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江西昆山钼钨铜多金属矿床辉钼矿 Re-Os 年代及其对成矿地球动力学演化的启示

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提要:【研究目的】昆山钼钨铜多金属矿床行政区划分属于江西修水县境内, 大地构造位置位于江南造山带中段之九岭隆起东段。本次研究试图对昆山 W-Cu-Mo 多金属矿床成矿年龄进行精准约束, 为该矿床成矿作用研究、成矿规律总结提供年代学证据, 同时对该矿床成岩成矿关系、成矿物质来源作简要探讨, 并对大湖塘钼钨铜多金属矿集区成矿期次进行划分。【研究方法】本文对昆山钼钨铜多金属矿床开展了辉钼矿 Re-Os 同位素年代学研究。【研究结果】实验获得的 10 件辉钼矿等时线年龄为 (148.8 ± 1.0) Ma, MSWD=0.42, Re-Os 加权平均年龄为 (148.7 ± 0.7) Ma, MSWD=0.10, 加权平均年龄与等时线年龄基本一致, 因此本次所采集的样品 Re-Os 体系是封闭的。 ^{187}Re 和 ^{187}Os 沿等时线距离拉开的较远, 表明本次定年数据较为可靠。昆山矿床成矿作用发生稍滞后于似斑状黑云母花岗岩侵位时限, 与晚侏罗世这期岩浆活动有成因联系。昆山钼钨铜多金属矿床的成矿物质以壳源为主, 考虑到存在较高 Re 两个样品 ($n \times 10^{-5}$), 说明可能存在部分幔源的成矿物质参与了该矿床的形成。【结论】软流圈物质自晚侏罗世开始, 沿着 NNE 向深断裂与先期(晋宁晚期)形成的 NEE 向古断裂的交汇部位上涌, 通过热传导效应引起地壳物质发生部分熔融, 于大湖塘地区形成了似斑状黑云母花岗岩(岩株)、细粒黑云母花岗岩(小岩株)、花岗斑岩(岩脉或岩枝)等一系列中酸性侵入岩。

关 键 词: 钼钨铜多金属矿床; Re-Os 年龄; 同位素; 成矿作用; 江南造山带; 矿产勘查工程; 昆山; 江西

创 新 点: (1) 采用辉钼矿 Re-Os 同位素年代学, 对昆山 Mo-W-Cu 多金属矿床进行了成矿时代约束; (2) 依据辉钼矿石中 Re 的含量确定了成矿作用过程中有幔源物质的参与; (3) 总结归纳了大湖塘地区的两期成矿作用与成岩成矿地球动力学演化模式。

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Molybdenite Re-Os isotopic age of the Kunshan Mo-W-Cu polymetallic deposit in the Jiangxi Province and its evolution of metallogenic geodynamics

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Abstract: This paper is the result of mineral exploration engineering.

[Objective] Kunshan molybdenum polymetallic deposit is located in Xiushui County, Jiangxi Province, and belongs to eastern segment of the Jiuling uplift from mid-segment of the Jiangnan orogen. This study attempts to precisely constrain the metallogenetic age of the Kunshan W-Cu-Mo polymetallic deposit, provide geochronological evidence for the study of mineralization and the summary of the metallogenetic regularity of the deposit, study the metallogenetic relationship of the deposit and the source of metallogenetic material, and divide the metallogenetic periods of the Dahutang tungsten polymetallic ore concentration area. **[Methods]** The Re-Os isotopic geochronology for the molybdenite from the Kunshan molybdenum polymetallic deposit was carried out in this paper. **[Results]** An isochron age of (148.8 ± 1.0) Ma, and a weighted mean age of (148.7 ± 0.7) Ma for 10 molybdenite samples were obtained, which were approximately simultaneous with each other, and the Re-Os system was closed accordingly. The points of ^{187}Re and ^{187}Os are far away along the isochron, which indicates that the geochronology data are reliable. Metallogenesis lags behind the emplacement of porphyritic-like biotitic granite in the Kunshan deposit, caused by Late Jurassic granitic magmatism. Metallogenetic materials of the Kunshan molybdenum polymetallic deposit mainly came from crust materials, in addition, there are two samples with higher Re, it shows that partial mantle source metallogenetic materials participated in the formation of the deposit. **[Conclusions]** The intrusion of asthenospheric materials began from the Late Jurassic epoch, through the intersection part of NNE deep faults and NEE ancient faults formed from Late Jinningian. Partial melting of crustal material was caused by the heat conduction effect, which forms a series of intermediate-acid intrusive rocks in the Dahutang area, such as porphyritic-like biotitic granite, fine-grained biotitic granite, and granite-porphyry.

Key words: Mo-W-Cu polymetallic deposit; Re-Os age; isotope; metallogenesis; Jiangnan orogen; mineral exploration engineering; Kunshan; Jiangxi Province

Highlights: (1) The molybdenite Re-Os isotope chronology of Kunshan Mo-W-Cu polymetallic deposit metallogenetic epoch constraints; (2) Based on the content of Re in molybdenite ore, it is determined that mantle-derived materials are involved in the process of mineralization; (3) The two-stage mineralization and the evolution model of diagenetic geodynamics in the Dahutang area are summarized.

