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鄂尔多斯盆地延川东地区本溪组储层特征与成岩过程

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Reservoir characteristics and diagenetic processes of the Benxi Formation in the East Yanchuan Block, Ordos Basin

Abstract: [Objective] The Benxi Formation in the East Yanchuan Block of the southeastern Ordos Basin has emerged as a significant exploration target for tight gas reservoirs. Despite its strategic importance, the sandstone reservoirs in this interval are typically characterized by low porosity, poor permeability, strong heterogeneity, and complex diagenetic histories. These characteristics limit their development potential and make accurate reservoir prediction challenging. This study aims to thoroughly investigate the reservoir characteristics and diagenetic evolution of the Benxi Formation sandstones, with the goal of clarifying how multi-stage diagenetic processes control pore evolution and reservoir quality differentiation. [Methods] An integrated analytical approach was applied, including casting thin section petrography, scanning electron microscopy (SEM), and high-pressure mercury intrusion (HPMI). These techniques were used to examine pore types, mineral assemblages, diagenetic alterations, and pore-throat distributions, allowing for a detailed reconstruction of the diagenetic evolution and its influence on petrophysical properties. [Results] The study reveals: (1) The Benxi Formation sandstones have undergone a multi-stage diagenetic evolution involving three principal phases: mechanical compaction, cementation enhancement, and dissolution modification. Each stage had a distinct influence on the pore structure and reservoir performance. (2) Quartz-rich sandstones showed relatively high resistance to mechanical compaction but experienced substantial porosity loss during the middle diagenetic stage due to the pervasive precipitation of quartz and carbonate cements. This stage led to the formation of a rigid framework, where the intergranular pores were severely occluded, significantly reducing effective porosity and permeability. (3) Lithic quartz sandstones exhibited a much more favorable diagenetic trajectory. These rocks experienced strong dissolution of unstable components, such as feldspar and lithic fragments, resulting in the formation of secondary intragranular and intergranular pores. This process markedly improved pore connectivity and storage capacity, making this lithofacies the most favorable for tight gas accumulation. (4) Lithic sandstones suffered from intense early-stage compaction and cementation, which drastically reduced both primary and secondary porosity. As a result, these rocks exhibit extremely poor reservoir quality and are considered the least favorable reservoir type. (5) Based on the observed differences in diagenetic intensity and pore evolution characteristics, four distinct diagenetic facies types were identified. [Conclusion] (1) The reservoir quality of the Benxi Formation sandstones is primarily controlled by the intensity and sequence of compaction, cementation, and dissolution. (2) Different

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mineral compositions lead to divergent diagenetic pathways and contrasting pore evolution patterns. (3) Lithic quartz sandstones hold the highest potential due to extensive dissolution, while quartz sandstones are more susceptible to cementation. (4) Lithic sandstones, dominated by early-stage densification, represent poor-quality reservoirs. (5) The proposed four-type diagenetic facies framework effectively reflects the reservoir heterogeneity and offers practical criteria for facies-based reservoir prediction. [Significance] This study advances the understanding of diagenetic control on pore structure in tight sandstone reservoirs and provides a robust geological foundation for classifying and predicting reservoirs and targeting sweet spots in future exploration efforts.

Keywords: tight sandstone; diagenesis; diagenetic facies; Benxi Formation; Ordos Basin

摘 要: 延川东地区本溪组是鄂尔多斯盆地东南部天然气勘探的重要目标层系, 储层物性较差、非均质性强且成岩演化过程复杂。文章以本溪组砂岩为研究对象, 基于铸体薄片、扫描电镜及高压压汞等多种实验手段, 系统分析了其储层特征及成岩作用差异, 明确了多阶段成岩过程对储层质量的控制作用。研究表明, 本溪组砂岩储层演化主要经历了压实致密、胶结加固与溶蚀改造3个阶段, 不同矿物组成砂岩在储集性能与成岩演化路径上存在显著差异: 石英砂岩抗压能力强但中成岩阶段易致密化; 岩屑石英砂岩溶蚀强烈, 是储层成岩改造潜力最大的岩石类型; 岩屑砂岩早期致密化强烈, 仅在局部发育数量有限次生孔隙。在此基础上, 建立了由弱压实弱胶结相、中等压实中等胶结相、胶结致密相及胶结溶蚀复合相组成的4类成岩相划分体系, 揭示了储层物性差异的微观形成机制。研究成果深化了致密砂岩成岩作用阶段性与非均质性成因的认识; 同时, 成岩相划分体系为致密砂岩储层分类评价与有利成岩相带的预测提供了可靠的地质依据。

关键词: 致密砂岩; 成岩作用; 成岩相; 本溪组; 鄂尔多斯盆地

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0 引言

近年来, 致密砂岩气藏作为非常规天然气资源的重要组成部分, 已成为中国油气勘探开发的重点方向(贾承造等, 2012; 李建忠等, 2025)。相较于常规储层, 致密砂岩普遍具有孔隙度低、渗透性差、非均质性强等特点, 其储集性能不仅取决于原始沉积条件, 更深受后期成岩作用的持续改造影响(刘雨航等, 2025; 刘志达等, 2025; 岳怀海等, 2025)。压实、胶结与溶蚀等成岩过程在不同演化阶段交替主导, 直接控制有效储集空间的保存程度, 并间接影响储层的渗流能力与改造潜力(邵鑫笛等, 2025; 王子龙等, 2025)。因此, 厘清致密砂岩储层的成岩演化机制, 识别主控成岩作用类型, 是实现储层质量精细评价的基础。在此基础上, 建立反映成岩过程—孔隙结构—储层质量耦合关系的成岩相划分体系, 已成为致密砂岩储层非均质性刻画的重要手段(魏兆胜等, 2024; 刘强和张莹, 2025; 于森等, 2025)。成岩相不仅可反映不同岩性在多阶段成岩作用中的物性响应特征, 还能够指示孔隙演化的阶段性与改造潜力, 为预测有利储层分布提供地质约束(刘强和张莹, 2025)。当前, 成岩相研究已广泛应用于塔

里木盆地、四川盆地、准噶尔盆地及三塘湖盆地中的多个非常规气藏中(程传捷等, 2020; 魏兆胜等, 2024; 杨杰等, 2025; 于森等, 2025), 在常规和非常规储层的分类评价、甜点识别与开发方案优化中发挥着日益重要的作用。

