

doi: 10.12029/gc20210313

周传芳,杨华本,蔡艳龙,张元厚,姜丽莉,魏小勇,段明新,冯嘉,王博超,赵喜东. 2021. 漠河盆地西缘漠河组形成时代及物源区构造环境判别[J]. 中国地质, 48(3): 832–853.

Zhou Chuanfang, Yang Huaben, Cai Yanlong, Zhang Yuanhou, Jiang Lili, Wei Xiaoyong, Duan Mingxin, Feng Jia, Wang Bochao, Zhao Xidong. 2021. Stratigraphic age of the Mohe Formation in the western margin of Mohe Basin and tectonic environment discrimination of its provenance[J]. Geology in China, 48(3): 832–853(in Chinese with English abstract).

漠河盆地西缘漠河组形成时代及物源区构造环境判别

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提要:大兴安岭北部漠河盆地广泛发育漠河组,其形成时代、物源区特征及大地构造背景等对研究盆地的形成与演化具有重要意义。本文在漠河盆地西缘恩和哈达地区漠河组内发现了2层凝灰岩夹层。采用LA-ICP-MS U-Pb测年方法进行了锆石U-Pb年代学研究,分别获得(158 ± 1)Ma、(157 ± 2)Ma的年龄,表明漠河组形成于晚侏罗世早期。砂岩岩石地球化学分析结果显示,漠河组砂岩 SiO_2 含量高(70.20%~76.13%), $\text{SiO}_2/\text{Al}_2\text{O}_3$ 值为4.99~5.87, $\text{K}_2\text{O}/\text{Na}_2\text{O}$ 值为0.59~1.62;在稀土元素球粒陨石标准化图解上,曲线与活动大陆边缘的稀土元素曲线吻合较好;样品富集高场强元素(Th,Hf)和大离子亲石元素(Rb,K),亏损高场强元素(Nb,Ta,Ti,P)和大离子亲石元素(Ba,Sr),微量元素蛛网图曲线与活动大陆边缘相似。地球化学物源分析图解和大地构造背景判别表明:漠河组砂岩的母岩主要来源于石英质沉积岩和酸性火山岩的混合区。砂岩碎屑锆石LA-ICP-MS U-Pb测年结果显示,碎屑锆石年龄与盆地南缘前寒武系基底和各期花岗岩类岩石的时代一致。综合研究认为,漠河组沉积物主要来自盆地南缘额尔古纳、兴安等地块,物源区既有古老陆壳物质,也有年轻的火成岩,大地构造背景以活动大陆边缘为主。

关 键 词:漠河组;形成时代;物源;构造环境判别;地质调查工程;漠河盆地;黑龙江

中图分类号:P597.3;P595;P588.21 文献标志码:A 文章编号:1000-3657(2021)03-0832-22

Stratigraphic age of the Mohe Formation in the western margin of Mohe Basin and tectonic environment discrimination of its provenance

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Abstract: The Mohe Formation is widely developed in Mohe Basin which is located in the northern part of the Da Hinggan

收稿日期:2019-10-23;改回日期:2020-03-23

基金项目:中国地质调查局项目(DD2016007803、DD20191014)资助。

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Mountains. Its formation age, provenance characteristics and tectonic background are of great geological significance for the study of the formation and evolution of the basin. Two interbeds of tuff were discovered in the Mohe Formation in the western margin of Mohe basin, Enhe–Hada area. Zircon U–Pb dating by LA–ICP–MS U–Pb method yields the weighted average ages of (158 ± 1) Ma and (157 ± 2) Ma respectively, indicating that the Mohe Formation rocks were formed in early Late Jurassic. The geochemical analysis of Mohe Formation sandstone shows that the content of SiO_2 is high (70.20%–76.13%), the value of $\text{SiO}_2/\text{Al}_2\text{O}_3$ is 4.99–5.87, and the value of $\text{K}_2\text{O}/\text{Na}_2\text{O}$ is 0.59–1.62. In the chondrite–standardized diagram of rare earth elements, the curve is consistent with that of the active continental margin. In the high field strong elements and large-ion lithophile elements, the samples are enriched in Th, Hf, Rb and K, and deficit in Nb, Ta, Ti, P, Ba and Sr. The cobweb curve of trace elements is similar to that of the active continental margin. The geochemical source analysis and tectonic background discrimination show that the parent rocks of the Mohe Formation sandstone were mainly originated from the mixing of the quartzite sedimentary rocks and the acidic volcanic rocks. The ages of the detrital zircons in the sandstone coincide with the ages of the Precambrian basement and granitic rocks in the southern margin of the basin. The formation age, provenance characteristics and tectonic setting of the Mohe Formation, which is widely developed in the Mohe Basin in the north of the Great Hinggan Mountains, are of great significance to the study of the formation and evolution of the basin.

Key words: Mohe formation; formation age; provenance; tectonic environment discrimination; geological survey engineering; Mohe Basin; Heilongjiang

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Fund support: Funded by China Geological Survey Project (No. DD2016007803, No. DD20191014).

1 引言

漠河盆地位于大兴安岭北段,呈东西向展布,东西长约300 km,南北宽约80 km,盆地面积约21500 km²,其地理坐标为 $52^{\circ}20' \sim 53^{\circ}33'N, 121^{\circ}03' \sim 126^{\circ}10'E$ 。盆地绝大部分隶属于黑龙江省大兴安岭地区,仅西侧小部分隶属于内蒙古自治区额尔古纳市,向西延入蒙古国境内,北面和东面过黑龙江延入俄罗斯境内,与俄罗斯境内的上阿穆尔盆地(也称乌舒蒙盆地)(17000 km²)相连为同一盆地(李春雷,2007;侯伟等,2010;李健等,2014;王少铁等,2015;王凯,2015;肖传桃等,2015;赵书跃等,2016;李良等,2017;甄森,2017;郭知鑫,2018),其北侧为蒙古—鄂霍茨克造山带,是中国地质学家直接了解蒙古—鄂霍茨克造山带中生代地质状况及其对中国北方中生代构造演化影响的窗口(和政军等,2003;侯伟等,2010;王少铁等,2015),因此,其沉积序列、盆地形成与演化、物源区特征等对蒙古—鄂霍茨克造山带的演化具有“探针”作用(李良等,2017)。但由于该区森林覆盖,地层出露差,研究程度低,因此对于漠河盆地性质、盆地成因与演化、形成时代、物源区特征等的认识还不统一。漠河盆地

的性质和动力学成因机制仍然存在争议较大(郭知鑫,2018),主要有前陆盆地(和政军等,2003;He et al., 2005;李春雷,2007;侯伟等,2010;王凯,2015)、磨拉石盆地(吴根耀等,2016)、活动大陆边缘盆地(赵寒冬等,2007)、挤压挠曲盆地或山间盆地(和钟铧等,2008)、由早期伸展断陷盆地演化为晚期磨拉石盆地(郭知鑫,2018)等。地层划分和时代限定观点不一,盆地形成时代从早侏罗世到早白垩世均有报道(王少铁等,2015;赵书跃等,2016;郭知鑫,2018)。盆地性质和形成时代等的不确定,严重制约了盆地基础地质调查和油气勘探等工作。

本文在前人研究基础上,对漠河组内凝灰岩夹层进行了锆石LA–ICP–MS U–Pb测年,对漠河组形成时代进行了限定,并结合漠河组砂岩岩相学、主微量元素地球化学和碎屑锆石U–Pb年龄特征,探讨了漠河盆地沉积物来源及物源区大地构造环境,为漠河盆地形成演化和蒙古—鄂霍茨克造山带构造演化提供了新的证据。

2 地质概况

漠河盆地位于兴蒙造山带东段额尔古纳地块北部,蒙古—鄂霍茨克造山带东南缘。盆地基底主

要为中新元古代兴华渡口岩群黑云钠长变粒岩、黑云石英片岩、二云(石英)片岩、斜长角闪岩等,南华系佳疙瘩组(倭勒根岩群)二云(石英)片岩、绿泥绢云片岩、板岩、千枚岩夹大理岩、石英岩等,震旦系额尔古纳河组灰岩、大理岩、千枚岩夹少量绿泥(石英)片岩等,古生界泥盆系下统泥鳅河组大理岩、千枚岩等地层及前中生代各期侵入岩。盆地盖层主要由三部分组成,第一部分为中晚侏罗世额木尔河群陆源碎屑沉积岩盖层,自下而上依次出露绣峰组、二十二站组、漠河组、开库康组(黑龙江省地质矿产局,1997);第二部分为早白垩世火山岩盖层,划分为塔木兰沟组玄武粗安岩、粗安岩、粗面玄武岩熔岩及火山碎屑岩(李春雷,2007),龙江组安山岩、角闪安山岩、安山质角砾熔岩夹安山质凝灰岩、玄武岩,光华组流纹岩、流纹质凝灰岩,九峰山组砂岩、含砾砂岩、(含集块)沉火山角砾岩夹粉砂岩、泥岩及煤线,甘河组玄武岩、玄武质安山岩及少量玄武安山质角砾熔岩等;第三部分为新生代古近系孙吴组及第四系堆积物。区内构造以漠河推覆构造为主,受其影响,盆地北部地质体多发生韧性变形。断裂以北西、北东向为主,近东西向、近南北向断裂次之(图1)。

漠河组最早由中国地质科学院1:100万地质图编组图于1963年在漠河县乌苏里沿江一带创建,指位于绣峰组或二十二站组之上,开库康组之下的碎屑岩组合,代表剖面位于漠河县乌苏里沿江(黑龙江省地质矿产局,1997),主要分布于漠河盆地中西部盆地中心地带(王少轶等,2015),岩性以灰黑色—灰黄色细、中、粗粒(岩屑)长石砂岩、长石岩屑砂岩,夹泥质粉砂岩、泥岩及少量凝灰岩、煤线,底部可见砾岩,向上变细,地层结构为退积型,为盆地扩张期产物(王少轶等,2015;赵书跃等,2016),该组属河流、沼泽相沉积,岩性在横向变化不大,沉降幅度各地相似,岩层厚度亦较稳定。产植物和孢粉化石,最大厚度3979.5 m(李春雷,2007)。研究区内漠河组底部与震旦系额尔古纳河组角度不整合接触,与上覆开库康组整合接触(图1b)。

3 样品特征

3.1 凝灰岩样品特征

本文测年样品采自额尔古纳市恩和哈达镇公

路旁,漠河组实测剖面PM05-3LT117(图2),该人工露头剖面为额尔古纳市恩和哈达南648高地剖面(PM05)上部层序,自下而上划分为27层,由中粒长石砂岩、粗粒长石砂岩、细粒长石砂岩、粉砂岩、泥岩、凝灰岩等,其中第3层和第22层为凝灰岩夹层(图2),镜下鉴定为流纹质晶屑熔结凝灰岩和流纹质晶屑玻屑凝灰岩(图3)。本次采取两层凝灰岩样品进行锆石LA-ICP MS测年,采样点地理坐标分别为 $51^{\circ}13'29.56''N, 121^{\circ}30'06.42''E$ 和 $51^{\circ}13'21.97''N, 121^{\circ}30'01.98''E$ 。

流纹质晶屑玻屑凝灰岩(WJL161 (PM05-3LT117-22)):浅灰绿色,凝灰结构,块状构造。镜下观察其成分主要有晶屑、玻屑火山尘等。晶屑:主要为石英、钾长石、斜长石、黑云母等。其中,石英呈棱角状一次棱角状,粒径0.05~0.2 mm,含量约8%;钾长石,呈棱角状一次棱角状,粒径0.1~0.2 mm,含量约8%;斜长石,粒径0.1~0.2 mm,含量约4%;黑云母呈不规则片状,少量。玻屑:为半塑性玻屑,呈鸡骨状、棒状、弓形状等,粒径0.1~0.5 mm,均脱玻化为镶嵌粒状或霏细状,含量约35%。火山尘:为粒径<0.05 mm的火山碎屑物,均脱玻化为隐晶质或霏细状,具弱绢云母化、方解石化,以集合体的形式分布于岩屑、晶屑和玻屑之间,含量约45%(图3a,b)。

