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## 北山造山带风雷山地区二长花岗岩LA-ICP-MS锆石U-Pb年龄及其构造背景

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**提要:**风雷山二长花岗岩位于北山北部的红石山—百合山—蓬勃山与芨芨台子—小黄山两条蛇绿混杂岩带之间,构造位置特殊。二长花岗岩LA-ICP-MS锆石U-Pb年龄加权平均值为( $320\pm1$ )Ma,代表了其成岩年龄。岩石具有富SiO<sub>2</sub>、Al<sub>2</sub>O<sub>3</sub>的特征,里特曼指数( $\sigma$ )为1.58~2.74(<3.3),碱度率( $AR$ )为2.35~4.05,属于钙碱性花岗岩类,A/NCK=0.92~1.15,显示出偏铝质、过铝质岩浆的特征。稀土元素总量( $\Sigma$ REE)较高,稀土配分模式呈轻稀土富集的右倾型[(La/Yb)<sub>n</sub>=4.60~8.83],轻稀土分异程度大于重稀土,所有样品均具有Eu负异常( $\delta$ Eu=-0.56~0.88,平均为0.72);强烈富集强不相容元素Th、U和大离子亲石元素Rb、K,亏损高场强元素Nb、Ta、Ti、P。曲线形态与上地壳的相似,表明岩浆作用过程有上地壳物质的参与。高的Th含量(8.58~21.62 μg/g)及低的Nb/Ta值(8.75~15.72)指示岩浆源区主要为地壳。在Rb-(Nb+Y)及 $R_1-R_2$ 等构造判别图解中样品大多数落入板块碰撞前火山弧花岗岩区,根据岩石共生组合特征,认为风雷山地区二长花岗岩为碰撞前陆缘弧岩浆作用的产物,与红石山—百合山—蓬勃山有限洋的洋壳俯冲作用有关。结合区域地质背景,推测红石山—百合山—蓬勃山有限大洋的闭合时间应晚于晚石炭世早期。

**关 键 字:**北山造山带;锆石U-Pb年龄;地球化学;活动大陆边缘弧;地质调查工程;风雷山;内蒙古

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## LA-ICP-MS zircon U-Pb dating and tectonic setting of the monzogranites in the Fengleishan area of Beishan orogenic belt, Inner Mongolia

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**Abstract:** Monzogranites in Fengleishan are located in northern Beishan and situated between two ophiolitic mélange belts of

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Hongshishan–Baiheshan–Pengboshan and Jijitaizi–Xiaohuangshan, having a special tectonic affinity. Zircons from the monzogranites yielded a weighted mean LA-ICP-MS U-Pb age of  $(320\pm1)$  Ma, representing the intrusion age. The samples are characterized by high content of  $\text{SiO}_2$  and  $\text{Al}_2\text{O}_3$  with Rittmann indexes of 1.58–2.74 ( $<3.3$ ) and alkaline ratios of 2.35–4.05, belonging to calc-alkaline granitoids. Their A/NCK values range from 0.92 to 1.15 and display metaaluminous to peraluminous affinities. They have relatively high content of total REE and are characterized by LREE-enriched patterns ( $(\text{La/Yb})_{\text{N}}=4.60–8.83$ ) and negative Eu anomalies ( $\delta\text{Eu}=0.56–0.88$ , 0.72 in average). They also show strong enrichment of incompatible elements (e.g., Th and U) and large ion lithophile elements (e.g., Rb and K) but depletion of high field strength elements (e.g., Nb, Ta, Ti and P). Their normalized patterns are similar to those of upper continental crust (UPC), indicating participation of materials from the UPC in their magma source. High values of Th (8.58–21.62  $\mu\text{g/g}$ ) and low ratios of Nb/Ta (8.75–15.72) point to a crust-dominated magma source. On the tectonic discriminant diagrams of  $\text{Rb}/(\text{Nb}+\text{Y})$  and  $R_1-R_2$ , most of the samples are plotted in the pre-collisional volcanic arc granite area. Based on associated lithological characteristics, it is proposed that monzogranites in the Fengleishan were formed by pre-collisional volcanic arc magmatism, which was related to oceanic crust subduction in the Hongshishan–Baiheshan–Pengboshan belt. Combined with regional geological background, it is inferred that the closure of the small ocean in the Hongshishan–Baiheshan–Pengboshan belt should postdate the early period of late Carboniferous.

**Key words:** Beishan orogenic belt; zircon U–Pb age; geochemistry; active continental margin; geological survey engineering; Fengleishan; Inner Mongolia

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## 1 引言

北山造山带作为中亚造山带的重要组成部分,位于西伯利亚板块、塔里木板块和华北板块交汇处的关键位置(图1a),属中亚造山带的中南缘,总体属古亚洲构造域的一部分,是近年来研究热点和焦点之一(李锦轶等,2006a,b,c; Windley et al., 2007; 肖文交等,2006,2008; Xiao et al., 2009,2010; 李舢,2009; 李小菲,2013)。北山造山带发育复杂而又独特的中亚型构造,大地构造格局具有盆山耦合、多块体多缝合带镶嵌的特征(肖序常等,1991,1992; 王国强,2015),经历了多次“开”、“合”的构造演化历史(王玉往等,1997)。由于其地质演化过程漫长,地质构造极其复杂,对北山古生代洋陆转化时限及古生代构造格局的探讨是该区研究的热点。左国朝等(左国朝等,1990a,2003; 杜玉良等,2009)认为该区古生代石板井—小黄山洋盆开启于早寒武世中期或更早的震旦纪,最终闭合于志留纪末—泥盆纪,晚古生代全区总体处于板内的开合构造活动时期,将明水—石板井—小黄山蛇绿混杂岩带作为北山地区旱山微板块与塔里木板块东端北缘的缝合带(左国朝等,1990a; 李锦春等,1996)、哈萨克

斯坦板块与塔里木板块的缝合带(左国朝等,2003; 郑荣国等,2012);何世平等认为红柳河—牛圈子—洗肠井(图1b)洋盆开启于寒武纪—早奥陶纪而闭合于志留纪末,红石山—百合山—蓬勃山有限大洋开启于早石炭世而结束于石炭纪末,认为晚古生代红石山—百合山—蓬勃山大洋双向俯冲,为哈萨克斯坦板块与塔里木板块最终的板块缝合带(何世平等,2002,2004,2005; 龚全胜等,2003);而另一种观点则认为洋盆闭合于二叠纪(聂凤军等,2002a; 江思宏等,2003; 苗来成等,2014)。刘雪亚等将红石山—黑鹰山—六驼山深大断裂作为晚古生代西伯利亚板块与哈萨克斯坦板块的分界,而将柳园—大奇山深大断裂作为晚古生代哈萨克斯坦板块与塔里木板块的分界(张新虎,1993; 刘雪亚等,1995; 聂凤军等,2002b,2004; 江思宏,2006);杨合群等以红柳河—牛圈子—洗肠井蛇绿混杂岩带代表早古生代最终板块缝合带,将北山地区北带划归哈萨克斯坦板块、南带划归塔里木板块,晚古生代为陆内裂谷演化时期(任秉琛等,2001; 杨合群等,2008,2009,2010; 徐学义等,2008);而谢春林等则认为代表古亚洲洋南缘消减带的实际位置应在雀儿山—英安山一线以北的蒙古境内,而北山岛弧带实属南侧东