延川东地区位于鄂尔多斯盆地东南部, 是典型的多源—多层致密气富集区(贺亚维等, 2019; 黄杏雨等, 2025)。本溪组作为延川东地区下古生界的重要储层发育层位之一, 受控于构造稳定背景与深埋埋藏环境, 具有压实、胶结与溶蚀等多阶段耦合改造特征, 成岩机制复杂, 储层非均质性较强(王集等, 2023; 梁状等, 2025; 万永平等, 2025a)。然而, 已有研究多聚焦于本溪组沉积相特征与宏观物性评价(冯娟萍等, 2021; 郭艳琴等, 2024), 而针对不同岩性在成岩演化过程中所体现出的作用类型差异、孔隙响应方式及成岩相组合特征缺乏系统研究, 制约了对致密储层质量主控因素的深入理解。为揭示本溪组砂岩储层的成岩演化过程及其对储层品质的作用机制, 文章以延川东地区为研究区, 基于铸体薄片、扫描电镜及高压压汞等分析数据, 系统开展了不同矿物组成砂岩的成岩作用识别、孔隙结构分析与成岩相划分。通过识别主要成岩作用类型, 厘清压实、胶结与溶蚀作用的时序关系与演化路

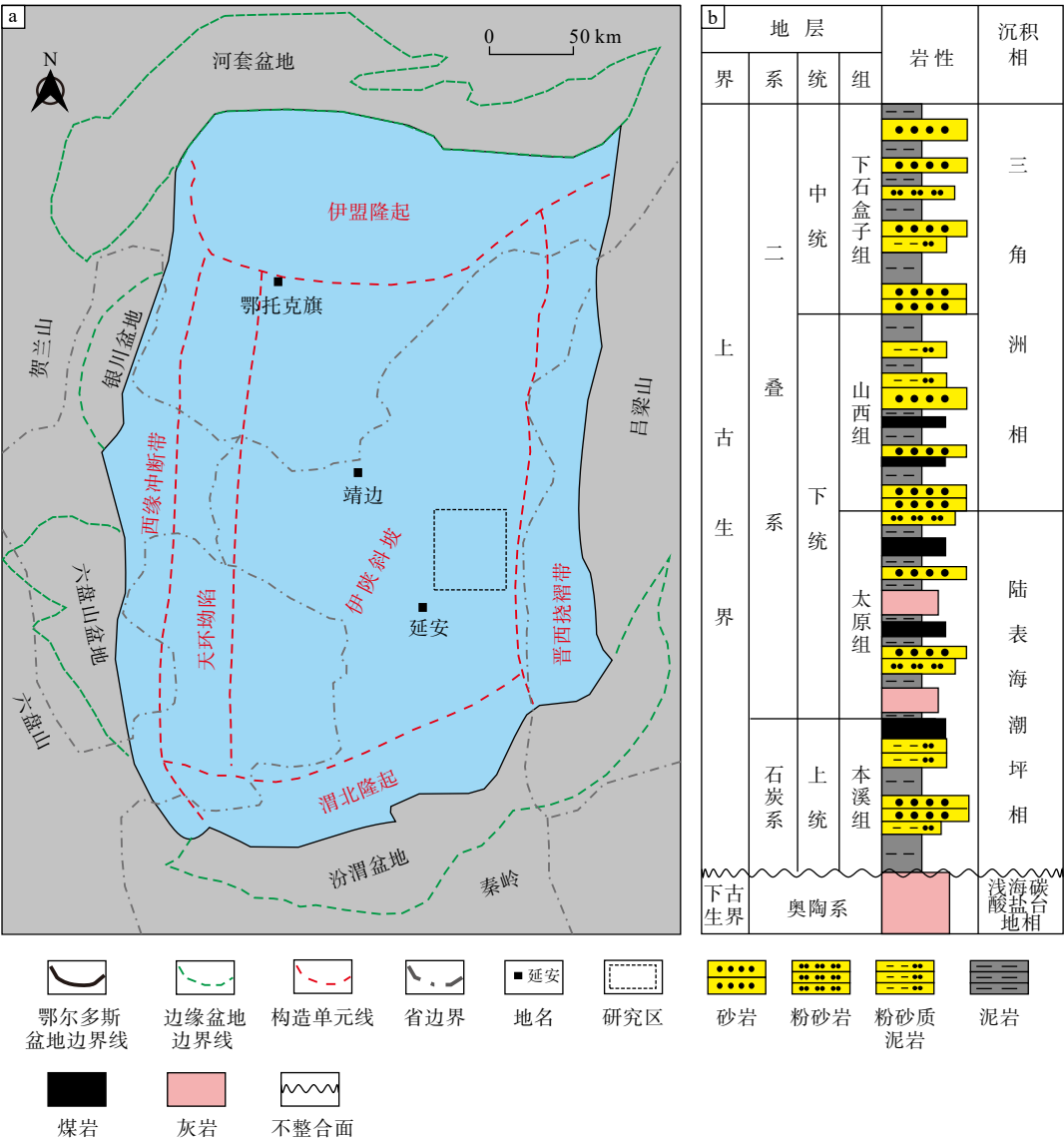
径, 建立成岩相划分体系, 并进一步探讨成岩相组合对储层非均质性及有效性发育的影响机制。研究成果可为延川东地区本溪组致密砂岩储层评价提供理论支撑, 也为同类非常规储层成岩过程分析与成岩相划分提供借鉴。

1 地质背景

鄂尔多斯盆地位于华北板块西部, 面积约为 $2.5\times 10^5\text{ km}^2$, 是在华北克拉通太古宙—古元古代结晶基底之上发育的一个多旋回叠合盆地, 发育多套生储盖组合, 呈现多层系油气分布格局(刘建等,

2021; 赵喆等, 2024; Huang et al., 2025; 伍岳等, 2025)。依据构造特征及演化, 盆地可划分为伊陕斜坡、晋西挠褶带、渭北隆起、伊盟隆起、天环坳陷与西缘冲断带 6 个一级构造单元(图 1a), 其中伊陕斜坡作为晚古生代—中生代沉积中心, 控制着鄂尔多斯盆地主力烃源岩展布, 是盆地主要的油气富集区(牛小兵等, 2024; 王冠民等, 2024; 岳怀海等, 2025; 何希鹏等, 2025)。

延川东地区位于伊陕斜坡中段, 构造形态整体平缓, 断裂稀少, 长期处于构造应力场低转换区, 表现出典型的沉积继承性+构造稳定性叠合特征, 是



a—鄂尔多斯盆地构造单元划分图; b—鄂尔多斯盆地地层柱状图

图 1 鄂尔多斯盆地构造单元划分与地层柱状图

Fig. 1 Structural units and stratigraphic column of the Ordos Basin

(a) Tectonic division map of the Ordos Basin; (b) Stratigraphic column of the Ordos Basin