流纹质晶屑熔结凝灰岩(WJL162 (PM05-3LT117-3)):灰色,晶屑熔结凝灰结构,假流纹构造。镜下观察其成分主要有岩屑、玻屑、晶屑及火山尘等。岩屑:刚性岩屑成分为玄武岩、流纹岩,粒径0.2~0.8 mm,含量约3%;塑性岩屑成分为流纹岩,呈透镜状、条带状,有压扁拉长现象,多为玻璃质,少部分脱玻化为隐晶质,多发生弯曲,粒径0.2~1.0 mm,含量约2%。玻屑:多为塑性,呈细条状、细纹状,部分脱玻化为隐晶质,具绢云母化,遇刚性岩屑、晶屑明显弯曲变薄,粒径0.05~0.6 mm,含量约15%。晶屑:为石英、钾长石和斜长石等。钾长石,呈次棱角状一次圆状,边缘具熔蚀现象,粒径0.1~1.0 mm,含量约5%;斜长石,粒径0.1~1.0 mm,含量约20%;石英,多熔蚀成港湾状,不规则状,晶面具熔蚀孔,少数呈棱角状,粒径0.1~1.0 mm,含量约15%。火山灰:粒度<0.05 mm的火山碎屑物,脱玻化为隐晶质,多绢云母化,与塑性岩屑、塑性玻屑等熔结在一起,形

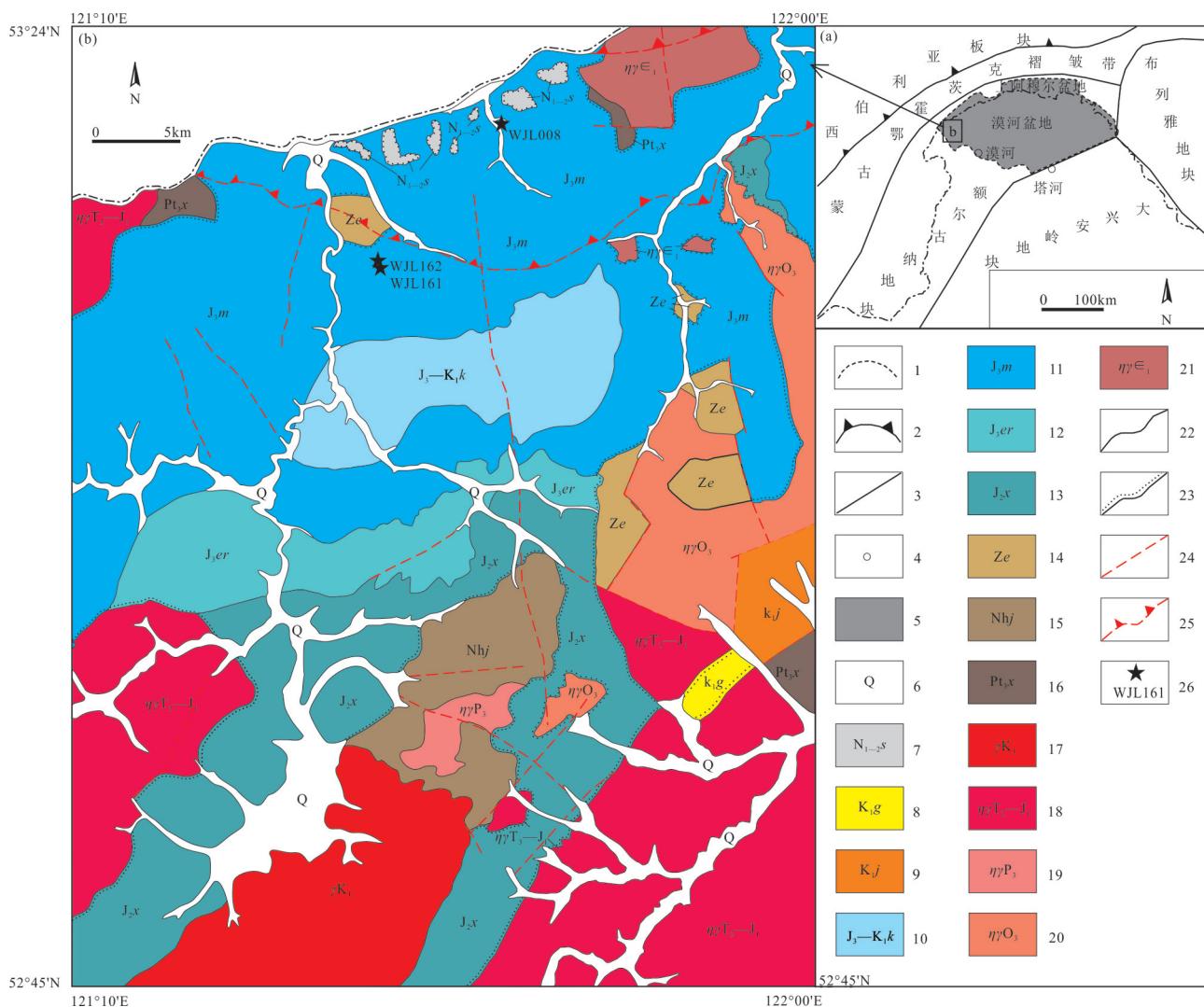


图1 大兴安岭北部大地构造分区图(a)和漠河盆地西缘地质简图(b)

(图a底图据1:25万恩和哈达幅、齐乾幅^①、1:25万漠河幅、漠河县幅^②和1:5万砂宝斯幅^③,在本文工作区调查基础上修改)

1—盆地边界;2—缝合带;3—构造边界;4—城镇;5—盆地范围;6—第四系;7—孙吴组;8—甘河组;9—九峰山组;10—开库康组;11—漠河组;12—二十二站组;13—绣峰组;14—额尔古纳河组;15—佳疙瘩组;16—兴华渡口群;17—早白垩世花岗岩;18—晚三叠世—早侏罗世二长花岗岩;19—晚二叠世二长花岗岩;20—晚奥陶世二长花岗岩;21—早寒武世二长花岗岩;22—整合界线;23—角度不整合界线;24—遥感解译断层;25—推覆构造;26—测年样品位置及编号

Fig. 1 Tectonic division of the northern part of the Da Hinggan Mountains (a) and geological map of the western margin of Mohe Basin (b)

(revised from the base maps of 1:250,000 Enhe, Hada, Qiqian, Mohe sheets and 1:50,000 Shabaoshi sheet)

1—Basin Boundary; 2—Suture Zone; 3—Tectonic Boundary; 4—Town; 5—Basin Scope; 6—Quaternary; 7—Sunwu Formation; 8—Ganhe Formation; 9—Jiufengshan Formation; 10—Kaikukang Formation; 11—Mohe Formation; 12—Ershierzhan Formation; 13—Xiufeng Formation; 14—Erguna Formation; 15—Jiageda Formation; 16—Xinghuadukou Group; 17—Early Cretaceous granite; 18—Late Triassic-Early Jurassic monzogranite; 19—Late Permian monzogranite; 20—Late Ordovician monzogranite; 21—Early Cambrian monzogranite; 22—Comfortable boundary; 23—Boundary of angular unconformity; 24—Remote sensing interpreting fault; 25—Nappe structure; 26—Location and number of dating sample

成假流纹构造,含量约40%(图3c,d)。

3.2 砂岩样品特征

主微量元素采自漠河县洛古河及兴华沟林场附近,样品均采自漠河组实测地质剖面,岩性以中粗粒长石砂岩和中粗粒岩屑长石砂岩为主,大部分

岩石受韧性剪切作用影响,发生片理化及糜棱岩化。岩石具陆源碎屑结构,块状构造,颗粒支撑,接触式—孔隙式胶结。碎屑物的磨圆度中等,分选性中等,成分成熟度中等。碎屑物约占85%~90%,填隙物占10%~15%。碎屑物主要由石英(30%~

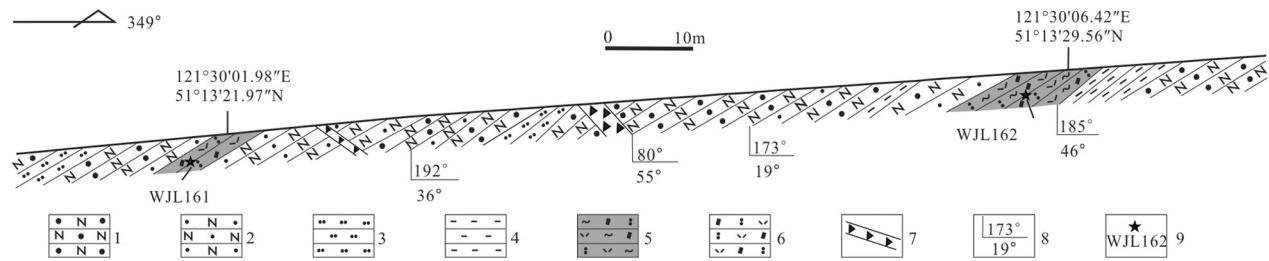


图2 漠河盆地西缘漠河组实测剖面图(PM05-3LT117)

1—粗粒长石砂岩;2—中细粒长石砂岩;3—粉砂岩;4—泥质岩;5—流纹质晶屑玻屑凝灰岩;6—流纹质晶屑熔结凝灰岩;7—断裂带;8—地层/断层产状;9—取样位置及编号

Fig 2 The measured profile of the Mohe Formation in the western margin of Mohe Basin (PM05-3LT117)

1—Coarse arkose;2—Medium fine arkose;3—Siltstone;4—Argillaceous rock;5—Rhyolitic glassy clastic tuff;6—Rhyolitic clastic fused tuff;7—Fault zone;8—Formation/fault occurrence;9—Location and number of sample

35%)、长石(40%~55%)、岩屑(10%~20%)，少量黑云母和白云母(<5%)等组成，粒径为0.25~2 mm。石英碎屑以单晶石英为主呈次棱角状一次圆状。长石分为钾长石和斜长石，钾长石含量高于斜长

石，K/F(钾长石/长石总量)约为0.6，斜长石为更一中长石，细而密的聚片双晶明显，钾长石为正长石、条纹长石、微斜长石。岩屑主要为安山岩、流纹岩、花岗岩、硅质岩、泥质岩、大理岩等。黑云母呈不规

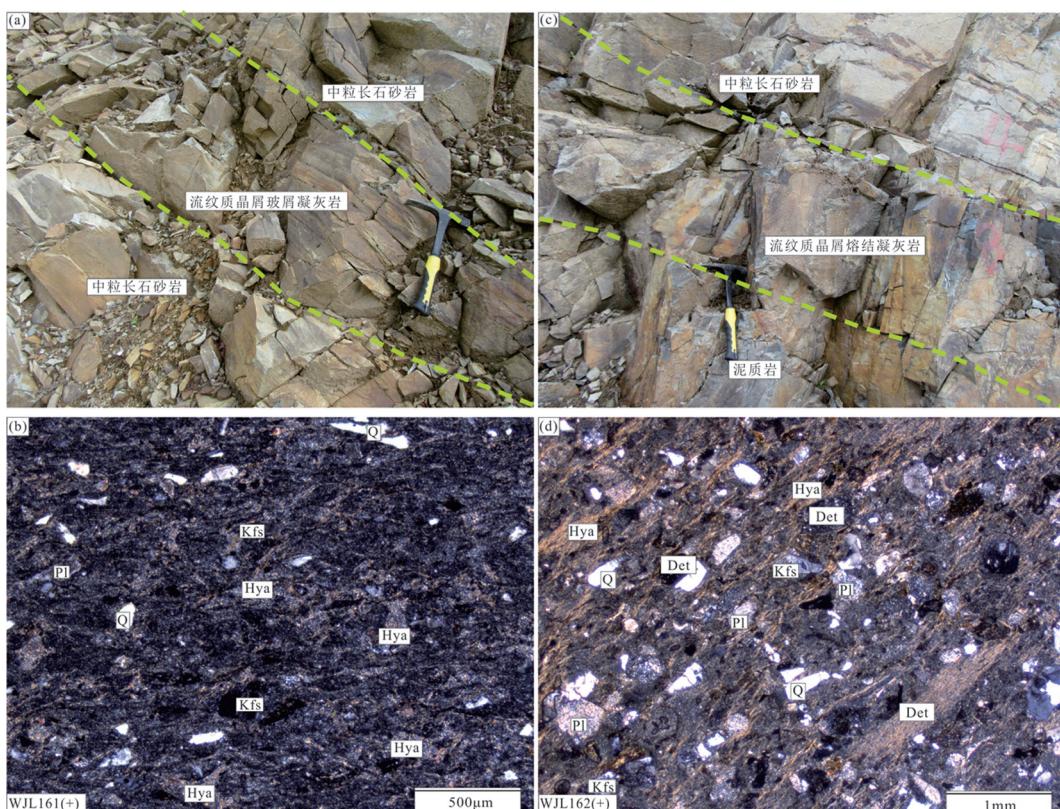


图3 漠河盆地西缘漠河组凝灰岩层露头及显微照片(+)

a—漠河组砂岩内流纹质晶屑玻屑凝灰岩;b—流纹质晶屑玻屑凝灰岩显微照片(+);c—漠河组砂岩内流纹质晶屑熔结凝灰岩;d—流纹质晶屑熔结凝灰岩显微照片(+);Q—石英;Kfs—钾长石;Pl—斜长石;Hya—玻屑;Det—岩屑

