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苔藓植物对川西亚高山区砂岩与碳酸盐岩风化特征的影响

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摘 要: 川西亚高山地区地形破碎、气候寒冷, 裸露岩石边坡生态修复困难, 探究苔藓生长对岩石风化过程的影响有利于寻找促进生态恢复的途径。文章通过对中国川西亚高山区不同尺度、不同风化条件下的砂岩和碳酸盐岩风化体和基岩性质进行分析, 以明确其风化微观特征, 并进一步探究苔藓植物生长对岩石风化的影响。砂岩和碳酸盐岩随着风化的进行, 矿物晶型遭到破坏, 岩体均出现裂隙, 出现次生矿物, 并伴随着铁氧化物和有机质等物质浸染, 随时间增加而加剧, 但砂岩在风化早期受到的风化作用比碳酸盐岩更强烈。苔藓生长对岩石风化的影响比较复杂, 在风化初期, 苔藓的作用不强, 随着风化的进行, 生物侵蚀作用加剧; 苔藓生长在砂岩表面, 则在风化初期就会加剧砂岩原本强烈的物理风化, 并发生物质成分的转变。总体而言, 苔藓植物的生长会加快岩石的风化进程, 但在风化初期的影响不明显, 若对碳酸盐岩边坡利用苔藓植物进行生态修复, 相比砂岩效果更好, 在一定时间内能起到保护作用。

关键词: 风化; 苔藓; 砂岩; 碳酸盐岩; 川西

创新点: 亚高山区苔藓植物对不同岩性基岩生物风化作用尚不明确。本研究通过多尺度技术手段, 发现苔藓对砂岩和碳酸盐岩风化微观特征的差异影响, 为山区裸露岩石边坡受损生态系统的恢复与演替研究提供理论依据。

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0 引 言

四川省西部山区是青藏高原与四川盆地的过渡区域, 气候复杂, 降雨充沛, 地貌以山地为主, 地势高差悬殊, 地质条件复杂, 构造活动频繁, 地质灾害频发, 加之人类活动强度逐渐增强, 裸露边坡和岩体分布广泛。岩石碎块风化脱落易引发次生灾害, 有着潜在的威胁; 岩石自出露地表开始风化, 着生微生物开始进行生态演替, 至岩面出现细土-砾质层^[1], 是土

壤生成的初级阶段^[2], 对土壤发生与形成有着重要的影响。因此有必要分析这些岩石边坡的风化机理和特征, 以及与生物, 尤其普遍分布的苔藓的关联作用, 可有助于寻求合适的生态修复方法, 维护岩石边坡生态系统稳定性, 降低山区裸坡其潜在生态风险与隐患。

岩石风化受岩性(母质)、气候、生物、微地貌和时间等因素控制, 关于气候与岩性对风化的影响已经有了许多研究结果^[3-4]。研究者们普遍认为气候会

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显著地影响风化岩石表面的结构特征,不同类型岩石风化过程中结构、化学成分等方面表现特征也不尽相同^[5]。而岩石风化成土过程也与岩生生物的演替紧密相连,目前已引起广泛的重视,有越来越多的研究者开始探索这一领域。微生物、地衣和苔藓通过附着在岩石表面生长,可使岩石进一步崩解^[6],或利用自身细胞结构,从外界吸收水分及养分,同时释放出糖类、氨基酸、脂肪酸及其衍生物,萜类和黄酮类化合物以及 CO₂ 等,其中生物酸和 CO₂ 对岩石均有一定的溶蚀作用^[7],改变风化岩石表面化学成分。苔藓对于岩石生境塑造十分重要,其生长于岩石表面,与微生物一起形成结皮层,可减小岩石基质遭受的风力侵蚀、降水冲刷等外力作用^[8],也能提高下伏土壤的微生物群落丰度,增强微生物活动,促进植被演替^[9]。目前已有研究尝试利用苔藓植物的生态功能进行裸露岩石边坡的生态修复^[10-11],有研究证明苔藓植物具有很高的成土率,可以有效缓解石漠化地区土壤稀少和养分贫瘠的现状^[12],也有研究通过数据分析表明某些苔藓种可用于石漠化治理^[13]。而不同岩性的基岩受苔藓活动的影响又有所不同,如石灰类岩石最容易遭受到苔藓的生物破坏^[14],砂岩则更容易因苔藓活动加剧受到的机械侵蚀^[15],也有研究表明苔藓植物生长在砂页岩上起到了固土持水的作用^[16]。因此,明确苔藓生长对不同岩石风化特征的具体影响,以评估利用苔藓植物进行岩石边坡生态修复的效果是很有必要的。

岩石风化过程是众多因素相互耦合的结果,而目前关于苔藓植物在岩石风化过程中的作用关注不足,有研究者统计并划分了苔藓植物对碳酸盐岩的溶蚀形态并在此之上分析其与环境因子的相关性^[17-18],也有研究表明苔藓植物会影响岩溶沉积的形态变化如加强钙化和石灰华现象^[19],这些研究目前大多停留在苔藓植物对岩石风化影响的描述性研究上,而关于山区裸露岩石在苔藓生长覆盖影响下的具体微观结构和成分变化尚不明确。因此本研究旨在通过分析四川西部亚高山地区,碳酸盐岩和砂岩在有或无苔藓覆盖情况下,经过不同的风化年限,其风化表面和新鲜基岩所展现的结构及成分特征,以此研究苔藓在岩石风化过程中发挥的作用以及对岩石风化微观特征的影响,这有助于明确在多因素影响情况下各单因素对岩石风化的作用效果,从而为人为控制风化成土进程、利用苔藓促进山区岩石边坡的生

态恢复提供理论依据。

1 材料与方法

1.1 研究区域概况

选择位于川西夹金山地区雅安市宝兴县、龙门山地区的阿坝藏族羌族自治州汶川县、理县以及成都市彭州市和德阳市什邡市山区为研究区域(图1)。

汶川县、理县位于岷江上游,区域构造活动强,山高坡陡,气候具有立体分布特征,植被和土壤垂直分布差异明显。河谷底部,气候干凉,年均气温 11.0 ℃,年降水量 400~600 mm,蒸发量为降水量的 2~4 倍^[20]。宝兴县气候同样垂直变化明显,具有亚热带到永冻带的垂直气候谱^[21]。彭州和什邡属于亚热带湿润气候区,气候温和湿润,雨量充沛,四季分明;境内气温由东南向西北逐渐降低,日照逐渐递减,降水量逐渐增多,丘陵、低山区、高山区间的气候差异明显(表1)。上述研究区地质条件复杂,构造活动频繁,气候环境变异大,因自然因素或人类活动(如道路修建)形成的岩石边坡面积广泛。