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1 引言

2012年江西省地质矿产勘查开发局916大队于大湖塘钨多金属矿集区石门寺钨铜钼多金属矿床取得WO₃金属量取得突破之后,打破了江西“南钨北铜”的分布格局,并成就了其迄今为止世界级第二大钨矿的地位,学者们从多种地质维度采用了多种实

验方法对其开展了一系列研究,并取得了丰硕的成果。项新葵等(2012a, 2013a,b,c)对石门寺矿床的地质特征进行了总结,约束了各类成矿条件。黄兰椿和蒋少涌(2012, 2013)对狮尾洞矿床似斑状白云母花岗岩和富钨花岗斑岩成因进行了探讨(Huang and Jiang, 2014)。项新葵等(2015a, b, c)对石门寺矿床同构造花岗斑岩成因及流体迁移规律进行了总结。

随着Re-Os同位素测年技术被应用于本区,先后获得了一系列成矿年龄数据,如石门寺矿床成矿年龄(139.2 ± 1.0)Ma(Mao et al., 2013)、(143.7 ± 1.2)Ma(丰成友等,2012)和(149.6 ± 1.4)Ma(项新葵等,2013a);狮尾洞矿床成矿年龄(140.9 ± 3.6)Ma(丰成友等,2012);昆山矿床成矿年龄(151.0 ± 1.3)Ma(张明玉等,2016);梅子坑矿床成矿年龄(148.9 ± 4.4)Ma(张志辉等,2013)等。这些研究工作为研究大湖塘钨多金属矿集区成矿作用、金属矿质来源、年代学格架搭建、地球动力学演化提供了重要支撑。

成矿年龄时限精准约束是认识成矿规律的基石,是建立成矿(找矿)地质勘查模型的重要参考资料,是反演成矿动力学演化和搭建年代学格架的基础资料(杨斌等,2021;丰晓路等,2022)。本文通过详细研究昆山钼钨铜多金属矿床基础资料,运用辉钼矿Re-Os同位素测年方法,试图对其成矿年龄进行精准约束,为该矿床成矿作用研究、成矿规律总结提供年代学证据,同时对该矿床成岩成矿关系、成矿物质来源作简要探讨,并对大湖塘钨多金属矿集区成矿期次进行划分,对该区成岩成矿地球动力学演化规律进行进一步总结。

2 成矿地质背景

昆山钼钨铜多金属矿床行政区划分属于江西修水县境内,大地构造位置位于江南造山带中段之九岭隆起东段(图1),东邻鄱阳凹陷,修水—武宁滑覆拗褶带以南,九岭南缘推覆构造带北侧(杨明桂和王昆,1994;项新葵等,2012a,b)。成矿区隶属于江南地块钨铜(金、钼等多金属)成矿带(项新葵等,2013c)。自古元古代以来,九岭隆起经历了扬子结晶基底的形成,随后其褶皱的隆升运动主要发生于古生代时期,至燕山期,则以板内构造运动(挤压-伸展)为主,伴随多期岩浆活动,并诱发一系列金属成矿作用(杨明桂和王昆,1994;林黎等,2006a,b;张志辉等,2021)。双桥山群安乐林组(中新元古界)在矿区大面积出露(图1),为一套含变余结构的粉(细)砂岩、千枚岩和板岩的火山-碎屑岩沉积建造。控制区内岩浆岩带和矿体展布的断裂主要有两组,一组呈近E-W向展布,另外一组为NE-NNE向。区内岩浆活动主要为晋宁期(花岗闪长岩)和燕山期(似斑状黑云母花岗岩、细粒白云母花

岗岩、细粒二云母花岗岩、花岗斑岩)。晋宁期花岗闪长岩于九岭地区呈规模较大的岩基产出,主要出露于矿区西部(项新葵等,2015a,b)。燕山期中酸性花岗岩出露较少,多呈岩株、岩脉(岩枝)侵入于花岗闪长岩岩基和双桥山群地层中(项新葵等,2013b,c,2015b,2017)。

3 矿床地质特征

昆山钼钨铜多金属矿床大面积出露的地层为以泥砂质浅变质岩和火山碎屑岩为主的双桥山群(图1),仅少数地段被白垩系红层覆盖(林黎等,2006b;张明玉等,2016,2018a)。晋宁期黑云母花岗闪长岩作为九岭地区的花岗岩基,隐伏于矿床深部,仅于西部和东南部地区部分出露(叶少贞等,2016;张明玉,2016)。燕山期花岗斑岩在区内出露数量较多,矿床北部、中部、南部均有花岗斑岩脉分布。区内隐伏燕山期似斑状黑云母花岗岩(图2),似斑状黑云母花岗岩主要呈岩株产出,为本区的主要成矿岩体,未出露地表,经钻孔验证,其侵入高度最高至650 m海拔附近(图2)。昆山钼钨铜多金属矿床区断裂主要分为两类,一类为成矿前断裂,成群成组呈带状分布,并被(含矿)石英脉不同程度的充填,主要呈NEE、NE、NWW向展布;另一类为规模较小的成矿后断裂,以切断(含矿)石英脉为特征,延伸距离较为有限(汪国华等,2015;王西荣等,2016;张明玉等,2018b)。

区内围岩蚀变现象较为普遍,主要包括云英岩化、硅化和绿泥石化等,基本呈线性分布。其中,云英岩化仅于脉壁间或近矿围岩中的小(或微小)裂隙中较为发育;硅化遍布整个矿区,主要发育部位为矿体两侧($n\sim n\times 10$)cm范围内。本矿床矿石矿物较丰富,其中辉钼矿、黑钨矿、黄铜矿占主导,白钨矿和黄铁矿次之。矿石结构主要为自形一半自形、他形粒状和交代残余等结构;矿石构造包含细脉状、浸染状、鳞片状构造等(叶少贞等,2016)。