Fig 3 Outcrop and microscopic photographs of Mohe Formation tuff in western margin of Mohe Basin (+)

a—Mohe formation sandstone rhyolitic crystalline glass clastic tuff; b—Photomicrograph of rhyolitic crystalline glass clastic tuff (+); c—Mohe formation sandstone rhyolitic crystal fusion tuff; d—Photomicrograph of rhyolitic clinking tuff (+); Q—Quartz; Kfs—Potassium feldspar; Pl—

Plagioclase; Hya—Vitric Pyroclast; Det—Detritus

则片状,均压弯变形,白云母呈不规则片状,粒径大小为0.1~1 mm。填隙物主要由黏土矿物,部分显微鳞片状绢云母,少量细小石英、长石碎屑等组成,以混晶集合体充填于碎屑之间(图4)。

4 样品测试方法

样品全岩主量元素和微量元素分析测试工作在河北省区域地质矿产调查研究所完成。其中主量元素采用熔片法X射线荧光光谱法(XRF)测定,利用X射线激发样品产生二次X射线,对仪器误差和样品成分的二次X射线进行适当校准,通过射线的强度最终确定元素的含量,分析准确度和精度优于2%~3%;微量元素和稀土元素是用Teflon熔样罐进行熔样,然后采用Finnigan MAT公司生产的双聚焦高分辨等离子体质谱仪ICP-MS进行测定,分析流程采用AGV-1和BHVO-1等国际标准物质进行质量监测,分析过程中加入重复样以监测分析流程的精密度。准确度和精度优于10%。主量元素和微量元素分析测试具体实验条件和步骤参考文献(Liang et al.,2000)。

本文测年样品在河北省区域地质矿产调查研究所进行粉碎,并用浮选和电磁选法进行分选,在双目镜下选出晶形较好的锆石,然后将锆石黏贴在环氧树脂表面,打磨抛光后露出锆石的表面,制成样靶对测试样靶中锆石进行透射光、反射光和阴极

发光(CL)照相之后,采用LA-ICP MS仪器对锆石进行U-Pb测年分析。锆石U-Pb分析在天津地质矿产研究所同位素实验室LA-ICP MS仪器上完成,利用193 nm激光器对锆石进行剥蚀,采用的激光剥蚀的斑束直径为35 μm ,激光能量密度为13~14 J/cm²,频率为8~10 Hz,激光剥蚀物质以氦为载气送入Neptune,利用动态变焦扩大色散可以同时接收质量数相差很大的U-Pb同位素,从而进行锆石U-Pb同位素原位测定。采用TEMORA作为外部锆石年龄标准。利用NIST610玻璃标样作为外标计算锆石样品的Pb、U、Th含量。样品信号采集时间60 s(其中20 s为空白的测定)。采用ICP-MS DataCal和Isoplot程序进行数据处理,数据处理方法同文献(Liu et al.,2008)。

5 测试结果

5.1 锆石U-Pb年代学

对洛古河地区漠河组内两个凝灰岩夹层样品和一个糜棱岩化砂岩样品进行了LA-ICP-MS锆石U-Pb同位素分析,分析结果见表1。

凝灰岩样品WJL161、WJL162和砂岩样品WJL008,其代表性锆石特征如图5。锆石CL图像显示,样品WJL161、WJL162锆石呈柱状,长宽比1:1~2:1,多数锆石具有清晰的振荡环带,显示岩浆成因锆石,样品WJL161中U含量为 47×10^{-6} ~ $544 \times$

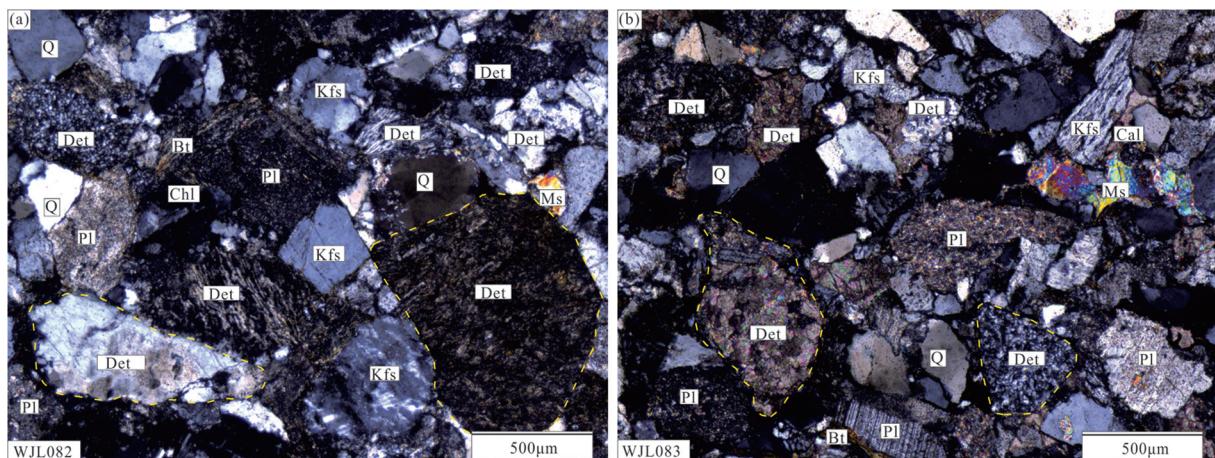


图4 漠河盆地西缘漠河组砂岩(a—样品WJL082;b—样品WJL083)显微照片(+)

Q—石英;Kfs—钾长石;Pl—斜长石;Chl—绿泥石;Bt—黑云母;Ms—白云母;Cal—方解石;Det—岩屑

Fig.4 Microscopic photographs of sandstone (a—sample WJL082; b—sample WJL083) of Mohe Formation in the western margin of Mohe Basin (+)

Q= Quartz; Kfs= Potassium feldspar; Pl= Plagioclase; Chl= Chlorite; Bt= Biotite; Ms= Muscovite; Cal= Calcite; Det—Detritus

表1 漠河盆地西缘漠河组凝灰岩层 LA-ICP-MS 锆石 U-Pb 分析结果

Table 1 LA-ICP-MS Zircon U-Pb assay data of tuff layer of the Mohe Formation in the western margin of Mohe Basin

测点号	同位素比值										年龄/ Ma				
	Pb	U	Th/U	$^{207}\text{Pb}/^{206}\text{Pb}$	1 σ	$^{207}\text{Pb}/^{235}\text{U}$	1 σ	$^{206}\text{Pb}/^{238}\text{U}$	1 σ	$^{207}\text{Pb}/^{206}\text{Pb}$	1 σ	$^{207}\text{Pb}/^{235}\text{U}$	1 σ	$^{206}\text{Pb}/^{238}\text{U}$	1 σ
WLJ161 流纹质晶屑玻屑凝灰岩															
1	135	446	0.324	0.05248	0.0020	0.30268	0.0114	0.04184	0.0002	306	85	268	9	264	1
2	142	328	0.804	0.05701	0.0017	0.21844	0.0063	0.02781	0.0001	500	60	201	5	177	1
3	156	289	0.955	0.04933	0.0059	0.17076	0.0206	0.02466	0.0003	165	256	160	18	157	2
4	166	424	0.205	0.05716	0.0006	0.57136	0.0069	0.07248	0.0004	498	22	459	4	451	2
5	127	421	0.464	0.04923	0.0024	0.18744	0.0095	0.02746	0.0002	167	145	174	8	175	1
6	133	471	0.273	0.05110	0.0033	0.24612	0.0487	0.03090	0.0021	256	152	223	40	196	13
7	118	478	0.303	0.05264	0.0016	0.30647	0.0092	0.04223	0.0001	322	73	271	7	267	1
8	92	544	0.245	0.05081	0.0022	0.22303	0.0094	0.03189	0.0001	232	98	204	8	202	1
9	168	338	0.606	0.04922	0.0074	0.16777	0.0252	0.02480	0.0002	167	309	157	22	158	2
10	170	257	1.106	0.04910	0.0026	0.16802	0.0092	0.02473	0.0002	154	119	158	8	158	1
11	230	47	8.117	0.04888	0.0018	0.16763	0.0061	0.02492	0.0002	143	90	157	5	159	1
12	95	511	0.319	0.05003	0.0024	0.20425	0.0097	0.02960	0.0001	198	109	189	8	188	1
13	157	470	0.081	0.07391	0.0005	0.90184	0.0071	0.08845	0.0004	1039	12	653	4	546	2
14	165	281	0.954	0.05000	0.0034	0.17204	0.0118	0.02488	0.0001	195	157	161	10	158	1
15	113	489	0.303	0.05088	0.0025	0.21761	0.0105	0.03114	0.0002	235	115	200	9	198	1
16	118	452	0.394	0.04987	0.0010	0.19522	0.0041	0.02837	0.0001	187	46	181	3	180	1
17	169	213	1.552	0.04898	0.0019	0.16745	0.0065	0.02478	0.0001	146	91	157	6	158	1
18	117	415	0.533	0.04948	0.0015	0.16965	0.0053	0.02484	0.0001	172	70	159	5	158	0
19	114	478	0.333	0.05009	0.0014	0.20435	0.0057	0.02957	0.0001	198	31	189	5	188	0
20	146	345	0.692	0.04956	0.0039	0.16970	0.0131	0.02493	0.0002	176	170	159	11	159	1
21	159	265	1.108	0.09118	0.0036	0.33917	0.0134	0.02697	0.0001	1450	75	297	10	172	1
22	168	273	0.990	0.05077	0.0048	0.18177	0.0244	0.02480	0.0007	232	204	170	21	158	5
23	112	473	0.354	0.05066	0.0020	0.21448	0.0085	0.03068	0.0001	233	95	197	7	195	1
24	141	297	0.999	0.04943	0.0096	0.18828	0.0362	0.02745	0.0003	169	400	175	31	175	2
25	177	241	1.185	0.04926	0.0028	0.16763	0.0094	0.02468	0.0002	167	133	157	8	157	1
26	176	251	1.107	0.04960	0.0012	0.16943	0.0044	0.02475	0.0002	176	57	159	4	158	1
27	186	378	0.270	0.05620	0.0009	0.55288	0.0110	0.07120	0.0004	461	31	447	7	443	3
28	128	442	0.393	0.04932	0.0021	0.16905	0.0072	0.02487	0.0002	161	98	159	6	158	1
29	139	417	0.421	0.05019	0.0061	0.21009	0.0255	0.03037	0.0002	211	254	194	21	193	1
30	131	376	0.618	0.05157	0.0042	0.21269	0.0176	0.02985	0.0002	265	182	196	15	190	1
31	142	330	0.794	0.05012	0.0054	0.17245	0.0189	0.02479	0.0002	211	220	162	16	158	1
32	141	326	0.823	0.04965	0.0044	0.16892	0.0149	0.02470	0.0002	189	183	158	13	157	1
WLJ162 流纹质晶屑熔结凝灰岩															
1	13	602	1.751	0.0810	0.0025	0.2961	0.0095	0.0265	0.0003	1221	61	263	8	169	2
2	8	239	0.807	0.0519	0.0023	0.1798	0.0081	0.0251	0.0003	283	100	168	8	160	2
3	9	368	1.321	0.0542	0.0020	0.1838	0.0068	0.0246	0.0003	381	82	171	6	157	2
4	14	705	1.724	0.0489	0.0013	0.1650	0.0046	0.0245	0.0003	143	63	155	4	156	2
5	7	198	0.785	0.0566	0.0021	0.1871	0.0072	0.0240	0.0003	475	83	174	7	153	2
6	14	220	0.670	0.0554	0.0012	0.2931	0.0068	0.0383	0.0004	430	49	261	6	243	3
7	9	300	0.991	0.0497	0.0018	0.1628	0.0060	0.0237	0.0003	183	86	153	6	151	2
8	7	157	0.597	0.0550	0.0022	0.1820	0.0075	0.0240	0.0003	411	91	170	7	153	2