天然气富集与保存的重要区带(陈刚和胡宗全, 2018; 姚红生等, 2022; 刘世民等, 2025)。自石炭纪以来,延川东地区经历了滨浅海、潮坪、三角洲、湖泊等多期沉积环境更替,本溪组—山西组发育煤系烃源岩和致密砂岩储层(李国永等, 2024; 李明瑞等, 2024)。印支期—燕山期,区域未发生剧烈褶皱改造,断裂活动以隐伏小断层为主,未见大规模逆冲或推覆构造,为后期气藏保存提供了稳定构造背景(王建民和张三, 2018)。石炭系本溪组在延川东地区呈区域稳定展布,与下伏奥陶系马家沟组不整合接触,与上覆二叠系太原组为连续沉积(图1b)。本溪组沉积环境属于陆表海潮坪沉积体系,具有典型的潮汐韵律层理,岩性组合表现为砂泥岩互层夹薄煤层,构成海陆交互相含煤建造序列。其中,本溪组砂岩作为延川东地区下古生界中重要的致密砂岩储层,为该区域天然气勘探的关键目标层段(孟祥振等, 2024; 李辉等, 2024; 万永平等, 2025b; 张蕊等, 2025)。

2 储层基本特征

2.1 矿物组成与岩石类型

本溪组砂岩碎屑成分以石英类(包括石英与燧石)为主,含量范围为2.5%~100.0%(平均为77.5%);岩屑含量为0%~97%(平均为22.3%),长石含量极低,平均仅为0.3%。砂岩成分成熟度较高(陈波等, 2012)。此外,岩屑组分主要为火山岩岩屑和中—低级变质岩岩屑,表现为火山物质与区域构造改造产物的混合来源。从砂岩碎屑组分三角图来看(图2),样品主要分布于石英砂岩(I类)、岩屑石英砂岩(III类)和岩屑砂岩(VII类)区。

本溪组砂岩粒度概率曲线整体呈S形陡升状(图3),具有典型的三段式结构,反映了沉积物粒径组成的分异特征与水动力条件的变化过程(郑浚茂等, 1980)。曲线前段平缓,粗颗粒所占比例较小,中段迅速上升,主粒径分布集中,说明搬运和沉积过程中能量适中、分选较好,末段细粒悬移组分呈平台状展布,比例较低(袁静等, 2011)。整体曲线特征揭示了本溪组砂岩形成于较稳定的搬运与沉积环境中,具有较好的分选性和成熟度。

2.2 物性与储集空间类型

本溪组砂岩储层孔隙度介于0.32%~13.82%(平均为6.30%),渗透率介于 0.026×10^{-3} ~ 1.091×10^{-3}

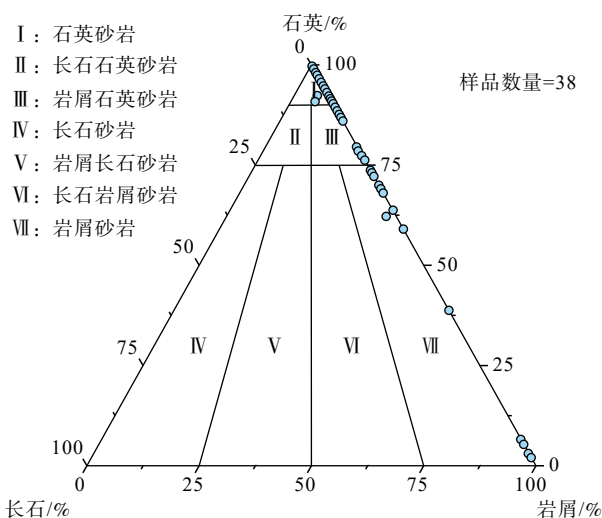


图2 延川东地区本溪组砂岩碎屑组分三角图

Fig. 2 Ternary QFL diagram of detrital sandstone components of the Benxi Formation in the East Yanchuan Block

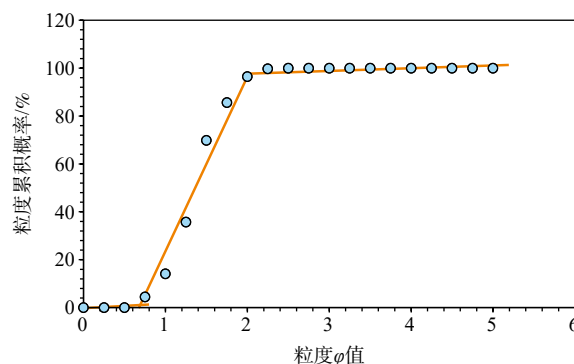


图3 延川东地区本溪组砂岩样品粒度概率累积曲线

Fig. 3 Cumulative grain size probability curve of sandstone samples from the Benxi Formation in the East Yanchuan Block