昆山矿床工业矿体类型主要为含Mo、Wu、Cu矿质的热液充填于先期断裂形成的石英脉型矿体,Mo和Wu共生、交替沉淀,呈近EW向分布(叶少贞等,2015),沿倾向(较陡)深部延伸较远(图2)。Mo-Wu矿体呈透镜状或薄层状,分支复合现象较为普遍,其产状与石英脉体走向(一般 $65^\circ\sim 80^\circ$)、倾

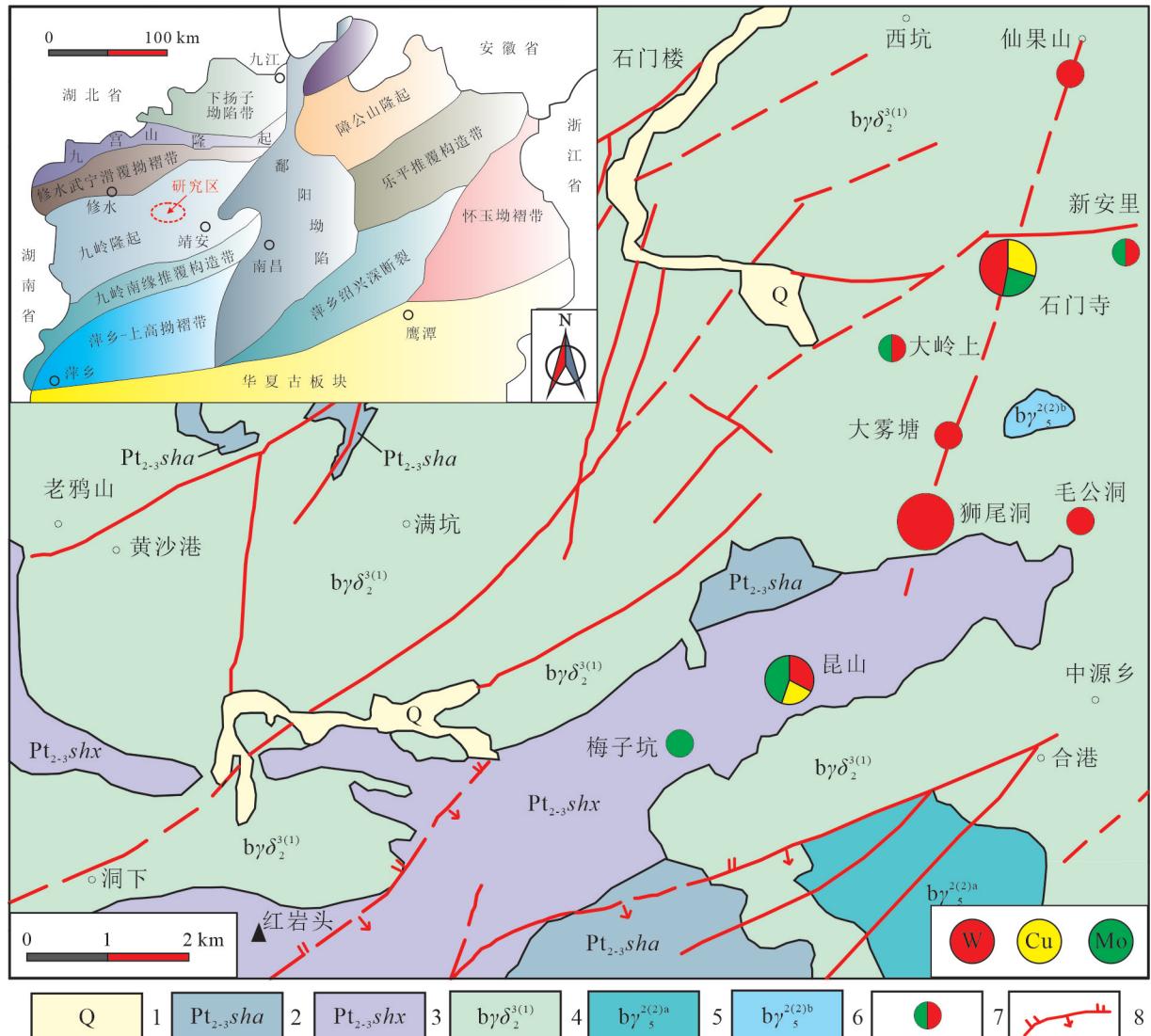


图1 江西九岭成矿带地质略图(据项新葵等 2012a, b修改)

1—第四系;2—安乐林组;3—修水组;4—黑云母花岗闪长岩(晋宁期);5—似斑状黑云母花岗岩(燕山期);6—细粒黑云母花岗岩(燕山期);7—矿床;8—断层

Fig.1 Regional geological sketch maps of Jiuling mineralization belt, Jiangxi (modified from Xiang Xinkui et al., 2012a, b)

1—Quaternary; 2—Anlelin Formation; 3—Xiushui Formation; 4—Biotite granodiorite (Jingningian); 5—Porphyritic biotitic granite (Yanshanian); 6—

Fine-grained biotitic granite (Yanshanian); 7—Ore deposit; 8—Fault

向(以南倾为主)相耦合(张明玉等, 2016)(图2)。钼矿体厚度一般3~26 m, 有的厚达60 m以上, Mo品位一般介于0.06%~0.32%, 平均品位0.11%; 钨矿体厚度1.4~6.3 m, 有的厚达44 m; WO₃品位0.12%~1.02%, 平均品位为0.30%(汪国华等, 2015)。