续表1

测点号	同位素比值										年龄/Ma					
	Pb	U	Th/U	$^{207}\text{Pb}/^{206}\text{Pb}$	1σ	$^{207}\text{Pb}/^{235}\text{U}$	1σ	$^{206}\text{Pb}/^{238}\text{U}$	1σ	$^{207}\text{Pb}/^{206}\text{Pb}$	1σ	$^{207}\text{Pb}/^{235}\text{U}$	1σ	$^{206}\text{Pb}/^{238}\text{U}$	1σ	
9	19	986	1.791	0.0605	0.0012	0.2085	0.0045	0.0250	0.0003	620	44	192	4	159	2	
10	14	838	2.048	0.0503	0.0018	0.1710	0.0063	0.0247	0.0003	207	83	160	6	157	2	
11	10	401	1.212	0.0531	0.0017	0.1766	0.0061	0.0241	0.0003	331	75	165	6	154	2	
12	15	807	1.744	0.0501	0.0016	0.1655	0.0053	0.0239	0.0003	201	73	156	5	153	2	
13	11	262	0.668	0.0798	0.0017	0.2661	0.0057	0.0242	0.0003	1192	41	240	5	154	2	
14	139	520	0.552	0.0682	0.0008	1.2922	0.0178	0.1374	0.0015	874	24	842	12	830	9	
15	15	179	0.467	0.0565	0.0012	0.2968	0.0064	0.0381	0.0004	471	45	264	6	241	3	
16	13	444	1.008	0.0529	0.0014	0.1771	0.0051	0.0243	0.0003	324	62	166	5	155	2	
17	6	76	0.317	0.0503	0.0029	0.1792	0.0105	0.0258	0.0003	209	132	167	10	164	2	
18	9	343	1.291	0.0496	0.0019	0.1799	0.0072	0.0263	0.0003	175	91	168	7	167	2	
19	10	372	1.053	0.0521	0.0016	0.1774	0.0055	0.0247	0.0003	290	69	166	5	157	2	
20	5	14	0.063	0.0532	0.0020	0.1825	0.0071	0.0249	0.0003	337	87	170	7	158	2	
21	8	326	1.278	0.0496	0.0021	0.1757	0.0077	0.0257	0.0003	176	100	164	7	164	2	
22	9	274	0.840	0.0541	0.0016	0.1858	0.0056	0.0249	0.0003	374	65	173	5	159	2	
23	21	187	1.659	0.0688	0.0011	1.3332	0.0233	0.1406	0.0015	891	33	860	15	848	9	
24	5	112	0.544	0.0499	0.0055	0.1695	0.0185	0.0246	0.0003	191	258	159	17	157	2	
25	18	204	0.479	0.0644	0.0014	0.3431	0.0077	0.0386	0.0004	756	47	300	7	244	3	
26	33	558	0.731	0.0520	0.0008	0.2752	0.0046	0.0384	0.0004	287	34	247	4	243	3	
27	8	209	0.744	0.0545	0.0032	0.1793	0.0108	0.0239	0.0003	392	131	167	10	152	2	
28	7	237	0.973	0.0523	0.0026	0.1748	0.0091	0.0242	0.0003	299	115	164	8	154	2	
29	6	177	0.835	0.0514	0.0024	0.1745	0.0084	0.0246	0.0003	258	107	163	8	157	2	
30	6	197	0.921	0.0550	0.0024	0.1883	0.0083	0.0248	0.0003	414	96	175	8	158	2	
31	45	607	0.557	0.0524	0.0007	0.2762	0.0043	0.0382	0.0004	303	32	248	4	242	3	
32	9	425	1.539	0.0537	0.0041	0.1847	0.0145	0.0249	0.0003	359	174	172	13	159	2	
WJL008糜棱岩化粗粒长石砂岩																
1	61	894	0.020	0.0557	0.0008	0.5734	0.0088	0.0746	0.0008	441	30	460	7	464	5	
2	33	815	0.525	0.0520	0.0007	0.2956	0.0046	0.0413	0.0004	284	33	263	4	261	3	
3	9	219	0.613	0.0535	0.0012	0.3071	0.0073	0.0416	0.0004	351	52	272	6	263	3	
4	48	636	0.290	0.0568	0.0008	0.6057	0.0092	0.0773	0.0008	485	31	481	7	480	5	
5	14	334	0.586	0.0525	0.0009	0.2828	0.0054	0.0391	0.0004	307	41	253	5	247	2	
6	25	299	0.644	0.0573	0.0012	0.5958	0.0153	0.0754	0.0008	503	48	475	12	469	5	
7	38	457	0.374	0.0581	0.0008	0.6465	0.0100	0.0807	0.0008	533	31	506	8	500	5	
8	4	124	0.779	0.0517	0.0027	0.1852	0.0097	0.0260	0.0003	271	121	173	9	165	2	
9	10	212	0.772	0.0532	0.0013	0.3125	0.0079	0.0426	0.0004	337	55	276	7	269	3	
10	22	271	0.519	0.0568	0.0009	0.6030	0.0102	0.0770	0.0008	483	34	479	8	478	5	
11	16	380	0.432	0.0527	0.0010	0.2998	0.0058	0.0413	0.0004	315	42	266	5	261	3	
12	9	108	0.470	0.0574	0.0014	0.6210	0.0160	0.0785	0.0008	506	55	491	13	487	5	
13	59	442	0.349	0.0658	0.0011	1.1893	0.0227	0.1310	0.0014	802	34	796	15	794	8	
14	8	190	0.739	0.0521	0.0014	0.2966	0.0084	0.0413	0.0004	292	63	264	7	261	3	
15	38	149	0.601	0.0878	0.0012	2.8769	0.0492	0.2377	0.0030	1378	27	1376	24	1375	17	
16	56	480	0.059	0.0652	0.0009	1.1158	0.0169	0.1241	0.0013	781	29	761	12	754	8	
17	53	673	0.052	0.0579	0.0008	0.6796	0.0107	0.0851	0.0009	527	31	527	8	526	6	
18	46	1071	0.628	0.0535	0.0008	0.3032	0.0049	0.0411	0.0004	349	34	269	4	260	3	
19	52	235	0.337	0.1014	0.0014	2.9962	0.0467	0.2143	0.0023	1650	25	1407	22	1252	13	

续表1

测点号	同位素比值												年龄/Ma			
	含量/ 10^{-6}		Th/U	$^{207}\text{Pb}/^{206}\text{Pb}$	1σ	$^{207}\text{Pb}/^{235}\text{U}$	1σ	$^{206}\text{Pb}/^{238}\text{U}$	1σ	$^{207}\text{Pb}/^{206}\text{Pb}$	1σ	$^{207}\text{Pb}/^{235}\text{U}$	1σ	$^{206}\text{Pb}/^{238}\text{U}$	1σ	
20	40	161	0.694	0.0960	0.0013	2.9818	0.0466	0.2253	0.0024	1547	26	1403	22	1310	14	
21	85	487	0.405	0.0732	0.0010	1.7265	0.0256	0.1712	0.0018	1018	27	1018	15	1019	10	
22	34	211	0.181	0.0716	0.0010	1.6065	0.0248	0.1627	0.0017	975	28	973	15	972	10	
23	34	439	0.400	0.0578	0.0008	0.6016	0.0093	0.0755	0.0008	522	31	478	7	469	5	
24	16	60	0.411	0.0906	0.0014	3.0601	0.0506	0.2450	0.0025	1438	29	1423	24	1413	14	
25	10	39	0.959	0.0839	0.0015	2.5086	0.0490	0.2169	0.0022	1290	36	1275	25	1265	13	
26	2	65	0.988	0.0521	0.0048	0.2045	0.0188	0.0284	0.0003	291	210	189	17	181	2	
27	12	270	0.613	0.0533	0.0011	0.3065	0.0065	0.0417	0.0004	340	46	271	6	264	3	
28	8	104	0.387	0.0579	0.0013	0.6436	0.0157	0.0806	0.0008	527	51	505	12	500	5	
29	15	476	1.008	0.0498	0.0009	0.1872	0.0036	0.0273	0.0003	185	42	174	3	173	2	
30	7	85	0.572	0.0576	0.0016	0.5875	0.0172	0.0740	0.0008	514	62	469	14	460	5	
31	29	367	0.526	0.0579	0.0008	0.6120	0.0094	0.0767	0.0008	525	31	485	7	476	5	
32	24	93	0.685	0.0880	0.0012	2.9202	0.0462	0.2406	0.0026	1383	27	1387	22	1390	15	
33	16	466	0.977	0.0518	0.0009	0.2254	0.0043	0.0316	0.0003	277	41	206	4	200	2	
34	18	88	0.293	0.0795	0.0012	2.1889	0.0378	0.1996	0.0023	1185	30	1178	20	1173	14	
35	78	466	0.074	0.0750	0.0010	1.8381	0.0268	0.1777	0.0018	1069	27	1059	15	1054	10	
36	18	237	0.708	0.0571	0.0009	0.5749	0.0100	0.0731	0.0007	494	36	461	8	455	5	
37	15	345	0.479	0.0525	0.0009	0.3251	0.0060	0.0449	0.0005	307	41	286	5	283	3	
38	30	410	0.092	0.0577	0.0008	0.6129	0.0096	0.0770	0.0008	519	32	485	8	478	5	
39	52	225	0.608	0.0951	0.0013	2.8699	0.0436	0.2189	0.0022	1530	26	1374	21	1276	13	
40	9	211	0.458	0.0512	0.0013	0.3004	0.0082	0.0425	0.0005	252	60	267	7	268	3	
41	8	237	0.721	0.0509	0.0013	0.2215	0.0060	0.0315	0.0003	238	60	203	5	200	2	
42	18	526	0.878	0.0501	0.0009	0.2167	0.0040	0.0314	0.0003	199	41	199	4	199	2	
43	6	223	0.732	0.0503	0.0017	0.1961	0.0067	0.0283	0.0003	210	78	182	6	180	2	
44	34	132	0.471	0.0914	0.0013	3.1278	0.0471	0.2481	0.0025	1456	26	1440	22	1429	14	
45	46	543	0.890	0.0573	0.0008	0.6146	0.0095	0.0778	0.0008	502	31	486	8	483	5	
46	39	297	0.300	0.0662	0.0009	1.2113	0.0182	0.1327	0.0013	813	29	806	12	803	8	
47	18	447	0.612	0.0520	0.0008	0.2724	0.0046	0.0380	0.0004	285	37	245	4	240	3	
48	109	763	0.638	0.0672	0.0009	1.2951	0.0192	0.1399	0.0014	842	29	844	13	844	8	
49	4	113	0.767	0.0518	0.0024	0.2159	0.0102	0.0303	0.0003	275	106	199	9	192	2	
50	13	289	0.781	0.0530	0.0010	0.2862	0.0061	0.0392	0.0004	327	44	256	5	248	3	
51	37	293	0.176	0.0766	0.0011	1.3467	0.0214	0.1276	0.0013	1110	29	866	14	774	8	
52	52	366	0.768	0.0662	0.0009	1.2169	0.0183	0.1333	0.0014	813	29	808	12	807	8	
53	42	545	0.509	0.0570	0.0008	0.6054	0.0092	0.0770	0.0008	493	31	481	7	478	5	
54	23	455	1.647	0.0513	0.0009	0.2855	0.0056	0.0404	0.0004	253	43	255	5	255	3	
55	2	68	0.637	0.0507	0.0036	0.2217	0.0158	0.0317	0.0004	228	165	203	15	201	2	
56	13	388	1.878	0.0500	0.0012	0.1831	0.0044	0.0266	0.0003	195	54	171	4	169	2	
57	109	1507	0.109	0.0583	0.0008	0.6237	0.0093	0.0777	0.0008	539	30	492	7	482	5	
58	91	818	0.273	0.0632	0.0009	0.9937	0.0193	0.1140	0.0017	716	29	701	14	696	10	
59	27	358	0.570	0.0559	0.0008	0.5785	0.0094	0.0750	0.0008	449	34	464	8	466	5	
60	24	298	0.933	0.0570	0.0009	0.5893	0.0097	0.0750	0.0008	490	34	470	8	466	5	
61	12	363	0.803	0.0514	0.0009	0.2274	0.0044	0.0321	0.0003	259	41	208	4	204	2	
62	11	268	0.785	0.0525	0.0012	0.2839	0.0067	0.0393	0.0004	305	51	254	6	248	3	
63	5	112	1.180	0.0530	0.0020	0.3135	0.0122	0.0429	0.0004	329	87	277	11	271	3	
64	6	114	1.174	0.0519	0.0020	0.3442	0.0136	0.0481	0.0005	282	87	300	12	303	3	