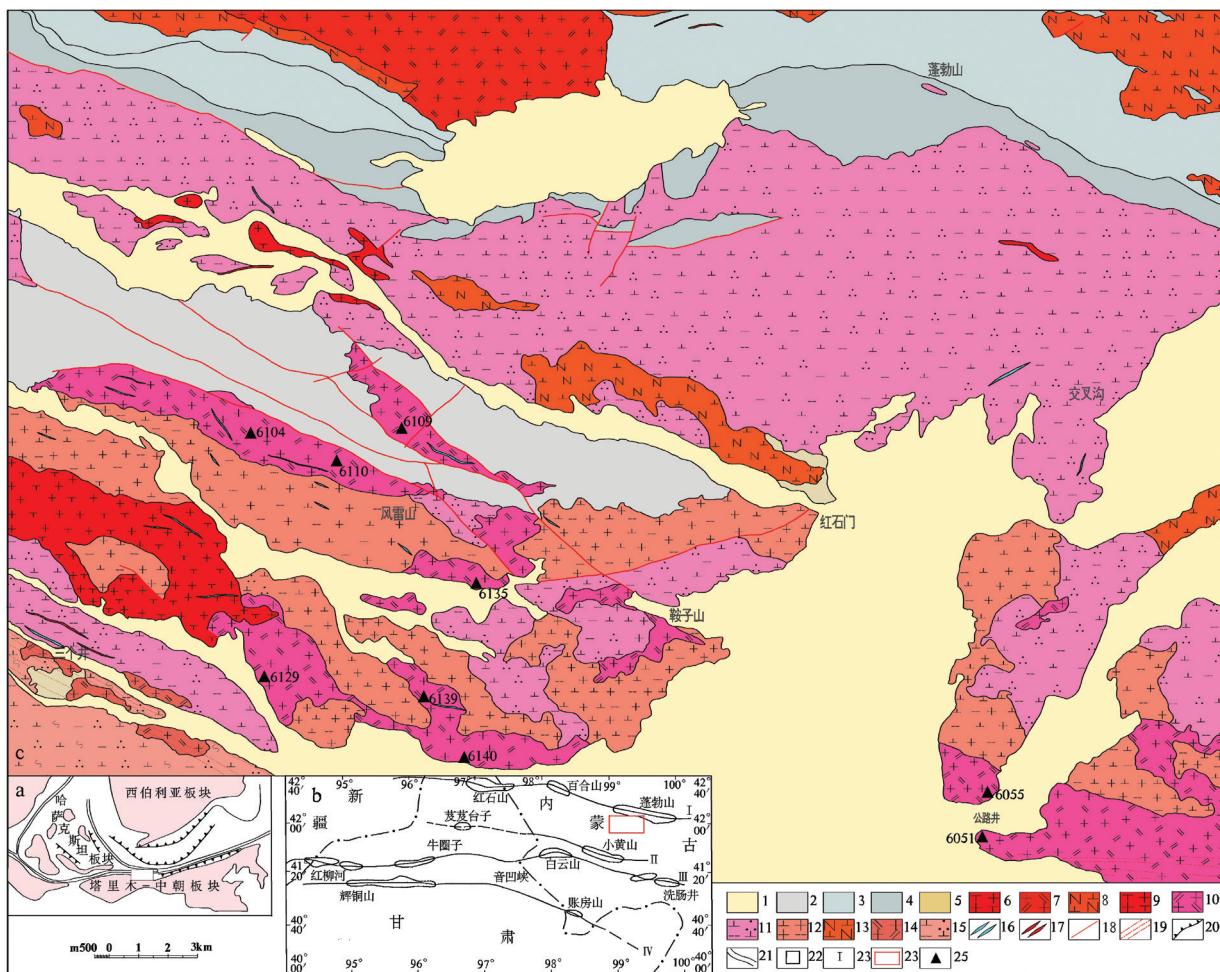


图1 风雷山地区地质简图(a据王国强,2015;b据杨合群等,2010)

1—中新生界;2—上石炭统火山岩;3—绿条山组二段;4—绿条山组一段;5—北山岩群;6—二叠纪正长花岗岩;7—早二叠世似斑状二长花岗岩;8—早二叠世英云闪长岩;9—晚石炭世正长花岗岩;10—晚石炭世二长花岗岩;11—晚石炭世石英闪长岩;12—早石炭世黑云花岗闪长岩;13—晚泥盆世英云闪长岩;14—早泥盆世二长花岗岩;15—早志留世黑云石英闪长岩;16—闪长玢岩/闪长岩脉;17—二长花岗岩/花岗岩脉;18—断层;19—糜棱岩化带;20—板块俯冲带;21—板块缝合带;22—北山造山带位置;23—蛇绿岩带;24—研究区;25—采样点位

Fig. 1 Geological sketch map of the Fengleishan area (a after Wang Guoqiang, 2015; b after Yang Hequn et al., 2010)

1—Mesozoic-Cenozoic; 2—Upper Carboniferous volcanic rock; 3—The second member of Lütiaoshan Formation; 4—The first member of Lütiaoshan Formation; 5—Beishan Group complex; 6—Permian syenogranite; 7—Early Permian porphyritic monzogranite; 8—Early Permian tonalite; 9—Late Carboniferous syenogranite; 10—Late Carboniferous monzogranite; 11—Late Carboniferous quartz diorite; 12—Early Carboniferous biotite granodiorite; 13—Late Devonian tonalite; 14—Early Devonian monzogranite; 15—Early Silurian biotite quartz diorite; 16—Diorite porphyrite vein and diorite vein; 17—Monzogranite vein and granite vein; 18—Fault; 19—Mylonitization zone; 20—Subduction zone; 21—Suture; 22—Beishan orogenic belt location; 23—Ophiolite; 24—Research area; 25—Sampling location

天山古陆陆缘增生地体的一部分(谢春林等,2009),古亚洲洋的闭合与增生造山时间发生在中泥盆世与早石炭世之间(卢进才等,2013)。最新的一种观点认为古生代时期整个北山地区属于古亚洲洋中的弧增生系统的组成部分(Xiao et al., 2010),划分为北部的北山弧盆系、南部的敦煌陆块北缘(潘桂棠等,2009)。显然,这些观点有着明显的差异,这也意味着北山地区古生代构造演化及构

造单元划分均存在明显的争议。

古生代是北山地区不同洋盆和陆块形成演化的重要时期,伴随着弧陆拼贴碰撞和大陆增生,该区古生代期间发生了多期次岩浆作用,其中以花岗岩类分布尤为广泛。北山地区花岗岩类总出露面积约30000 km<sup>2</sup>,占整个北山地区面积的30%,其侵位时间可以从前寒武纪一直延续到中生代,其中以华力西期,尤其是华力西晚期的花岗岩类分布最为

广泛(甘肃省地质矿产局,1989;内蒙古自治区地质矿产局,1991;江思宏等,2006;聂凤军等,2006;杜玉良等,2009;)。左国朝等(1990b)将北山地区发育的花岗岩类自北向南划分为3个带:北带指明水—石板井一线以北、中蒙边界以南的区域,花岗岩类约占该区总面积的28%;中带为白云山—东七一山一带,岩性主要是英云闪长岩及花岗闪长岩;南带指南起疏勒河,北达马鬃山—横岱山隆起带以南的狭长地带,该带内花岗岩类分布广泛,构成北山地区两大花岗岩带之一。近年来,在公婆泉一带、雀儿山一带及狼娃山—黑鹰山一带取得的花岗岩锆石U-Pb年龄、Ar-Ar、K-Ar和Rb-Sr年龄将古生代花岗岩类大致划分为两个阶段:早—中古生代、晚古生代(李小菲,2013;李舢,2013)。但仍需要精确的锆石U-Pb定年资料支撑及构造背景研究。本文通过对风雷山地区二长花岗岩的岩石学、锆石U-Pb年代学及地球化学等系统研究,分析其成因,探讨其形成的构造背景,以期为北山北部晚古生代区域大地构造演化提供新的约束。

