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# 鲁西莱芜三岔河矽卡岩型铁金矿床岩体成因:锆石 U-Pb年代学和岩石地球化学证据

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**提要:**【研究目的】莱芜三岔河矿床是鲁西地区首次发现的矽卡岩型铁金矿床(不含铜),以往研究相对薄弱,本文旨在对成矿岩体形成时代、源区性质、岩石成因和构造背景进行讨论。【研究方法】本文开展了LA-ICP-MS锆石U-Pb测年、岩体主微量测试工作。【研究结果】结果显示辉石闪长岩侵位年龄为( $138.4\pm1.2$ )Ma,属早白垩世。辉石闪长岩里特曼指数 $\sigma=2.20\sim2.54$ ,为钙碱性岩石;辉石闪长岩具高Mg<sup>#</sup>、富Na特征,Sc、Cr、Co、Ni含量较高,表明其具有幔源岩浆的属性;富集轻稀土元素(LREE),亏损重稀土元素(HREE),Eu异常不明显,富集大离子亲石元素(Cs、Ba、Sr),亏损高场强元素(Nb、Ta、Zr),暗示有地壳物质的涉入。【研究结论】三岔河辉石闪长岩可能起源于受华北克拉通古老地壳物质改造的富集岩石圈地幔的部分熔融,形成于岩石圈减薄的强烈伸展构造背景,可能与太平洋板块俯冲后撤引起的伸展构造环境有关。

**关 键 词:**铁金矿;矽卡岩型;岩体成因;U-Pb年代学;岩石地球化学;矿产勘查工程;莱芜;山东

**创 新 点:**(1)首次报道了三岔河辉石闪长岩侵位年龄为( $138.4\pm1.2$ )Ma;(2)岩体起源于富集岩石圈地幔的部分熔融,同时岩浆源区受到了华北克拉通古老地壳物质的改造。

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## Petrogenesis of the Sanchahe skarn type iron-gold deposit in Laiwu area, western Shandong: Evidences from zircon U-Pb geochronology and geochemistry

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**Abstract:** This paper is the result of mineral exploration engineering.

**[Objective]** Laiwu Sanchahe iron–gold deposit is the first skarn type iron–gold deposit (without copper) discovered in western Shandong. Previous research on the Sanchahe deposit is weak. The purpose of this paper is to discuss the age of ore-forming pluton, magma sources, petrogenesis and tectonic setting. **[Methods]** We carried out LA-ICP-MS zircon U-Pb dating, major and trace elements analysis. **[Results]** Zircon U-Pb age of pyroxene diorite is  $(138.4 \pm 1.2)$  Ma in the Early Cretaceous. Pyroxene diorite is calc-alkalic (Rittmann index  $\sigma = 2.20\text{--}2.54$ ) and characterized by high Mg<sup>#</sup> and high contents of Na<sub>2</sub>O, Sc, Cr, Co and Ni, which demonstrates that it has geochemical affinity with mantle-derived magma. It is enriched in LREE and depleted in HREE, with no obvious Eu anomaly, and shows enrichment in Cs, Ba, Sr and depletition in high field strength elements (Nb, Ta, Zr), indicating involvement of some crustal materials. **[Conclusions]** Geochemical features of pyroxene diorite combined with geochronological data imply that it may be generated by partial melting of the enriched lithospheric mantle which had been metasomatized by the ancient crust materials of the North China Craton, and formed in the strong extensional tectonic background of lithospheric thinning, which may be caused by the retreat of the subducting Paleo-Pacific plate.

**Key words:** iron–gold deposit; skarn type; petrogenesis; U–Pb dating; geochemistry; mineral exploration engineering; Laiwu; Shandong Province

**Highlights:** (1) It is reported for the first time that the age of Sanchahe pyroxene diorite is  $(138.4 \pm 1.2)$  Ma. (2) pyroxene diorite may be generated by partial melting of enriched lithospheric mantle, which had been metasomatized by the ancient crust materials of the North China Craton.

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## 1 引言

山东是黄金大省,胶东地区发现了多个大型、超大型金矿床,是中国最重要的金矿资源基地(宋明春等,2022)。鲁西地区与胶东地区同处沂沭断裂带,具有相似的构造背景和岩体形成时代,具有一定找金潜力,区内发现了多处金矿床,但数量和规模远远不及胶东地区。鲁西地区金矿床具有规模小、类型多、分散分布的特点。三岔河矿床为矽卡岩型铁金同体共生矿床(不含铜),查明金矿石金属量6941 kg,平均品位4.54 g/t,矿床规模为中型,是鲁西地区找新发现,对山东尤其鲁西地区金矿勘查具有一定指示作用。

20世纪80年代以来,矽卡岩型金矿床的勘查和研究工作在国内外均取得了较大进展,发现了一大批矽卡岩型金矿床,贡献了1000 t以上的黄金产量(陈衍景等,1997;赵一鸣等,2017)。据不完全统计,我国已探明矽卡岩型金矿储量约700 t,约占中国黄金总储量的15%,但已发现的矽卡岩型金矿床中多与铜元素共生,铁金共生矿床(不含铜)极为少见,因此三岔河矿床具有重要研究价值。矽卡岩型矿床多形成于中基性–中酸性侵入岩与碳酸盐岩地层的接触带及其附近,含矿高温热液与围岩发生接触交代作用形成(赵一鸣等,1997;彭聪等,1999),矿化类型与岩浆岩的性质有着密切的联系,因此探讨岩体的成因和来源对了解矿床成因和指导下一

步矿产勘查工作具有重要意义。众多学者从不同角度对鲁西岩体进行了研究,认为岩浆起源于上地幔物质的部分熔融,并受到地壳物质的混染作用,但对地壳物质的来源还存在多种认识,主要观点有:①拆沉的华北下地壳(许文良等,2003;杨承海等,2007; Lan et al., 2013);②扬子地壳物质的部分加入(巫祥阳等,2003;徐义刚等,2007;③古太平洋俯冲的洋壳物质(陈立辉等,2003;宁培松等,2013)。三岔河矿床以往研究程度较低,缺少对矿床形成时代、岩体地球化学性质等方面的研究,本文简要介绍三岔河铁金矿床地质特征,对与成矿关系密切的岩体进行详细的岩石学研究,报道岩体锆石U-Pb年龄和岩石地球化学特征,结合以往资料对岩体形成时代、源区性质、岩石成因和构造背景进行了讨论,以期约束三岔河成矿岩体的成因机

制,为鲁西相似地段寻找该类型金矿提供参考。

## 2 地质背景和矿床特征

### 2.1 区域地质背景

鲁西地区位于华北克拉通东南部,兰考—聊城断裂以东、齐河—广饶断裂以南、郯庐断裂以西、丰沛断裂以北,东南与苏鲁超高压变质带相接(Yang et al., 2005; Zhao et al., 2005; 张超等,2017; 杨浩田等,2018)(图1a,b)。研究区位于鲁西地块中部,泰莱凹陷东南部,结晶基底为新太古界泰山岩群花岗岩类,盖层主要为古生界寒武系—奥陶系灰岩、中生界碎屑岩以及新生界沉积物。区内构造以断裂为主,其展布方向主要为NNW向、NE向及近EW向3组,NNW向断裂是区域上的主干断裂,铜冶店—孙祖断裂规模最大,该断裂具有多期活动的特点,