4 分析方法

4.1 样品分解和蒸馏分离 Os

首先将待测定辉钼矿样品进行准确称重,加

入至 Carius 管，然后配置乙醇和液氮黏稠体，用来保温盛好样品的 Carius 管（-50~ -80℃）。利用适量 HCl 溶液将 ¹⁹⁰Os 和 ¹⁸⁵Re 稀释剂导入 Carius 管底部位置，再加入 4~5 mL 浓度为 16 mol/L 的 HNO₃ 溶液，当 Carius 管底部冷冻冻结后，加热密封，置入不锈钢套管内保存，待其恢复室内温度后，便均匀加温至 200℃ 保温存放 24 h。将 Carius 管开口，在冰水浴中升温后，加入 1.5 倍样品种积的超纯水，把连接两根 Teflon 管的胶头固定于 Carius 管顶

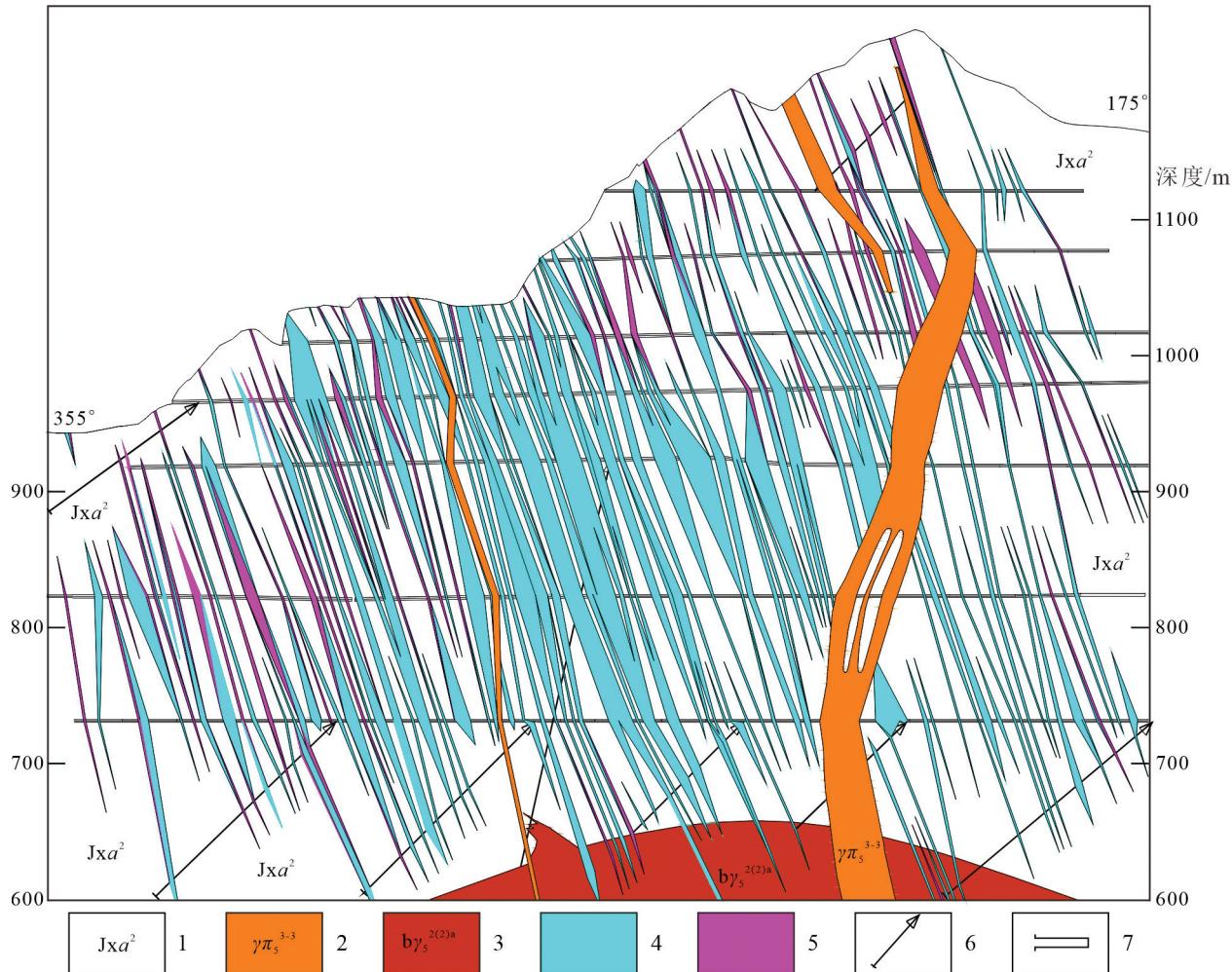


图2 昆山钼钨铜多金属矿床0线勘探SE—NW向剖面图(据叶少贞等, 2015, 2016修改)

1—双桥山群安乐林组中段;2—燕山期花岗斑岩;3—燕山期似斑状黑云母花岗岩;4—钼矿体;5—钨矿体;6—钻孔;7—平硐

Fig.2 Geological section along the NE-SW trending explosion line of Kunshan Mo-W-Cu polymetallic deposit (modified from Ye Shaozhen et al., 2015, 2016)