## 2 区域地质概况

风雷山地区地处红石山—百合山—蓬勃山与芨芨台子—小黄山两条蛇绿混杂岩带之间,旱山地块中北部(左国朝等,2003;杨剑洲,2019)。研究区内构造以近EW向或NNW向的一系列逆断层为主,后期叠加了NE向的次一级断裂。区内前寒武纪变质基底主要由经受了中—低级变质作用的古元古界北山岩群组成(图1c),其构造样式复杂,经历了多期变质、变形作用的改造,主要表现为近水平顺层剪切作用下形成的片理、石香肠构造等,记录了北山地区早期地壳形成演化的历史(内蒙古自治区地质矿产局,1996;甘肃省地质矿产局,1997)。古生代时期,研究区所在北山地区在前寒武纪统一古陆基础上,先后经历了两期板块构造体制下的“开”(裂解)“合”(俯冲—碰撞造山)作用(何世平等,2002,2005),然而地层记录在本区保留甚少,其中早古生代地层在区内未见出露,而晚古生代地层也仅限于分布在本区北部的下石炭统绿条山组碎屑岩及上石炭统白山组中酸性火山岩。中新生代以来,风雷山地区主要以断陷作用为主,形成了白垩纪至新近纪的沉积盆地,孕育了赤金堡

组、苦泉组及第四系沉积。区内岩浆活动强烈,侵入岩分布广,侵入期次多,时限由志留纪岩浆侵入开始至二叠纪岩浆侵入结束,出露岩性主要有石英闪长岩、英云闪长岩、花岗闪长岩、二长花岗岩及正长花岗岩等。风雷山二长花岗岩呈北西西向展布,与区域构造线走向一致,呈岩株状产出,侵入早石炭世花岗岩及上石炭统白山组火山岩,被中新生代地层覆盖。

## 3 岩石学特征

风雷山二长花岗岩岩体出露于风雷山北、风雷山南及公路井等地,总体呈北西西向带状展布,侵入早石炭世花岗闪长岩、上石炭统白山组、晚石炭世石英闪长岩等,被晚石炭世正长花岗岩侵入,局部可见闪长质暗色包体。岩石具中细粒花岗结构,块状构造(图2)。岩石中斜长石多呈1~3.5 mm半自形板状—半自形粒状,普遍绢云母化;钾长石多数为1~4.5 mm半自形粒状,为微斜长石、微斜条纹长石;石英呈0.5~3.5 mm他形粒状;黑云母呈0.3~1 mm鳞片状,多数绿泥石化。斜长石(绢云母化)含量25%~35%,钾长石为25%~45%,石英为20%~27%,黑云母小于5%,局部含少量角闪石,磁铁矿少量,磷灰石微量。

## 4 样品分析测试

本次在二长花岗岩岩体的不同部位共采集了9件新鲜岩石样品进行岩石地球化学分析(图1c),选取其中的1件样品(TW6104)进行锆石U-Pb测年。主量元素、微量及稀土元素的测定均由河北省区域地质矿产调查研究所实验室完成。主量元素采用X射线荧光光谱仪(Axiosmax)测定,微量及稀土元素采用等离子体质谱仪(X-serise2)测定,元素含量精度优于5%。

LA-ICP-MS锆石U-Pb测年样品的锆石挑选由河北省区域地质矿产调查研究所实验室完成。经机械粉碎至80~100目,用浮选、电磁选方法进行分离,然后在双目镜下挑选出晶形和透明度较好的锆石颗粒用于年龄测定。锆石制靶及阴极发光(CL)图像拍照在北京锆年领航科技有限公司进行,将完整的锆石单矿物用无色透明的环氧树脂固定,待环氧树脂充分固结后抛磨,使锆石内部充分暴露,然后进行锆石的反射光、投射光及阴极发光

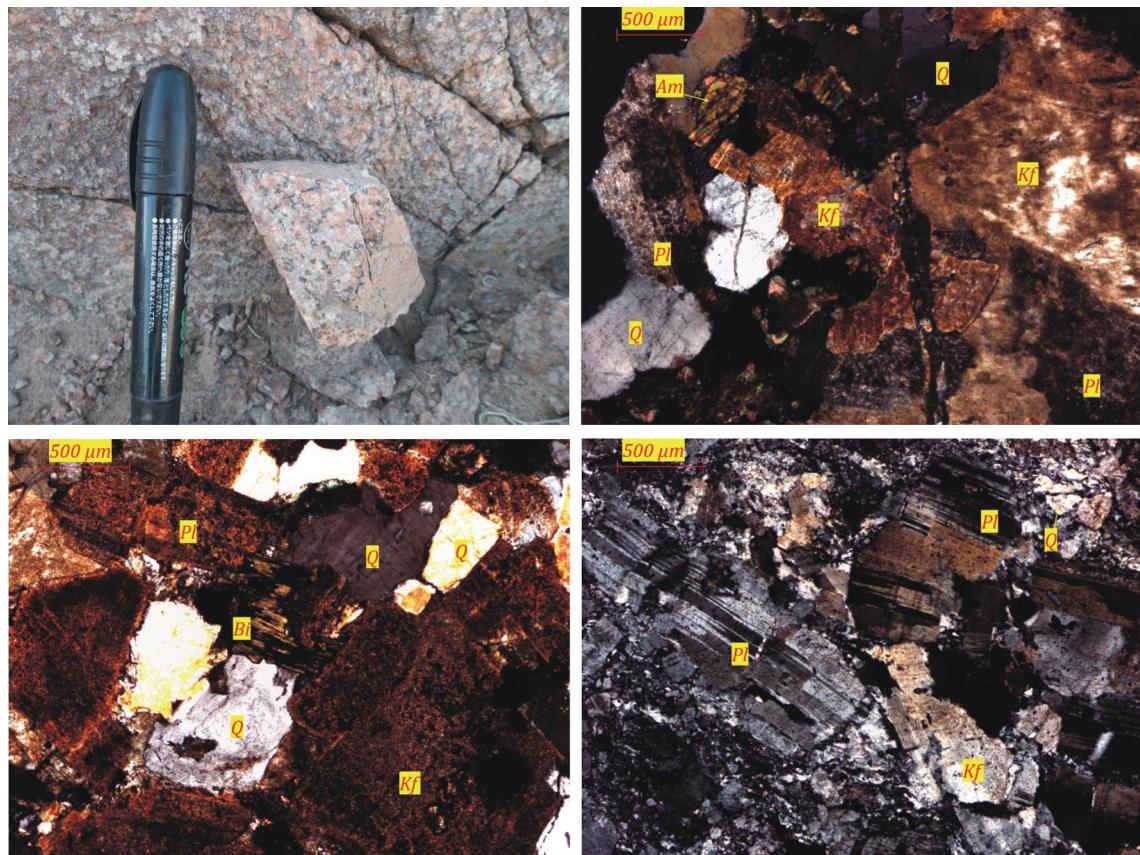


图2 北山风雷山二长花岗岩野外露头及显微(正交偏光)照片  
Kf—钾长石;Pl—斜长石;Q—石英;Bi—黑云母;Am—角闪石

Fig.2 Photographs of monzogranites in Fengleishan, Beishan (crossed nicols)  
Kf—K-feldspar; Pl—Plagioklase; Q—Quartz; Bi—Biotite; Am—Amphibole

