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关于溶蚀试片法计算溶蚀速率公式的一点认识

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摘要: 测算碳酸盐岩的溶蚀速率对研究岩溶发育规律和提高岩溶碳汇估算精度至关重要。溶蚀试片法是得到广泛认可和使用的计算研究区碳酸盐岩溶蚀速率的方法, 但目前利用其计算溶蚀速率的公式却并不统一, 还存在日溶蚀量和年溶蚀量计算错误、时间尺度外推过大、单位不统一等问题。文章对溶蚀试片法的起源与方法原理进行了简要介绍, 并就当前利用溶蚀试片法测算溶蚀速率的部分已发表公式进行了分析, 发现相关公式存在忽略试验时间长度、日溶蚀量与年溶蚀量转换系数(方法)错误、短期试验结果进行了长时间尺度外推等问题, 导致溶蚀速率测算误差较大, 区域间无法进行对比。基于溶蚀试片法的基本原理, 本文提出了用溶蚀试片法测算溶蚀速率的建议公式, 以期进一步提高就溶蚀试片法测算溶蚀速率的规范性及结果的对比性, 从而提高岩溶碳汇估算精度提供数据支持。

关键词: 碳酸盐岩; 溶蚀试片法; 溶蚀速率; 岩溶; 问题讨论

创新点: 系统梳理现有利用溶蚀试片法测算溶蚀速率的公式, 分析部分公式的应用误区, 提出基于溶蚀试片法计算溶蚀速率的建议公式, 以提高测算规范性及结果对比性。

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0 引言

大量研究表明, 碳酸盐岩溶蚀在全球千年—万年时间尺度碳循环中发挥了重要作用, 是全球重要的碳汇, 对于平衡全球碳收支和明确“剩余陆地汇”的归趋具有重要意义^[1-4], 这主要起因于如下两项化学反应式: $\text{CaCO}_3 + \text{CO}_2 + \text{H}_2\text{O} \rightleftharpoons \text{Ca}^{2+} + 2\text{HCO}_3^-$ (石灰岩溶蚀) 和 $\text{CaMg}(\text{CO}_3)_2 + 2\text{CO}_2 + 2\text{H}_2\text{O} \rightleftharpoons \text{Ca}^{2+} + \text{Mg}^{2+} + 4\text{HCO}_3^-$ (白云岩溶蚀)。因此, 碳酸盐岩溶蚀速率的测算是岩溶碳循环调查和碳汇效应评价的重要方法之一, 也是研究岩溶发育和岩溶水文地质问

题的关键性指标^[5]。碳酸盐岩溶蚀试片法是目前定量评价岩溶作用及其碳汇强度的主要方法之一^[6-7], 在近来发布的相关行业标准中也得到了推荐^[8-9]。此外, 国内外众多学者研究了大气或土壤中 CO_2 ^[10-11]、土地利用类型^[12-14]、岩性及岩石成分^[15-17]、地形地貌^[18-20]、气候条件等影响因子对不同碳酸盐岩的溶蚀速率及其时空变化的影响^[19, 21-22]。

岩溶碳汇量可通过参与反应的原材料与反应产物之间的质量关系进行计算^[23]。溶蚀试片法是通过参与反应的原材料即量化碳酸盐岩溶蚀量来估算岩溶碳汇, 其独特优势是可以对气候、岩性、土壤理化性质等影响碳酸盐岩溶蚀速率的因素进行对比分析,

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且可通过在土下不同剖面埋设试片来进一步探究土下岩溶作用的时空变化特征。

然而,当前在利用溶蚀试片法测算溶蚀速率的相关工作中,溶蚀速率的计算公式还存在不同版本,其应用还存在一定问题,导致同一监测数据应用不同公式计算出的溶蚀速率值存在较大差异,对准确揭示区域溶蚀过程及评估岩溶碳汇效应带来不利影响。本文根据溶蚀速率计算的基本原理和国内外使用的相关公式所存在的问题,对相关公式进行了分析和讨论,并进一步就计算公式的使用提出了建议,对今后更好地利用该方法开展碳酸盐岩溶蚀速率对比研究及区域岩溶碳汇估算有一定的参考价值。