1—The middle member of Anlelin Formation; 2—Granite-porphyry of Yanshanian period; 3—Porphyritic biotitic granite of Yanshanian period; 4—Molybdenum orebody; 5—Tungsten orebody; 6—Drilling; 7—Adit

部位置, 将 Carius 管放置于电加热蒸笼中, 利用 100°C 恒温蒸汽持续加热 40~60 min。通过以上操作获得的水吸收液(吸附 OsO₄)便可以通过实验测试获得其 Os 同位素比值, 为方便下一步萃取分离 Re, 需将蒸馏获得的残液导入至 Teflon 烧杯中。详细步骤参见文献(Du et al., 2007; 屈文俊等, 2008; 李超等, 2009)。

4.2 萃取分离 Re 和质谱测定

将残液在电热板上烘至近干后, 再少量水烘至近干。碱化过程中需加入 4~10 mL、浓度为 10 mol/L 的 NaOH 溶液, NaOH 溶液用量视沉淀量多少来调节。为使样品达到碱性条件, 可对样品溶液进行微

加热, 再置入 Teflon 离心管中, 在振荡萃取 Re 之前, 先加入 4~10 mL 二甲基酮, 持续离心振荡约 1 min 后, 将上层二甲基酮转移至盛有少量水的 Teflon 烧杯中, 为促使样品溶液中二甲基酮挥发, 将其加热 50°C 附近, 持续 30~60 min, 并继续加热升温至 120°C 烘干。为分离其余的 Os, 再次烘干之前, 需滴入适量 30% 的 H₂O₂ 溶液和浓度较高的 HNO₃ 溶液。将烘干的残渣用适量硝酸溶解稀释后, 用于测定 Re 同位素比值的样品溶液便制备完成。将上述分离的 Os 和 Re 两组样品通过 X-series ICP-MS 测定。详细步骤参见文献(杜安道, 2012; 周利敏等, 2012; 王礼兵等, 2013)。

5 分析结果

为避免辉钼矿Re-Os同位素因失耦影响定年数据准确性,本次在辉钼矿采集过程中遵循“多取样,细磨碎”的原则(Stein et al., 2003; Selby and Creaser, 2004)。用于本次研究的10件样品均采自昆山钼钨铜多金属矿床PD920平硐内(采集间距10 m以上),采样时主要从辉钼矿较富集的地段采集,所采集的辉钼矿样品不同程度的伴生少量黄铁矿或黄铜矿,辉钼矿单矿物分选后每件样品重量都大于1 g,考虑到测试数据Re-Os等时线上分布过于集中而引起等时线年龄误差变大的情况,在PD920平硐采样过程中每个样品采集位置均距离较远,并分别于不同的石英脉矿体采集样品,以达到Re含量或Re/Os比值变化较大要求。实验测试由国家地质测试中心完成,通过Isoplot程序分别对10

组¹⁸⁷Re-¹⁸⁷Os年龄数据(表1)做等时线,并计算出年齡加权平均值(图3)。年齡数据介于(148.0±2.4)~(149.6±2.4) Ma,分布比较集中。在¹⁸⁷Re-¹⁸⁷Os等时线图上数据分布近似一条直线,且¹⁸⁷Re和¹⁸⁷Os沿等时线距离拉开的较远(图3),因此等时线年齡(148.8±1.0) Ma (MSWD=0.42)较为可靠。模式年齡加权平均值为(148.7±0.7) Ma (MSWD=0.10)。本次测得的辉钼矿¹⁸⁷Re-¹⁸⁷Os等时线年齡与加权平均年齡基本一致,考虑到统计学MSWD指标,选取加权平均年齡(148.7±0.7) Ma (MSWD=0.10)作为本区辉钼矿的形成年齡。

6 讨论

6.1 成矿时限厘定

昆山矿床精细成矿年齡报道较少(张明玉等,2016),为进一步针对本区成矿年齡进行有效精准

表1 昆山钼钨铜多金属矿床辉钼矿Re-Os同位素模式年齡及含量

Table 1 Re-Os isotopic model age and contents of molybdenite in the Kunshan Mo-W-Cu polymetallic deposit

样品编号	Re/10 ⁻⁶		普 Os/10 ⁻⁹		¹⁸⁷ Re/10 ⁻⁶		¹⁸⁷ Os/10 ⁻⁹		模式年齡/Ma	
	测定值	不确定度	测定值	不确定度	测定值	不确定度	测定值	不确定度	测定值	不确定度
KS6-1	10.26	0.080	0.091	0.038	6.448	0.053	16.06	0.100	149.3	2.1
KS6-3	25.72	0.250	0.019	0.019	16.16	0.150	40.21	0.300	149.1	2.3
KS6-7	10.12	0.080	0.018	0.120	6.360	0.048	15.73	0.120	148.3	2.1
KS6-8	5.794	0.042	0.018	0.022	3.642	0.026	9.021	0.082	148.5	2.3
KS6-9	7.971	0.058	0.018	0.027	5.010	0.037	12.40	0.090	148.3	2.1
KS6-10	8.022	0.058	0.018	0.027	5.042	0.036	12.46	0.120	148.2	2.3
KS6-11	5.059	0.041	0.018	0.025	3.179	0.026	7.847	0.077	148.0	2.4
KS6-13	43.51	0.470	0.010	0.022	27.35	0.290	68.23	0.420	149.6	2.4
KS6-15	2.947	0.019	0.059	0.017	1.852	0.012	4.603	0.033	149.0	2.1
KS6-16	3.307	0.024	0.000	0.012	2.078	0.015	5.165	0.037	149.0	2.1