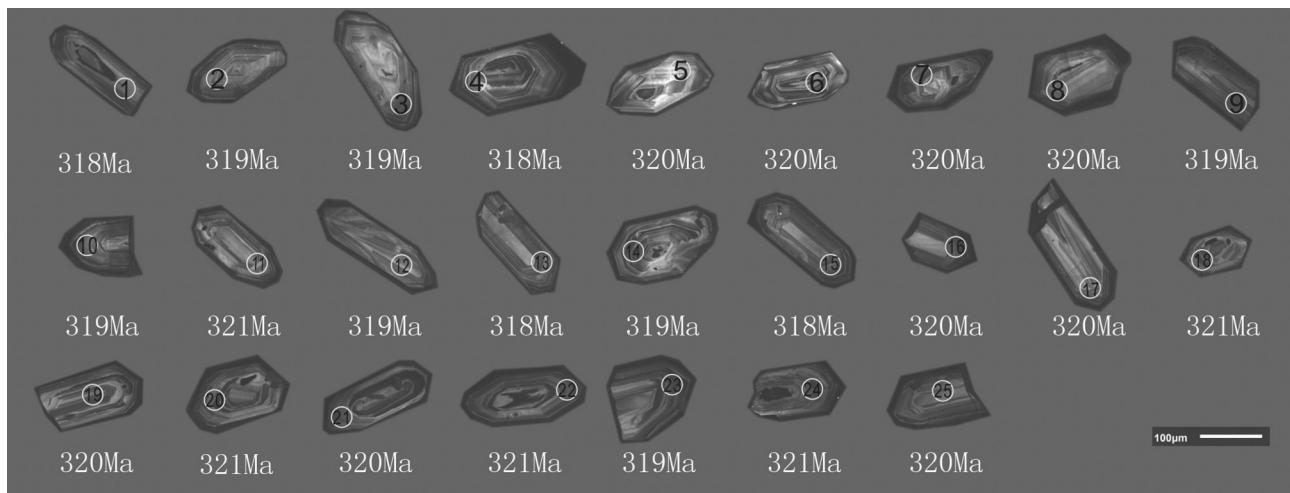


图3 北山风雷山二长花岗岩(TW6104)中锆石的阴极发光图  
Fig.3 CL images of representative zircons of monzogranites (TW6104) in Fengleishan, Beishan

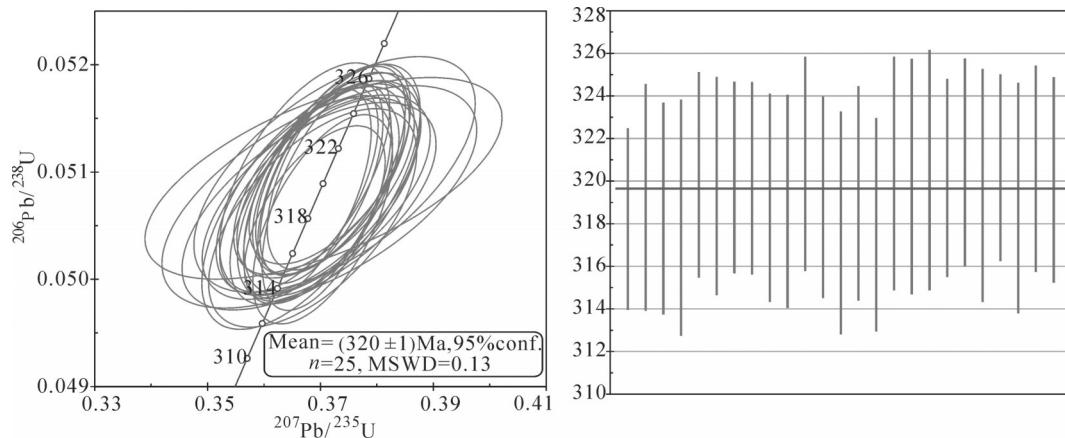


图4 北山风雷山二长花岗岩锆石U-Pb同位素谐和图  
Fig.4 U-Pb isotopic concordia diagram for zircons of monzogranites in Fengleishan, Beishan

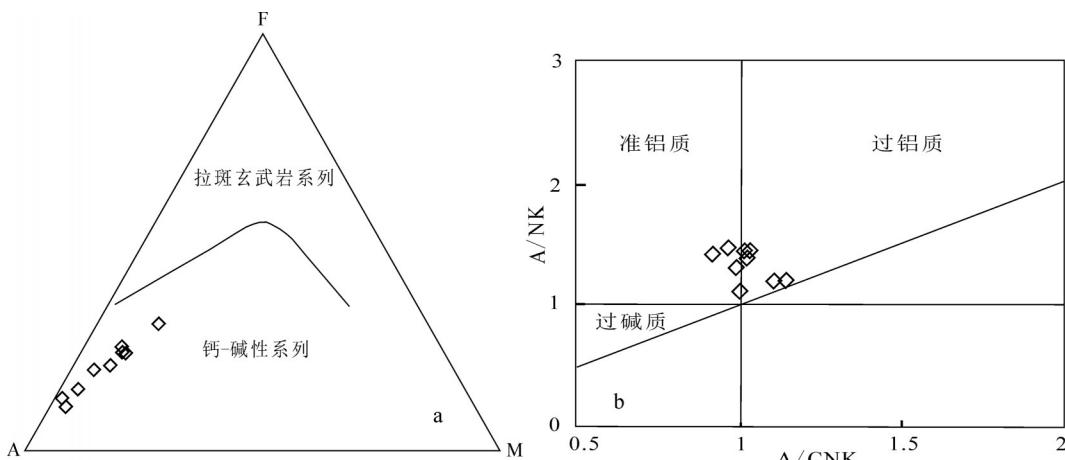


图5 AFM图解和A/CNK-A/NK图解(据Irvine and Barage, 1971; Peccerillo and Taylor, 1976; Maniar and Piccoli, 1989)  
Fig. 5 AMF and A/CNK-A/NK diagrams (after Irvine and Barage, 1971; Peccerillo and Taylor, 1976; Maniar and Piccoli, 1989)

(CL)照相。LA-ICP-MS锆石微区原位单点定年在天津地质调查中心实验室完成,使用仪器为Neptune多接收电感耦合等离子体质谱仪和193 nm激光取样系统(LA-MC-ICP-MS)。激光束斑直径为35 μm,能量密度为13~14 J/cm<sup>2</sup>,频率为8~10 Hz。实验室采用He作为激光剥蚀物质的载气,利用动态变焦扩大色散使质量数相差很大的U-Pb同位素可以同时接收从而进行U-Pb同位素测定,锆石标样采用TEMORA标准锆石。数据处理采用中国地质大学Liu et al.(2008)编写的ICPMsDataCal程序,加权平均年龄计算及谐和图的绘制采用ISPLT(ver3.0)程序(Ludwig, 2001)。采用<sup>208</sup>Pb对普通铅进行校正,利用NIST 612玻璃标样作为外标计算锆石样品的Pb、U、Th含量。具体测试方法与

详细分析步骤见文献(李怀坤等,2009),采用<sup>206</sup>Pb/<sup>238</sup>U年龄的加权平均值作为岩体的结晶年龄,同位素比值和年龄误差1σ,可信度95%。

## 5 同位素年代学

本次工作的二长花岗岩采自风雷山北,地理坐标:东经99°05'59",北纬41°57'45"。对其锆石进行了25个点的测试,分析结果见表1。所测锆石呈无色,长柱状,锆石粒径多为150~200 μm,长宽比多为2:1~3:1。在阴极发光CL图像(图3)上,显示出锆石内部具典型的明暗相间的振荡环带结构,表明其属于岩浆结晶的产物(吴元保等,2004)。25个点的测试结果显示锆石的Th/U含量分别为78×10<sup>-6</sup>~432×10<sup>-6</sup>(平均值207×10<sup>-6</sup>)和107×10<sup>-6</sup>~458×10<sup>-6</sup>(平