1 溶蚀试片法简要介绍

碳酸盐岩溶蚀速率测定方法主要有溶蚀试片法、微侵蚀计法^[24-26]、水化学法^[27-28]、宇宙成因核素法^[29-30]、灰岩基座法^[31-32]等。微侵蚀计(Micro-Erosion Meter, MEM)是一种直接测定岩石表面相对高度的高精度装置,由固定的不锈钢基座、活动的不锈钢基架和数字式微尺组成,由 Hanna 于 20 世纪 60 年代首次应用^[33],随着技术的不断进步,微侵蚀计法在石刻文物、建筑石材等领域也得到了广泛应用^[34-35]。水化学法是通过监测岩溶水径流量和溶质含量来估算溶蚀速率,最早由 Corbel 于 1959 年提出^[27],水化学法所需数据量较大,但可计算出具体流域的溶蚀速率^[28]。宇宙成因核素法是根据地表暴露岩石或埋藏物中宇宙成因核素浓度估算溶蚀或抬升速率^[29-30],这一方法可估算研究区十几至几十万年的平均溶蚀速率,但测试成本较高,存在仪器误差、影响因素判断较难等问

题。灰岩基座法是通过覆盖于地下灰岩基座与裸露岩面的高度差及距末次冰期时长来推算溶蚀速率^[31],该方法多用于高寒地区溶蚀速率的计算^[32],存在一定的地域限制。

溶蚀试片是由岩石样本切割制成的规格统一的石片,为便于对比研究,常使用统一成分和结构成因类型的石灰岩制作。国内学者一般使用广西桂林七星岩附近上泥盆统融县组(D₃r)石灰岩制成标准试片,试片通常为圆形,直径 4 cm,厚 3~5 mm,表面积 28~31 cm²。国外常使用斯洛文尼亚白垩纪灰岩制成标准试片,通常为圆形,直径 4.1 cm,厚 2~3 mm,重 9~12 g。部分学者为探究岩性对溶蚀速率的影响或获取更准确的当地溶蚀速率也采用试验点当地岩石制作试片^[17, 36-38],也有部分研究采用更容易切割的正方体或长方体^[39]。试片切割后,通常会进行清洗、干燥、称重等处理,为减小切割痕迹对试片溶蚀的影响,部分研究人员还会对试片进行打磨、酸蚀^[39-41]。溶蚀试片法根据试片在埋放前和埋放一定时间后的重量损失来计算溶蚀速率。其方法一般为将试片放置于空中 1~1.5 m、土壤表面和土壤不同深度(通常为土下 20 cm、50 cm)四个层位,试片需平放,以便水、气能均匀地作用于试片上,可用木桩或金属圆盘标记埋放点,以便在回收试片时定位。根据试验需要于一定时间后(通常为一个水文年)取出、洗净、烘干后用万分之一精度天平称重(试片取出后处理步骤应与放置前完全一样),从试片重量的减少来计算溶蚀速率及其消耗的大气或土壤 CO₂ 量。根据各种常用碳酸盐岩溶蚀速率测算方法的原理,表 1 总结了各方法的优点与局限。

表 1 常用溶蚀速率测算方法的特点

Table 1 Characteristics of commonly used methods for measuring dissolution rates

方法	优点	缺点
溶蚀试片法	不需要长期监测;埋放方法成熟;接近自然状态;影响因子可控;试验成本低	难以获得历史数据;需多个代表性测试点;时间分辨率不高;动态变化不易掌握
微侵蚀计法	应用范围广;相对短时间内获得数据;限制条件少	难以获得历史数据;不适合大尺度区域研究
水化学法	可得出流域溶蚀速率;适合较大尺度区域研究	所需数据量较大
宇宙成因核素法	应用范围广;可估算长时间尺度数据	测试成本较高;精确度有待提升
灰岩基座法	原理简单;可估算长时间尺度数据	存在地域限制;误差较大