1.2 样品采集

选取位于宝兴县泽根村和黄店子沟、汶川县萝卜寨村、理县毕棚沟和古尔沟以及彭州市小鱼洞镇、什邡市蓥华镇等地的亚高山地区道路边坡岩石创面布设样点(图1),并在各样点分别选取砂岩或碳酸盐岩(石灰岩和白云岩)进行采样。岩石风化时间依据道路修筑年代确定。同时选取生长苔藓的岩石样点,其上覆盖的苔藓植物为以丛藓科、真藓科、羽藓科等苔藓植物为主的丛集型、平铺型混生苔藓植物群落(不包含地衣等其他植物)。在各采样点,分别采集苔藓覆盖生长、无苔藓生长的风化岩石和新鲜岩石样品(表1)。

1.3 分析与研究方法

用 20 cm×20 cm 的金属方框于岩石表面布设样方,并利用 GIS 软件的监督分类提取法提取样方内的苔藓覆盖度^[24]。实验选取不同岩性不同风化程度的代表试样进行薄片制作和岩石与矿物类型和特征鉴定。采用 BM-58XCS 型号偏光显微镜分析, Nikon 相机进行薄片照相鉴定。针对岩石试样,实验前将样品处理成面积为 2 cm²、厚度不超过 5 cm 的块状样,并将于导电纸接触的一面打磨平整,采用日本电

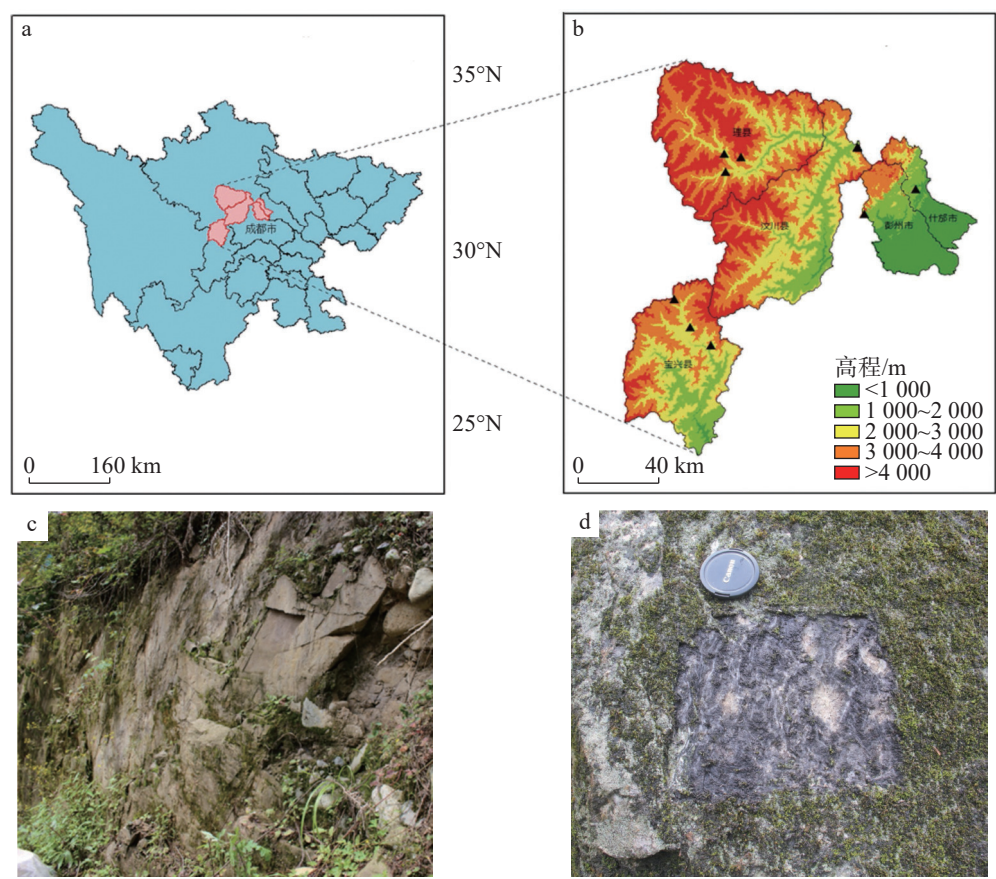


图 1 川西亚高山区研究样点情况
a, b.样点分布情况图 c.XY1 样地苔藓覆盖的风化砂岩 d.HDZ1 样地苔藓覆盖的风化碳酸盐岩表面
Fig. 1 Conditions of sampling points in the subalpine region of western Sichuan
a,b.distribution of sampling points c.moss-covered weathered sandstone in Plot XY1 d.moss-covered weathered carbonate rock surface in Plot HDZ1

表 1 采样点基本情况
Table 1 Basic information of sampling points

采样区域	采样地点/编号	海拔/m	风化时间/a	岩性	样方典型苔藓物种	苔藓盖度/%	年均温度/℃ ^[20]	年降水量/mm ^[22]
理县	古尔沟 (GEG)	2 215	2	砂岩	长叶提灯藓 <i>Mnium lycopodioides</i>	60	11.4	650 ^[23]
	毕棚沟 (BPG3)	2 290	49	砂岩	细叶小羽藓 <i>Haplocladium microphyllum</i>	90	11.4	650 ^[23]
宝兴县	夹金山 (JJSB)	3 880	3	砂岩	反扭藓 <i>Timmiella anomala</i>	30	19.3	1 275 ^[22]
什邡	莹华山 (YHSZ1)	1 003	1~2	砂岩	细叶小羽藓 <i>Haplocladium microphyllum</i>	10	14.9	770 ^[22]
彭州	小鱼洞镇 (XY1)	1 043	2.5	砂岩	拟灰羽藓 <i>Thuidium glaucinoides</i>	20	15.9	865 ^[22]
汶川	萝卜寨 (LBZ1)	2 013	2	灰岩	大羽藓 <i>Thuidium cymbifolium</i>	50	12.9	525 ^[23]
	萝卜寨 (LBZ2)	1 942	9	灰岩	硬叶对齿藓 <i>Didymodon rigidulus</i>	30	12.9	525 ^[23]
宝兴县	泽根村 (ZGX3)	2 430	5	灰岩	真藓 <i>Bryum argenteum</i>	15	15.3	990 ^[21]
	黄店子沟 (HDZ1)	2 353	27	白云岩	细叶真藓 <i>Bryum capillare</i>	90	15.3	990 ^[21]