μm^2 (平均为 $0.275 \times 10^{-3} \mu\text{m}^2$),以低孔—特低孔、超低

渗储层为主(图4)。其中石英砂岩储层孔隙度介于

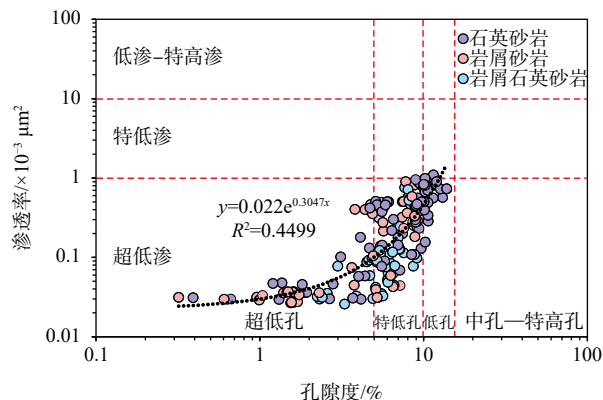


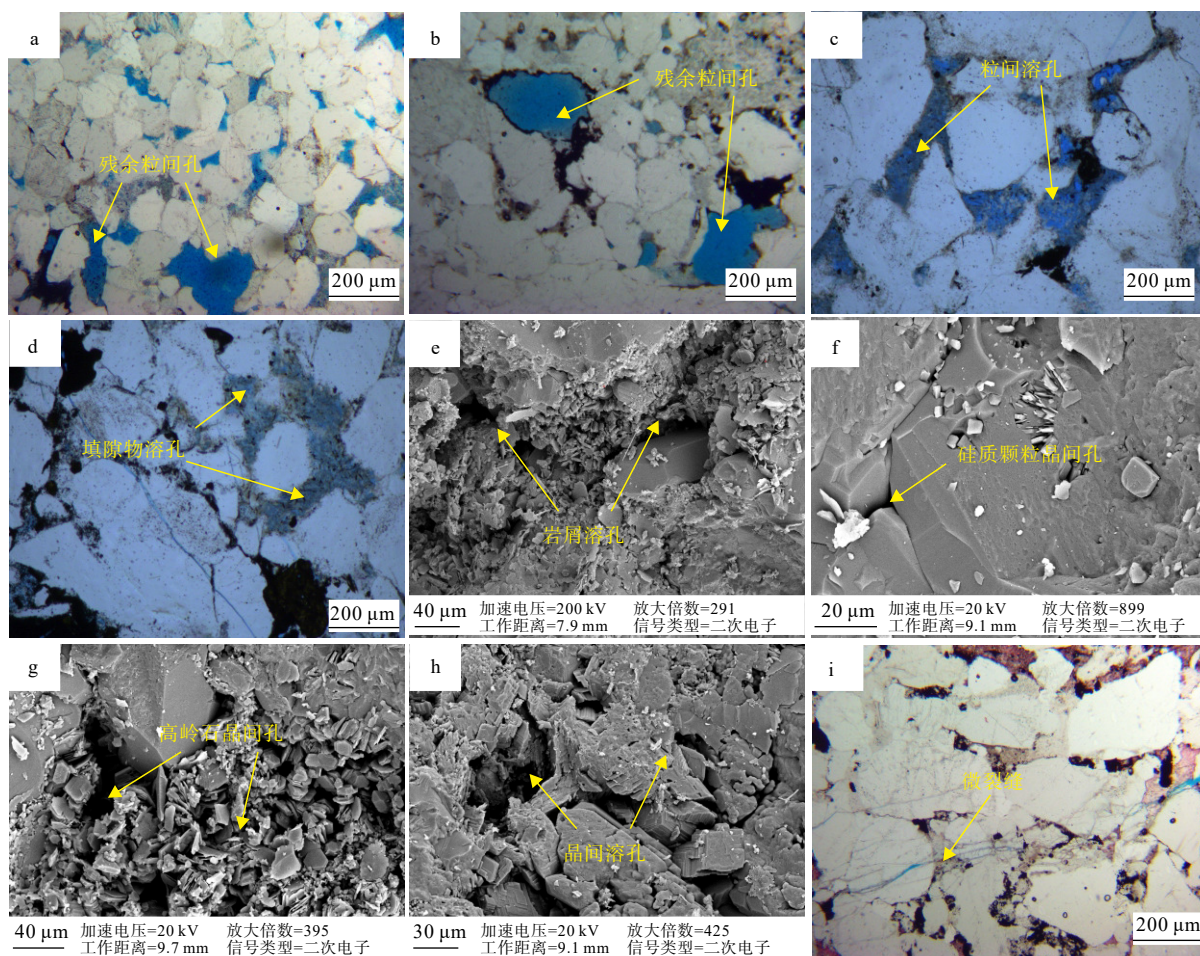
图4 延川东地区本溪组砂岩储层孔隙度与渗透率关系图

Fig. 4 Relationship between porosity and permeability of sandstone reservoirs in the Benxi Formation, East Yanchuan Block

1.40%~13.82% 之间(平均为 4.26%), 渗透率介于 $0.027 \times 10^{-3} \sim 1.059 \times 10^{-3} \mu\text{m}^2$ (平均为 $0.070 \times 10^{-3} \mu\text{m}^2$); 岩屑石英砂岩储层孔隙度介于 2.00%~8.50% (平均为 3.73%), 渗透率介于 $0.031 \times 10^{-3} \sim 0.763 \times 10^{-3} \mu\text{m}^2$ (平均为 $0.060 \times 10^{-3} \mu\text{m}^2$); 岩屑砂岩储层孔隙度介于 0.32%~9.10% (平均为 4.28%), 渗透率介于 $0.026 \times 10^{-3} \sim 1.091 \times 10^{-3} \mu\text{m}^2$ (平均为 $1.020 \times 10^{-3} \mu\text{m}^2$)。可见,

本溪组砂岩储层非均质性较强, 且岩性对储层物性控制作用并不明显, 不同矿物组成的砂岩均可形成优质储层。

延川东地区本溪组面孔率介于 0.30%~5.30% (平均为 3.60%), 孔隙类型以粒间孔为主, 局部可见次生溶孔及晶间微孔, 少量样品中发育微裂缝(图 5)。残余粒间孔面孔率为 0.30%~4.30% (平均 0.55%),



a—延 304 井, 2671.51—2671.61m, 本溪组, 残余粒间孔, 铸体薄片; b—延 149 井, 2453.49—2453.54m, 本溪组, 残余粒间孔, 铸体薄片; c—延 272 井, 2767.85—2767.91m, 本溪组, 粒间溶孔, 铸体薄片; d—延 272 井, 2767.85—2767.91m, 本溪组, 填隙物溶蚀孔, 铸体薄片; e—延 336 井, 2583.67—2583.86m, 本溪组, 岩屑溶孔, 扫描电镜; f—延 330 井, 2699.66—2699.81m, 本溪组, 硅质颗粒晶间孔, 扫描电镜; g—延 149 井, 2455.20—2455.25m, 本溪组, 高岭石晶间孔, 扫描电镜; h—延 330 井, 2699.66—2699.81m, 本溪组, 晶间溶孔, 扫描电镜; i—延 272 井, 2769.01—2769.15m, 本溪组, 微裂缝, 铸体薄片

图 5 延川东地区本溪组储层储集空间类型特征

Fig. 5 Characteristics of reservoir space types in the Benxi Formation, East Yanchuan Block

(a) Well Yan 304, 2671.51—2671.61 m, Benxi Formation, residual intergranular pores, cast thin section; (b) Well Yan 149, 2453.49—2453.54 m, Benxi Formation, residual intergranular pores, cast thin section; (c) Well Yan 272, 2767.85—2767.91 m, Benxi Formation, intergranular dissolution pores, cast thin section; (d) Well Yan 272, 2767.85—2767.91 m, Benxi Formation, interstitial dissolution pores, cast thin section; (e) Well Yan 336, 2583.67—2583.86 m, Benxi Formation, lithic dissolution pores, scanning electron microscope (SEM) image; (f) Well Yan 330, 2699.66—2699.81 m, Benxi Formation, inter-crystalline pores in siliceous particles, SEM image; (g) Well Yan 149, 2455.20—2455.25 m, Benxi Formation, inter-crystalline pores in kaolinite, SEM image; (h) Well Yan 330, 2699.66—2699.81 m, Benxi Formation, inter-crystalline dissolution pores, SEM image; (i) Well Yan 272, 2769.01—2769.15 m, Benxi Formation, microfractures, cast thin section