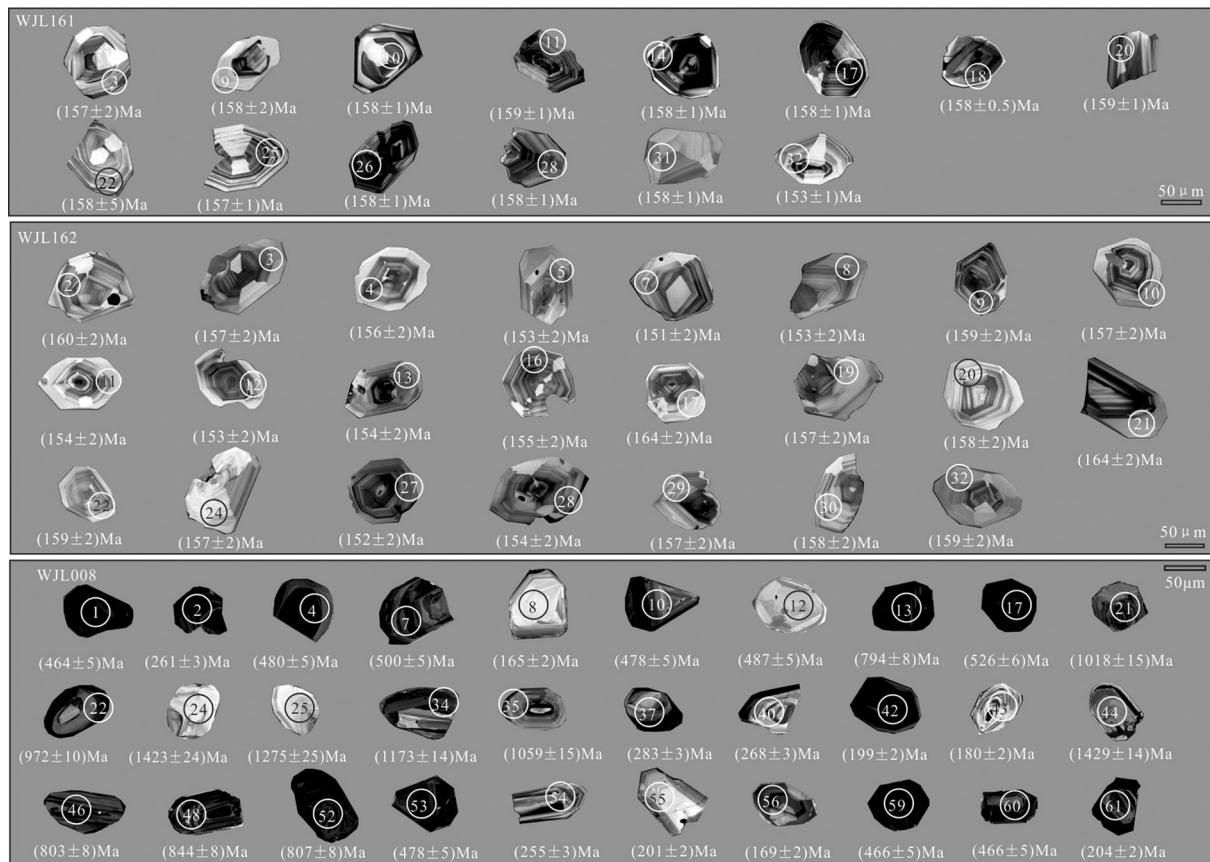


图5 漠河盆地西缘漠河组凝灰岩和糜棱岩化砂岩代表性锆石CL图像与表面年龄

Fig.5 Typical zircon CL images and surface ages of tuff and mylonitized sandstone in Mohe Formation, western margin of Mohe Basin

10^{-6} , Pb含量为 $92 \times 10^{-6} \sim 230 \times 10^{-6}$, Th/U比值为在0.08~8.12(多数为0.2~1.2),样品WJL162中U含量为 $113 \times 10^{-6} \sim 1090 \times 10^{-6}$, Pb含量为 $4 \times 10^{-6} \sim 65 \times 10^{-6}$, Th/U比值在0.06~2.05(多数为0.3~1.8),锆石均显示为岩浆锆石成因(Koschek, 1993),样品WJL008中U含量为 $39 \times 10^{-6} \sim 1507 \times 10^{-6}$, Pb含量为 $2 \times 10^{-6} \sim 109 \times 10^{-6}$, Th/U比值在0.02~1.88,多数锆石显示为岩浆锆石成因,其中44号锆石,有明显的变质增生边、34号锆石有面状分带结构、35号锆石有椭圆形残留核,增生边有面状分带结构(图5),Th/U比值0.07,应为变质成因锆石(李亚萍等,2007)。

样品WJL161、WJL162两个样品均测试了32个测点,测点主要选择在具有明显震荡环带的区域,部分测点年龄偏离谐和线,为凝灰岩中碎屑锆石,少数测点谐和度低,可能为后期铅丢失或者有非放射性成因铅混入的影响,其年龄无意义。样品WJL161谐和线上13个分析点的 $^{206}\text{Pb}/^{238}\text{U}$ 加权平均年龄为 $(158 \pm 1)\text{ Ma}$ (图6a),样品WJL162谐和线上

23个分析点的 $^{206}\text{Pb}/^{238}\text{U}$ 加权平均年龄 $(157 \pm 2)\text{ Ma}$ (图6b),代表凝灰岩形成时代为晚侏罗世。因此,洛古河地区漠河组沉积年龄应在157~158 Ma,属晚侏罗世早期。样品WJL008中64个测点锆石U-Pb年龄为165~1429 Ma,时间跨度较大,表明漠河组砂岩源区复杂。按年龄及频率分布特征划分为7个年龄组:(1)1019~1429 Ma:该组共11粒锆石,占总数的17.18%,属于中元古代;(2)696~972 Ma:该组共8粒锆石,占总数的12.5%,峰值年龄为800 Ma,属于新元古代;(3)455~526 Ma:该组共18粒锆石,占总数的28.13%,峰值年龄为479 Ma,属于早奥陶世;(4)303 Ma:该组共1粒锆石,占总数的1.56%,属于晚石炭世;(5)240~283 Ma:该组共15粒锆石,占总数的23.43%,峰值年龄为260 Ma,属于中二叠世;(6)199~204 Ma:该组共5粒锆石,占总数的7.81%,峰值年龄为201 Ma,属于晚三叠世;(7)165~192 Ma:该组共6粒锆石,占总数的9.39%,峰值年龄为180 Ma,属于早侏罗世。最年轻锆石U-Pb年龄为

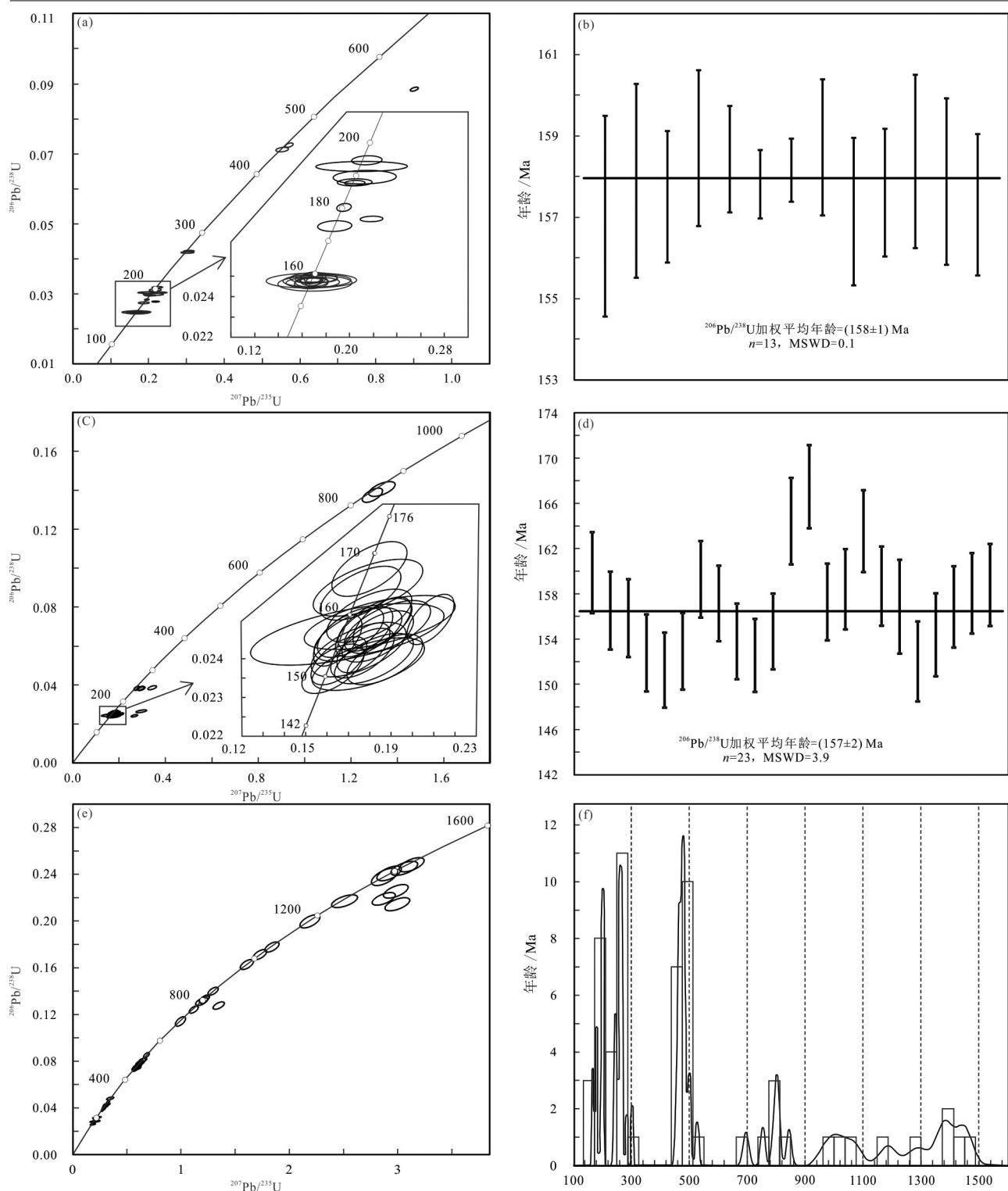


图6 漠河盆地西缘漠河组凝灰岩锆石U-Pb谐和年龄图(a、c)、加权平均年龄图(b、d)和糜棱岩化砂岩锆石U-Pb谐和年龄图(e)、直方图(f)

Fig.6 Zircon U-Pb concordance age (a, c), weighted average age (b, d) and zircon U-Pb concordance age (e) and histogram (f) of mylonitized sandstone of Mohe Formation in the Western Margin of Mohe Basin

表2 漠河盆地西缘漠河组砂岩岩石地球化学特征

Table 2 Geochemical characteristics of sandstone in the Mohe Formation, western margin of Mohe Basin