1953 年,法国岩溶学家 P. Chevalier 在第一届国际洞穴学大会上首次提出了用溶蚀试片测量岩溶区

碳酸盐岩溶蚀速率的方法,其研究目标是了解溶蚀和侵蚀效应在地下管道形成和扩大中的作用^[42]。随

后, 1957 年, I. Gams 于斯洛文尼亚 Podpeška 洞穴使用该方法研究洞穴溶蚀过程^[43]。1964 年, R. Rebek 也在斯洛文尼亚 Ljubljana 洞穴利用薄矩形样式的溶蚀试片(3 mm 厚)研究了洞内静水和洞外土壤层中的溶蚀速率, 并首次给出了计算溶蚀速率的公式^[44]。随后, I. Gams 和 S. T. Trudgill 等对该方法进行了广泛的运用, 研究了欧洲多个岩溶区的溶蚀速率及其控制因素问题^[34, 45-48]。1985 年, I. Gams 系统分析了全球 60 多个点的溶蚀试片研究资料, 初步建立降水与不同深度土壤层中溶蚀速率的线性关系, 发现了溶蚀速率在全球尺度上的规律性, 成为溶蚀试片法的一个划时代的研究成果^[40, 49-50]。相较而言, 我国开展相关研究的时间较晚, 20 世纪 80 年代袁道先等将其引入我国^[51-52], 对我国不同气候带不同海拔的溶蚀速率进行了研究, 得出我国西南岩溶区的石灰岩溶蚀速率较华北半干旱区和西北干旱区更快, 气候特别是年降雨量与石灰岩的溶蚀速度有很好的对应关系, 生物作用、土层中的 CO₂ 浓度及土壤的渗透性对溶蚀速率有重要影响, 该方法后来在我国主持的国际地质对比计划 IGCP299 项目中得到广泛应用。近几十年来, 相关学者利用溶蚀试片法开展了大量研究, 研究发现降雨量、气温等气候要素是碳酸盐岩溶蚀速率的主要决定因素之一^[19, 53], 雨季溶蚀速率显著大于旱季^[54-56]; 不同土地利用类型下土壤理化性质如 pH、含水率、孔隙度等会影响溶蚀速率的大小^[57-59]; 植被的正向演替可促进岩溶作用的进行^[18, 60-61];

岩石组成、结构等也对溶蚀速率产生影响^[16, 36], 石灰岩溶蚀速率普遍较白云岩更高^[4, 37-38]; 我国存在溶蚀速率由地下>地表至地表>地下、碳酸盐岩由溶解向沉积转变的过渡带^[22, 62-63]。然而, 目前利用溶蚀试片法计算溶蚀速率的公式还存在不同版本, 部分公式存在一定问题, 不利于减小试验结果与实际溶蚀速率的误差, 为提高对区域岩溶碳汇的估算精度, 有必要对利用溶蚀试片法计算溶蚀速率的公式进行讨论。

2 利用溶蚀试片法计算溶蚀速率的讨论

2.1 使用的相关公式及其问题

在溶蚀试片法研究溶蚀速率的单位使用上, 部分学者以试验期间试片的重量损失百分比(试片溶蚀量/试片原重×100%)来衡量溶蚀作用的强度^[45, 48, 64-65], 也有以单位面积和时间的重量损失(即 mg·cm⁻²·a⁻¹)来表示溶蚀强度。若已知岩石密度, 则可转换为 μm·a⁻¹ 或 mm·ka⁻¹, 即单位时间内岩石被剥蚀的厚度^[5, 66]。