多分布于石英颗粒之间, 形态不规则(图 5a、5b), 部分被伊利石、高岭石胶结物及次生石英加大充填, 该类孔隙虽较为普遍, 但受压实与胶结改造影响明显, 含量有限。次生溶孔主要包括粒间溶孔(图 5c)、填隙物溶孔(图 5d)与岩屑溶孔(图 5e), 其中岩屑溶孔更为常见, 平均面孔率为 2.80%, 其形态多呈蜂窝状、残角状或解理面延伸式, 孔径一般介于 20~70 μm 之间。晶间孔主要分布于硅质颗粒以及自生伊利石、高岭石等黏土矿物晶簇之间(图 5f—5h), 孔径较小(<0.01 mm), 结构分散, 连通性弱, 平均面孔率仅为 0.30%。此外, 少数岩屑石英砂岩样品中发育微裂隙(图 5i), 其数量有限, 但对局部孔隙连通具有一定贡献(王鹏威等, 2021)。

2.3 孔隙结构特征

本溪组储层排驱压力介于 0.01~1.54 MPa(平

均为 0.71MPa), 孔喉分选系数介于 0.05~4.43(平均为 2.43), 最大进汞饱和度介于 62.55%~96.76%(平均为 79.21%), 详见表 1, 孔喉非均质性总体较强。依据孔喉结构划分标准, 可将本溪组储层划分为 I 类至 IV 类 4 种孔隙结构类型。其中, I 类和 II 类结构表现出较大的孔喉尺寸、较高的进汞饱和度及相对较好的连通性(图 6a、6b, 表 1), 主要与粒间孔与溶蚀孔复合发育有关(Zhu et al., 2024; 李琦等, 2025); III 类和 IV 类结构则以中—高排驱压力、细至微细喉为特征(图 6c、6d, 表 1), 为胶结加剧或孔隙封闭背景下的致密型结构(宋泽章等, 2023; Zhou et al., 2024; 刘硕等, 2024)。总体来看, 本溪组砂岩储层具有明显的多结构复合性特征, 不同类型的孔隙结构广泛分布于石英砂岩、岩屑石英砂岩和岩屑砂岩中, 结构类型多样且组合关系复杂。

表 1 延川东地区本溪组砂岩高压压汞参数分布特征

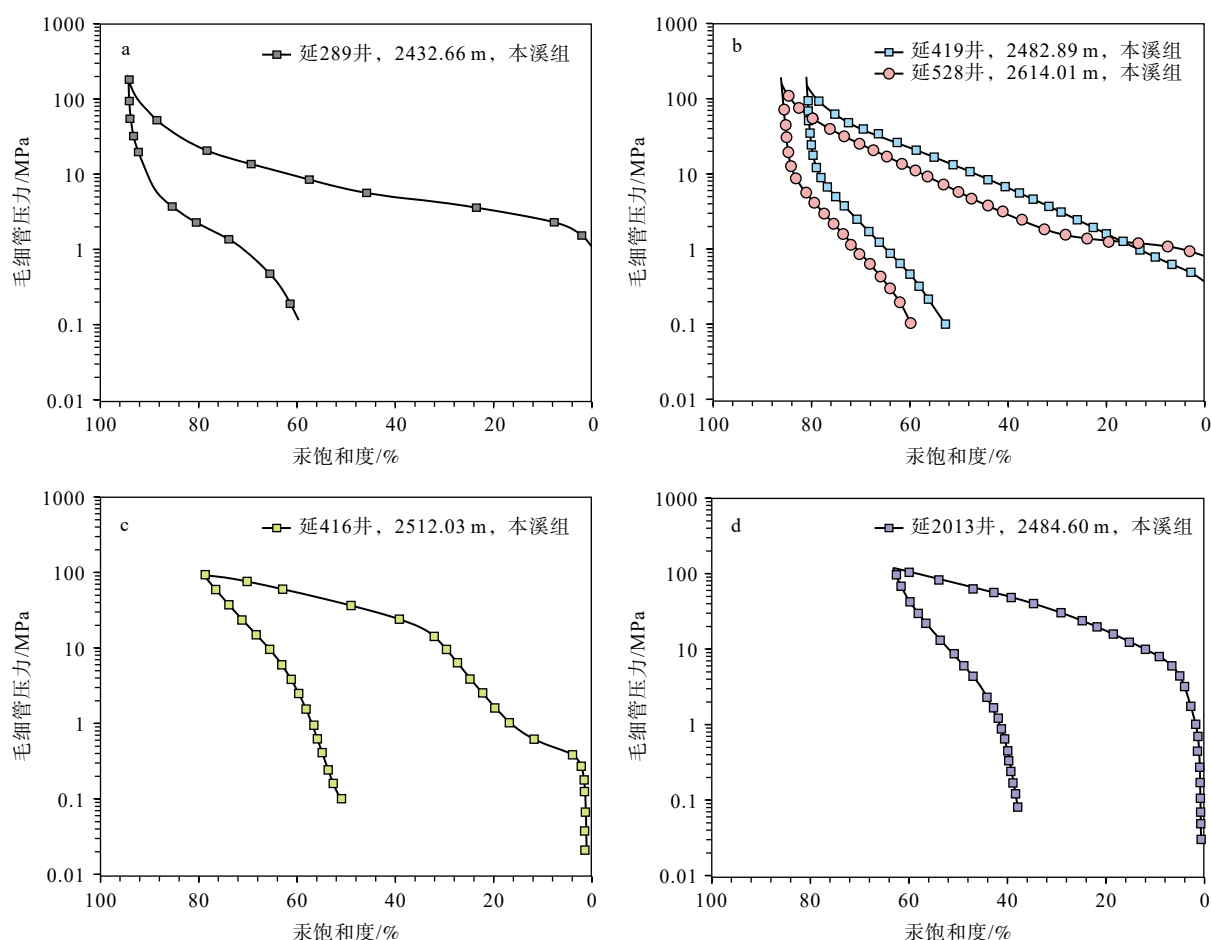
Table 1 Distribution of high-pressure mercury intrusion parameters of Benxi Formation sandstones from the East Yanchuan Block												
参数		物性		孔喉大小			孔喉分选特征			孔喉连通性/%		
层位	样品数	孔隙度/ %	渗透率/ ×10 ⁻³ μm ²	排驱压 力/MPa	中值压 力/MPa	中值半 径/μm	孔喉分 选系数	变异 系数	歪度 系数	最大进汞 饱和度	退出 效率	
本溪组	28	最小值	2.10	0.023	0.01	1.98	0.01	0.05	0.11	-0.22	62.55	18.25
		最大值	7.24	2.409	1.54	69.46	0.37	4.43	0.77	2.23	96.76	41.23
		平均值	4.28	0.267	0.71	20.29	0.14	2.43	0.31	1.31	79.21	30.36
Ⅰ类储层	8	最小值	4.64	0.355	0.01	1.98	0.19	0.05	0.11	-0.11	85.24	31.46
		最大值	7.24	1.249	0.56	5.48	0.37	0.97	0.34	0.57	96.75	41.23
		平均值	6.21	0.543	0.24	2.66	0.26	0.56	0.18	0.21	88.78	35.65
Ⅱ类储层	6	最小值	4.21	0.237	0.26	4.79	0.11	0.54	0.16	0.36	80.16	30.40
		最大值	6.52	2.409	0.98	11.34	0.24	1.58	0.49	1.09	87.19	35.85
		平均值	3.97	0.431	0.49	8.62	0.21	1.01	0.22	0.87	84.21	28.70
Ⅲ类储层	8	最小值	3.31	0.052	0.58	13.37	0.09	1.76	0.28	-0.22	73.15	23.17
		最大值	5.47	0.424	1.37	43.29	0.15	3.31	0.57	1.57	82.78	30.34
		平均值	3.56	0.215	1.01	19.91	0.11	2.57	0.36	1.04	79.26	24.59
Ⅳ类储层	6	最小值	2.1	0.023	0.79	29.64	0.01	2.55	0.45	0.96	62.55	18.25
		最大值	3.24	0.264	1.54	69.46	0.09	4.43	0.77	2.33	75.43	25.14
		平均值	2.47	0.189	1.15	35.76	0.04	2.93	0.39	1.55	71.11	21.34