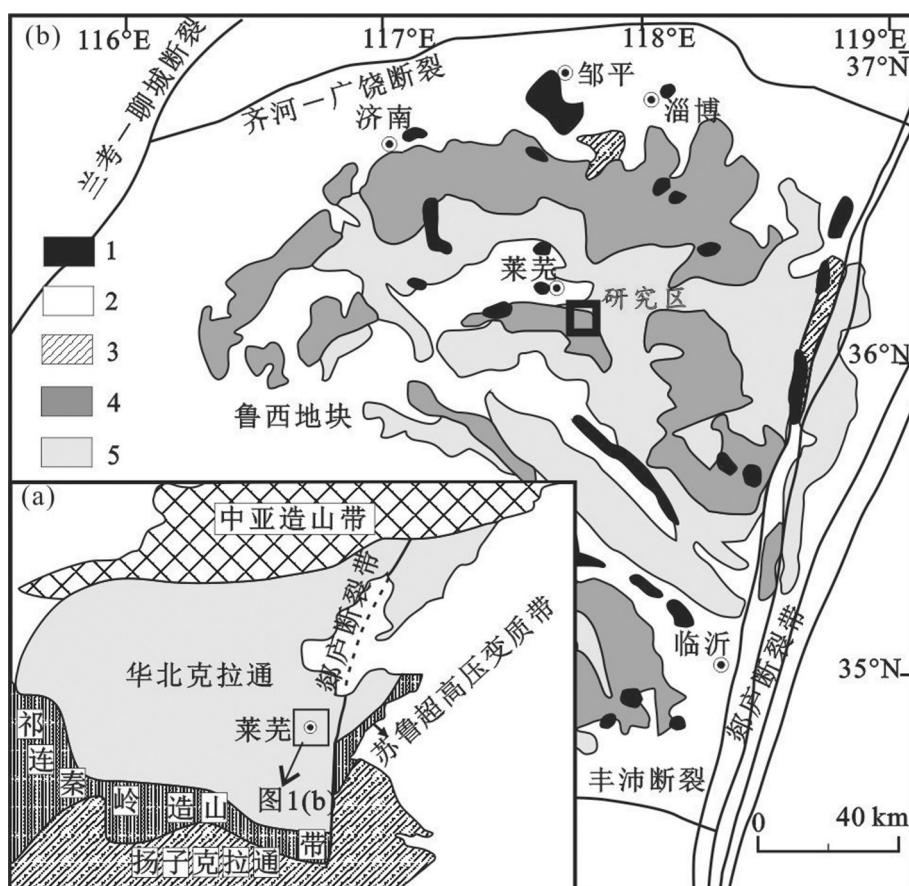


图1 华北克拉通地质简图(a)和鲁西地块地质简图(b)(a据Zhang et al., 2005;b据Lan et al., 2011修改)  
1—中生代岩浆岩;2—新生界;3—中生界;4—古生界;5—前寒武纪基底

Fig.1 Geological map of the North China Craton (a) and geological map of the Luxi terrane(b) (a after Zhang et al., 2005; b modified from Lan et al., 2011)

1—Mesozoic magmatic rocks; 2—Cenozoic; 3—Mesozoic; 4—Paleozoic; 5—Precambrian basement

为岩浆的上侵提供了空间和通道。区内岩浆活动较为强烈,主要有新太古代晚期花岗岩和燕山晚期铁铜沟杂岩体,铁铜沟杂岩体与成矿关系密切。铁铜沟杂岩体位于莱芜东南铁铜沟一带,产于铜冶店—孙祖断裂与青泥沟断裂交汇部位,岩体大致呈“十”字型展布,出露面积约7.5 km<sup>2</sup>(图2)。铁铜沟杂岩体由早到晚岩性依次为铁铜沟单元二长花岗岩、铜汉庄单元石英闪长玢岩、靳家桥单元角闪长玢岩、西杜单元细粒辉石闪长岩、东明生单元中粒辉石闪长岩、凤凰峪中粒角闪石岩、药山橄榄苏长辉长岩,是一个由中基性向中酸性演化的系列。

## 2.2 矿床地质特征

三岔河矿床位于铁铜沟杂岩体的北部,矿体隐伏于第四系之下(图2),与成矿关系密切的东明生单元辉石闪长岩为铁铜沟杂岩体的组成部分,是铁铜沟杂岩体早期岩浆活动的产物。由于岩浆沿区域性断裂上侵,大量灰岩被吞食、断陷、倒置,形成捕掳体接触带构造,在捕掳体附近,侵入体与围岩发生强烈的化学反应,形成矽卡岩型铁金矿床。

矿床由一个矿体组成,主矿种为金,铁矿体与其同体共生,伴生银和硫。矿体赋存于辉石闪长岩内的奥陶系灰岩捕掳体附近,赋矿岩性主要为矿化矽卡岩,其次为蚀变辉石闪长岩。矿体形态产状较稳定,呈似层状、透镜状产出,局部形态受捕掳体的形态发生变化,矿体走向近南北,倾向东,倾角平均值为33°。金矿体平均真厚度4.65 m,中部矿体较厚,两侧矿体较薄,出现尖灭现象,平均品位4.54×10<sup>-6</sup>。铁矿体被金矿体包裹(图3),平均厚度2.81 m,中部矿体较厚,两侧矿体较薄,出现尖灭现象,矿体内无夹石。矿区内的矿石类型主要为磁铁矿化矽卡岩(磁性弱,只含金)和黄铁矿化磁铁矿化矽卡岩(金铁共生),矿石矿物主要为磁铁矿、黄铁矿、自然金,脉石矿物主要为透辉石、透闪石、蛇纹石、金云母等矽卡岩矿物。

## 3 样品描述及分析方法

### 3.1 样品描述

锆石U-Pb年龄样采自ZK0003钻孔,全岩主微量元素测试样品采自ZK803、ZK0004、ZK0005钻孔,岩性皆为辉石闪长岩,样品新鲜,未受风化。岩石(图

4)为灰黑色,粒状结构,块状构造,晶形较好,主要由斜长石、普通辉石组成,其次为黑云母、普通角闪石、石英、磷灰石和金属矿物等,镜下可见多条石膏细脉穿插分布。斜长石(55%~60%),为自形一半自形板状,聚片双晶发育,局部可见环带结构,晶形较完整,表面具泥化蚀变而显得浑浊不净,粒径一般为0.4~1.2 mm;普通辉石(25%~30%),局部可见两组近相互垂直的解理,可见普通角闪石的反应边结构,粒径一般为0.2~1.2 mm;黑云母(5%~10%),褐色,为半自形柱状、鳞片状,粒径一般为0.3~1.0 mm;普通角闪石(±5%),为半自形柱状,常常以普通辉石的反应边形式存在,粒径一般为0.2~0.8 mm;石英(<5%),为他形粒状,表面光滑干净,充填在上述矿物之间,含量较少。