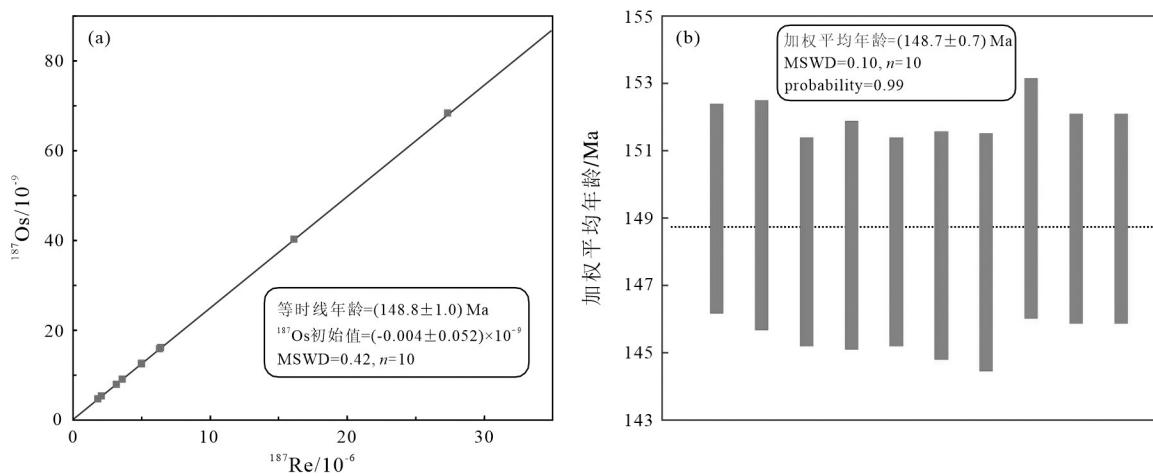


图3 昆山钼钨铜多金属矿床PD920平硐中辉钼矿Re-Os同位素等时线年龄(a)和加权平均年龄(b)

Fig.3 Re-Os isochrone (a) and the weighted average (b) ages of molybdenite in PD920 adit of Kunshan Mo-W-Cu polymetallic deposit

厘定。通过实验获得的10件辉钼矿等时线年龄为 (148.8 ± 1.0) Ma (MSWD=0.42); Re-Os加权平均年龄为 (148.7 ± 0.7) Ma (MSWD=0.10)。其加权平均年龄与等时线年龄基本一致,因此本次所采集的样品Re-Os体系是封闭的,在采集足量的辉钼矿样品以满足实验测试条件的基础上, ^{187}Re 和 ^{187}Os 沿等时线距离拉开的较远,表明本次定年数据较为可靠。由于本次用于定年的辉钼矿样品采自于辉钼矿、黑钨矿、黄铜矿共生的含矿石英脉,因此MSWD值较小的Re-Os加权平均年龄 (148.7 ± 0.7) Ma (MSWD=0.10),可以直接代表昆山钼钨铜多金属矿床的成矿年龄,该矿床为晚侏罗世岩浆期后热液活动的产物。

6.2 成岩成矿关系

总结认识规模较大的客观成岩成矿事件是以较为丰富的精细定年数据为基础的(丰成友等,2012; 黄兰椿和蒋少涌, 2012, 2013; 项新葵等, 2013a, 2015c; 张明玉, 2016; 张明玉等, 2016; 江丽, 2017)。项新葵等(2012a, b, 2013c, 2017)认为大湖塘钨多金属矿集区燕山期岩浆活动具有多期多次性,从似斑状黑云母花岗岩→细粒黑云母花岗岩→花岗斑岩(成矿期)→花岗斑岩(成矿后),岩石粒度粒径具有递减变细的特点,侵位顺序由早到晚(项新葵等, 2012a, b, 2013a, b, c, 2015a, b, c, 2017)。昆山矿床似斑状黑云母花岗岩、花岗斑岩的锆石单矿物U-Pb年龄分别为 (151.7 ± 1.3) Ma、 (136.6 ± 2.5) Ma(张明玉, 2016),侵入时间由早到晚,时空关系上,基本符合矿集区内岩浆侵位规律。昆山矿床燕山期至少存在两期成岩作用,第一期是152 Ma前后的晚侏罗世(以似斑状黑云母花岗岩为代表)(张明玉, 2016),略早于成矿(148.7 ± 0.7 Ma),暗示昆山矿床成矿作用发生稍滞后于似斑状黑云母花岗岩侵位时限,因此昆山钼钨铜多金属矿床为典型的岩浆期后热液矿床,与晚侏罗世岩浆活动有成因联系;第二期岩浆活动为137 Ma前后的早白垩世(以花岗斑岩为代表)(张明玉, 2016),滞后于成矿(148.7 ± 0.7 Ma)约12 Ma,表明第二期岩浆活动可能不具有成矿意义,经过野外地质调查发现,早白垩世花岗斑岩产出状态主要为岩枝或岩脉,与围岩的接触关系截然,未见明显热液矿化蚀变现象,切穿似斑状黑云母花岗岩和Mo-W-Cu工业矿体,对矿体形态