子(JEOL)生产的JSM-7500F场发射扫描电子显微镜测试。样品矿物组成在四川大学分析测试中心利用X射线衍射仪(荷兰帕纳科公司, EMPYREAN)测定。

2 结果与讨论

2.1 风化碳酸盐岩与砂岩基本微观特征

2.1.1 风化碳酸盐岩微观特征

新鲜石灰岩中方解石晶体层理清晰, 晶型结构

完整, 呈块状或鲕粒状, 晶体之间粘着紧密、无缝隙(图2a, 图2d, 图2g); 而风化作用使碳酸盐岩表面微观结构发生明显变化, 结构变差, 晶体结构的完整性遭到破坏, 岩石表面出现明显裂缝、裂隙, 出现溶蚀孔洞和有机质浸染(图2b, 图2e, 图2h)。风化时间为2a的LBZ1灰岩样品, 除溶蚀孔洞和缝隙等现象外, 难以观察到黏土矿物(图2b)。风化时间为5a的ZGX3灰岩样品, 除方解石外, 能观察到少量石英以及隐晶质黏土矿物(图2e)。暴露27a的HDZ1白云岩样品, 晶体结构变化更明显, 孔隙更大, 溶蚀孔洞更多, 有机质浸染现象明显, 除原有矿物白云石、方

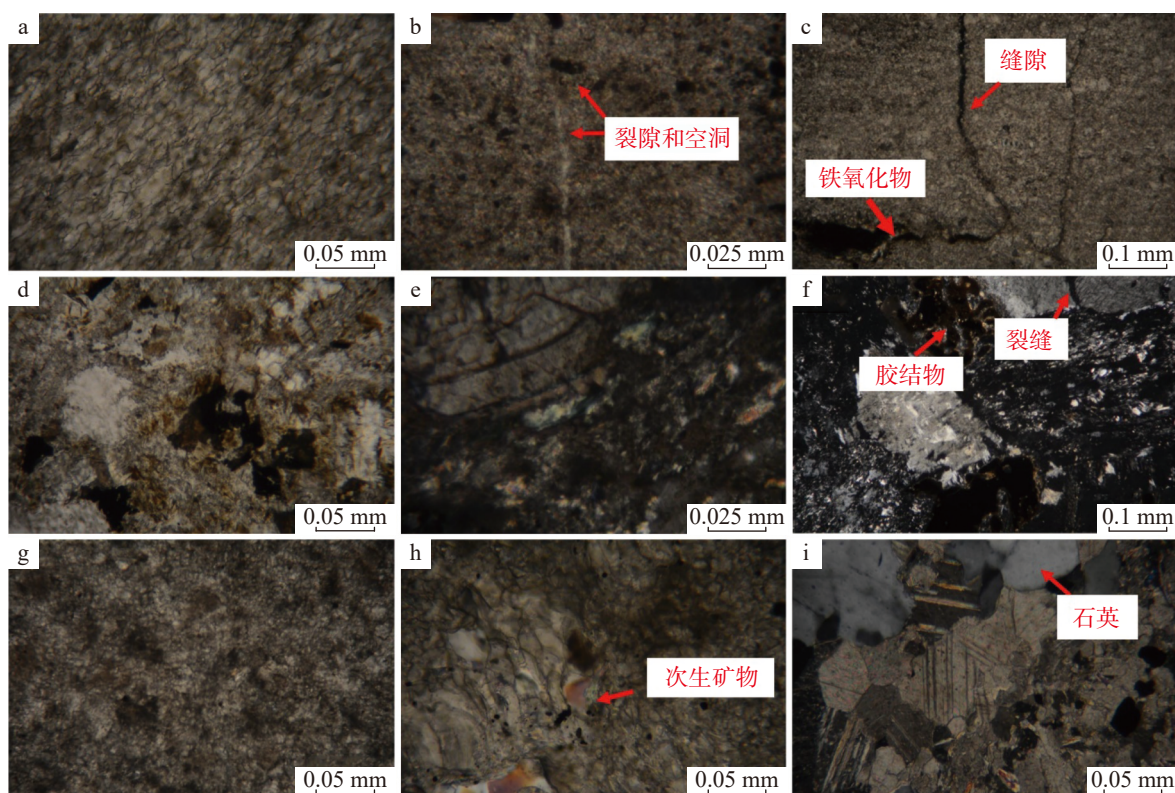


图2 新鲜、无苔藓覆盖、有苔藓覆盖碳酸盐岩表面微形态特征

a.LBZ1样地石灰岩新鲜基岩, 方解石晶形完整, 颗粒均一 b.LBZ1样地无苔藓覆盖的石灰岩表面, 风化时间2a, 方解石晶形开始不完整, 有明显的裂隙和孔洞 c.LBZ1样地有苔藓覆盖的石灰岩表面, 风化时间2a, 有较深较长的裂隙, 出现铁氧浸染现象 d.ZGX3样地石灰岩新鲜基岩, 方解石晶形较完整, 成块状 e.ZGX3样地无苔藓覆盖的石灰岩表面, 风化时间5a, 表面有机质浸染, 可见次生矿物 f.ZGX3有苔藓覆盖的石灰岩表面, 风化时间5a, 有明显裂缝, 可见较多杂质 g.HDZ1样地白云岩新鲜基岩, 白云石晶形较为完整, 以粒状或块状为主 h.HDZ1样地无苔藓覆盖的白云岩表面, 风化时间27a, 白云石晶形发生改变, 可见次生矿物 i.HDZ1样地有苔藓覆盖的白云岩表面, 风化时间27a, 白云石变少, 可见斜长石、石英, 有较多溶蚀孔洞

Fig. 2 Micro-morphological features of weathered surfaces of fresh, moss-free, and moss-covered carbonate rocks

a.fresh bedrock of limestone in Plot LBZ1, with intact crystal forms of calcite and uniform grain size b.moss-free limestone surface of Plot LBZ1, weathered for 2 years, showing partial degradation of calcite crystal forms with visible cracks and pores c.moss-covered limestone surface of Plot LBZ1, weathered for 2 years, featuring deeper and longer cracks with iron oxide staining d.fresh bedrock of limestone in Plot ZGX3, with relatively intact calcite crystals in a massive form e.moss-free limestone surface of Plot ZGX3, weathered for 5 years, exhibiting organic matter staining and visible secondary minerals f.moss-covered limestone surface of Plot ZGX3, weathered for 5 years, showing prominent cracks and an increase in impurities g.fresh bedrock of dolomite in Plot HDZ1, characterized by relatively intact dolomite crystals primarily in a granular or block form h.moss-free dolomite surface of Plot HDZ1, weathered for 27 years, displaying altered dolomite crystal forms with visible secondary minerals i.moss-covered dolomite surface of Plot HDZ1, weathered for 27 years, showing a decrease in dolomite crystals with visible plagioclase and quartz, and numerous solution pores.