### 3.2 LA-ICP-MS测年

将辉石闪长岩人工粉碎后,采用重选及电磁选分离出锆石,在双目镜下挑选出晶形和透明度较好的锆石颗粒,将其用双面胶黏在载玻片上并套上靶环,注入环氧树脂后,冷却和打磨并露出锆石表面,用于透射光、反射光和阴极发光(CL)图像的采集。锆石U-Pb测年使用193 nm ArF准分子激光剥蚀系统(型号为New Wave)和电感耦合等离子质谱仪(ICP-MS,型号为ICAP RQ)。实验中采用He作为剥蚀物质的载气,分析时采用8 Hz的激光频率,能量为5 J/cm<sup>2</sup>,束斑直径为32 μm,激光剥蚀时间为40 s。测试结果通过ICPMsDataCal软件离线处理完成,年龄计算时以国际标准锆石91500为外标进行同位素比值校正,元素含量以国际标样NIST 610为外标,以Si为内标标定锆石中的元素含量。详细的实验分析步骤和数据处理方法见文献(Yuan et al., 2004; Liu et al., 2010)。样品加工和分析测试在广州拓岩检测技术有限公司进行。

### 3.3 岩石地球化学测试

主量元素测试方法为X射线荧光光谱法(XRF),所用的仪器为荷兰PANalytical生产的 Axios;微量元素采用德国生产的电感耦合等离子质谱仪(ICP-MS,型号为ICAP RQ)进行分析测试。主量元素的分析精度和准确度优于5%,微量元素的分析精度和准确度为5%~10%。样品加工和分析测试在广州拓岩检测技术有限公司进行。

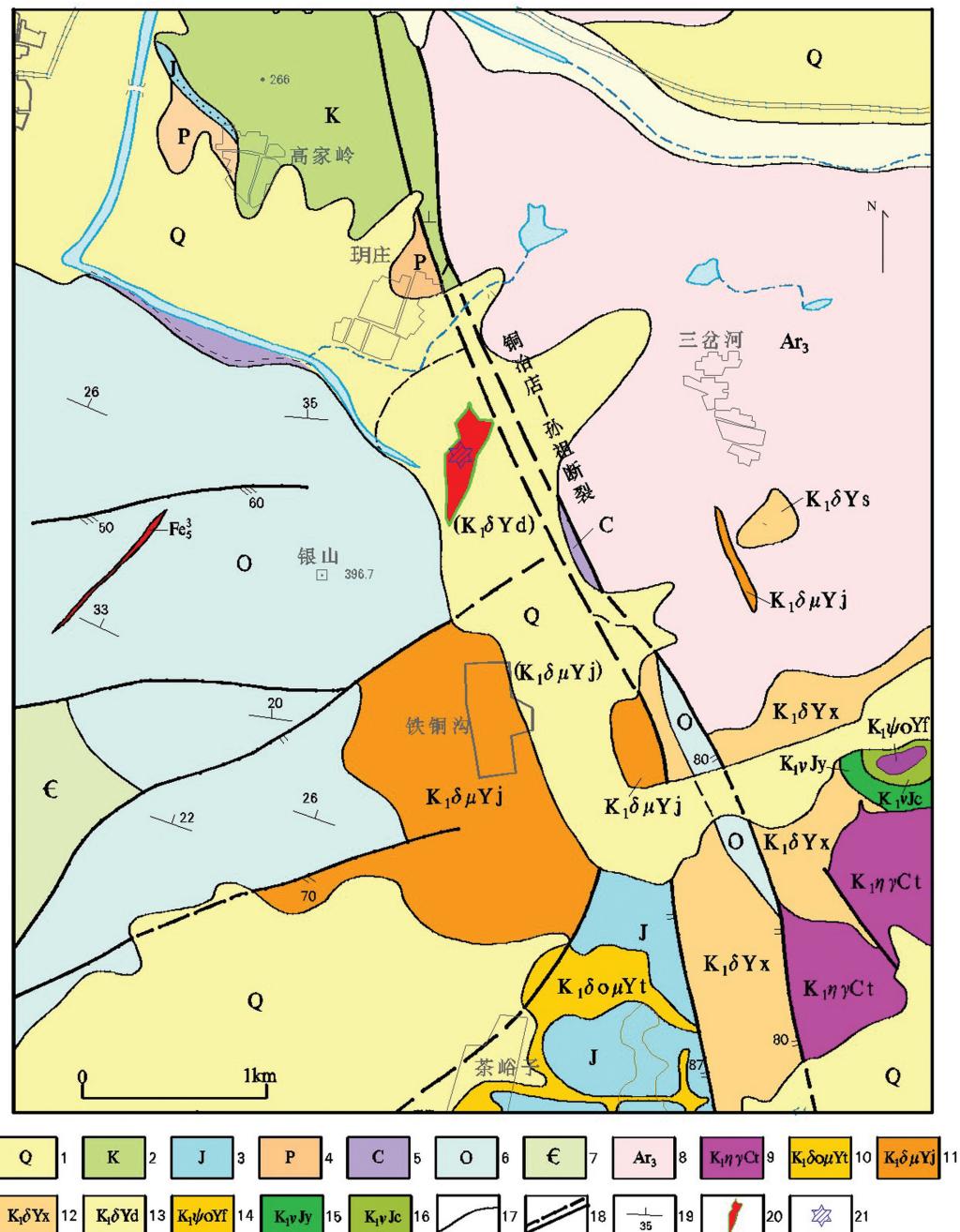


图2 三岔河矿区地质简图

1—第四系;2—白垩系;3—侏罗系;4—二叠系;5—石炭系;6—奥陶系;7—寒武系;8—新太古代花岗岩;9—铁铜沟单元二长花岗岩;10—铜汉庄单元石英闪长玢岩;11—靳家桥单元角闪闪长玢岩;12—西杜单元细粒辉石闪长岩;13—东明生单元中粒辉石闪长岩;14—凤凰峪中粒角闪岩;15—药山橄榄苏长辉长岩;16—茶叶山角闪苏长辉长岩;17—地质界线;18—断层;19—地层产状;20—矿体范围;21—取样位置

Fig.2 Geological sketch of Sanchahe mining area

1—Quaternary; 2—Cretaceous; 3—Jurassic; 4—Permian; 5—Carboniferous; 6—Ordovician; 7—Cambrian; 8—Neoarchean granite; 9—Monzogranite of Tietonggou unit; 10—Quartz diorite porphyry of Tonghanzhuang unit; 11—Hornblende diorite porphyry of Jinjiaqiao unit; 12—Finegrained pyroxene diorite of Xidu unit; 13—Medium grained pyroxene diorite of Dongmingsheng unit; 14—Medium grained hornblendite of Fenghuangyu unit; 15—Olivine norite gabbro of Yaoshan unit; 16—Hornblende norite gabbro of Chayeshan unit; 17—Geological boundary; 18—Fault; 19—Attitude of strata; 20—Orebody range; 21—Sampling location