3 储层成岩作用类型

3.1 压实作用

本溪组砂岩经历了较强的压实作用, 颗粒普遍紧密接触(图 7a、7b)。石英砂岩以刚性石英颗粒为主, 抗压能力强, 早期压实主要表现为颗粒滑动与

转动, 形成线接触, 部分原生粒间孔得以保留(于海跃等, 2024; 岳怀海等, 2025)。但在深埋条件下, 石英颗粒之间发生压溶重结晶, 导致次生加大和孔隙丧失。岩屑石英砂岩具有一定抗压性, 也含一定量的塑性组分, 在压实过程中, 易形成凹凸接触, 塑性颗粒如云母、泥质岩屑则发生压扁变形, 充填粒间孔隙形成假杂基结构(丛森等, 2017; 邵鑫笛等,



a—I类储层压汞曲线特征；b—II类储层压汞曲线特征；c—III类储层压汞曲线特征；d—IV类储层压汞曲线特征

图6 延川东地区本溪组砂岩高压压汞曲线特征

Fig. 6 Characteristics of high-pressure mercury intrusion (MIP) curves of Benxi Formation sandstones from the East Yanchuan Block

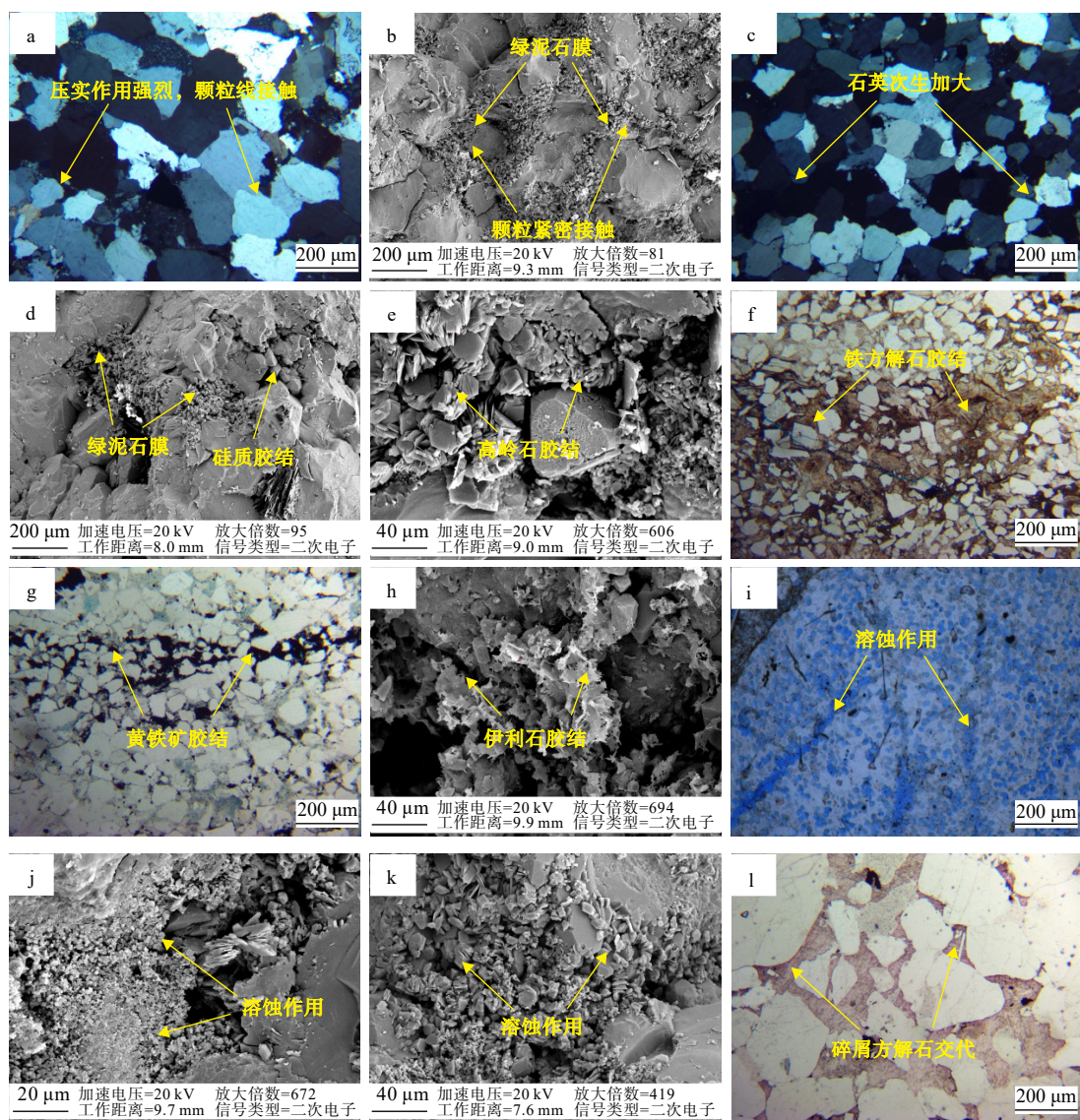
(a) Mercury intrusion curves of a Type I reservoir; (b) Mercury intrusion curves of a Type II reservoir; (c) Mercury intrusion curves of a Type III reservoir; (d) Mercury intrusion curves of a Type IV reservoir

2025; 王子龙等, 2025)。岩屑砂岩抗压性差, 在早期埋藏阶段即发生剧烈压实, 颗粒破碎、结构紧密, 原生粒间孔遭受严重破坏, 尤其在杂基含量较高时 (>12%), 压实更为敏感, 孔隙度显著降低(于海跃等, 2024)。