解石外,也可观察到黏土矿物(图 2 h)。

2.1.2 风化砂岩微观特征

未风化砂岩内部石英等矿物晶型完整、结构较为清晰,物质组成不均一,骨骼颗粒矿物之间由碳酸盐岩等微晶或隐晶矿物胶结,微观结构呈块状和网状,结构完整性较高(图 3a, 图 3d)。风化时间为 2.5 a 的 XY1 砂岩样品,岩石风化表面出现碎裂和片状龟裂,有片状和颗粒状剥落现象,岩石风化体有明显裂缝裂隙,杂质增多,填充孔隙,结构疏松,连结不紧密,石英颗粒变小,表面微观结构不规则,呈现块状或板状,有明显断裂痕迹,孔隙较多(图 3b); 风化时间为 2 a 的 GEG 砂岩样品,其表面层状结构发生明显改变,呈现不规则的蜂窝状,杂质明显增多,孔隙变多(图 3e)。

2.2 生长苔藓的风化岩石微观特征

2.2.1 生长苔藓的风化碳酸盐岩微观特征

对比生长苔藓的风化岩石表面和未生长苔藓的风化岩石表面可以发现,生长苔藓的风化岩石比没

有苔藓生长的岩石出现的裂隙更深、更宽,有机质和铁氧化物填充孔隙现象更明显,岩石面孔率增大(图 2c, 图 2f)。这是由于苔藓植物对岩石有较大的生物溶蚀作用,侵入岩石表面凹凸不平处或者微小裂隙开始生长,使得裂纹变宽加深^[25]。苔藓植物在干湿交替作用下发生膨胀、收缩和卷曲,分泌有机酸溶蚀基岩表面大量碳酸盐岩颗粒,从而扩大了岩石参与风化的面积^[26],使灰岩中的方解石晶形发生改变。

新鲜石灰岩中方解石晶体微观形状相对完整(图 4a),而风化 2 a、有苔藓生长的石灰岩整体微观结构改变不大,但可见絮状杂质,晶体粒径略微变小,具有孔隙(图 4c); 风化和新鲜石灰岩主要化学组成相同,均以 C、O、Ca 为主(图 4a, 图 4d)。风化 27 a 的 HDZ1 白云岩样品矿物以白云石为主,但岩石出现许多缝隙和溶蚀孔洞,白云石晶体粒径显著减小,晶体形状遭到破坏,有块状、板状次生矿物分布(图 4d)。另外,风化时间短的 LBZ1 样地灰岩,各元素含量变化不显著,风化灰岩表面矿物成分与新鲜基岩基本一致,均以方解石为主,含量都在 90% 以上

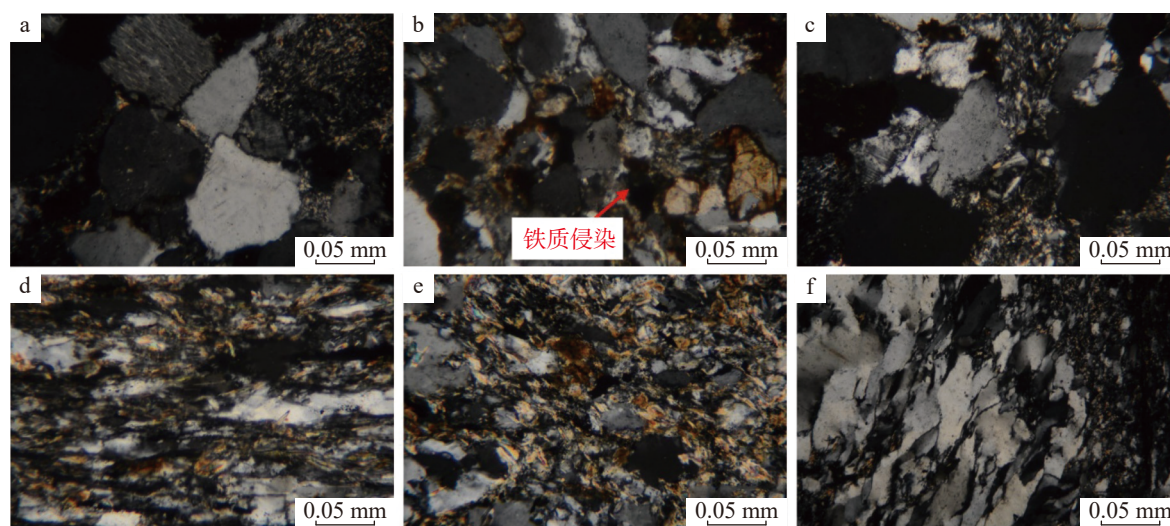


图 3 新鲜、无苔藓覆盖、有苔藓覆盖碳酸盐岩风化表面微形态特征

a.XY1 样地砂岩新鲜基岩,晶型完整,结构清晰,呈网状 b.XY1 样地无苔藓覆盖的砂岩表面,风化时间 2.5 a,晶体粒径变小,杂质变多,有次生矿物,结构不完整 c.XY1 样地有苔藓覆盖的砂岩表面,风化时间 2.5 a,结构破坏严重,孔隙变多,有大量铁质浸染,片状剥落痕迹明显 d.GEG 样地砂岩新鲜基岩,结构清晰紧致,杂质较少,呈层状 e.GEG 样地无苔藓覆盖的砂岩表面,风化时间 2 a,杂质增多,有机质和铁氧化物填充,结构松散 f.GEG 样地有苔藓覆盖的砂岩表面,风化时间 2 a,结构破坏严重,孔隙增多,片状破坏痕迹明显