表1 北山风雷山二长花岗岩LA-ICP-MS锆石U-Pb同位素测年数据  
Table 1 LA-ICP-MS zircon U-Pb isotopic data of monzogranites in Fengleishan, Beishan

测点	含量/ $10^{-6}$				Th/U	同位素比值						表面年龄/Ma					
	Pb	Th	U	$^{206}\text{Pb}/^{238}\text{U}$		$1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb}$	$1\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb}$	$1\sigma$	
1	22	432	353	1.22	0.0506	0.0003	0.3690	0.0051	0.0529	0.0008	318	2	319	4	324	33	
2	7	78	120	0.65	0.0508	0.0004	0.3698	0.0114	0.0528	0.0015	319	3	320	10	322	66	
3	9	113	176	0.64	0.0507	0.0004	0.3680	0.0092	0.0527	0.0013	319	2	318	8	314	55	
4	11	128	203	0.63	0.0506	0.0004	0.3659	0.0069	0.0524	0.0010	318	3	317	6	304	41	
5	6	87	107	0.81	0.0509	0.0004	0.3705	0.0129	0.0528	0.0019	320	2	320	11	318	80	
6	13	159	238	0.67	0.0509	0.0004	0.3697	0.0074	0.0527	0.0010	320	3	319	6	317	44	
7	13	165	249	0.66	0.0509	0.0004	0.3698	0.0062	0.0527	0.0009	320	2	320	5	315	37	
8	25	301	458	0.66	0.0509	0.0004	0.3698	0.0043	0.0527	0.0006	320	2	319	4	315	26	
9	10	150	168	0.89	0.0508	0.0004	0.3693	0.0083	0.0528	0.0011	319	2	319	7	318	49	
10	13	153	238	0.64	0.0507	0.0004	0.3690	0.0061	0.0527	0.0008	319	3	319	5	318	36	
11	15	225	272	0.83	0.0510	0.0004	0.3723	0.0056	0.0529	0.0009	321	3	321	5	325	38	
12	13	214	224	0.95	0.0508	0.0004	0.3697	0.0072	0.0528	0.0010	319	2	319	6	321	44	
13	17	221	288	0.77	0.0506	0.0004	0.3686	0.0075	0.0529	0.0009	318	3	319	7	323	37	
14	20	302	350	0.86	0.0508	0.0004	0.3677	0.0046	0.0525	0.0007	319	3	318	4	307	29	
15	14	163	243	0.67	0.0506	0.0004	0.3704	0.0051	0.0531	0.0007	318	3	320	4	335	31	
16	14	136	263	0.52	0.0510	0.0004	0.3672	0.0062	0.0523	0.0009	320	3	318	5	297	38	
17	18	260	315	0.82	0.0509	0.0004	0.3721	0.0064	0.0530	0.0008	320	3	321	6	328	35	
18	15	164	264	0.62	0.0510	0.0004	0.3725	0.0102	0.0530	0.0012	321	3	322	9	329	53	
19	14	204	233	0.87	0.0509	0.0004	0.3709	0.0064	0.0528	0.0009	320	2	320	6	322	40	
20	21	308	363	0.85	0.0510	0.0004	0.3728	0.0050	0.0530	0.0006	321	2	322	4	328	26	
21	14	165	262	0.63	0.0509	0.0004	0.3715	0.0066	0.0530	0.0009	320	3	321	6	328	37	
22	16	207	287	0.72	0.0510	0.0003	0.3716	0.0058	0.0528	0.0008	321	2	321	5	322	35	
23	22	267	404	0.66	0.0508	0.0004	0.3711	0.0047	0.0530	0.0006	319	3	320	4	330	27	
24	18	257	306	0.84	0.0510	0.0004	0.3719	0.0057	0.0529	0.0007	321	2	321	5	325	31	
25	22	308	377	0.82	0.0509	0.0004	0.3700	0.0045	0.0527	0.0006	320	2	320	4	317	26	

均值 $270\times10^{-6}$ ), Th/U为0.52~1.22, 均大于0.5, 也说明了锆石属于典型岩浆成因(李志昌等, 2004)。在锆石U-Pb年龄 $^{206}\text{Pb}/^{238}\text{U}-^{207}\text{Pb}/^{235}\text{U}$ 谐和图(图4)中, 分析数据均分布在谐和线上,  $^{206}\text{Pb}/^{238}\text{U}$ 加权平均年龄为( $320\pm1$ ) Ma(MSWD=0.13), 属于晚石炭世早期, 代表了该岩体的侵位年龄。

## 6 地球化学特征

### 6.1 主量元素

岩石SiO<sub>2</sub>含量为65.90%~75.98%, 属于中酸性岩、酸性岩类。Al<sub>2</sub>O<sub>3</sub>含量较高, 为12.50%~14.85%, 具有富硅、铝的特征(表2)。TiO<sub>2</sub>含量较低, 为0.21%~0.61%。Na<sub>2</sub>O+K<sub>2</sub>O含量高, 为6.75%~9.05%, K<sub>2</sub>O/Na<sub>2</sub>O值为0.65%~1.58, 体现出岩石中钾长石、斜长石含量的变化。岩石里特曼指数( $\sigma$ )为1.58~2.74(<3.3), 碱度率( $AR$ )为2.35~4.05, 属于钙

碱性花岗岩类, 在SiO<sub>2</sub>-(Na<sub>2</sub>O+K<sub>2</sub>O)图解(图略)中样品落入亚碱性系列区, 在AFM图解(图5a)中样品呈现钙-碱性系列岩石的演化趋势, 在SiO<sub>2</sub>-K<sub>2</sub>O图解中(图略)大部分样品落入高钾钙碱性系列区域。A/NCK=0.92~1.15(>1), A/NK=1.13~1.49(>1), 显示出偏铝质、过铝质岩浆的特征, 在A/NCK-A/NK图(图5b)中样品也分布于准铝质和过铝质区。分异指数( $DI$ )为76.44~94.69, 固结指数( $SI$ )为2.18~13.02, 反映岩浆分异程度较强; 长英指数( $FL$ )为70.25~95.44, 镁铁指数( $MF$ )为70.89~85.90, 揭示岩浆分离结晶作用较强。指数变化范围较大, 体现出岩浆演化的不均一性。

### 6.2 稀土元素与微量元素

从稀土分析结果表(表2)中可以看出二长花岗岩稀土总量( $\Sigma\text{REE}$ )较高(平均为 $126.35\times10^{-6}$ ), 变化范围较大( $86.29\times10^{-6}$ ~ $158.50\times10^{-6}$ ), 轻稀土总量

表2 北山风雷山二长花岗岩主量元素(%)和微量元素( $10^{-6}$ )组成Table 2 Major (%) and trace elements ( $10^{-6}$ ) content of monzogranites in Fengleishan, Beishan