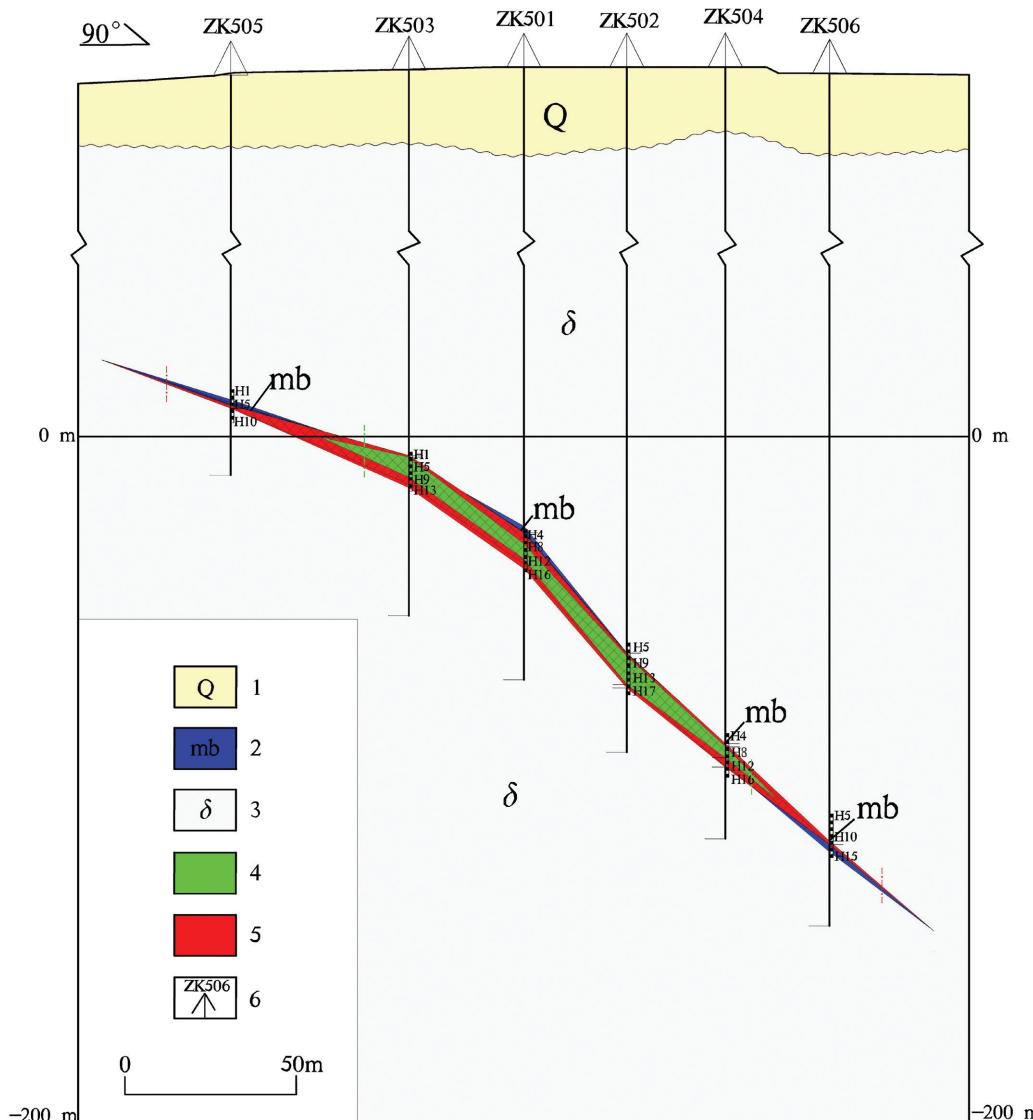


图3 三岔河矿区5勘查线地质剖面图

1—第四系;2—大理岩;3—辉石闪长岩;4—铁金矿体;5—金矿体;6—钻孔位置及孔号

Fig.3 Geological profile of No.5 exploration line in Sanchahe mining area

1—Quaternary; 2—Marble; 3—Pyroxene diorite; 4—Iron gold ore body; 5—Gold ore body; 6—Position and number of drill bore

## 4 分析结果

### 4.1 锆石U-Pb年龄

三岔河辉石闪长岩锆石阴极发光图像(图5)显示,辉石闪长岩中锆石晶型较好,多为长柱状、短柱状,少数粒状,颗粒大小不均,70~270 μm,表现出典型的岩浆锆石环带结构(吴元保等,2004;李长民,2009)。这些锆石Th、U含量较高,分别为 $187.83 \times 10^{-6}$ ~ $633.29 \times 10^{-6}$ 、 $176.42 \times 10^{-6}$ ~ $456.30 \times 10^{-6}$ ,变化范围较小,Th/U比值为0.99~1.43,表明锆石均具有岩

浆成因锆石的特征(Hoskin and Black, 2000;雷玮瑛等, 2013)。本次挑选25粒锆石进行LA-ICP-MS测年,结果见表1。7个锆石样品谐和度小于90%,未参与最终年龄计算,18个谐和度大于90%样品 $^{206}\text{Pb}/^{238}\text{U}$ 年龄加权平均值为( $138.4 \pm 1.2$ ) Ma (MSWD=1.10)(图6),属早白垩世,锆石年龄可以代表三岔河辉石闪长岩岩体的结晶年龄。

### 4.2 元素地球化学组成

#### 4.2.1 主量元素

三岔河铁金矿区辉石闪长岩地球化学分析结

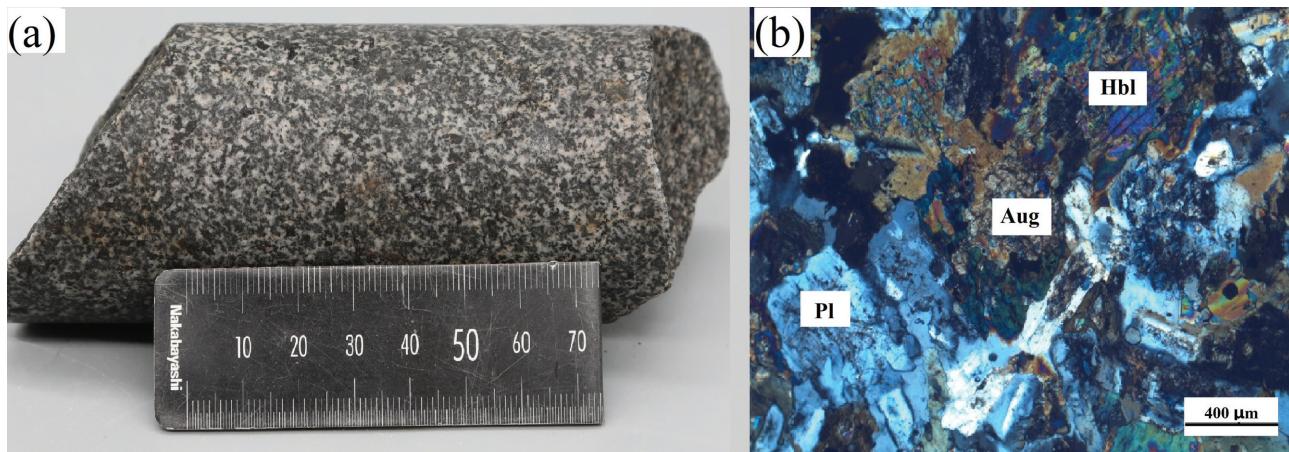


图4 三岔河辉石闪长岩岩心照片(a)和显微镜下照片(b)