Rebek 于 1964 年给出了第一个计算溶蚀速率的公式^[31], Trudgill 在 1975 年提出以单位 cm·yr⁻¹ 来表示的计算公式^[34]。随后, 在相关的研究中一些学者提出了许多相似的公式用于计算溶蚀速率(表 2)。然而, 由于对溶蚀作用过程、溶蚀速率计量的理解存在一些偏差, 部分公式的使用还存在一定问题, 导致出现了多种计算公式, 且同一数据计算结果误差较

表 2 利用溶蚀试片法计算溶蚀速率的公式
Table 2 Formulas used to calculate dissolution rates using rock tablets

公式	注释	问题	出处
$d = \frac{(G_0 - G_1)}{A \times 2.71}$	d : 溶蚀速率/mm G_0 、 G_1 : 试验前后试片重量/mg A : 试样的表面积/mm ² 2.71: 石灰岩比重/mg·mm ⁻³	缺少试验时间, 不利于对比研究	[44]
$W = \frac{wm}{mt} \times 365; v = \frac{w}{d}; h = \frac{v}{a}$	W : 年损失量/g w : 试片重量损失/g mt : 代表测量的时间/d v : 损失体积/cm ³ d : 密度/g·cm ⁻³ h : 岩石剥蚀厚度/cm·yr ⁻¹ a : 试片表面积/cm ²	无误, 单位面积溶蚀量可据此转换为剥蚀厚度	[47]
$Sr = \frac{(W_2 - W_1)}{2.7A \cdot t}$	Sr : 溶蚀速率/mm·ka ⁻¹ W_1 、 W_2 : 试验前后试片重量 t : 试验时间 A : 试样的表面积/mm ² 2.7: 石灰岩比重/g·cm ⁻³	式中 Sr 应为 mm·a ⁻¹	[11]