样品编号	WJL063	WJL064	WJL067	WJL068	WJL069	WJL070	WJL082	WJL083	WJL190	WJL193	WJL194	WJL195
SiO ₂	73.63	75.04	73.21	71.74	73.44	76.13	70.54	70.20	73.54	71.95	72.90	75.61
TiO ₂	0.20	0.27	0.31	0.28	0.25	0.25	0.35	0.38	0.45	0.41	0.32	0.20
Al ₂ O ₃	13.77	13.69	13.88	14.39	13.98	13.40	13.70	12.92	13.29	13.74	12.87	12.88
Fe ₂ O ₃	0.62	0.91	1.17	0.98	0.62	1.26	0.34	0.28	0.84	0.96	0.76	0.85
FeO	1.10	0.95	0.86	1.13	0.92	0.38	1.77	2.27	1.59	1.51	1.48	0.58
MnO	0.03	0.02	0.10	0.05	0.03	0.03	0.05	0.07	0.059	0.046	0.043	0.026
MgO	0.91	0.78	0.91	0.89	0.73	0.41	1.17	1.27	1.07	1.30	1.27	0.62
CaO	0.63	0.21	0.17	1.25	0.26	0.17	1.79	2.37	0.88	1.06	1.25	0.74
Na ₂ O	3.22	4.16	3.10	4.31	3.23	3.78	3.92	3.53	4.03	2.81	3.23	2.81
K ₂ O	4.00	2.44	3.86	3.59	5.23	2.64	3.54	3.28	3.11	3.88	3.54	3.79
P ₂ O ₅	0.06	0.07	0.07	0.06	0.07	0.06	0.09	0.08	0.092	0.090	0.067	0.055
LOI	1.76	1.37	1.75	1.21	1.12	1.41	2.62	3.26	0.92	2.12	2.16	1.72
Total	98.17	98.54	97.64	98.66	98.74	98.52	97.26	96.66	99.88	99.87	99.90	99.90
F ₁	-4.85	-2.48	-4.73	-2.87	-6.27	-2.26	-4.13	-4.37	-4.00	-5.36	-5.09	-4.66
F ₂	0.32	-1.59	-0.31	0.47	2.12	-1.18	0.21	-0.25	-0.83	-0.47	-0.61	0.09
Y	18.64	18.98	16.00	18.66	17.79	16.91	21.11	24.38	27.6	24.9	21.4	15.7
La	25.55	25.19	21.98	24.17	25.53	23.98	29.65	36.79	47.2	37.4	27.7	21.2
Ce	36.66	36.14	59.64	40.37	41.38	48.17	49.48	66.29	96.4	72.0	52.3	38.2
Pr	4.47	4.66	4.92	4.54	4.73	5.48	6.08	7.91	10.6	8.62	6.43	4.58
Nd	17.08	17.95	18.03	17.35	17.99	20.38	23.03	29.03	40.8	34.2	25.3	19.0
Sm	3.28	3.47	3.36	3.36	3.42	3.75	4.14	5.08	6.53	5.78	4.37	3.40
Eu	0.79	0.71	0.73	0.86	0.85	0.72	0.79	0.87	1.01	1.08	0.84	0.80
Gd	2.95	3.20	3.30	3.12	3.10	3.55	3.96	4.66	5.83	4.91	3.96	2.90
Tb	0.50	0.52	0.50	0.51	0.50	0.53	0.65	0.73	0.91	0.75	0.64	0.48
Dy	3.04	3.13	2.99	3.05	2.95	3.07	3.89	4.02	5.08	4.16	3.75	2.71
Ho	0.58	0.61	0.57	0.58	0.57	0.58	0.73	0.78	1.03	0.82	0.75	0.55
Er	1.69	1.78	1.70	1.70	1.70	1.71	2.05	2.18	3.08	2.39	2.20	1.58
Tm	0.26	0.28	0.27	0.27	0.27	0.26	0.33	0.35	0.57	0.43	0.40	0.29
Yb	1.63	1.79	1.73	1.65	1.72	1.63	2.01	2.21	3.41	2.50	2.36	1.70
Lu	0.28	0.30	0.27	0.27	0.29	0.26	0.32	0.35	0.50	0.37	0.34	0.25
δEu	0.77	0.66	0.67	0.82	0.80	0.60	0.60	0.55	0.50	0.62	0.62	0.78
δCe	0.84	0.82	1.41	0.94	0.92	1.03	0.90	0.95	1.06	0.98	0.96	0.95
ΣREE	98.76	99.73	119.99	101.80	105.00	114.07	127.11	161.25	222.95	175.41	131.34	97.64
LREE/HREE	2.97	2.88	3.97	3.04	3.25	3.60	3.23	3.68	4.22	3.85	3.27	3.34
(La/Yb) _N	11.23	10.12	9.09	10.48	10.64	10.54	10.56	11.91	9.94	10.71	8.43	8.96
(δEu) _{UCC}	1.10	0.93	0.96	1.16	1.13	0.85	0.85	0.78	0.71	1.10	0.93	0.96
(La/Yb) _{UCC}	1.01	0.91	0.82	0.94	0.96	0.95	1.07	0.89	0.96	0.76	0.81	
Li	13.58	15.03	17.09	13.27	12.71	16.46	15.81	14.59	22.7	19.3	16.2	14.1
Be	2.83	2.49	2.78	2.83	2.61	2.47	2.50	2.43	2.45	2.70	2.38	2.34
Sc	3.29	4.26	5.82	5.63	4.41	4.60	4.86	5.78	6.05	6.08	4.85	3.50
V	26.79	31.80	38.97	33.48	32.60	26.82	43.63	49.44	37.0	38.1	31.5	21.2
Cr	10.80	17.37	21.89	16.09	12.43	13.74	25.51	20.90	21.3	24.9	21.2	14.1
Co	2.77	3.22	5.09	7.18	8.36	2.14	4.26	4.95	5.73	3.50	2.29	1.68
Ni	4.22	6.37	13.84	8.77	5.05	4.72	9.01	7.42	8.68	8.71	8.44	3.73
Cu	5.27	1.76	4.25	4.86	15.28	3.06	5.81	4.61	4.65	6.88	2.66	3.74
Zn	34.31	36.78	41.34	44.02	36.38	28.30	39.82	47.21	54.3	51.5	41.2	34.7
Ga	16.97	16.51	16.96	17.00	17.90	15.61	17.02	16.96	16.9	18.8	16.7	16.1
Rb	125.33	98.19	136.66	112.28	147.05	102.45	111.72	105.06	111	148	134	129
Sr	174.81	126.05	100.88	207.52	127.75	91.55	185.75	180.76	145	184	133	96.3
Zr	115.81	112.55	111.90	101.03	132.29	114.18	131.65	156.45	216	167	122	113
Nb	7.00	8.73	9.99	8.69	8.77	8.06	10.68	11.87	14.8	12.2	10.6	6.91
Mo	0.18	0.11	0.20	0.31	0.19	0.14	0.16	0.12	0.15	0.10	0.10	0.20
Cs	3.77	7.34	4.36	3.27	4.18	3.37	3.70	3.64	4.75	6.29	4.97	5.79
Ba	704.60	416.67	541.60	631.74	780.85	313.00	582.77	514.46	562	689	532	659
Hf	3.44	3.28	3.42	2.96	3.89	3.37	3.95	4.86	6.64	4.47	3.65	3.31
Ta	0.56	0.66	0.74	0.68	0.68	0.63	0.80	1.07	1.28	1.17	0.77	0.55
Pb	19.87	14.17	19.63	20.00	23.38	20.43	19.55	20.30	22.1	24.3	17.9	28.7
Bi	0.14	0.14	0.16	0.17	0.12	0.13	0.15	0.13	0.18	0.26	0.24	0.22
Th	7.11	7.05	9.09	7.96	8.60	8.88	9.55	13.74	18.7	12.2	10.7	7.39
U	1.20	1.02	1.53	2.04	1.60	1.44	1.74	1.96	3.30	1.92	1.75	1.13
Rb/Sr	0.72	0.78	1.35	0.54	1.15	1.12	0.60	0.58	0.76	0.80	1.01	1.34
Rb/Nb	17.91	11.25	13.68	12.93	16.77	12.72	10.46	8.85	7.51	12.12	12.61	18.62

注:主量元素单位为%,稀土、微量元素单位为 10^{-6} ;F₁=-1.773ω(Ti₂O)+0.607ω(Al₂O₃)+0.76ω(Fe₂O₃)-1.5ω(MgO)+0.616ω(CaO)+0.509ω(Na₂O)-1.224ω(K₂O)-9.09;F₂=-0.445ω(Ti₂O)+0.07ω(Al₂O₃)-0.25ω(Fe₂O₃)-1.142ω(MgO)+0.438ω(CaO)+0.475ω(Na₂O)+1.426ω(K₂O)-6.861,据Roser and Korsch(1988)和李良(2017);“N”表示相对于球粒陨石标准化值,全文同;UCC据(Rudnick and Gao, 2014)。

(165 ± 2) Ma, 说明漠河组沉积年龄不早于 165 Ma, 与凝灰岩的年龄 157~158 Ma 接近。

5.2 地球化学特征

5.2.1 主量元素

漠河组砂岩样品主量元素测试结果(表 2)显示, 砂岩中 SiO_2 含量较高, 为 70.20%~76.13%, 平均含量为 73.16%, Al_2O_3 含量为 12.87%~14.39%, 平均含量为 13.54%, $\text{SiO}_2/\text{Al}_2\text{O}_3$ 值为 4.99~5.87, 平均值为 5.41, K_2O 含量为 2.44%~5.23%, Na_2O 含量为 2.81%~4.31%, $\text{K}_2\text{O}/\text{Na}_2\text{O}$ 值为 0.59~1.62, 平均值为 1.05。 Fe_2O_3 含量为 0.28%~1.26%、 FeO 含量 0.38%~2.27%、 MgO 含量为 0.41%~1.30%, 含量明显较低。结合镜下特征漠河组源区岩石主要以长英质岩石为主, 少量为黑云母等, 各氧化物的含量与活动大陆边缘环境的碎屑岩接近。

5.2.2 稀土元素

稀土元素(REE)具非迁移性, 沉积岩中稀土元素含量主要受源区岩石的元素丰度及源区风化条件的控制, 沉积过程和沉积后成岩过程对稀土元素的含量改变比较微小(Bhatia, 1985; Taylor et al., 1985; Gu et al., 2002; 杨志新等, 2018)。因此, 可以用沉积岩中稀土元素特征来反映源区岩石的性质。

漠河组砂岩 ΣREE 值在 97.64×10^{-6} ~ 222.95×10^{-6} , 平均值为 129.59×10^{-6} , 与平均大陆上地壳的 146.4×10^{-6} 相当(Taylor et al., 1985), 总体表现为轻稀土富集, 重稀土亏损的特征, LREE/HREE 值在 2.88~4.22, 平均值为 3.44, $(\text{La/Yb})_{\text{N}}$ 值在 8.43~11.91, 平均值为 10.22, 表明轻、重稀土分异明显, δEu 值为 0.50~0.82, 平均值为 0.67, 具中等偏弱的负 Eu 异常, Ce 异常不明显。稀土元素球粒陨石标准化图解上, 呈右倾模式, 重稀土元素分配曲线平坦, 在图 6a 中可见, 在样品的稀土曲线上轻稀土曲线与漠河盆地南部砂岩(和政军等, 2003)曲线相似, 重稀土含量高于漠河盆地砂岩, 与活动大陆边缘的稀土元素曲线吻合较好(图 6a)。样品全球平均大陆上地壳(UCC)标准化稀土元素配分曲线较为平缓, 具弱的轻稀土亏损, 重稀土富集的左倾模式, $(\text{La/Yb})_{\text{UCC}}$ 值在 0.76~1.07, 平均值为 0.92, 多数样品呈现出轻微的 Eu 负异常, $(\delta\text{Eu})_{\text{UCC}}$ 值为 0.71~1.16, 平均值为 0.96, 与全球平均大陆上地壳的稀土元素含量非常接近(图 7b)。