Pl—斜长石; Hbl—角闪石; Aug—辉石

Fig.4 Microphotographs of the Sanchahe pyroxene diorite

Pl—Plagioclase; Hbl—Hornblende; Aug—Augite

果见表2。辉石闪长岩中  $\text{SiO}_2$  含量为 55.32%~58.73%，表现中性岩浆岩特征； $\text{Al}_2\text{O}_3$  为 14.60%~15.31%，具有高铝质特点； $\text{Na}_2\text{O}$ 、 $\text{K}_2\text{O}$  分别为 3.40%~3.52% 和 2.07%~2.36%， $\text{Na}_2\text{O}/\text{K}_2\text{O}$  值为 1.49~1.66，显示相对富 Na 的特征；在 TAS 图解中(图 7a)中，样品落在了闪长岩范围内；在  $\text{K}_2\text{O}-\text{SiO}_2$  图解(图 7b)中，样品落入了高钾钙碱性系列；样品的里特曼指数  $\sigma=2.20\sim2.54$ ，小于 3.30，具有钙碱性系列岩浆演化趋势； $\text{MgO}$  含量 5.89%~6.80%， $\text{Mg}^#$  为 59.66~62.83，具高 Mg 闪长岩的特点；A/CNK 为 0.75~0.76，为准铝质岩石。

#### 4.2.2 微量元素

辉石闪长岩样品在球粒陨石标准化的稀土元

素配分曲线上(图 8a)，显示富集轻稀土元素，相对亏损重稀土元素，稀土元素总量( $\Sigma \text{REE}$ )为  $118.05 \times 10^{-6}\sim127.81 \times 10^{-6}$ ，轻稀土总量( $\Sigma \text{LREE}$ )为  $90.54 \times 10^{-6}\sim102.08 \times 10^{-6}$ ，重稀土总量( $\Sigma \text{HREE}$ )为  $25.73 \times 10^{-6}\sim27.52 \times 10^{-6}$ ， $(\text{La/Yb})_{\text{N}}=10.29\sim11.84$ ，轻重稀土分馏程度较高，具有较平坦的 HREE 配分模式， $\delta \text{Eu}$  为 0.88~0.98，具有微弱的 Eu 负异常，与鲁西沂南、金岭等地区同时代岩体具有相似的稀土配分模式(王永等, 2011；金子梁, 2017；张超等, 2017)。在原始地幔标准化蛛网图(图 8b)中，样品富集大离子亲石元素(Rb、Cs、Ba、Sr)，相对亏损高场强元素(Nb、Ta、Zr)，暗示了有地壳物质的涉入；样品具有较高的 Rb/Sr 值(0.04~0.1)和较低的 Sm/Nd 值(0.2)也反映

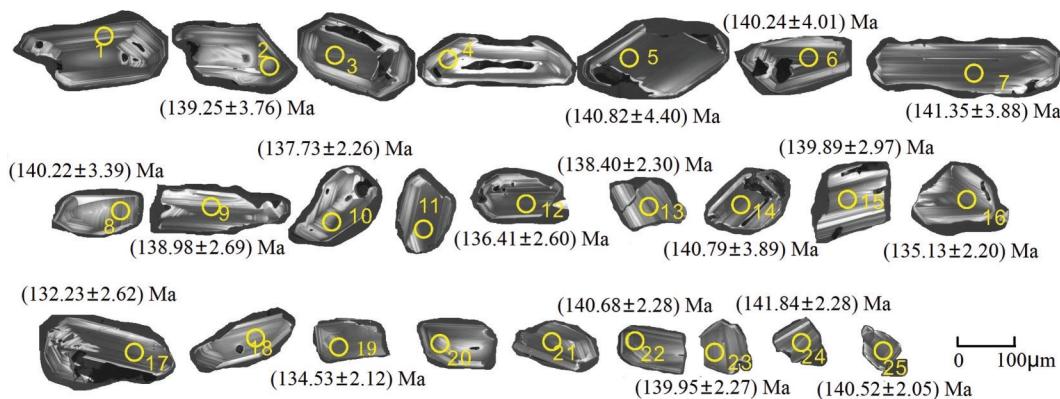


图5 三岔河辉石闪长岩锆石阴极发光图像

Fig.5 Cathodoluminescence images of zircons of the Sanchahe pyroxene diorite

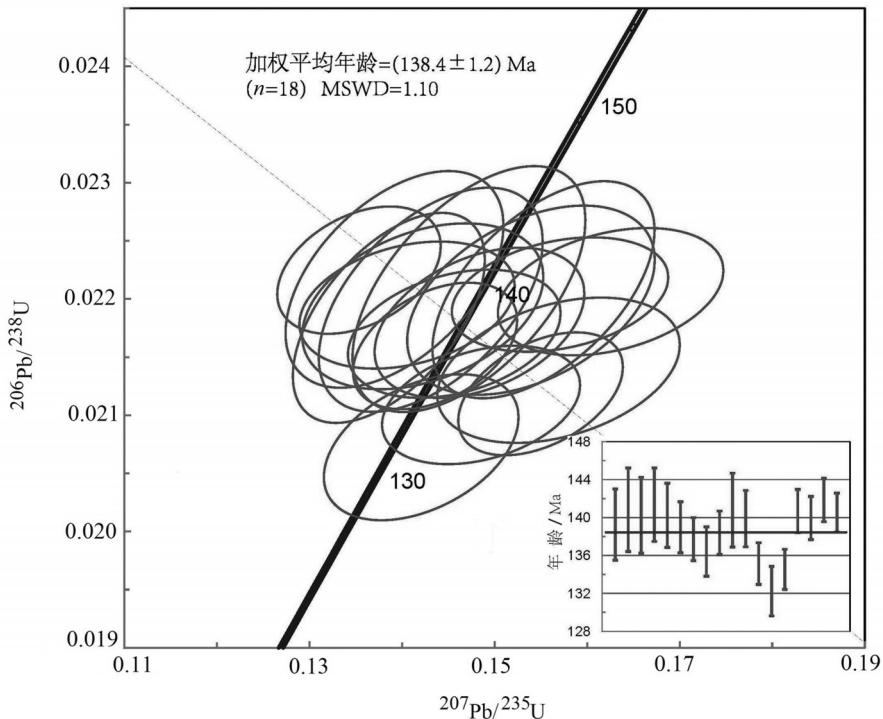


图6 三岔河辉石闪长岩锆石LA-ICP-MS U-Pb年龄谐和图  
Fig. 6 Concordia diagrams showing LA-ICP-MS zircon U-Pb dating of the Sanchahe pyroxene diorite

了地壳物质的参与。样品具有较高的Sr含量和较低的Yb含量, Sr为 $668 \times 10^{-6} \sim 853 \times 10^{-6}$ , Yb为 $1.31 \times 10^{-6} \sim 1.35 \times 10^{-6}$ , 具有典型埃达克岩的地球化学特征; 样品具有较高的Sc、Cr、Co、Ni含量, Sc为 $18.60 \times 10^{-6} \sim 24.5 \times 10^{-6}$ , Cr为 $289 \times 10^{-6} \sim 334 \times 10^{-6}$ , Co为 $24.2 \times 10^{-6} \sim 27.7 \times 10^{-6}$ , Ni为 $93.4 \times 10^{-6} \sim 104 \times 10^{-6}$ , 暗示其岩浆源区为地幔; Nb/Ta值为14.64~16.72, Zr/Hf值为38.39~87.28, Ba/Rb值为12.38~21.91, 其相关比值也暗示辉石闪长岩具有地幔橄榄岩部分熔融的岩浆成分特征。