3.2 胶结作用

胶结作用既导致本溪组砂岩储层进一步致密, 也起到了局部保孔的作用(丛森等, 2017; Guo et al., 2024)。镜下观察显示, 研究区砂岩普遍发育硅质胶结、高岭石胶结、铁质胶结、绿泥石胶结和伊利石胶结等多种胶结类型(图7c—7h)。其中, 绿泥石常以定向排列的包膜产出于颗粒边缘, 在早期成岩阶段形成, 有效抑制了后期胶结物的进一步沉淀, 对孔隙具有良好的保护作用, 这一现象在鄂尔多斯盆地其他地区亦有报道(梁状等, 2025; 岳怀海等,

2025)。此外, 石英砂岩中常见多期硅质胶结, 多由压溶作用形成的次生石英充填粒间孔隙, 增强了骨架颗粒的稳定性, 但同时封堵孔喉, 降低了渗透性。然而, 在部分样品中发育有早期的边缘式或桥连式硅质胶结, 能够在颗粒之间形成支撑, 有效抑制压实作用进程, 促进了原生粒间孔的局部保存(Xi et al., 2015; Guo et al., 2024)。岩屑石英砂岩由于成分复杂、胶结类型多样, 发育高岭石与伊利石的复合胶结。高岭石呈细粒状充填粒间孔隙, 当分布不连续或充填不完全时, 部分残余粒间孔得以保留, 而伊利石常以片状或纤维状附着于颗粒表面, 对孔喉具有一定的支撑作用(王艳忠等, 2024)。岩屑砂岩中广泛发育碳酸盐类胶结, 常沿颗粒边缘或孔隙内沉淀, 形成致密的块状充填结构, 显著影响了储层物性(Lai et al., 2018; 宋昆鹏等, 2020; 林潼等, 2025)。



a—延 349 井, 2830.35—2830.46 m, 本溪组, 颗粒线接触, 少量火山岩屑变形, 铸体薄片; b—延 349 井, 2831.28—2833.09 m, 本溪组, 压实作用强烈, 颗粒紧密接触, 扫描电镜; c—延 176 井, 2779.87—2780.06 m, 本溪组, 石英次生加大, 铸体薄片; d—延 347 井, 2889.09—2889.25 m, 本溪组, 硅质胶结, 扫描电镜; e—延 149 井, 2457.37—2457.66 m, 本溪组, 高岭石胶结, 扫描电镜; f—延 149 井, 2890.59—2890.8 m, 本溪组, 铁方解石胶结, 铸体薄片; g—延 272 井, 2765.31—2765.39 m, 本溪组, 黄铁矿胶结, 铸体薄片; h—延 176 井, 2777.56—2777.68 m, 本溪组, 伊利石胶结, 扫描电镜; i—延 149 井, 2457.80—2457.85 m, 本溪组, 溶蚀作用, 铸体薄片; j—延 349 井, 2829.49—2831.28 m, 本溪组, 溶蚀作用, 扫描电镜; k—延 336 井, 2580.91—2581.01 m, 本溪组, 溶蚀作用, 扫描电镜; l—延 272 井, 2769.01—2769.15 m, 本溪组, 碎屑方解石被铁方解石交代, 铸体薄片

图 7 延川东地区本溪组砂岩典型成岩作用类型

Fig. 7 Typical diagenetic types of Benxi Formation sandstones in the East Yanchuan Block

(a) Well Yan 349, 2830.35—2830.46 m, Benxi Formation, point-line contact between grains with slight deformation of volcanic lithic fragments, cast thin section; (b) Well Yan 349, 2831.28—2833.09 m, Benxi Formation, strong compaction with tight grain contacts, scanning electron microscope (SEM) image; (c) Well Yan 176, 2779.87—2780.06 m, Benxi Formation, quartz overgrowth, cast thin section; (d) Well Yan 347, 2889.09—2889.25 m, Benxi Formation, siliceous cementation, SEM image; (e) Well Yan 149, 2457.37—2457.66 m, Benxi Formation, kaolinite cementation, SEM image; (f) Well Yan 149, 2890.59—2890.80 m, Benxi Formation, ferrocalcite cementation, cast thin section; (g) Well Yan 272, 2765.31—2765.39 m, Benxi Formation, pyrite cementation, cast thin section; (h) Well Yan 176, 2777.56—2777.68 m, Benxi Formation, illite cementation, SEM image; (i) Well Yan 149, 2457.80—2457.85 m, Benxi Formation, dissolution features, cast thin section; (j) Well Yan 349, 2829.49—2831.28 m, Benxi Formation, dissolution features, SEM image; (k) Well Yan 336, 2580.91—2581.01 m, Benxi Formation, dissolution features, SEM image; (l) Well Yan 272, 2769.01—2769.15 m, Benxi Formation, detrital calcite replaced by ferrocalcite, cast thin section

3.3 溶蚀作用

本溪组储层在埋藏过程中经历了广泛的溶蚀作用(图 7i—7k)。石英砂岩抗蚀性较强,仅在局部弱胶结或含少量不稳定矿物(如燧石、重矿物)的部位,酸性流体沿粒间缝隙对胶结物及颗粒边缘产生溶蚀现象(岳怀海等, 2025)。岩屑石英砂岩富含如高岭石、伊利石、铁质胶结物及不稳定岩屑,溶蚀作用广泛且强烈。酸性流体沿原始孔隙、裂缝及粒间缝隙持续运移,形成粒间—粒内复合溶孔,并建立局部的次生孔隙网络(王宏博等, 2023)。岩屑砂岩中广泛分布的火山岩屑稳定性较差,易形成岩屑溶孔,但由于该类岩石普遍经历了强烈的早期压实和胶结,孔隙连通性差,导致反应产物迁移受限,储层质量改善并不明显(张顺, 2018; 邵鑫笛等, 2025; 杨杰等, 2025)。总体来看,溶蚀作用在不同岩性中的响应强度与作用结果受到矿物成分、孔隙结构和裂缝网络等多重因素共同控制。其中,反应产物能否及时从体系中迁出,是决定溶蚀是否能够转化为有效孔隙的关键因素(于森等, 2022; 梁状等, 2025; 刘志达等, 2025)。

3.4 交代作用

交代作用在本溪组致密砂岩中局部发育,显微镜下可观察到少量铁方解石对早期碳酸盐胶结物的交代(图 7l)、伊利石对高岭石的交代,以及黄铁矿对不稳定颗粒的替代现象。这些交代作用多发生于中成岩阶段,反映了成岩流体性质变化所引起的局部矿物重组(王艳忠等, 2024)。

交代作用对本溪组储层的改造作用较为有限,部分交代产物如铁方解石、黄铁矿等呈致密块状沉淀,易封闭孔隙,加剧致密化(王宏博等, 2023; 于海跃等, 2024)。然而,个别样品中交代过程破坏原有胶结结构或颗粒外缘,形成较小的次生孔隙,改善局部储集性能。总体来看,交代作用在研究区砂岩中发育强度较弱,对储层物性演化的控制作用较小,是成岩演化中的一种次要改造机制。