Fig. 3 Micro-morphological features of weathered surfaces of fresh, moss-free, and moss-covered carbonate rocks

a.fresh bedrock of sandstone in Plot XY1, exhibiting intact crystal forms and clear, mesh-like structures b. moss-free sandstone surface of Plot XY1, weathered for 2.5 years, showing reduced crystal grain sizes, increased impurities, and the presence of secondary minerals, with an incomplete structure c.moss-covered sandstone surface of Plot XY1, weathered for 2.5 years, displaying significant structural damage, increased porosity, abundant iron staining, and clear signs of flaking d.fresh bedrock of sandstone in Plot GEG, characterized by a clear, compact structure with minimal impurities and a layered appearance e.moss-free sandstone surface of Plot GEG, weathered for 2 years, showing increased impurities with organic matter and iron oxides filling the structure, resulting in a loose texture f.moss-covered sandstone surface of Plot GEG, weathered for 2 years, exhibiting severe structural damage, increased porosity, and prominent signs of flaking.

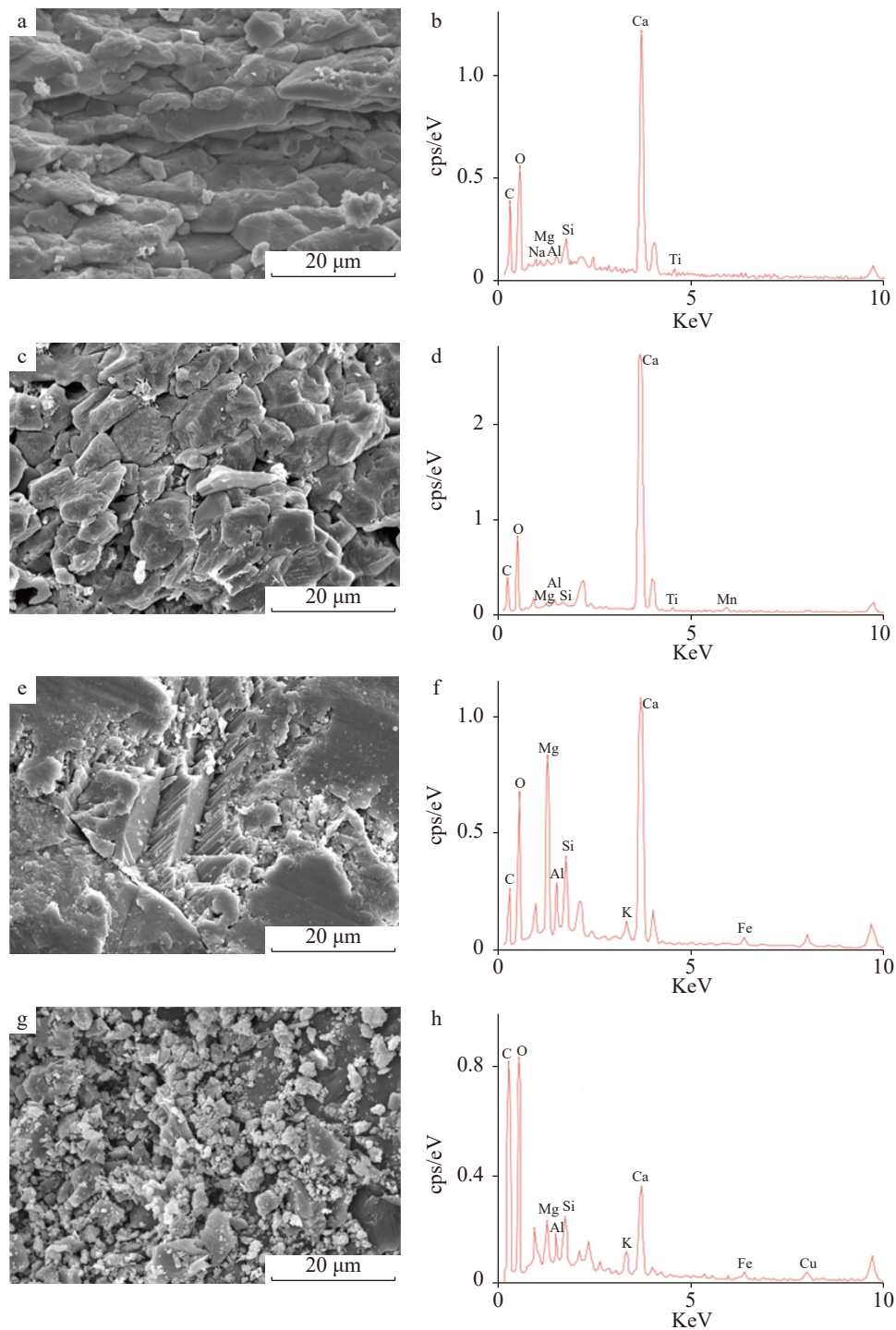


图 4 新鲜与苔藓生长碳酸盐岩样品扫描电镜及 EDS 图

a, b. LBZ1 样地新鲜基岩, 结构完整清晰 c, d. LBZ1 样地苔藓覆盖风化表面, 晶体粒径变小, 有絮状杂质, 产生孔隙 e, f. HDZ1 样地新鲜基岩, 结构完整, 有断裂痕迹 g, h. HDZ1 样地苔藓覆盖风化表面, 晶体破坏严重, 形状发生改变, 杂质增多, 孔隙较多。

Fig. 4 SEM and EDS images of fresh and moss-growth carbonate rock samples

a, b. fresh bedrock of Plot LBZ1, exhibiting an intact and clear structure c, d. moss-covered weathered surface of Plot LBZ1, characterized by reduced crystal grain size, the presence of flaky impurities, and increased porosity e, f. fresh bedrock of Plot HDZ1, displaying an intact structure with signs of fracturing g, h. moss-covered weathered surface of Plot HDZ1, showing significant crystal degradation, altered shapes, increased impurities, and greater porosity.