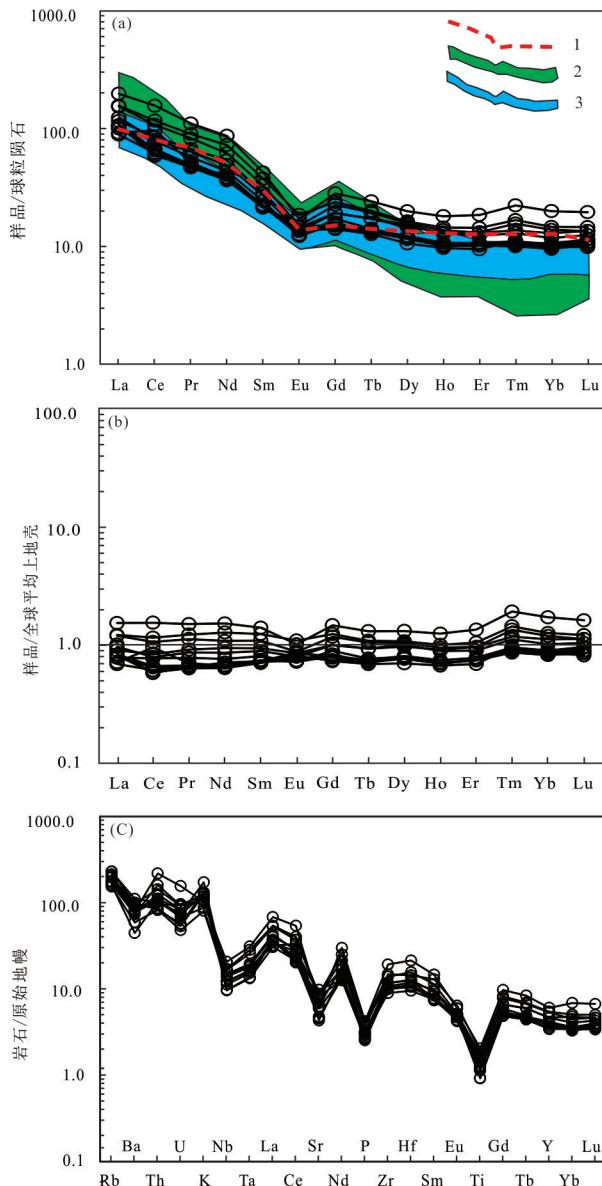


图 7 漠河盆地西缘漠河组砂岩球粒陨石(a)、全球平均大陆上地壳(b)标准化稀土元素配分曲线和原始地幔标准化微量元素蛛网图(c)(a、c据 Sun and McDonough, 1989; b据 Rudnick and Gao, 2014; d据和政军等, 2003)
1—活动大陆边缘砂岩的平均值;2—漠河盆地北部砂岩的平均值;
3—漠河盆地南部砂岩的平均值

Fig7 Standardized REE partition curves of sandstone chondrites (a), global average continental upper crust (b) and standard trace element cobweb maps of the primitive mantle on the western margin of Mohe Basin (c) (a, c after Sun and McDonough, 1989; b after Rudnick and Gao, 2014; d after He Zhengjun et al., 2003)

Average value of sandstone in active continental boundary;2—Average value of sandstone in the north of Mohe basin;3—Average value of sandstone in the southern of Mohe Basin

5.2.3 微量元素

漠河组砂岩微量元素含量见表2,在原始地幔标准化微量元素蛛网图中(图6c),各样品曲线高度一致,亏损高场强元素Nb、Ta、Ti、P,大离子亲石元素Ba和Sr,富集高场强元素Th、Hf,大离子亲石元素Rb、K,说明具有相同的物质来源,曲线与活动大陆边缘相似。Sr/Ba比值介于0.14~0.35,平均值为0.26,显示为陆源淡水沉积环境;V/Ni比值介于2.82~6.66,平均值4.97,V/(V+Ni)比值介于0.74~0.87,平均值为0.83,反映其为缺氧沉积环境(朱志军等,2010)。

6 讨 论

6.1 漠河组形成时代

漠河组的形成时代,乃至漠河盆地的形成时代一直以来都悬而未决。地质科学院1:100万地质图编图组建立漠河组时将其时代定为中或晚侏罗世;赵书跃等(2016)通过对漠河组内顺层侵入的糜棱岩化微晶黑云母石英闪长岩锆石U-Pb测年(182.8 Ma),并结合地层中化石组合,认为漠河组形成于(早侏罗世晚期—)中侏罗世;王少铁等(2015)对漠河组碎屑锆石测年,最小年龄为170 Ma,认为漠河组沉积下限不早于中侏罗世早期;肖传桃等(2015)对额木尔河群的古植物群进行了系统分析与研究,将漠河组划分为中侏罗世中晚期;吴河勇等(2003)通过对额木尔河群中的孢粉化石组合的研究,认为漠河组(原文称额木尔河组)应形成于晚侏罗世;郭知鑫(2018)通过系统的碎屑锆石测年分析研究,确定漠河组(原文称额木尔河组)沉积年龄为早白垩世(145~137 Ma)。但由于漠河盆地额木尔河群缺少标准化石,而碎屑锆石测年仅能确定地层形成的下限,因此,漠河组的形成时代一直未得到准确厘定,时代归属还存在争议。

对漠河盆地西部额尔古纳市恩和哈达镇漠河组内两个凝灰岩层进行锆石LA-ICP-MSU-Pb测年,锆石震荡环带结构发育,具岩浆锆石成因特征。年代学研究表明,两个凝灰岩样品加权平均年龄分别为(158±1)Ma、(157±2)Ma,与大兴安岭满克头鄂博组;火山岩年龄相当(刘凯等,2018;臧延庆等,2018;何鹏等,2020),该期岩浆事件与中国东北及邻区在晚侏罗世处于伸展拉张构造环境,发育与伸展性构造环境相关的中—酸性岩浆活动(郭知鑫,2018)相一致,说明

了该时期漠河盆地存在火山喷发活动,并形成了凝灰岩层。因此,158~157 Ma代表了漠河盆地西部漠河组形成的年龄,即晚侏罗世牛津阶。

6.2 物源区分析

稀土元素在指示沉积物源区中非常可靠(Cullers et al., 2000;宋春彦等,2013),其配分模式是判定盆地沉积岩源区性质的重要手段之一(McLennan, 1989;谢国梁等,2013;张建军等,2017)。漠河盆地西缘漠河组砂岩稀土总量平均值为 129.59×10^{-6} ,LREE/HREE平均值为3.44, δEu 值为0.50~0.82,平均值为0.67,具中等偏弱的负 δEu 异常,(La/Yb)_N值在8.43~11.91,平均值为10.22,在ΣREE-La/Yb判别图解上(图8c),样品多落于沉积岩和花岗岩区。稀土元素球粒陨石标准化图解上,呈右倾模式,重稀土元素分配曲线平坦,样品全球平均大陆上地壳(UCC)标准化稀土元素配分曲线较为平缓,具弱的轻稀土亏损,重稀土富集的左倾模式,稀土元素特征与全球平均大陆上地壳的稀土元素含量非常接近,表明漠河组砂岩物源区岩石主要来自于上地壳。沉积岩中微量元素La、Ce、Nd、Y、Th、Zr、Hf、Ti、Sc等活动性较弱,在风化、搬运和沉积过程中能定量转移到碎屑沉积物中,可以较好地反映源区母岩性质及形成时的构造环境(Bhatia, 1985;Gu et al., 2002;王丛山等,2016;杨志新等,2018)。La/Th比值在2.41~3.59,平均值为2.93,Hf含量为 2.96×10^{-6} ~ 6.64×10^{-6} ,平均值为 3.94×10^{-6} ,在Hf-La/Th图解上(图8d),样品多落于长英质岩石和酸性岛弧物源区。样品的Rb/Sr比值在0.54~1.35,平均值为0.90,高于地幔和下地壳相应元素的比值,而与上地壳平均值大致一致。在图La/Yb-La/Th图解中(图8a)显示物源区母岩与平均上地壳一致。在主量元素F1~F2判别方程图解(图8b)中显示,漠河组砂岩碎屑物来源于石英质沉积岩区和酸性火山岩区。综上所述,结合镜下特征,漠河盆地西缘漠河组碎屑岩的碎屑物母岩主要为长英质砂岩和花岗岩类。

前人对沉积相和古水流方向研究,判断漠河组砂岩物源主要来自盆地东南部(和钟铧等,2008;侯伟等,2010),通过对漠河组碎屑锆石U-Pb年龄的研究,可以反映碎屑物来源和基底信息,7组碎屑锆石年龄与盆地南缘(额尔古纳地块、兴安地块)出露的地质体时代吻合较好。第一组1019~1429 Ma年

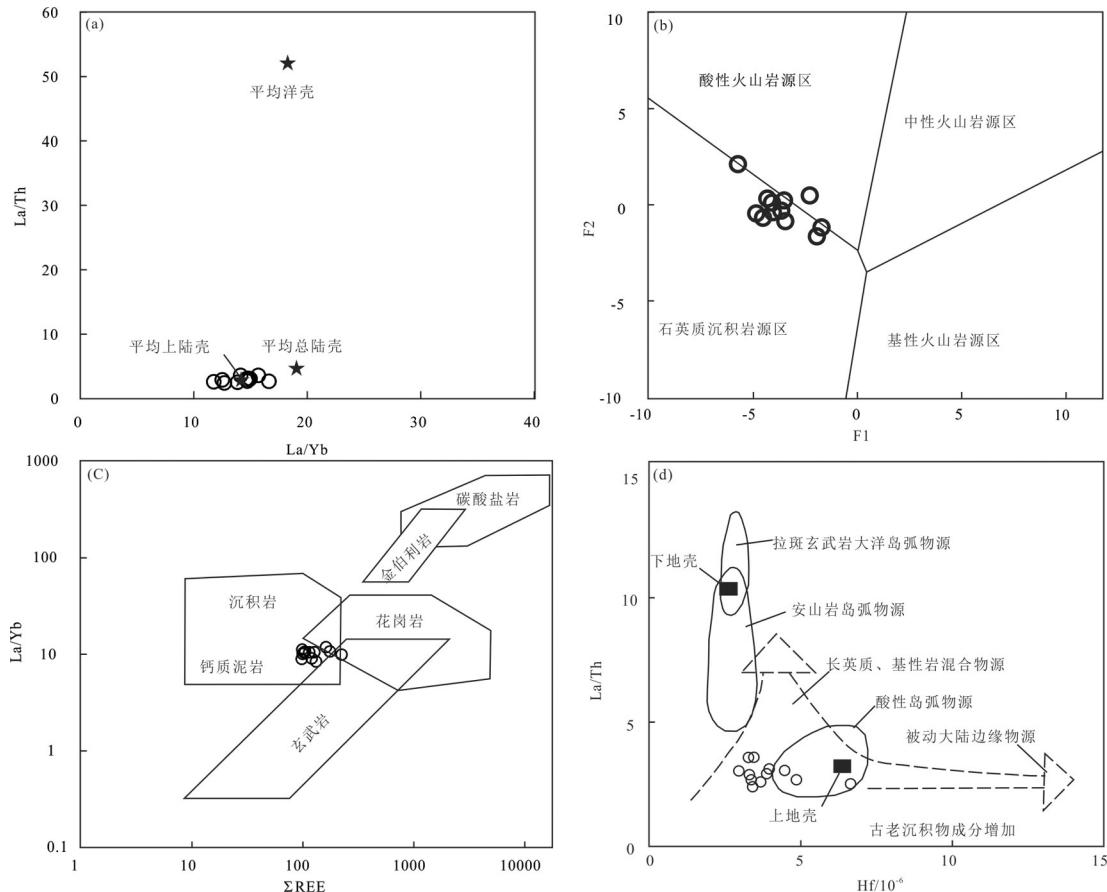


图8 漠河盆地西缘漠河组砂岩La/Yb-La/Th(a,据Shao et al.,2001)、F1-F2(b,据Roser et al.,1988)、ΣREE-La/Yb(c,据Allègre et al.,1978)、Hf-La/Th判别图解(d,据Floyd et al.,1987)

Fig8 Sandstone La/Th-La/Yb(a,after Shao et al.,2001), F1-F2(b,after Roser et al.,1988), REE-La/Yb(c,after Allègre et al.,1978), Hf-La/Th of Mohe Formation in the western margin of Mohe basin(d,after Floyd et al.,1987)