续表 2

公式	注释	问题	出处
$ER = (W_1 - W_2) \times 1000 \times T / 365 / S$	ER: 溶蚀速率/ $\text{mg} \cdot \text{cm}^{-2} \cdot \text{a}^{-1}$ W_1 、 W_2 : 试验前后试片重量/g T : 埋放时间/d S : 试片表面积/ cm^2	误将试片绝对溶蚀量与埋放时间相乘; 日溶蚀速率转换为年溶蚀速率计算错误	[14, 67–68]
$ER = (W_1 - W_2) \times 10^7 / T / S$	ER: 溶蚀速率/ $\text{mg} \cdot \text{m}^{-2} \cdot \text{d}^{-1}$ W_1 、 W_2 : 试验前后试片重量/g $W_1 - W_2$: 绝对溶蚀量/mg T : 埋放天数/d S : 试片表面积/ cm^2	无误	[36, 54, 56–57, 69–71]
$ER = (W_1 - W_2) \times 1000 \times 365 / (T \times S)$	ER: 溶蚀速率/ $\text{mg} \cdot \text{cm}^{-2} \cdot \text{a}^{-1}$ W_1 、 W_2 : 试验前后试片重量/g $W_1 - W_2$: 试片绝对溶蚀量/g T : 埋放时间 S : 试片表面积/ cm^2	无误	[58, 72]
$ER = (W_1 - W_2) \times 1000 \times T \times 365^{-1} \times S^{-1}$	ER: 溶蚀速率/ $\text{mg} \cdot \text{cm}^{-2} \cdot \text{a}^{-1}$ W_1 、 W_2 : 试验前后试片重量/g T : 埋放时间/d S : 试片表面积/ cm^2	误将试片绝对溶蚀量与埋放时间相乘; 日溶蚀速率转换为年溶蚀速率计算错误	[12, 18, 20, 60, 73–74]
$E = (W_1 - W_2) / S$	E: 溶蚀速率/ $\text{g} \cdot \text{cm}^{-2} \cdot \text{a}^{-1}$ W_1 、 W_2 : 试验前后试片重量/g $W_1 - W_2$: 试片绝对溶蚀量/g S : 溶蚀试片的表面积/ cm^2	省略了溶蚀时间, 易造成误导	[75]
$D = (W_1 - W_2) \times 1000 \times S^{-1}$	D: 喀斯特溶蚀量/ $\text{mg} \cdot \text{cm}^{-2}$ W_1 、 W_2 : 试验前后试片重量/g S : 试片表面积/ cm^2	省略了溶蚀时间, 易造成误导	[76]
$R = \frac{(W_1 - W_2) \times T}{365 \times S}$	R: 溶蚀速率/ $\text{mg} \cdot \text{cm}^{-2} \cdot \text{a}^{-1}$ W_1 、 W_2 : 试验前后试片重量/mg T : 埋放时间/d S : 试片表面积/ cm^2	误将试片绝对溶蚀量与埋放时间相乘; 日溶蚀速率转换为年溶蚀速率计算错误	[7, 62, 77]
$ER = \frac{W_1 - W_2}{T \cdot S}$	ER: 溶蚀速率/ $\text{mg} \cdot \text{cm}^{-2} \cdot \text{a}^{-1}$ W_1 、 W_2 : 试验前后试片重量/mg T : 溶蚀时间/a S : 试片表面积/ cm^2	无误	[21, 78–79]
$ER = (W_1 - W_2) \times 10000 / (T \times S)$	ER: 溶蚀速率/ $\text{mg} \cdot \text{m}^{-2} \cdot \text{d}^{-1}$ W_1 、 W_2 : 试验前后试片重量/mg T : 埋放天数/d S : 试片表面积/ cm^2	无误	[38]
$RS = 3.65 \times 10^{10} n (W_1 - W_2) / (A \times T)$	RS: 碳酸盐岩溶蚀或碳汇强度/ $\text{mol} \cdot \text{km}^{-2} \cdot \text{a}^{-1}$ n : 试片中碳酸盐的含量 W_1 、 W_2 : 试验前后试片重量/g T : 埋放时间/d A : 试片表面积/ cm^2	日溶蚀速率转换为年溶蚀速率计算错误	[28]
$E = 365 \times 10^4 (W_1 - W_2) / (A \times T)$	E: 溶蚀量/ $\text{g} \cdot \text{km}^{-2} \cdot \text{a}^{-1}$ W_1 、 W_2 : 试验前后试片重量/g T : 埋放时间/d A : 试片表面积/ cm^2	面积单位换算错误	[23]
$DR = (W_1 - W_2) \times 1000 \times T / S$	DR: 溶蚀速率/ $\text{mg} \cdot \text{a}^{-1} \cdot \text{cm}^{-2}$ W_1 、 W_2 : 试验前后试片重量/g T : 埋放时间/a S : 试片表面积/ cm^2	误将试片绝对溶蚀量与埋放时间相乘	[80]

大,具体见表2。

根据表2所示,部分学者所用公式误将试片绝对溶蚀量与试验时间相乘,或在将单位面积日溶蚀量转换为年溶蚀量时误把“ $\times 365$ ”计算为“ $\div 365$ ”,这将导致所计算出的溶蚀速率与实际相差较大,此外部分研究还存在单位换算错误的问题。已有相关研究表明,由于环境条件(如气候、土壤、植被覆盖等)的变化,溶蚀速率并不是恒定不变的,因此虽然溶蚀试片所得溶蚀速率便于对比空间差异,但要避免对试验结果的长时间尺度外推^[5,48,81]。

除溶蚀速率计算公式外,部分研究为探究不同岩性试片的溶蚀速率差异采用不同岩性岩石制作试片,但对其岩石密度未提及,也不利于开展对比研究^[38]。受地层、地形等作用,溶蚀试片埋放位置也会对溶蚀速率产生影响^[18],建议选择在不同地层岩性及地貌部位(如山顶、山腰、洼地、垭口等)埋放试片,同时考虑不同植被覆盖类型、土地利用方式,试片埋放点应与岩样、土样、水样等采样点配套^[8]。我国北方干旱半干旱岩溶区土壤无机碳含量较高,易产生沉积物阻碍溶蚀作用的进行,导致溶蚀试片法所得结果较小甚至出现重量增加的情况^[22,82],应结合试验点土壤无机碳数据与水化学法所得结果对比分析。在岩溶发育强烈、地势起伏较大的西南地区,需要考虑实际发生岩溶作用的面积与岩溶区面积的比值^[83]。