造成一定的破坏,为成矿后花岗斑岩。

6.3 成矿物质来源

Re-Os同位素组成不仅在约束成矿作用时限方面发挥重要作用,而且可以反演成矿物质初始来源的深度,进一步示踪不同深度混源的程度。Re元素化学习性表现为分散元素和不相容元素,在地壳中表现为强烈亏损,结晶为独立的自然矿物较为困难(涂光炽等, 2004)。壳源→壳幔混源→幔源,辉钼矿中Re相差一个数量级,即 $n\times10^{-6}\rightarrow n\times10^{-5}\rightarrow n\times10^{-4}$ 呈现递增趋势(Mao et al., 1999; Stein et al., 2001)。单一矿种的钼金属矿床Re一般波动较小,并且较多矿种的钼多金属矿床而言,Re表现更低(叶会寿等, 2006)。当然,有时在同一个矿床中,也存在辉钼矿Re波动范围较大的情况(代军治等, 2007)。辉钼矿中Re高低通常与成矿主岩不均一的源物质组成,经历部分熔融作用后,在岩浆期后热液成矿演化过程中,析出挥发份的能力有关(Selby and Creaser, 2001),期间,成矿主岩的差异性演化,通常会导致Cu(或W)多金属矿床中辉钼矿Re显著高于Mo多金属矿床对应的含量(Berzina et al., 2005)。

昆山钼钨铜多金属矿床中辉钼矿Re为 $2.95\times10^{-6}\sim43.51\times10^{-6}$,仅KS6-3和KS6-13两个样品Re稍高(分别为 25.72×10^{-6} 和 43.51×10^{-6}),其他8个样品Re较低(平均 6.69×10^{-6}),综合以上分析结果,认为昆山钼钨铜多金属矿床的成矿物质以壳源为主,考虑到存在较高Re两个样品($n\times10^{-5}$),说明可能存在部分幔源的成矿物质参与了该矿床的形成。这一结果与张明玉等(2018a)获得的Hf同位素认识一致,作为昆山矿床成矿主岩的似斑状黑云母花岗岩,其20个锆石原位Hf同位素测点 $\epsilon_{\text{Hf}}(t)$ 值介于-8.8~4.48,存在两个大于0(0.03和4.48)的测点,表明似斑状黑云母花岗岩以壳源为主,并存在少量幔源物质参与其形成。昆山钼钨铜多金属矿床含矿石英脉和黑钨矿氧同位素研究表明,12件石英 $\delta^{18}\text{O}_{\text{V}-\text{SMOW}}$ 值介于 $11.2\text{\textperthousand}\sim16.9\text{\textperthousand}$,平均为 $12.5\text{\textperthousand}$;黑钨矿 $\delta^{18}\text{O}_{\text{V}-\text{SMOW}}$ 值为 $6.2\text{\textperthousand}\sim8.1\text{\textperthousand}$,平均为 $7.1\text{\textperthousand}$,显示成矿流体为再平衡岩浆水,因此本区成矿流体应为成矿主岩母岩演化晚期分异而成,因此昆山矿床属于岩浆期后热液矿床,成矿物质来源相应继承了成矿主岩的源区特征,说明本区辉钼矿Re-Os同位素体系对昆山矿床成矿物质来源的示踪较为可靠。

表2 大湖塘钨铜多金属矿集区成矿定年结果汇总

Table 2 A summary of mineralization ages of Dahutang W–Cu–Mo polymetallic ore-concentrated area

矿床名称	测试对象	测试方法	成岩、成矿年龄/Ma	数据来源
石门寺	辉钼矿	Re-Os法	139.2±1.0	等时线年龄(Mao et al., 2013)
	辉钼矿	Re-Os法	143.7±1.2	等时线年龄(丰成友等, 2012)
	辉钼矿	Re-Os法	149.6±1.4	等时线年龄(项新葵等, 2013a)
狮尾洞	辉钼矿	Re-Os法	140.9±3.6	等时线年龄(丰成友等, 2012)
	辉钼矿	Re-Os法	148.7±0.7	等时线年龄(本文)
昆山	辉钼矿	Re-Os法	151.0±1.3	等时线年龄(张明玉, 2016)
	辉钼矿	Re-Os法	148.9±4.4	加权平均年龄(张志辉等, 2013)

6.4 地球动力学演化

昆山矿床所属的大湖塘钨铜多金属矿集区已报道了一批与成矿相关的年代学数据(表2),总结后推测本区燕山期成矿事件至少存在两期。其中第一期晚侏罗世成矿事件时限为151.0~148.7 Ma,昆山辉钼矿Re–Os等时线年龄(151.0±1.3) Ma(张明玉, 2016),梅子坑辉钼矿Re–Os加权平均年龄(150.0±1.0) Ma(张志辉等, 2013),石门寺辉钼矿Re–Os等时线年龄(149.6±1.2) Ma(项新葵等, 2013a),本次研究测得的昆山辉钼矿Re–Os加权平均年龄(148.7±0.7) Ma;第二期早白垩世成矿事件(140.9±3.6) Ma(丰成友等, 2012),石门寺辉钼矿

Re–Os等时线年龄时限约为143.7~139.2 Ma,石门寺、狮尾洞辉钼矿Re–Os等时线年龄分别为(143.7±1.2) Ma、(139.2±1.0) Ma(Mao et al., 2013)。项新葵等(2015c)将江南造山带中段的“挤压–伸展旋回”时限约束为146~135 Ma,因此本区第一期晚侏罗世成矿事件(151.0~148.7 Ma)形成于碰撞挤压(陆内造山)环境,为“挤压成矿”;第二期早白垩世成矿事件(143.7~139.2 Ma)发生于陆内挤压向伸展转换的背景,为“挤压至伸展过渡成矿”。大湖塘矿集区在伸展拉张转换之前便已有成矿作用发生,说明伸展拉张引发的地壳减薄环境对本区的成矿作用发生不是所必需的。结合上述区域成矿