(图 5), 这说明有苔藓覆盖的岩石风化表面, 在风化时间不长的情况下, 其生物风化作用并不强烈, 矿物

成分的改变并不明显。经过长时间风化的白云岩样品, C 元素相对含量增加 20%, O、Ca、Mg 元素含量

下降 8% 左右,指示主要矿物白云石($\text{CaMg}(\text{CO}_3)_2$)在苔藓分泌的有机酸作用下,遭受长期溶蚀,Ca、Mg 元素淋失^[27],而硅酸盐矿物为主的黏土矿物相对增加,有机物质残留在岩石表面(图 4b,图 4d)。不同

风化时间的碳酸盐岩风化特征表现出较大的差异,这可能也应证了苔藓不同演替阶段对岩石的溶蚀效果表现不同的观点^[18]。

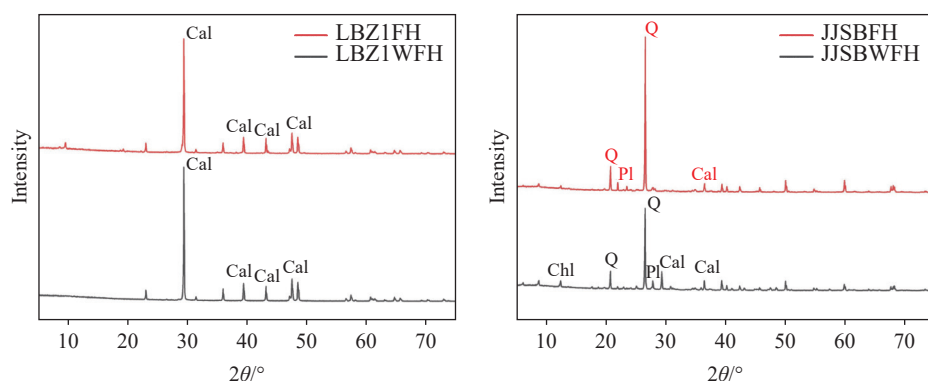


图 5 新鲜和苔藓覆盖风化石灰岩和砂岩 XRD 衍射图

a.LBZ1 灰岩样品,风化时间 2 a 样品(LBZ1FH),风化表面与新鲜基岩(LBZ1WFH)成分基本一致 b.JJSB 砂岩样品,风化时间 3a 样品(JJSBFH),风化表面相对于新鲜基岩(JJSBWFH),方解石、斜长石相对含量减少,石英增多

Q.石英 Cal.方解石 Dol.白云石 Pl.斜长石 Chl.绿泥石 Ms.白云母

Fig. 5 XRD diffraction patterns of fresh and moss-covered weathered limestone and sandstone

a.the LBZ1 limestone sample, weathered for 2 years (LBZ1FH), showing a composition that is largely consistent with the fresh bedrock (LBZ1WFH) b.the JJSB sandstone sample, weathered for 3 years (JJSBFH), exhibiting a relative decrease in the content of calcite and plagioclase compared to the fresh bedrock (JJSBWFH), while with the increase in amount of quartz Q.quartz Cal.calcite Dol.dolomite Pl.plagioclase Chl.chlorite Ms.Muscovite

2.2.2 苔藓覆盖的风化砂岩微观特征

有苔藓覆盖的砂岩风化表面对比新鲜基岩和无苔藓覆盖的风化表面,其晶体结构遭到明显改变,原有层状或块状结构遭到破坏,表面出现大量孔洞和鳞片状剥离脱落的痕迹(图 3c,图 3f),这是因为砂岩颗粒会因温度变化或周期性温差变化导致膨胀和收缩不均一,进而导致岩石表层碎裂和片状龟裂,最终导致片状酥裂。同时,苔藓植物在干湿交替作用下发生膨胀、收缩和卷曲,会使砂岩本就因温度、湿度变化发生的起翘、膨胀、脱落加剧^[28]。而岩石表面出现片状开裂酥解和微裂缝现象,表明在苔藓覆盖的砂岩风化早期,物理风化很强^[29],且苔藓会加剧砂岩受到物理风化强度。风化表面可见有机杂质增多,仍有次生矿物(图 3c,图 3f),总体呈现石英占比进一步增大,绿泥石、斜长石等被转化,矿物成分单一化的趋势。

从微观上看,新鲜砂岩晶体形状以显晶颗粒为主,晶形完整(图 6a,图 6c)。风化后的岩石尽管结构和矿物晶形变化基本不变,但大部分疏松,连结不紧密,疏松部可见较多风化碎屑以及溶缝、溶孔,有片

状和颗粒状剥落情况,还有较多絮状基质(图 6b,图 6d)。岩石的化学元素组成也出现相应变化,经 2a 左右风化的砂岩 C 元素含量百分比增加 8% 以上, Si 元素含量百分比降低 5%(图 6b,图 6h),表明在苔藓活动影响下, Mg、Al 等易活动元素相对淋失,含量减少,有机质充填岩石缝隙,导致 C 元素含量增加;因石英稳定性很高,抗风化能力强,风化过程中不容易发生化学风化^[25],使 Si 等元素相对富集^[30],但可能因为岩石碎块脱落造成含量降低。风化表面 Fe、Mg 元素含量降低, O 元素增加,这是因为砂岩风化中胶结物是最先风化的地方^[31],而随着风化程度的加深,岩石裂隙和孔隙增大,氧化物含量增多。

2.3 碳酸盐岩与砂岩风化特征差异

风化初期,砂岩整体结构受到的破坏更严重,碳酸盐岩的整体结构更完整,但容易产生溶蚀孔洞,这与灰岩主要矿物方解石在酸性风化条件下,易发生化学溶蚀作用有关^[32]。砂岩风化过程中,风化表面的断裂、剥落痕迹,使得风化砂岩表面显现许多杂色(图 4b,图 4e)。砂岩矿物成分复杂,含有石英、长石、云母以及方解石、黏土矿物、铁氧化物等^[33]。当风化

