的强岩溶渗漏带。

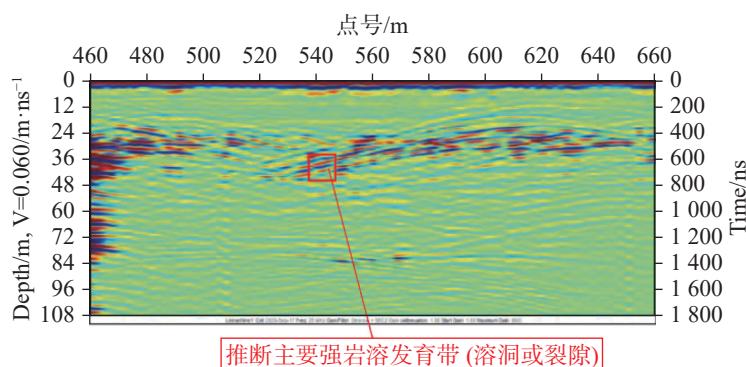


图 5 地质雷达法结果图

Fig. 5 Results by geological radar method

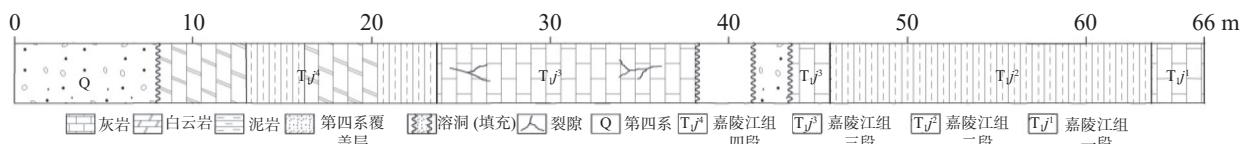


图 6 ZK01 钻孔岩性柱状图

Fig. 6 Lithology histogram of borehole ZK01

### 3.4 分析讨论

综合水文地质调查、示踪试验、物探和钻探结果,认为湖区渗漏带主要由岩溶管道构成,沿嘉陵江组三段灰岩和二段泥岩岩性界面发育,主要依据如下:

(1) 岩性接触带是岩溶强烈发育的地区之一,落水洞、裂隙、岩溶管道等常沿着岩性接触带分布。水文地质调查发现位于洼地西侧有一落水洞 X01,碧云湖水源源不断流入落水洞 X01,可能是碧玉湖水减少的原因之一。落水洞平面上呈长方形,长约 3.5 m, 宽约 1.5 m, 深约 2.0 m。落水洞垂直发育,底部形态与地表相似,渗漏点湖水汇入一侧可见基岩出露,而另一侧未见基岩,但见泥土,落水洞底部泥土堆积,未见大的裂隙和管道,表明渗漏点可能位

于岩性接触带上:嘉陵江组三段( $T_j^3$ )灰岩地层和嘉陵江组二段( $T_j^2$ )泥岩地层。

(2) 大功率充电法地球物理探测结果显示,测点 540 m、680 m 和 720 m 附近出现局部电位极大值,且电位梯度曲线上接近零值,680 m、720 m 也可能是渗漏部位,有充电异常、但没有雷达资料,钻探也没有查证,因此存疑,需开展后续工作。地质雷达地球物理探测结果显示,测点 540~550 m 在 30~40 m 深度区间电磁波呈现相对电磁波强反射带,推断测点 540 m 附近为测区近南北向主管道经过处,为碧云湖水库的主要渗漏通道,可能有大型溶洞存在。

(3) 水文地质钻探结果显示,测点 540 m 处 ZK01 号钻孔在 38~43.2 m 位置发生掉钻现象,为充泥溶洞,

位于嘉陵江组三段灰岩和嘉陵江组二段泥岩岩性界线附近。

## 4 结论与建议

(1) 据水文地质调查成果,碧云湖属于向斜构造盆地,碧云湖渗漏点和地下河管道主要发育在靠近八步复式向斜轴的位置。下部嘉陵江组二段泥岩构成岩溶渗漏带的隔水底板,湖水沿渗漏带渗漏后,向北西方向径流,最终于八步地下河出口处排泄。

(2) 综合水文地质调查、物探和钻探结果,认为岩溶渗漏带主要为岩溶管道型渗漏,主管道位于物探测线测点 540 m 处,管道深约 40 m;根据钻孔资料校正,地质雷达参数 V 取  $0.060 \text{ m} \cdot \text{ns}^{-1}$  适用于该地区。

(3) 对碧云湖水库的防渗治理建议采用帷幕灌浆,根据水文地质调查、地形地貌条件和地球物理探测结果,建议以测点 540 m 为中心向两边帷幕灌浆。

(4) 钻探结果证实了在岩溶区采用水文地质调查、示踪试验与物探相结合的技术方法的可靠性,可以很好地确定渗漏通道的走向以及渗漏带的空间位置,从而指导岩溶区湖库的渗漏治理。

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## Hydrogeological conditions of karst leakage and locating of leakage channels in Biyun lake, Guizhou

ZENG Jie<sup>1,2</sup>, PAN Xiaodong<sup>1,2</sup>, REN Kun<sup>1,2</sup>, LIU Wei<sup>1,2</sup>, PENG Cong<sup>1,2</sup>, ZHENG Zhijie<sup>1,2</sup>