样号	6051	6055	6104	6109	6110	6135	6139	6140	6129
SiO <sub>2</sub>	72.93	75.87	71.90	71.78	71.46	65.90	71.63	75.98	73.56
Al <sub>2</sub> O <sub>3</sub>	14.18	12.50	14.02	13.19	14.15	14.85	13.83	12.99	13.10
TiO <sub>2</sub>	0.24	0.27	0.34	0.38	0.37	0.61	0.45	0.21	0.31
Fe <sub>2</sub> O <sub>3</sub>	0.82	1.07	1.22	1.55	1.44	2.45	2.08	0.60	1.53
FeO	0.60	0.50	0.98	0.98	1.07	1.67	0.68	0.46	0.60
CaO	0.87	0.56	2.12	2.86	2.24	2.99	2.35	0.39	1.87
MgO	0.23	0.40	0.83	0.99	0.90	1.69	0.87	0.34	0.50
K <sub>2</sub> O	4.12	4.82	3.47	3.43	2.64	3.23	3.36	4.85	4.39
Na <sub>2</sub> O	4.93	3.05	3.71	3.32	4.09	3.95	3.48	3.23	3.10
MnO	0.010	0.017	0.101	0.068	0.095	0.094	0.051	0.012	0.036
P <sub>2</sub> O <sub>5</sub>	0.044	0.044	0.101	0.068	0.124	0.145	0.113	0.048	0.076
H <sub>2</sub> O <sup>+</sup>	0.49	0.75	0.74	0.89	1.00	1.50	0.73	0.57	0.63
灼失量	0.80	0.78	1.07	1.26	1.29	2.29	0.99	0.78	0.81
总和	99.77	99.88	99.87	99.87	99.87	99.86	99.88	99.88	99.87
Na <sub>2</sub> O+K <sub>2</sub> O	9.05	7.88	7.18	6.75	6.73	7.17	6.84	8.08	7.48
K <sub>2</sub> O/Na <sub>2</sub> O	0.84	1.58	0.94	1.03	0.65	0.82	0.97	1.50	1.42
A/CNK	1.00	1.11	1.02	0.92	1.04	0.96	1.01	1.15	0.99
$\sigma$	2.74	1.89	1.79	1.58	1.59	2.25	1.64	1.98	1.83
AR	4.01	4.04	2.61	2.45	2.39	2.35	2.47	4.05	3.00
DI	93.39	93.16	84.28	82.02	83.00	76.44	82.77	94.25	87.16
SI	2.18	4.01	8.14	9.63	8.92	13.02	8.29	3.58	4.93
FL	91.21	93.36	77.24	70.25	75.00	70.61	74.48	95.44	79.99
MF	85.90	79.92	72.61	71.93	73.50	70.89	76.06	75.8	81.03
Cs	0.41	9.80	2.15	1.46	1.39	1.86	5.37	6.17	1.90
Rb	31.4	181.9	116.4	152.9	72.4	99.8	121.9	190.1	140.2
Ba	1506.0	536.3	650.9	701.6	625.6	629.0	680.8	721.3	638.1
Th	10.54	18.05	12.98	22.13	8.85	8.58	10.4	21.11	21.62
U	1.52	1.04	1.14	3.31	1.47	1.98	1.70	1.49	2.29
K	34210	40033	28829	28445	21895	26781	27912	40274	36405
Nb	14.15	11.41	9.77	11.76	8.90	8.35	9.04	11.37	14.78
Ta	0.90	1.08	0.89	1.32	0.65	0.58	0.77	1.30	1.63
Sr	117.3	86.6	232.1	302.6	281.5	253.1	221.8	120.0	135.5
P	192	192	441	298	543	634	494	208	333
Zr	311.3	147.2	158.6	136.0	156.8	194.3	180.2	132.0	175.1
Hf	10.15	5.75	5.91	5.00	6.06	7.22	6.66	4.24	5.44
Ti	1426	1596	2016	2249	2218	3636	2717	1247	1885
Y	21.90	12.29	24.04	32.13	17.67	25.83	20.97	17.40	32.43
La	34.44	20.63	26.77	29.56	26.06	22.82	22.13	28.41	23.55
Ce	65.02	34.22	51.71	57.19	51.31	47.51	47.43	46.69	68.77
Pr	7.97	4.41	6.40	7.08	5.77	5.76	5.38	5.57	5.56
Nd	28.75	15.23	24.39	25.98	20.82	22.84	20.26	18.85	21.3
Sm	5.15	2.52	4.60	5.18	3.81	4.73	3.91	3.33	4.63
Eu	0.99	0.50	1.13	1.03	1.04	1.26	1.04	0.57	0.94
Gd	4.53	2.21	4.06	5.09	3.40	4.39	3.67	2.93	4.56
Tb	0.72	0.34	0.67	0.85	0.50	0.72	0.56	0.44	0.75
Dy	3.93	1.99	3.87	5.03	2.81	4.29	3.37	2.63	5.15
Ho	0.79	0.42	0.79	1.03	0.58	0.87	0.70	0.53	1.06
Er	2.37	1.35	2.52	3.15	1.77	2.66	2.13	1.83	3.16
Tm	0.41	0.24	0.45	0.58	0.32	0.45	0.37	0.34	0.57
Yb	2.89	1.89	2.89	3.82	2.12	2.91	2.53	2.46	3.67
Lu	0.54	0.33	0.46	0.62	0.34	0.44	0.35	0.43	0.67
$\Sigma$ REE	158.50	86.29	130.70	146.20	120.70	121.70	113.80	115.00	144.30
LREE	146.90	79.73	119.10	131.10	112.20	109.30	103.80	106.40	129.30
HREE	11.65	6.56	11.64	15.07	8.45	12.35	10.01	8.67	15.03
LREE/HREE	12.60	12.15	10.23	8.70	13.29	8.85	10.37	12.27	8.61
$\delta$ Eu	0.63	0.65	0.80	0.61	0.88	0.85	0.84	0.56	0.63
(La/Yb) <sub>N</sub>	8.55	7.85	6.65	5.55	8.83	5.62	6.26	8.28	4.60
(La/Sm) <sub>N</sub>	4.32	5.29	3.76	3.69	4.41	3.12	3.65	5.51	3.28
(Gd/Yb) <sub>N</sub>	1.30	0.97	1.16	1.10	1.33	1.25	1.20	0.99	1.03

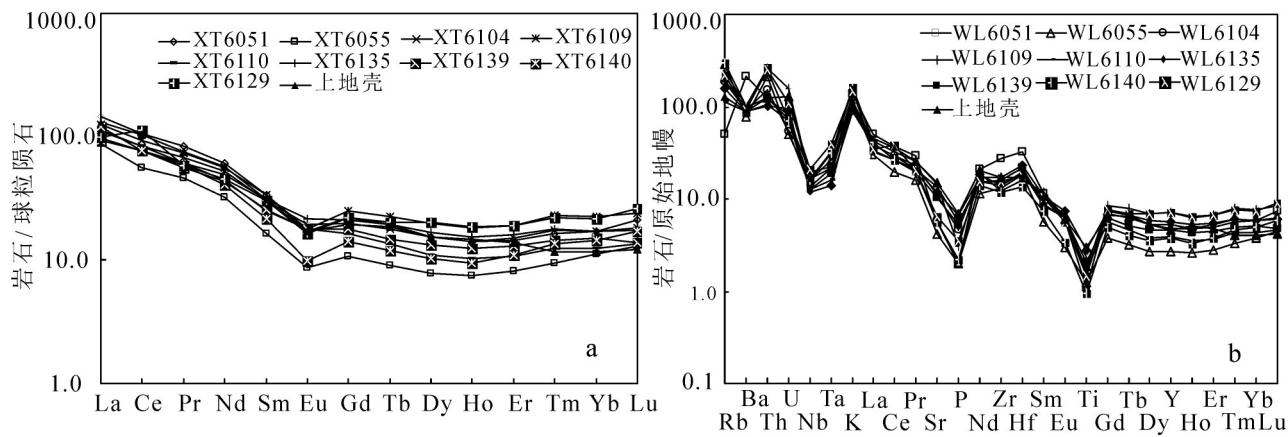


图6 北山风雷山二长花岗岩的稀土元素球粒陨石标准化配分型式(a)和微量元素原始地幔标准化蛛网图(b)  
(据Sun and McDonough, 1989)

Fig.6 Chondrite-normalized REE patterns (a) and primitive mantle normalized trace elements spider diagram (b) of monzogranites in Fengleishan, Beishan (after Sun and McDonough, 1989)