龄,存在变质锆石,年龄与盆地基底古老碎屑锆石相近(苗来成等,2007;表尚虎等,2012);第二组696~972 Ma年龄,与盆地南缘佳疙瘩组碎屑锆石年龄(变安山岩单颗粒锆石U-Pb蒸发法年龄 (723 ± 8) Ma(郭灵俊等,2005),碎屑锆石U-Pb加权平均年龄 (846.2 ± 3.9) Ma和 (721 ± 8) Ma,待发表)和额尔古纳北段新元古代花岗质侵入岩年龄相当(余宏全等,2012;赵硕等,2016a;赵硕等,2016b);第三组455~526 Ma年龄,与早寒武世—晚奥陶世花岗质岩石(葛文春等,2007;柴明春等,2018;杜兵盈等,2019;吴琼等,2019)、中基性火山岩(汪岩等,2017)和额尔古纳地块兴华渡口群年龄相对应(苗来成等,2007;赵院冬等2017;李锦轶等,2019);第四组303 Ma,与兴安地块和松嫩地块晚石炭世花岗质岩石年龄相一致(郑常青等,2009;崔芳华等,2013;赵院冬等2017;贾旭等,2020;周传芳等,2020);第五

组240~283 Ma年龄,与漠河盆地南缘晚二叠世二长花岗岩年龄($272\sim252$ Ma,加权平均年龄257 Ma,未发表)和大兴安岭中北段林西组(刘兵等,2014;秦涛等,2018)、老龙头组地层碎屑锆石年龄相当(刘兵等,2014);第六组199~204 Ma和第七组165~192 Ma年龄,与漠河盆地南缘晚三叠世—早中侏罗世花岗质岩石年龄相对应(李研等,2017;胡军海等,2018;李良等,2018;于跃江等,2019)。 $500\sim526$ Ma、 $455\sim487$ Ma、 $240\sim283$ Ma、 303 Ma、 $165\sim204$ Ma与(Wu et al.,2011)统计的中国东北显生宙花岗岩年龄谱系相一致。因此,漠河组沉积物主要来自盆地南缘额尔古纳、兴安等地块,物源区既有古老陆壳物质,也有年轻的火成岩。

6.3 物源区大地构造环境

不同构造环境形成的沉积岩的碎屑物质组成和地球化学成分不同,可据此判定沉积物源区的构

表3 不同大地构造环境砂岩的化学成分含量(%)平均值对比
Table 3 Comparison of average chemical composition (%) of sandstones in different geotectonic environments

	大洋岛弧	大陆岛弧	活动大陆边缘	被动大陆边缘	上地壳	研究区
SiO ₂	58.83	70.69	73.86	81.95	66	73.16
Al ₂ O ₃	17.11	14.04	12.89	8.41	15.20	13.54
Fe ₂ O ₃	1.95	1.43	1.30	1.32		0.80
CaO	5.83	2.68	2.48	1.89	4.20	0.90
MgO	3.65	1.97	1.23	1.39	2.20	0.94
Na ₂ O	4.10	3.12	2.77	1.07	3.90	3.51
K ₂ O	1.60	1.89	2.90	1.71	3.40	3.58
TiO ₂	1.06	0.64	0.46	0.49	0.50	0.30
Fe ₂ O ₃ +MgO	11.73	6.79	4.63	2.89	7.15	1.74
Al ₂ O ₃ /SiO ₂	0.29	0.20	0.18	0.10	0.23	0.19
K ₂ O/Na ₂ O	0.39	0.61	0.99	1.60	0.87	1.05
La	8.20	27	37	39	30	28.87
Ce	19.40	59	78	85	64	53.09
ΣREE	58	146	186	210	146	129.58
δEu	1.04	0.80	0.60	0.55	0.65	0.67
La/Yb	4.20	11	12.5	15.90	13.60	14.25
(La/Yb) _n	2.80	7.50	8.30	10.80	9.20	10.22
LREE/HREE	3.80	7.70	9.10	8.50	9.47	3.32
Sc	19.50	14.80	8	6	11	4.93
La/Th	4.26	2.36	1.77	2.20	2.80	2.93
La/Y	0.48	1.02	1.33	1.31	1.36	1.42
La/Sc	0.55	1.82	4.55	6.25	2.73	5.85

注:大洋岛弧、大陆岛弧、活动大陆边缘和被动大陆边缘据文献(Roser and Korsch.,1985; Bhatia and Crook, 1986);上地壳据文献(Taylor and McLennan,1985)。

造背景(陈小双等,2018)。盆地构造环境分为大洋岛弧、大陆岛弧、活动大陆边缘和被动大陆边缘4种类型(Bhatia,1984; Roser and Korsch,1986; 冯兴雷

等,2016)。漠河组砂岩主量元素与典型大地构造背景下的砂岩主量元素相比较其Al₂O₃、K₂O、Na₂O、TiO₂、Al₂O₃/SiO₂与上地壳砂岩主量元素较为相似(表3),SiO₂、Al₂O₃、Fe₂O₃、CaO、MgO、Na₂O、TiO₂、Al₂O₃/SiO₂、K₂O/Na₂O与活动大陆边缘和大陆岛弧吻合较好(表3)。在砂岩主量元素构造环境判别图解(Fe₂O₃+MgO)-(Al₂O₃/SiO₂)(图9a)和(Fe₂O₃+MgO)-TiO₂(图9b)上,样品全部落于活动大陆边缘。微量元素对构造背景的判别,主要是运用沉积岩中La、Th、Sc、Co、Zr等元素的比值或相关性构造环境判别(冯兴雷等,2016),在微量元素构造判别图上(图10),样品多落于活动大陆边缘,少量样品落于大陆岛弧和被动大陆边缘。

漠河盆地处于蒙古—鄂霍茨克造山带东南缘。蒙古—鄂霍茨克造山带自西向东呈剪刀式关闭,西部于晚三叠世开始闭合(余宏全等,2012;唐杰等,2018),而东部最终在晚侏罗世(Zorin,1999; Sorokin et al.,2004)或早白垩世初闭合(郭知鑫,2018)。蒙古—鄂霍茨克洋具南北双向俯冲的构造背景得到了越来越多的报道(Tomurtogoo et al.,2005; Orolmaa et al.,2008; 陈志广等,2010; Wu et al., 2011; 许文良等,2013; Tang et al., 2014; 唐杰等,2018)。额尔古纳地块晚三叠世—早侏罗世花岗岩类(流纹岩)岩石地球化学表明其具有活动大陆边缘性质(李研等,2017; 胡军海等,2018; 李良等,2018; 于跃江等,2019),这与该期岩浆活动形

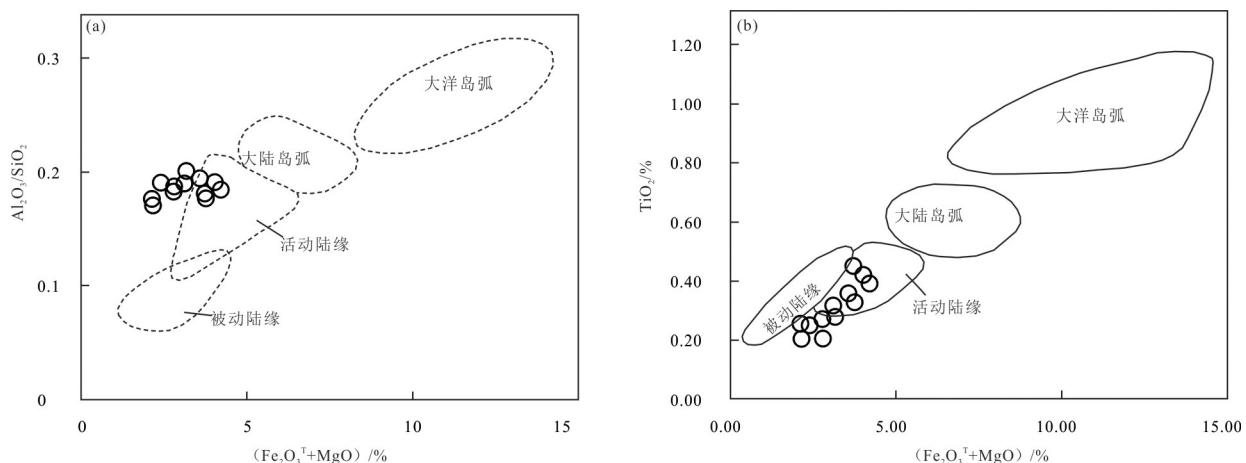


图9 漠河盆地西缘漠河组砂岩主量元素(Fe₂O₃+MgO)-Al₂O₃/SiO₂(a)、(Fe₂O₃+MgO)-TiO₂(b)构造环境判别图
(据Roser et al.,1985)

Fig9 Discriminant diagram of major elements (Fe₂O₃+MgO)-Al₂O₃/SiO₂(a)、(Fe₂O₃+MgO)-TiO₂ (b)tectonic environment of the Mohe Formation sandstone in western margin of Mohe basin (after Roser et al.,1985)

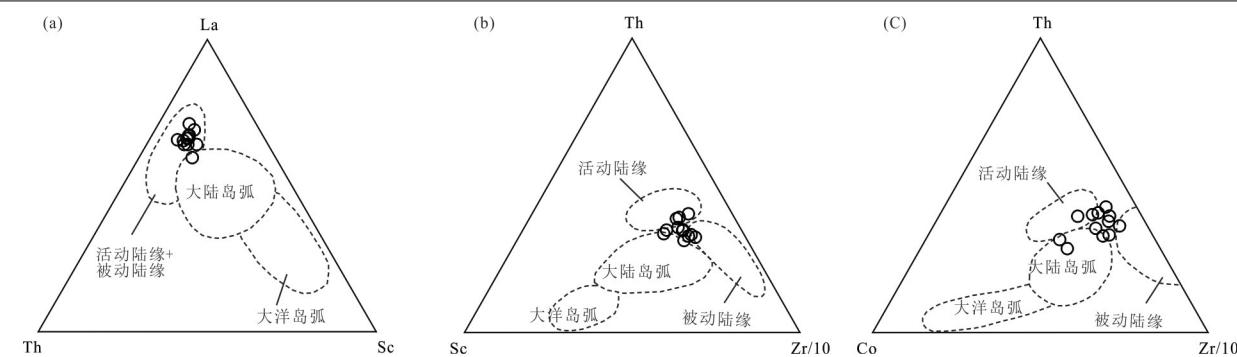


图10 漠河盆地西缘漠河组砂岩物源区微量元素La–Th–Sc、Th–Sc–Zr/10、Th–Co–Zr/10构造环境判别图解
(据Bhatia et al., 1986)

Fig10 La–Th–Sc、Th–Sc–Zr/10、Th–Co–Zr/10 diagrams for tectonic environment discriminant of Mohe Formation sandstone, western Margin of Mohe Basin (after Bhatia et al., 1986)

成于蒙古—鄂霍茨克洋板块向南俯冲于额尔古纳地块之下的活动大陆边缘环境相吻合(唐杰等, 2018)。

根据漠河组砂岩的碎屑物质组成、主量和微量元素地球化学研究,结合区域构造背景,认为漠河组砂岩主要物源区的大地构造背景以活动大陆边缘为主。

7 结 论

通过对洛古河地区漠河盆地漠河组凝灰岩层锆石U–Pb测年、砂岩地球化学特征和物源分析研究,可以得出以下结论:

(1) 漠河盆地西缘漠河组凝灰岩夹层锆石均为岩浆锆石,2个样品加权平均年龄为 (158 ± 1) Ma、 (157 ± 2) Ma。漠河盆地西缘漠河组形成时间为158~157 Ma,为晚侏罗世牛津阶。

(2) 漠河组砂岩的母岩主要来源于盆地南缘广泛分布的前中生代沉积–变质基底(兴华渡口岩群、佳疙瘩组、额尔古纳河组等)和各期中酸性火成岩岩石(寒武—奥陶纪、二叠纪、晚三叠—早侏罗世等)。

(3) 漠河组物源区大地构造背景以活动大陆边缘为主。漠河组砂岩形成于晚侏罗世蒙古—鄂霍茨克洋闭合造山过程中,漠河盆地南侧物源区快速剥蚀、快速搬运并快速沉积的环境。

致谢:原武警黄金第三支队总工程师李向文在成文过程给予悉心指导,怀宝峰、马靖、刘宁波、郑吉林、王大可、宋昊南、李贺参与了野外工作,中国地质调查局自然资源调查指挥中心王梁博士对本文进行指导,

匿名审稿专家及编辑老师对本文提出了许多建设性的意见和建议,在此一并表示衷心感谢!

注释:

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