2.2 建议公式

根据溶蚀试片法的原理,计算单位时间内单位面积的重量损失,应首先确定试片的绝对溶蚀量。这可由试片埋设前的初始重量减去埋设后的重量得出。随后,将试片绝对溶蚀量除以试片的表面积与埋放时间,即可得到试片的溶蚀速率。鉴于溶蚀试片通常较轻,为减小试验数据的误差,建议均采用万分之一精度的电子天平进行称量,并以mg作为重量单位。为在减少误差的同时有利于开展区域对比分析,溶蚀试片通常埋放一或多个水文年,部分研究为观测其季节性变化从而更好地揭示数据结果与环境因子之间的关系,按旱雨季进行埋放。为使获取的数据更符合实际情况,也便于对比,按季节尺度埋放试片或埋放时间不足一个水文年建议采用以下公式计算溶蚀速率:

$$ER = \frac{W_0 - W_1}{S \cdot T}$$

式中: ER 为碳酸盐岩日溶蚀速率($\text{mg} \cdot \text{cm}^{-2} \cdot \text{d}^{-1}$); W_0 为试片埋放前重量(mg); W_1 为试片埋放后取回烘干重量(mg); $W_0 - W_1$ 为试片绝对溶蚀量(mg); S 为试片表面积(cm^2); T 为试片埋放时间(d)。

按水文年尺度埋放试片可将上述公式所得溶蚀速率计算结果乘以365,得到单位面积的重量年损失量。即:

$$ER = \left(\frac{W_0 - W_1}{S \cdot T} \right) \times 365$$

式中: ER 为碳酸盐岩年溶蚀速率($\text{mg} \cdot \text{cm}^{-2} \cdot \text{a}^{-1}$); W_0 为试片埋放前重量(mg); W_1 为试片埋放后取回烘干重量(mg); $W_0 - W_1$ 为试片绝对溶蚀量(mg); S 为试片表面积(cm^2); T 为试片埋放时间(d)。

这样的数据处理方式能使试验分析结果更符合实际情况,为开展全球分析和对比研究,揭示岩溶发育规律提供可靠的数据支撑。

3 研究展望

由于溶蚀试片法无法直接获取历史溶蚀数据,相关研究只能通过对比分析来间接了解过去的岩溶作用情况。因此,在应用这一方法时,确保试验过程的规范性、计算统一性和可对比性显得尤为重要。碳酸盐岩溶蚀速率的测算方法多样且各具优缺点,应从研究区域特征、研究目的出发,选择更合适的方法。如溶蚀试片法在气候湿润的南方岩溶区应用效果较好,而在我国北方干旱半干旱地区,为减少土壤无机碳对溶蚀试片法造成的干扰,可适当延长试验时间,使试验结果更接近自然状态,根据北方流域边界清晰、流量稳定的特点,还可结合水化学法进行对比研究。

受客观条件的限制,溶蚀试片法很难考虑到岩层内部和岩层间的溶蚀作用。应加强各试验方法的关联研究,建立碳酸盐岩溶蚀试验模型,从而使得不同研究的数据可以相互参照,更全面地了解溶蚀过程的特征和规律,也能形成统一的行业规范,有助于提高估算岩溶碳汇通量的精度,为相关领域的研究和应用提供更为坚实的依据,推动岩溶学和相关学科的发展。

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Discussion on the formulas used to calculate the dissolution rate by the rock tablet method