## 5 讨 论

### 5.1 成岩成矿时代

三岔河铁金矿床为典型的矽卡岩型矿床, 成矿与三岔河辉石闪长岩关系密切, 侵入体与围岩地层发生强烈的化学反应, 促使Si、Al、Ca、Mg等物质成分交换, 形成矽卡岩矿床, 岩体形成时代大致可代表矿床形成时代。三岔河辉石闪长岩为隐伏岩体, 位于铁铜沟杂岩体的北部, 为铁铜沟杂岩体的组成部分。以往对铁铜沟杂岩体研究工作较多, 《山东省区域地质志》根据岩体与其他地质体的接触关

系, 即岩体侵入侏罗系三台组, 将杂岩体时代定为早白垩世; 许文良等(2004a)测得辉石闪长岩中黑云母的 $^{40}\text{Ar}/^{39}\text{Ar}$ 年龄为 $(132.79 \pm 0.27)$ Ma; 王冬艳等(2004)测得二辉闪长岩中黑云母的 $^{40}\text{Ar}/^{39}\text{Ar}$ 年龄 $(132.82 \pm 0.5)$ Ma; 杨承海等(2007)通过LA-ICP-MS锆石U-Pb方法对岩体侵位时代进行了测定, 其加权平均年龄为 $(134.5 \pm 2.3)$ Ma, 属于早白垩世。本文对莱芜三岔河辉石闪长岩体进行了LA-ICP-MS锆石U-Pb测年, 锆石的阴极发光图像、结构特征及Th/U比值显示其为典型的岩浆成因, 样品加权平均年龄为 $(138.4 \pm 1.2)$ Ma, 为早白垩世早期, 所测得的年龄可以代表岩体的形成时代, 鉴于矽卡岩型矿床成岩-成矿作用在时间上密切关系, 推测三岔河铁金矿床的成矿年龄应稍稍晚于成岩年龄。以上分析测试结果与杨承海等(2007)所测得的结果基本一致, 并具有相对偏老的特点, 同时与鲁西济南、莱芜、金岭、沂南、临朐铁寨、沂源等地区岩体形成时代测试结果相近(巫祥阳等, 2003; 刘金民等, 2006; 徐义刚等, 2007; 董树义, 2008; 杨承海等, 2008; 胡芳芳等, 2010; 李洪奎等, 2011; 王永等, 2011; 宇培松等, 2013; 田京祥等, 2015; 王浩等, 2015; Xie et al.,

表1 三岔河辉石闪长岩LA-ICP-MS锆石U-Pb同位素分析结果  
Table 1 Analyzed results of LA-ICP-MS zircons U-Pb of the Sanchaha pyroxene diorite

测点	含量/ $10^{-6}$		同位素比值						表观年龄/Ma			谐和度				
	Th	U	$^{207}\text{Pb}^{206}\text{Pb}$	$1\sigma$	$^{207}\text{Pb}^{205}\text{U}$	$1\sigma$	$^{206}\text{Pb}^{208}\text{U}$	$1\sigma$	$^{207}\text{Pb}^{206}\text{Pb}$	$1\sigma$	$^{207}\text{Pb}^{235}\text{U}$	$1\sigma$				
ZK0003B4@1*	265.27	278.43	0.95	0.0705	0.0042	0.2054	0.0110	0.0219	0.0006	942.59	122.23	189.67	9.23	139.86	3.54	69%
ZK0003B4@2	403.38	334.21	1.21	0.0479	0.0024	0.1386	0.0068	0.0218	0.0006	94.54	118.50	131.77	6.10	139.25	3.76	94%
ZK0003B4@3*	98.95	116.15	0.85	0.1075	0.0120	0.6463	0.1896	0.0378	0.0040	1766.67	206.64	506.18	116.94	239.33	25.07	28%
ZK0003B4@4*	145.93	159.77	0.91	0.0909	0.0066	0.2808	0.0197	0.0230	0.0008	1444.14	138.58	251.29	15.64	146.48	4.86	47%
ZK0003B4@5	276.83	236.77	1.17	0.0467	0.0028	0.1472	0.0083	0.0221	0.0007	31.58	137.02	139.46	7.35	140.82	4.40	99%
ZK0003B4@6	434.65	355.11	1.22	0.0469	0.0026	0.1431	0.0080	0.0220	0.0006	55.65	112.95	135.80	7.09	140.24	4.01	96%
ZK0003B4@7	363.57	311.31	1.17	0.0454	0.0023	0.1392	0.0078	0.0222	0.0006	—	—	132.32	6.93	141.35	3.88	93%
ZK0003B4@8	187.83	176.42	1.06	0.0507	0.0030	0.1539	0.0088	0.0220	0.0005	227.85	139.80	145.38	7.77	140.22	3.39	96%
ZK0003B4@9	248.19	231.50	1.07	0.0507	0.0031	0.1484	0.0075	0.0218	0.0004	233.40	140.72	140.46	6.61	138.98	2.69	98%
ZK0003B4@10	633.29	456.30	1.39	0.0486	0.0023	0.1436	0.0058	0.0216	0.0004	127.87	112.95	136.22	5.19	137.73	2.26	98%
ZK0003B4@11*	48.30	90.38	0.53	0.5532	0.0231	6.4256	0.4872	0.0715	0.0037	4389.82	61.05	2035.78	66.63	445.04	22.32	-29%
ZK0003B4@12	434.89	323.89	1.34	0.0533	0.0029	0.1569	0.0086	0.0214	0.0004	342.65	124.06	148.00	7.59	136.41	2.60	91%
ZK0003B4@13	332.32	297.40	1.12	0.0501	0.0030	0.1456	0.0076	0.0217	0.0004	198.23	145.35	138.06	6.72	138.40	2.30	99%
ZK0003B4@14	580.70	404.94	1.43	0.0532	0.0030	0.1553	0.0079	0.0221	0.0006	344.50	125.91	146.57	6.97	140.79	3.89	95%
ZK0003B4@15	206.02	189.29	1.09	0.0475	0.0031	0.1422	0.0086	0.0219	0.0005	72.32	148.13	135.04	7.65	139.89	2.97	96%
ZK0003B4@16	587.52	442.49	1.33	0.0540	0.0023	0.1550	0.0059	0.0212	0.0003	372.28	96.29	146.28	5.16	135.13	2.20	92%
ZK0003B4@17	540.94	428.90	1.26	0.0504	0.0025	0.1421	0.0069	0.0207	0.0004	216.74	116.65	134.88	6.17	132.23	2.62	98%
ZK0003B4@18*	103.98	98.11	1.06	0.2683	0.0215	1.4524	0.1351	0.0307	0.0012	3295.99	126.54	910.86	55.93	194.67	7.24	-30%
ZK0003B4@19	341.89	392.16	0.87	0.0483	0.0028	0.1485	0.0071	0.0211	0.0003	122.31	120.35	140.62	6.24	134.53	2.12	95%
ZK0003B4@20*	238.67	1.10	0.0819	0.0042	0.2442	0.0118	0.0217	0.0004	1242.59	100.46	221.88	9.63	138.17	2.34	53%	
ZK0003B4@21*	196.81	141.80	1.39	0.1766	0.0123	0.7495	0.0609	0.0280	0.0012	2621.30	115.59	567.95	35.32	177.77	7.39	-5%
ZK0003B4@22	423.66	359.00	1.18	0.0553	0.0026	0.1625	0.0080	0.0221	0.0004	338.95	111.10	152.92	7.02	140.68	2.28	91%
ZK0003B4@23	249.39	251.53	0.99	0.0465	0.0027	0.1405	0.0077	0.0219	0.0004	33.43	133.32	133.53	6.84	139.95	2.27	95%
ZK0003B4@24	425.52	329.49	1.29	0.0442	0.0021	0.1353	0.0058	0.0222	0.0004	—	—	128.89	5.21	141.84	2.28	90%
ZK0003B4@25	407.55	379.74	1.07	0.0517	0.0027	0.1571	0.0077	0.0220	0.0003	272.29	118.50	148.14	6.77	140.52	2.05	94%