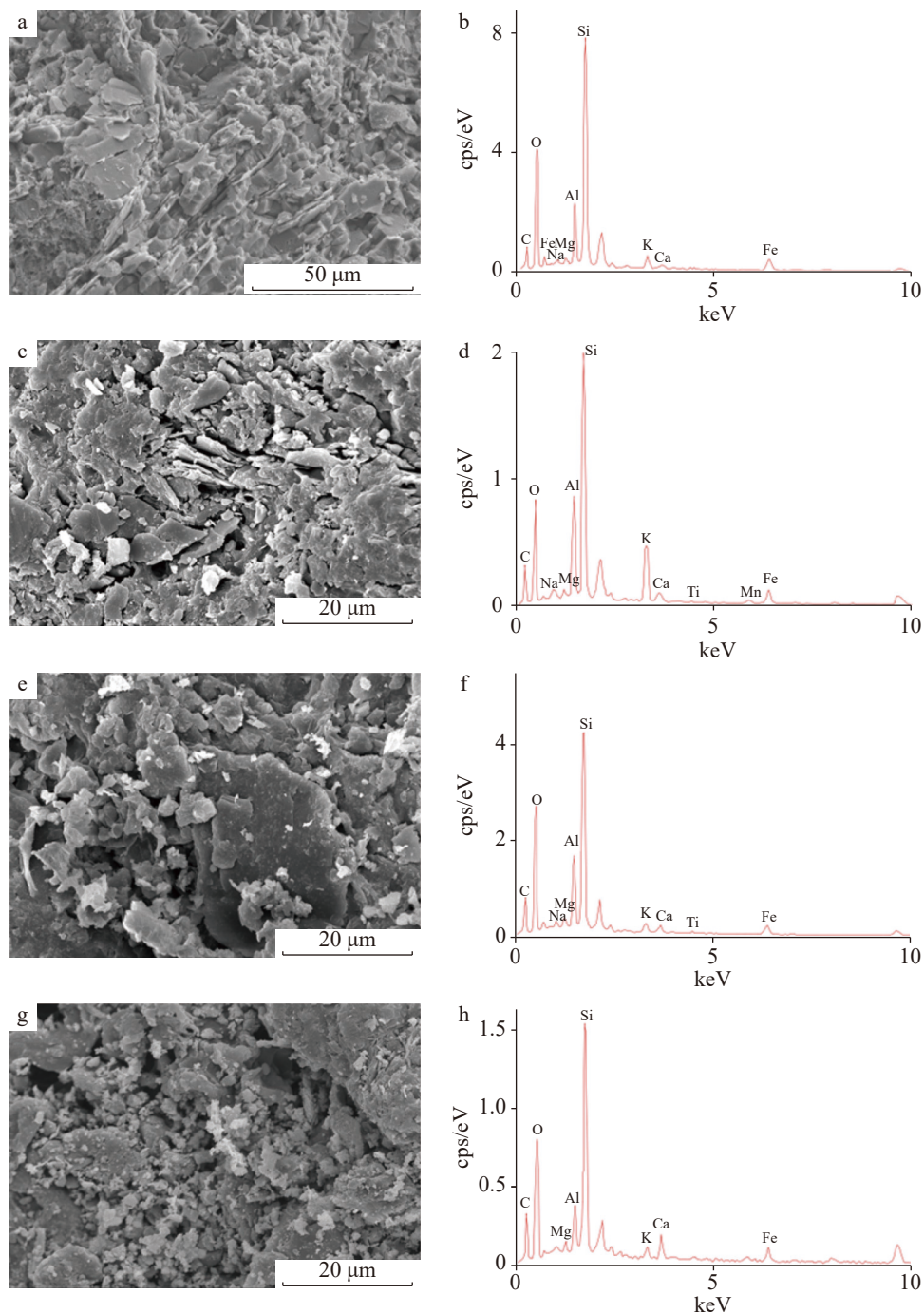


图6 砂岩新鲜基岩与苔藓覆盖风化表面扫描电镜及EDS线总谱图

a, b. XY1 样地, 新鲜基岩, 结构清晰紧致, 呈层状 c, d. XY1 样地, 风化表面, 结构破坏严重, 有许多片状碎片, 可见杂质和次生矿物 e, f. YHSZ1 样地, 新鲜基岩, 晶体颗粒较为完整, 呈块状, 杂质较少 g, h. YHSZ1 样地, 风化表面, 晶体结构明显变小, 结构松散, 孔隙增多, 杂质增多, 可见次生矿物

Fig. 6 SEM and EDS spectrum of fresh sandstone bedrock and moss-covered weathered surface

a, b. Plot XY1, fresh bedrock exhibiting a clear and compact structure with distinct layering c, d. Plot XY1, the weathered surface showing significant structural degradation, with numerous flaky fragments and visible impurities along with secondary minerals e, f. Plot YHSZ1, the fresh bedrock featuring relatively intact crystal grains and a blocky appearance with minimal impurities g, h. Plot YHSZ1, the weathered surface revealing a pronounced reduction in crystal structure, a loose texture, increased porosity, and a rise in impurities, along with the presence of secondary minerals.

发生时, 方解石等胶结物发生溶蚀、水解等作用, 最终除石英外只剩少量云母、长石等, 晶体孔隙扩大, 结构变松散^[34]。灰岩原生矿物较简单, 主要为方解

石, 含少量石英、长石、白云石及黏土矿物, 方解石化学性质不稳定, 易在外界环境因素影响下发生流失, 而石英化学性质十分稳定^[30], 长石、白云石等随

着风化其含量减少,甚至消失,石英则在风化过程中填充至岩石缝隙(图 2i)。随着风化进行,灰岩的矿物成分趋向于复杂,过程中原生矿物分解,出现其他黏土矿物,石英相对含量逐渐增加;而砂岩的主要矿物成分一直都是石英,矿物种类趋于减少。

苔藓生长的碳酸盐岩表面除缝隙扩大外其他变化不明显(图 2c, 图 2f),而风化时间 2 a 的 LBZ1 新鲜和苔藓覆盖风化石灰岩样品都只检测出方解石(图 5),说明在风化初期,苔藓植物对风化作用影响尚不明显。风化时间 5 年的 ZGX3 碳酸盐岩样品,其表面苔藓盖度相对于 LBZ1 样地更低(表 1),其结构破坏更为严重,受到的风化作用更强;随着风化时间增加,苔藓分泌更多有机酸,生物风化作用增强,风化 27 年的 HDZ1 碳酸盐岩样品表面的溶蚀孔洞明显增多,填充孔隙和溶洞的有机物质增多。这与已有研究中苔藓植物对碳酸盐岩的生物溶蚀形态相印证^[35]。可见在风化初期,苔藓生长对碳酸盐岩风化的影响不明显,随着风化进行,苔藓对碳酸盐岩风化特征的影响也明显加强^[36]。