(LREE平均为 $115.31 \times 10^{-6}$ )明显高于重稀土(HREE平均为 $11.05 \times 10^{-6}$ )。在球粒陨石标准化稀土配分图中(图6a),曲线呈右倾型,LREE/HREE变化于8.61~13.29,平均为10.79,(La/Yb)<sub>N</sub>变化于4.60~8.83,平均为6.91,说明轻重稀土分馏较为明显。(La/Sm)<sub>N</sub>变化于3.12~5.51,平均为4.11;(Gd/Yb)<sub>N</sub>变化于0.97~1.33,平均为1.15,表现出轻稀土分馏程度大于重稀土。所有样品的Eu负异常明显( $\delta\text{Eu}$ 为0.56~0.88,平均为0.72),表明可能发生了斜长石的分离结晶作用(Hugh, 2000)。曲线形态基本平行于上地壳曲线,表明源岩物质有上地壳物质。

从微量元素分析结果(表2)及微量元素原始地幔标准化的蛛网图(图6b)可以看出:岩石强烈富集了强不相容元素Th、U和大离子亲石元素Rb、K,相对亏损高场强元素Nb、Ta、P、Ti,呈现Rb、Th峰和Nb、Ta谷,可能为地幔物质部分熔融形成,但Nb、Ta、Ti的亏损,表明了二长花岗岩源岩物质与地壳密切相关(李再会等,2012)。曲线形态与上地壳的相似,也表明岩浆作用过程有上地壳物质的参与。曲线形态与Pearce et al.(1984)建立的用于花岗岩构造环境判别的微量元素图解中的火山弧型钙碱性系列相似,暗示岩浆活动与火山弧环境有关(Hugh, 2000)。

## 7 讨 论

### 7.1 岩浆侵位时代、成因及构造环境

本文对风雷山地区二长岩花岗岩锆石U-Pb定

年获得 $^{206}\text{Pb}/^{238}\text{U}$ 加权平均年龄为 $(320 \pm 1)\text{Ma}$ ,可以与同区域同时期形成的交叉沟石英闪长岩( $306.3 \pm 1.2\text{Ma}$ )(赵志雄等,2015)进行对比,同为红石山—百合山—蓬勃山有限洋南缘狼娃山—黑鹰山岩浆带华力西中期侵入岩的一员。因此,本文所测年龄数据可信,代表了岩浆结晶侵位年龄,岩体形成时代属于晚石炭世。

二长花岗岩矿物组合为斜长石、钾长石、石英及黑云母等,角闪石局部可见,有少量闪长质暗色包体发育。岩石地球化学特征显示出富 $\text{SiO}_2$ 、 $\text{Al}_2\text{O}_3$ ,富碱的特征, $\text{Rb}/\text{Sr}$ 比值高(0.26~2.10),平均为0.73,说明其分异程度较高,在 $(\text{Na}_2\text{O}+\text{K}_2\text{O})/\text{CaO}-(\text{Zr}+\text{Nb}+\text{Ce}+\text{Y})$ 图(图7a)中4个样品落入分异花岗岩区,5个样品落入非分异花岗岩区。 $\text{A/CNK}=0.92 \sim 1.15$ ,显示出偏铝质、过铝质高钾钙碱性岩浆的特征。对于I型、S型的划分,早期研究提出的铝饱和指数 $\text{A/NCK}=1.1$ 作为两者的分界值,但这一指标仅适用于未经强烈结晶分异的花岗岩。 $\text{P}_2\text{O}_5$ 在弱过铝质和强过铝质岩浆中随 $\text{SiO}_2$ 增加变化趋势不同,这种性质被成功地用于区分I型和S型花岗岩类。在 $\text{SiO}_2-\text{P}_2\text{O}_5$ 图解(图7b)上,数据点总体沿I型演化趋势分布,即 $\text{P}_2\text{O}_5$ 与 $\text{SiO}_2$ 含量呈负相关关系(Chappell et al., 1992)。所有样品中均未见S型花岗岩所特有的富铝特征矿物白云母、堇青石、石榴子石及电气石等,同样说明不具备S型花岗岩的特征,应为I型花岗岩。

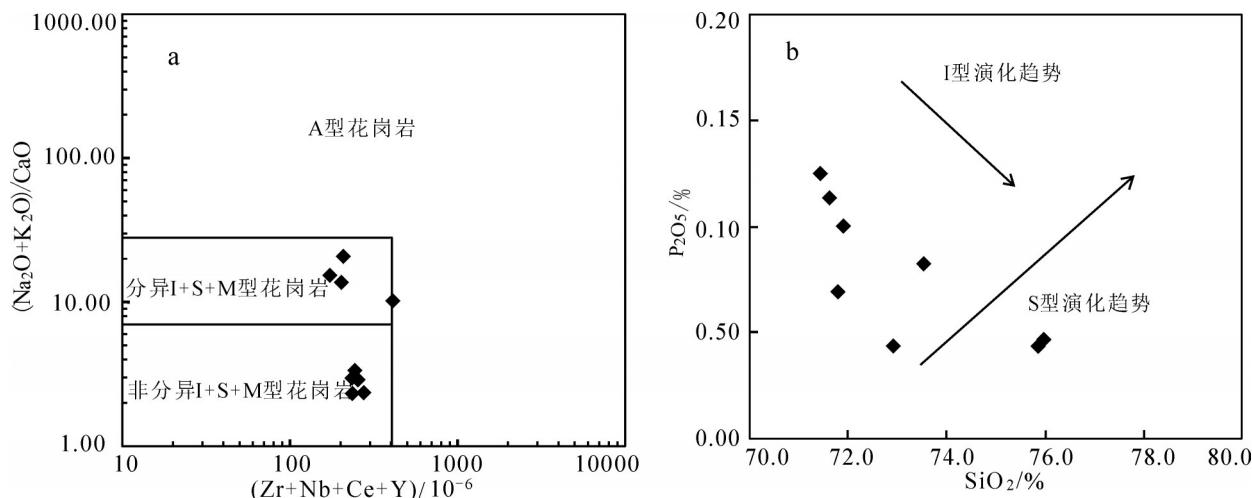


图7  $(\text{Na}_2\text{O}+\text{K}_2\text{O})/\text{CaO}-(\text{Zr}+\text{Nb}+\text{Ce})+\text{Y}$  图(据 Whalen et al., 1987) 及  $\text{SiO}_2-\text{P}_2\text{O}_5$  图解(据 Chappell and White, 1992)  
Fig.7  $(\text{Na}_2\text{O}+\text{K}_2\text{O})/\text{CaO}-(\text{Zr}+\text{Nb}+\text{Ce})+\text{Y}$  (after Whalen et al., 1987) and  $\text{SiO}_2-\text{P}_2\text{O}_5$  diagrams (after Chappell and White, 1992)

地幔中 Th 的丰度仅为  $0.05 \mu\text{g/g}$  (Sun, 1980), 而地壳(尤其是花岗岩)中的 Th 含量高达  $16\sim21 \mu\text{g/g}$  (Pitcher et al., 1985), 风雷山地区二长花岗岩的 Th 含量为  $8.58\sim21.62 \mu\text{g/g}$ , 平均为  $14.92 \mu\text{g/g}$ , 远高于地幔丰度均值而更接近于地壳。李昌年(1992)认为 Nb 的负异常能反映花岗岩具有大陆壳的特征, 风雷山地区二长花岗岩的  $\text{Nb}/\text{Ta}$  值为  $8.75\sim15.72$ , 均值为  $11.54$ , 远低于地幔平均值  $60$ , 而更接近地壳平均值  $11$  (Green et al., 1987)。且二长花岗岩的稀土元素球粒陨石标准化配分型式和微量元素原始地幔蛛丝网曲线形态与上地壳的相似, 也表明岩浆作用过程有上地壳物质的参与, 综合表明岩浆源区主要为地壳(丁坤等, 2020)。