注: \*为谐和度较低的锆石测点。

表2 三岔河辉石闪长岩主量元素(%)、微量元素和稀土元素( $10^{-6}$ )化学分析结果

Table 2 Chemical compositions of major elements, trace elements and REE elements of the Sanchahe pyroxene diorite

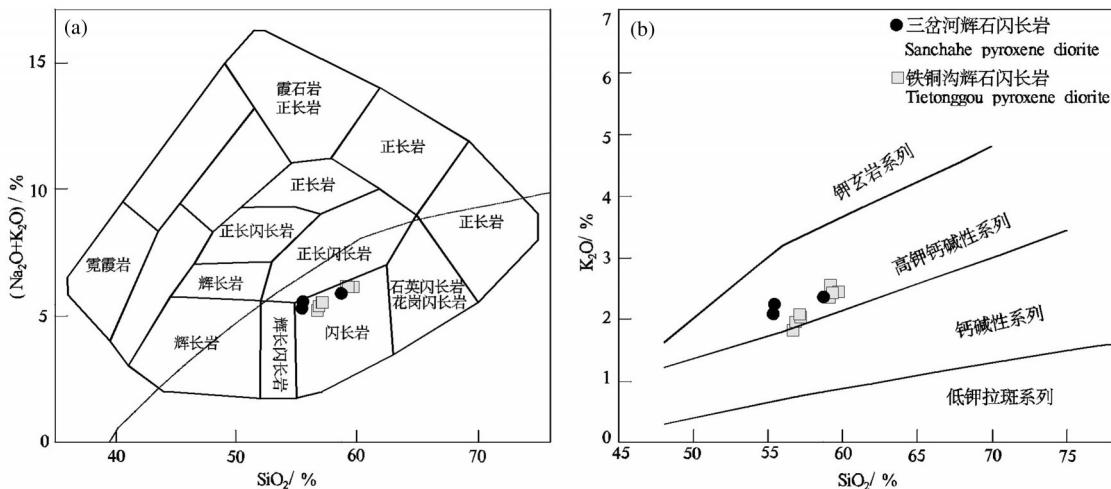
分析项目	ZK803	ZK0004	ZK0005	分析项目	ZK803	ZK0004	ZK0005
SiO <sub>2</sub>	58.73	55.41	55.32	Mo	1.20	3.19	2.59
TiO <sub>2</sub>	0.70	0.71	0.71	Cd	0.22	0.03	0.03
Al <sub>2</sub> O <sub>3</sub>	14.60	15.31	15.14	Sn	1.62	3.07	2.99
Fe <sub>2</sub> O <sub>3</sub>	6.97	8.75	8.90	Sb	0.57	0.33	0.35
MnO	0.12	0.12	0.12	Cs	3.59	2.02	1.94
MgO	5.89	6.47	6.80	Ba	842.00	816.00	802.00
CaO	5.92	6.70	6.64	La	23.00	20.60	21.20
Na <sub>2</sub> O	3.52	3.40	3.43	Ce	46.20	35.80	37.60
K <sub>2</sub> O	2.36	2.22	2.07	Pr	5.54	5.44	5.57
P <sub>2</sub> O <sub>5</sub>	0.21	0.27	0.27	Nd	21.80	22.90	23.80
烧失量	0.86	0.93	1.20	Sm	4.34	4.49	4.66
Total	99.88	100.28	100.59	Eu	1.20	1.33	1.35
K <sub>2</sub> O+Na <sub>2</sub> O	5.88	5.62	5.50	Gd	3.99	3.59	3.72
A/CNK	0.76	0.75	0.75	Tb	0.53	0.58	0.61
Mg <sup>#</sup>	62.83	59.66	60.45	Dy	2.73	3.11	3.11
Na <sub>2</sub> O/K <sub>2</sub> O	1.49	1.53	1.66	Ho	0.52	0.57	0.57
$\sigma$	2.20	2.54	2.46	Er	1.43	1.64	1.61
Li	23.80	31.40	28.70	Tm	0.21	0.21	0.21
P	963.00	1170.00	1157.00	Yb	1.31	1.35	1.34
Sc	18.60	24.50	23.90	Lu	0.20	0.21	0.20
Ti	3940.00	4286.00	4227.00	Hf	3.36	2.30	2.28
V	148.00	164.00	161.00	Ta	0.40	0.36	0.31
Cr	334.00	289.00	289.00	Pb	16.40	12.99	12.59
Mn	792.00	911.00	908.00	Th	6.26	2.80	2.90
Co	24.20	27.70	27.50	U	1.81	0.89	0.89
Ni	104.00	97.60	93.40	$\Sigma$ LREE	102.08	90.53	94.14
Cu	20.50	7.75	11.54	$\Sigma$ HREE	25.73	27.52	25.79
Zn	78.10	27.10	29.10	$\Sigma$ REE	127.81	118.05	119.93
Ga	20.50	15.60	16.30	Rb/Sr	0.10	0.05	0.04
Ge	1.60	0.99	0.89	Ba/Rb	12.38	21.14	21.91
Rb	68.00	38.60	36.60	Zr/Hf	38.39	77.83	87.28
Sr	668.00	853.00	844.00	La/Sm	5.30	4.59	4.55
Y	14.80	14.50	14.40	Nb/Ta	16.72	14.64	15.03
Zr	129.00	179.00	199.00	(La/Yb) <sub>N</sub>	11.84	10.29	10.67
Nb	6.72	5.27	4.66	$\delta$ Eu	0.88	0.98	0.96

2015;杨浩田等,2018),也与华北克拉通岩石圈减薄及其发生强烈岩浆活动的时间相同,应属于华北克拉通岩石圈减薄过程中岩浆活动产物。

## 5.2 岩浆源区性质及岩石成因

莱芜三岔河辉石闪长岩为铁铜沟杂岩体的组成部分,是铁铜沟杂岩体早期岩浆活动的产物,铁铜沟杂岩体的地球化学特征可为探讨三岔河辉石

闪长岩的岩浆源区性质和岩石成因提供参考。三岔河辉石闪长岩具有高镁(MgO=5.89%~6.80%)和高的Mg<sup>#</sup>值(59.66~62.83)以及富钠(Na<sub>2</sub>O=3.40%~3.52%)特征,结合相对高的Cr、Sc、Co、Ni含量,表明它们具有地幔的源区性质(Rapp and Watson, 1995)。此外,样品的Ba/Rb值为12.38~21.91接近于原始地幔值(11.00)(McDonough and Sun, 1995);