苔藓覆盖的砂岩在风化早期,就体现出明显的作用,使砂岩本就因气候变化发生的起翘、膨胀、脱落加剧,加强了砂岩受到的物理风化。苔藓盖度更低的 XY1 样地砂岩较风化时间相近的苔藓盖度更高的 GEG 样品的岩石风化表面结构破坏明显,岩石碎屑剥离崩解更明显以及产生更多隐晶质粘土矿物(图 3c, 图 3f),这可能与前者位于更高年均温和年降水量的龙门山东坡,风化作用更强有关^[37]。同时,因苔藓和微生物在岩石表层活动作用加强,生物残体等会淀积表面,是苔藓为主的生物风化作用标志^[38],生长苔藓的砂岩风化表面可见有机质增加(图 3c, 图 3f)以及碳含量提高(图 6d, 图 6h)。因此苔藓生长对砂岩风化的影响在风化早期就很明显,不仅加剧了砂岩的物理风化,苔藓本身的生物风化作用也很强烈,不过苔藓盖度更低的砂岩样品受到的结构破坏更严重。

3 结 论

本研究分析了川西亚高山区碳酸盐岩和砂岩风化表面的微观结构特征、矿物成分和元素变化,探究了苔藓植物群落生长对于山区岩石风化特征的具体微观影响,得到以下结论:

(1) 碳酸盐岩风化以化学风化为主,随着风化时间变长,除去结构特征的改变逐渐明显外,其化学成分由简单变复杂。砂岩新鲜基岩矿物组成非单一,在风化初期,以物理风化为主,结构上的变化相比碳酸盐岩更明显,矿物成分逐渐由复杂变简单,并伴有少量次生矿物。

(2) 生长苔藓的碳酸盐岩在短时间内风化作用并不强烈,各元素含量百分比的变化不超过 4%,风化时间长的碳酸盐岩,其结构特征和化学成分的变化都更明显,C 元素含量均增加超过 20%,O、Ca、Mg 等元素含量降低超过 8%,其受到了苔藓强烈的生物溶蚀。生长苔藓的砂岩,从风化初期开始,就加剧了物理风化的相关现象,使主要矿物成分受到损失,Si 元素含量下降超过 5%,C 元素含量增加超过 8%。植被盖度更高的碳酸盐岩风化岩石表面,在风化初期相对于盖度低的风化岩石表面,结构更完整,砂岩的表现与其类似。砂岩风化过程受苔藓生长的影响更强烈。

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Influence of mosses on the weathering features of sandstone and carbonate rocks in the subalpine region of western Sichuan

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Abstract Sichuan Province is located in the Southwest China, with its western mountainous area serving as a transitional zone between the Qinghai-Xizang Plateau and the Sichuan Basin. The climate is complex, characterized by abundant rainfall, and the terrain is mainly mountainous, featuring significant variations in elevation. The geological conditions are complex, tectonic activities are frequent, and geological disasters occur regularly. In addition, with the increasing intensity of human activities, exposed slopes and rock masses are widely distributed. The weathering and detachment of rock fragments can easily trigger secondary disasters, posing potential threats. The widespread presence of rocky slopes makes these areas prone to such events, underscoring the urgent need for ecological restoration. Mosses play a critical role in ecological succession and soil formation on rock surfaces, acting as pioneer species that promote the development of thin soil layers. Exploring the impact of moss growth on rock weathering processes is beneficial for finding ways to promote ecological restoration. This study aims to investigate the weathering processes and mechanisms in this region by analyzing weathered rock bodies and bedrock properties under varying weathering conditions to elucidate their microscopic characteristics. Additionally, the influence of moss growth on rock weathering was further investigated. We collected samples from carbonate rocks and sandstones at various sites within

several counties in western Sichuan Province. Our study focused on sites with varying degrees of moss coverage and differing weathering durations. The analyses included examining the microstructural features of weathered rock surfaces, assessing mineral composition, and analyzing elemental variations to understand how mosses interact with these geological materials. These investigations were conducted by using electron microscopy, scanning electron microscopy (SEM), X-ray diffraction (XRD) for mineral composition, energy-dispersive spectroscopy (EDS), and other analytical techniques.

The results indicate that as weathering progressed, the breakdown of mineral crystal structures, the development of fractures in the rock mass, and the formation of secondary minerals occurred in both sandstone and limestone. These processes were accompanied by the infiltration of iron oxides and organic matter, and they intensified over time. Notably, sandstone exhibited a more pronounced weathering response during the early stages compared to limestone. During the weathering of carbonate rocks, a large amount of secondary minerals was produced, and the mineral composition evolved from simple to complex. In contrast, the main component of sandstone weathering has consistently been quartz; as weathering progressed, secondary minerals disappeared, and the mineral composition shifted from complex to single. The elemental content in limestone did not change significantly during the early stages of weathering, while the elemental content in sandstone changed markedly. In total, results show that the weathering of carbonate rocks is predominantly chemical, whereas physical weathering is more pronounced in sandstone.

Furthermore, the impact of mosses on rock weathering is complex. For carbonate rocks, initial weathering stages exhibited minimal elemental changes with moss presence; however, with prolonged exposure, significant modifications in structural integrity and chemical composition occurred. The percentage of carbon and other elements, such as calcium and magnesium, changed markedly, due to the bioerosion effects of moss-generated organic acids, with carbon increasing and calcium and magnesium decreasing. Microstructural analysis revealed that moss-covered rocks exhibited deeper and wider fissures compared to their bare counterparts. This increase in fissure size is correlated with the mechanical and biochemical actions of mosses, which exploited small crevices to enhance their growth and further facilitate the weathering process. Over time, mosses create a conducive environment for microbial activity, which adds to the biochemical weathering of the rock substrate. When mosses grew on the surface of sandstone, it exacerbated the already intense physical weathering of sandstone during the early stages and resulted in compositional changes. Overall, moss growth accelerates the weathering process of rocks, but its influence is not significant during the early stages.

The findings indicate that mosses significantly contribute to both the physical and chemical weathering processes of rocks, which in turn influence soil formation and ecological stability. These insights have important implications for ecological restoration strategies on exposed rock slopes, particularly in areas susceptible to erosion and geological instability. By harnessing the natural weathering capabilities of mosses, we can develop sustainable methods for restoring degraded landscapes and mitigating ecological risks associated with exposed rock surfaces. For the ecological restoration of limestone slopes, the use of mosses is more effective than its application on sandstone, as it provides protective benefits over time. Therefore, it is essential to consider the differences in rock types when we select suitable plants for ecological rehabilitation. Future studies should explore the long-term effects of moss-induced weathering on soil development and ecosystem restoration, particularly in sensitive mountainous regions.

Key words weathering, moss, sandstone, carbonate rock, western Sichuan

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