(1. Institute of Karst Geology, CAGS/Key Laboratory of Karst Dynamics, MNR & GZAR/International Research Center on Karst under the Auspices of UNESCO, Guilin, Guangxi 541004, China; 2. Pingguo Guangxi, Karst Ecosystem, National Observation and Research Station, Pingguo, Guangxi 531406, China)

**Abstract** China is a country with the widest distribution area and the most complete karst development type in the world, and its karst area accounts for about one third of China's territorial area. Finding leakage channels of lakes and reservoirs in karst areas has been an urgent problem for the lake and reservoir construction and the control of leakage disasters. In this study, Biyun lake in Zhijin county of Guizhou Province was taken as an example to analyze the hydrogeological conditions around the lake in response to the problem of watershed shrinkage. Furthermore, with the combination of tracer tests and geophysical detection methods, the hydrogeological conditions of leakage in Biyun lake and the location of the main leakage channel were determined and verified by drilling test results.

The results of the tracer test show that Moion concentration increased significantly 32 hours after the tracer was put into Biyun lake, and the value peaked 68 hours after the setting of tracer, and then decreased. The tracer results confirm that Biyun lake flows along the karst leakage zone in the north-east direction after seeping through the sinkhole, and is finally discharged at the outlet of the Babu underground river. In the results of high-power charging method, the potential curve and gradient curve show a relatively maximum potential value and a relatively minimum value of potential gradient at the measuring point of 540 m, where the main leakage channel of Biyun lake was inferred to be located and a large-scale karst cave is likely developed in the vicinity. The rest of the anomalous sections of the curves is inferred to be a strong karst development zone, which can provide conditions for the storage and transport of groundwater. A relatively strong reflective surface of electromagnetic wave shown by geological radar exploration at the depth of about 20–28 m is inferred to be a contact surface between mudstone and greystone. A strong reflection surface of relative electromagnetic wave at the depth of 36–48 m is inferred to be a strong karst development zone. These results are basically in line with the location of the strong karst leakage zone detected by high-power charging method. The borehole drilled at the measuring point of 540 m has verified the strong karst leakage zone inferred from

the physical exploration. Based on the correction of the data from the borehole, the parameter of geological radar,  $V$ , is taken to be  $0.060 \text{ m}\cdot\text{ns}^{-1}$  which is applicable to the study area.

The study results indicate that Biyun lake belongs to the oblique tectonic basin, where leakage points and underground river pipes are mainly developed close to the axis of the Babu compound oblique, along the lithological interface of three sections of greystone and the second section of the mudstone of Jialingjiang Group. The mudstone constitutes a water separating plate of the karst leakage zone. After lake water seeps along the zone, it flows northwest, and is discharged in the end at the outlet of the Babu underground river. In addition, the leakage of Biyun lake is mainly of the karst pipeline type. The karst leakage zone is located at the measuring point of 540 m, with a 40-meter-deep of karst pipeline; therefore, the leakage section is relatively narrow, and it is feasible to plug the leakage by grouting. The drilling test results confirm the reliability of the technical method used in karst areas, which is combined by hydrogeological investigation, tracer test and physical exploration, because this method can well determine the direction of the leakage channel and the location of the leakage zone, and thus providing guidance on the leakage control of lakes or reservoirs in karst areas.

**Key words** Biyun lake, karst leakage, hydrogeology, leakage channel, geophysical exploration technology

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constructed an index system for evaluating the benefit of rocky desertification control from ecological, economic and social perspectives, and analyzed the effectiveness of the control in each county/district of southwest Guangxi in 2010, 2015 and 2020 by this index system.

The results show that, (1) Benefit indexes of rocky desertification control in southwest Guangxi during 2010–2020 are ranked as follows: ecological benefit > social benefit > economic benefit. Among these indexes, the ecological benefit increased the most; social benefit grew slowly; economic benefit increased at a lower rate. The regional ecological environment and socio-economic conditions have been improved to a certain extent, showing overall benign development. (2) Counties/districts of southwest Guangxi have experienced differences in the effectiveness of rocky desertification control over the past 10 years. The northwestern and central regions have been maintaining high ecological benefits, while the areas with high economic and social benefits are mainly located in the main urban area of Nanning City and the areas under the jurisdiction of Binyang county and Hengzhou City in the southeast under the context of the "strong capital strategy", indicating that there is still much room for improvement in the effectiveness of rocky desertification control. As a result, the southeastern region should strengthen ecological and environmental protection, while the northwestern region should improve its socio-economic development. (3) The average values of the coupling and coordination degree of "ecological–economic–social" benefits are between 0.47 and 0.51, which is on the verge of disorder and barely coordinated state. The coupling and coordination relationship among the benefits of rocky desertification control in each county/district in 2020 showed lagged effects in different degrees, and the lagged economic and social benefits are the key factors limiting the effects of rocky desertification control. This indicates that the level of socio-economic development lags behind ecological governance in the control of rocky desertification. In the subsequent control of rocky desertification, it is necessary for us to focus on ecological and environmental protection and to accelerate high-quality socio-economic development as well.

**Key words** at a scale of county/district, benefits of rocky desertification control, coupling and coordination, limiting factor, southwest Guangxi

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