图7 三岔河辉石闪长岩硅碱图解(a)和SiO<sub>2</sub>-K<sub>2</sub>O图解(b)

(a,底图据Cox et al., 1979; b,底图据Rickwood, 1989;铁铜沟数据据Yang et al., 2005)

Fig.7 TAS (a) and K<sub>2</sub>O-SiO<sub>2</sub> (b) diagram of the Sanchahe pyroxene diorite

(a, after Cox et al., 1979; b, after Rickwood, 1989; Tietonggou data after Yang et al., 2005)

Nb/Ta值为14.64~16.72,接近于幔源岩浆的Nb/Ta值(17.50),而明显高于壳源岩浆的Nb/Ta值(11.00)(Green, 1995);Zr/Hf值为38.39~87.28明显高于大陆地壳值(11.00),而接近于原始地幔值(36.27)(刘永江等,2016)。此外,铁铜沟岩体中含有丰富的橄榄岩包体也证明了岩浆源区为上地幔(Xu et al., 2003;王冬艳等,2004;杨承海等,2007)。综上,笔者认为三岔河辉石闪长岩岩浆源区为上地幔。

莱芜三岔河辉石闪长岩富集大离子亲石元素

(LILE)和轻稀土元素(LREE),亏损高场强元素(HFSE)和重稀土元素(HREE),具有较高的稀土丰度( $\Sigma$  REE=118.09~127.81),且样品有较高的Rb/Sr值(0.04~0.1)和较低的Sm/Nd(0.2)值,表明岩浆源区或演化过程中有地壳物质的参与;此外,铁铜沟岩体具有相对较高的初始<sup>87</sup>Sr/<sup>86</sup>Sr值(0.70672~0.70715)和相对较低的ε<sub>Na</sub>(t)值(Yang et al., 2005),也表明该类岩浆有地壳物质的参与,同时排除了俯冲大洋板片熔融物质对岩石圈地幔的交代,因为后

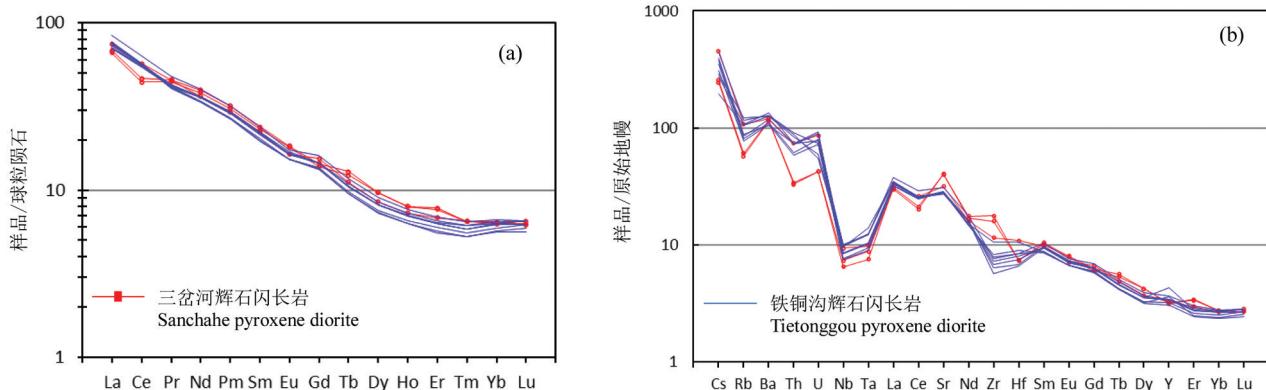


图8 三岔河辉石闪长岩稀土元素球粒陨石标准化配分图(a)和微量元素地幔标准化蛛网图(b)

(a底图据Boynton, 1984;b底图据Sun and McDonough, 1989;铁铜沟数据据Yang et al., 2005)

Fig.8 Chondrite normalized rare earth element pattern diagram (a) and primitive mantle normalized spider diagram (b) of the Sanchahe pyroxene diorite

(a after Boynton, 1984; b after Sun and McDonough, 1989; Tietonggou data after Yang et al., 2005 )

者形成的岩浆具有相对较低的初始 $^{87}\text{Sr}/^{86}\text{Sr}$ 值和相对较高的 $\varepsilon_{\text{Nd}}(t)$ 值(Defant and Drummond, 1990)。同时,铁铜沟岩体Nd模式年龄位于1555~1987 Ma,且存在少量具有新太古代年龄的浑圆状锆石(杨承海等,2007),这些特征暗示在原始岩浆起源或演化过程中有古老地壳物质的涉入。

### 5.3 构造背景

三岔河辉石闪长岩形成于早白垩世早期( $138.4 \pm 1.2$ ) Ma,这与华北克拉通大规模分布的早白垩世岩浆活动相一致,暗示它们具有相同的形成环境。鲁西地区早白垩世形成了大量的火山岩和侵入岩,早白垩世的岩浆活动在华北克拉通东部主要表现为双峰式的火成岩组合(许文良等,2004b),表明它们形成于伸展的构造背景,这与中国东部广泛分布的变质核杂岩、拉张的沉积盆地、A型花岗岩和碱性岩的形成时间相对应(王德滋等,1995; Wang et al., 2005; Wu et al., 2007; Wang et al., 2015; 李春林等,2020; 朱强等,2021),同时与华北克拉通东部岩石圈减薄的峰期时间相吻合(许文良等,2004a; 魏巍等,2020)。综上,莱芜三岔河辉石闪长岩形成于华北克拉通东部早白垩世岩石圈强烈减薄的伸展构造背景,而这一时期恰恰是古太平洋板块向欧亚大陆斜向俯冲的时期(邓军等,2006; 翟明国,2020),可能为太平洋板块俯冲后撤引起的伸展构造环境。

## 6 结 论

(1) 莱芜三岔河辉石闪长岩岩体侵位时间为早白垩世( $138.4 \pm 1.2$ ) Ma,与鲁西广泛分布的闪长质岩石侵位年龄基本一致。铁金矿床形成时间稍晚于岩体形成时间。

(2) 莱芜三岔河辉石闪长岩具有高Mg<sup>#</sup>值、富Na,富集轻稀土元素和大离子亲石元素,亏损重稀土元素和高强场元素特征,结合铁铜沟岩体Sr-Nd同位素特征,暗示它们起源于富集岩石圈地幔的部分熔融,同时岩浆源区受到了华北克拉通古老地壳物质的改造。

(3) 莱芜三岔河辉石闪长岩形成于华北克拉通东部早白垩世时期岩石圈强烈伸展的构造背景,而这一时期恰恰是古太平洋板块向欧亚大陆斜向俯冲的时期,可能为太平洋板块俯冲后撤引起的伸展构造环境。

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