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Abstract The dissolution rate of carbonate rocks is an important parameter for investigating the karst carbon cycle and evaluating the carbon sink effect; as well as a key indicator for studying karst development and karst hydrogeological problems. The rock tablet method is currently one of the main methods for quantitatively evaluating karstification and its carbon sink intensity, and has been widely recognized and used. The rock tablet method can not only analyze the factors affecting the dissolution rate of carbonate rocks (e.g., climate, lithology, soil properties), but also further explore the temporal and spatial variation characteristics of subsurface karstification by burying carbonate rock tablets at different depths. However, in current studies related to the calculation of carbonate dissolution rates using the rock tablet method, there are different versions of the dissolution rate calculation formula, and there are still some problems in its application, resulting in significant differences in the dissolution rate values calculated using different formulas for the same monitoring data, which adversely affects the accurate determination of the regional dissolution process and assessment of the karst carbon sink effect.

This study firstly reviewed the origin, operational procedure, and basic principles of the carbonate rock tablet method and clarified the key links such as the preparation of the rock tablets (the standard rock tablets were uniformly made of limestone in the same stratum, usually the Rongxian formation of the upper Devonian collected from Guilin in Guangxi, or the upper Cretaceous collected from Slovenia, with a diameter of 4 cm and a thickness of 3mm to 5 mm), burial layers (in the air, on the surface, and at the soil depths of 20 cm and 50 cm), burial duration (typically corresponding to one hydrological year), and weighing (using a balance with one-millionth precision). The advantages and limitations of commonly used carbonate dissolution rate calculation methods are summarized.

Then, the paper analyzed some of the published formulas for calculating the carbonate dissolution rate using the rock tablet method and discovered that some of the formulas have the following issues, (1) The coefficient (method) from the daily dissolution rate to the annual dissolution rate is incorrect, further intensifying the uncertainty of the results; (2) The short-term test results were extrapolated over a long time scale, disregarding the influence of the dynamic changes in environmental conditions on the carbonate dissolution rate; (3) Unit conversion errors. These problems can result in significant calculation errors in the carbonate dissolution rate and make it difficult to conduct comparisons among regions.

Finally, based on the fundamental principle of the rock tablet method, this study proposes a suggested formula for calculating the carbonate dissolution rate using the rock tablet method. The suggested formula can further enhance the

standardization of the carbonate dissolution rate using the rock tablet method, providing data support for conducting global analysis and comparative research, revealing the laws of karst development, and improving the accuracy of karst carbon sink estimation. This study also proposes that methods should be selected based on the characteristics of the study area and research purpose. At the same time, it is of great necessity to strengthen the correlations between various test methods.

Key words carbonate rock, rock tablets method, dissolution rate, karst, problem discussion

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this regions, further diminishing the carbon sequestration capacity of the basin and gradually shifting its carbon source-sink effect from a carbon sink to a carbon source.(3) The single-factor detection results from the geographic detector indicate that land use is the main driving factor influencing the spatial heterogeneity of carbon storage in the Hongshui River Basin, with a q value of 0.833. Additionally, the average annual NDVI has been shown to explain the spatial heterogeneity of carbon storage, with a q value of 0.545. The interactive detection results show that the interaction between land use and the annual average NDVI factor have the most significant effect on the change in carbon storage within the Hongshui River Basin, with an explanatory power of 0.833. This indicates that the specific combination of the land use interactions, annual average NDVI and other factors—such as annual average temperature, annual average rainfall, digital elevation model, and population density—will influence the spatial distribution of carbon storage. The land use change factor is the main contributor to the increase of carbon storage in the Hongshui River Basin, followed by the annual average NDVI factor. The findings of this study may provide significant theoretical and data support for the sustainable development of carbon storage within ecosystem services in the Hongshui River Basin. Furthermore, they will assist in the formulation of more effective ecological protection and resource management policies aimed at enhancing the carbon sink capacity of the ecosystem and promoting environmental health and sustainable development.

Key words carbon storage, driving factors, inVEST model, karst area, Hongshui River Basin

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