构造运动控制着岩浆活动, 而岩浆活动是构造运动的一种表现形式, 不同的岩石组合、岩石成分及特征微量元素组合特点对构造环境有着明显的指示作用(赵振华, 2007)。风雷山地区二长花岗岩明显富集了强不相容元素 Th、U 和大离子亲石元素 Rb、K, 而强烈亏损高场强元素 Nb、Ta、Ti、P, 具有明显的 Eu 负异常, 曲线均呈现右倾型, 具有火山弧花岗岩的特征(吕达鑫等, 2018; 杨延乾等, 2018)。而二长花岗岩在  $\text{Rb}-\text{Nb}+\text{Y}$  图解(图 8a)中也落入火山弧花岗岩区, 在  $R_1-R_2$  图解(图 8b)中主要落入板块碰撞前区域(Hugh, 2000; 张旗等, 2008; 邓晋福等, 2015b)。有学者认为 I 型花岗岩类与岛弧和活动大陆边缘有关(王德滋等, 1999), 李昌年认为大陆弧

背景下造山花岗岩均具有 P、Ti 等元素的亏损, 而 Nb 的负异常更能反映花岗岩具有大陆壳的特征(李昌年, 1992; 刘明强等, 2018)。风雷山地区与二长花岗岩在时空上共生的岩浆岩多为钙碱性花岗岩及低镁的钙碱性火山岩(贾元琴等, 2016), 花岗岩类具有 TTG 岩类的特征, 说明该地区晚古生代存在洋壳俯冲环境(冯艳芳等, 2011), 而这一岩石共生组合特征也符合活动大陆边缘的岩石组合类型(路凤香等, 2002; 邓晋福等, 2007), 且为活动大陆边缘弧靠内陆一侧的火成岩弧的内带的标志性火成岩组合, 说明了风雷山地区晚古生代花岗岩的构造环境为活动大陆边缘弧环境(邓晋福等, 2007, 2015a; 袁四化等, 2009)。根据上述特征, 结合区域构造演化, 认为风雷山地区二长花岗岩为碰撞前陆缘弧岩浆作用的产物, 与红石山—百合山—蓬勃山有限洋的洋壳俯冲作用有关。

## 7.2 构造意义

古生代时期, 北山地区在前寒武纪统一古陆基础上, 先后经历了两期板块构造体制下的“开”(裂解)“合”(俯冲—碰撞造山)作用(何世平等, 2002, 2005), 其中: 第一期板块构造体制下的开合主要出现在早古生代早期, 北山地区沿红柳河—牛圈子—洗肠井一带裂解形成洋盆, 晚奥陶世—志留纪发生由南向北俯冲, 志留纪末大洋封闭, 泥盆纪以碰撞造山作用为主(左国朝等, 2003)。第二期开合旋回出现在晚古生代中期, 这一时期北山地区经过泥盆

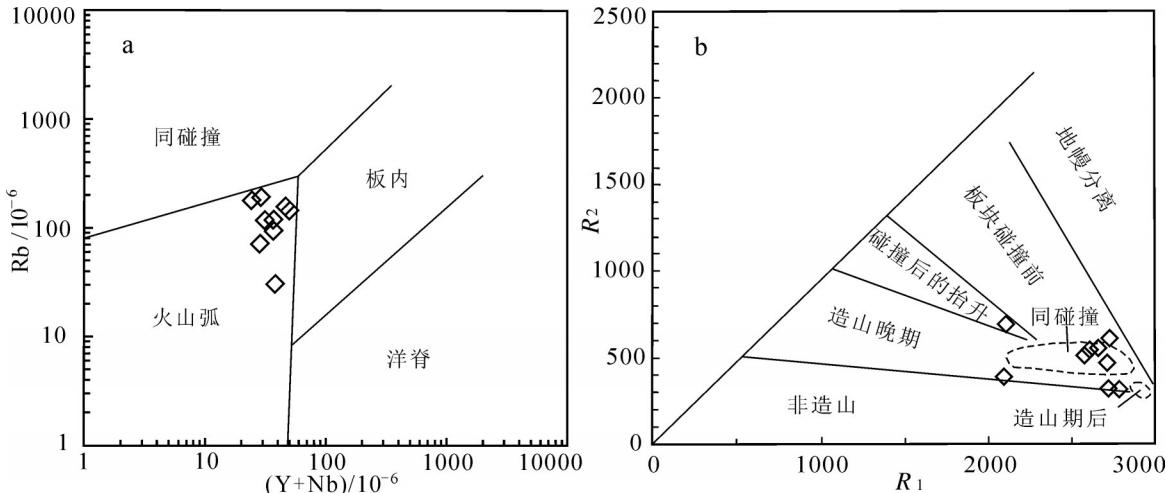


图8 北山风雷山二长花岗岩的构造环境判别图解(据Pearce et al., 1984; Batchelor and Bowden, 1985)

Fig. 8 Tectonic setting discrimination diagrams of monzogranites in Fengleishan, Beishan (after Pearce et al., 1984; Batchelor and Bowden, 1985)

纪碰撞造山作用后重新形成的统一大陆沿研究区以北的红石山—百合山—蓬勃山一带发生裂解,再度形成洋盆,可能由于扩张时间较短,洋盆扩张宽度不大,属“有限大洋”(何世平等,2005)或“红海型”海槽,即裂谷强烈拉张向初始小洋盆过渡(杨合群等,2010)。之后,洋壳在不断扩张的同时向南、向北发生双向俯冲,进而闭合(何世平等,2002,2005;牛文超等,2019)。在此期间出现了强烈的中酸性火山活动,形成了以白山组( $C_2b$ )为代表的喷出岩,同时伴随有华力西中期大量中酸性岩浆岩的侵入(郝增元,2018)。风雷山地区二长花岗岩即为该期侵入岩的重要组成单元,其构造背景为碰撞前陆缘弧,证实了风雷山地区该时期的构造环境,侵位时代为( $320\pm1$ )Ma,有效地限制了红石山—百合山—蓬勃山有限洋盆的闭合时限,即该有限洋盆的消解碰撞时间晚于( $320\pm1$ )Ma。这一认识为北山北部地区晚古生代的构造演化研究提供了新的资料。

## 8 结 论

基于对北山风雷山地区二长花岗岩的岩石学、岩石地球化学及锆石U-Pb年代学特征的系统研究,初步得出如下结论:

(1) 锆石LA-ICP-MS U-Pb年代学表明,风雷山地区二长花岗岩锆石 $^{206}\text{Pb}/^{238}\text{U}$ 的年龄加权平均值为( $320\pm1$ )Ma,代表岩体的成岩年龄,属于晚石炭世

早期。

(2) 主量元素富 $\text{SiO}_2$ 、 $\text{Al}_2\text{O}_3$ 及碱含量较高,稀土元素配分模式呈现轻稀土富集的右倾海鸥型,具有较为明显的负Eu异常,强烈富集大离子亲石元素Rb、Th、U、K,相对亏损高场强元素Nb、Ta、Ti、P,加之矿物组成,共同表明风雷山二长花岗岩为偏铝质、过铝质高钾钙碱性I型花岗岩。岩浆源区主要为地壳。

(3) 北山风雷山地区二长花岗岩形成于活动大陆边缘构造环境,与洋壳的俯冲、消减有关,推测红石山—百合山—蓬勃山有限大洋的闭合时间应晚于晚石炭世早期。

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