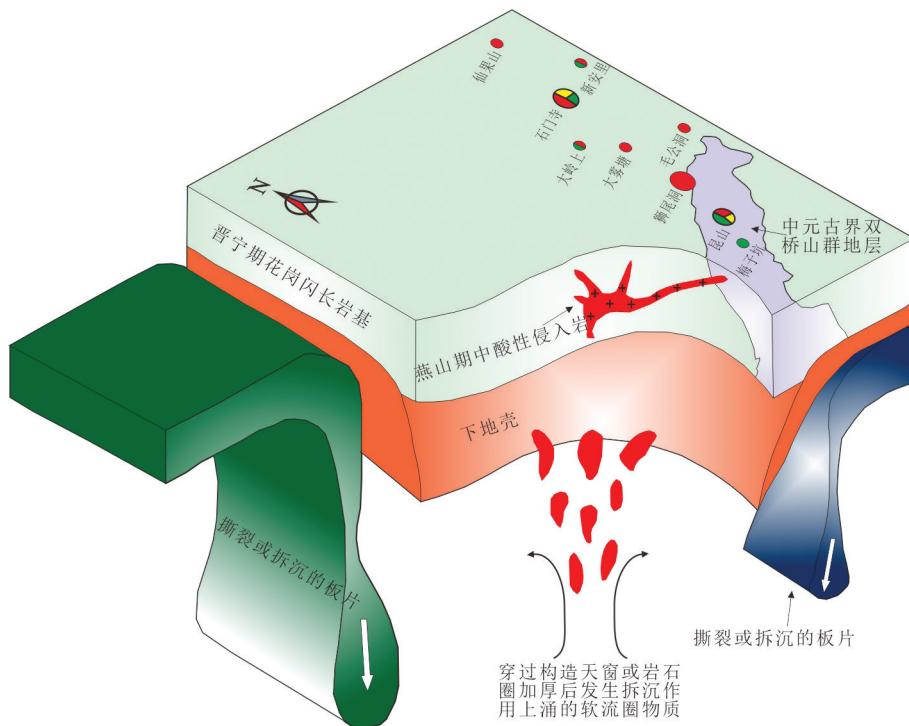


图4 大湖塘钨多金属矿集区成岩成矿地球动力学演化模式图

Fig.4 The geodynamical evolution pattern diagrams of ore-and rock-forming in W–Cu–Mo polymetallic ore-concentrated area

作用时限及构造旋回演化特征,本文推测热源供给途径有两种可能性,第一种诱发软流圈上涌途径,早侏罗世Izanagi板块(高浮力块体,如海山岛链、无震海岭等)俯冲至九岭地区时俯冲角度发生转换,诱发俯冲板片出现裂缝而撕裂,从而打开构造天窗(Mao et al., 2011);第二种诱发软流圈上涌途径,在陆内造山作用产生的碰撞挤压应力场控制下,引发岩石圈加厚而产生拆沉(项新葵等, 2015c)。通过以上途径之一导致在该地区形成一系列NNE向壳幔相互作用强烈的低压带,软流圈物质自晚侏罗世开始,沿着NNE向深断裂与先期(晋宁晚期)形成的NEE向古断裂的交汇部位上涌,通过热传导效应引起地壳物质发生部分熔融,于大湖塘地区形成了似斑状黑云母花岗岩(岩株)、细粒黑云母花岗岩(小岩株)、较早期花岗斑岩(岩脉或岩枝)等一系列中酸性侵入岩(图4)。当它们侵位于新元古界双桥山群地层时,则发生第一期成矿作用,即以石英脉型为主的Mo-W-Cu多金属矿床(如昆山矿床、梅子坑矿床,图4);当它们侵位于晋宁期黑云母花岗闪长岩时,第一期和第二期成矿作用均有发生,形成“一区三型”W-Cu-Mo多金属矿床(项新葵等, 2015a),即石英脉型、云英岩型(或细脉浸染型)和热液隐爆角砾岩型(如石门寺、新安里、大岭上等矿床,图4)。

7 结 论

(1)本次测得的辉钼矿Re-Os加权平均年龄(148.7 ± 0.7) Ma (MSWD=0.10),可以直接代表昆山钼钨铜多金属矿床的成矿年龄,该矿床为晚侏罗世岩浆期后热液活动的产物。

(2)昆山钼钨铜多金属矿床属于典型的岩浆期后热液矿床,成矿物质来源相应继承了成矿主岩的源区特征。成矿物质以壳源为主,可能存在部分幔源的成矿物质参与了该矿床的形成。

(3)大湖塘钨多金属矿集区燕山期成矿事件至少存在两期,第一期晚侏罗世成矿事件时限为151.0~148.7 Ma,第二期早白垩世成矿事件时限为143.7~139.2 Ma。

(4)当燕山期中酸性侵入岩侵位于新元古界双桥山群地层时,则发生第一期成矿作用,即以石英脉型为主的Mo-W-Cu多金属矿床;当侵位于晋宁

期黑云母花岗闪长岩时,第一期和第二期成矿作用均有发生,形成石英脉型、云英岩型和热液隐爆角砾岩型W-Cu-Mo多金属矿床。

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