4 储层成岩演化阶段

本溪组致密砂岩储层物性整体较差,主要储集空间以残余粒间孔、粒间溶孔和岩屑溶孔为主。储层成岩温度约 85 ~ 140 °C,主要处于中成岩 A 期向 B 期过渡阶段。该阶段压实作用基本终止,胶结作用广泛发育,溶蚀作用开始对孔隙结构产生主导性

改造。

采用 Beard and Weyl(1973)提出的公式对本溪组致密储层进行原始孔隙度恢复,结合粒度参数,计算得到石英砂岩、岩屑石英砂岩以及岩屑砂岩初始孔隙度分别为 39.87%、37.65% 和 34.24%。石英砂岩在早成岩阶段压实作用较弱,部分粒间孔得以保留,但随着埋藏加深,中成岩阶段压溶作用增强,次生石英加大及硅质胶结普遍发育,孔隙持续封闭,储层物性逐渐降低,表现为“中成岩阶段致密化”特征(图 8)。岩屑石英砂岩成分复杂、塑性组分丰富,早期压实作用较强,随后经历高岭石、伊利石等多期胶结,进入中成岩阶段后,受酸性流体影响,普遍发生胶结物和不稳定颗粒的选择性溶蚀,形成粒间溶孔和岩屑溶孔,孔隙演化呈“早成岩阶段降低—中成岩阶段回升”趋势(图 8),是区内容蚀改造潜力最大的岩石类型。岩屑砂岩压实敏感性强,早成岩阶段孔隙度迅速降低,叠加碳酸盐胶结作用进一步致密化,虽在部分样品经历了中成岩期较强的溶蚀作用,但由于孔隙结构普遍较差,对储层质量的改造作用有限,表现为“多期致密化特征”(图 8)。

5 储层成岩相

通过砂岩胶结物含量与粒间体积交会图分析发现(图 9),本溪组砂岩储层物性变差的主控因素为压实作用,减孔率介于 40%~80%。此外,胶结减孔率主要介于 5%~50%,是储层质量变差的次要因素。石英砂岩受压实作用较弱,虽然胶结物含量高,但多数已经历规模溶蚀,孔隙中仅残留部分溶蚀产物,岩屑砂岩受到压实和胶结等破坏性成岩作用最为强烈,而岩屑石英砂岩的压实和胶结作用强度介于上述两者之间。

基于不同样品的成岩作用类型、成岩演化过程与孔隙演化特征,可以划分出弱压实—弱胶结相、中等压实中等胶结相、胶结致密相和胶结—溶蚀复合相共 4 类成岩相。不同成岩相显微特征、储层质量及岩性分布具有明显差异(表 2)。

5.1 弱压实—弱胶结相

弱压实—弱胶结相常发育于潮坪砂坝与边缘三角洲等高能砂体沉积区,具有较强搬运与分选能力,砂体厚度大,泥质杂基含量低。多见于分选性好、杂基含量低的石英砂岩及部分结构均一的岩屑石英砂岩中(表 2, 表 3, 图 10)。其主要形成于早期

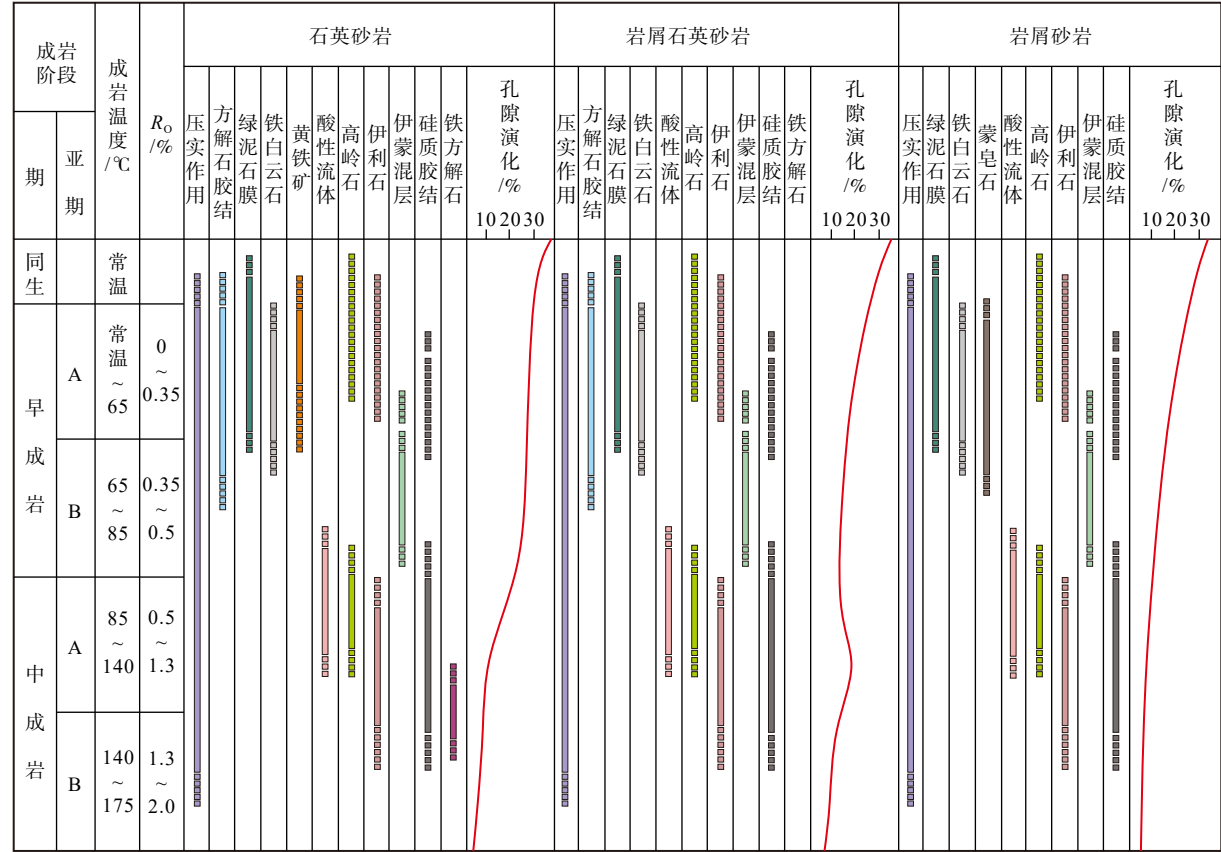


图 8 延川东地区本溪组不同岩性砂岩的成岩演化阶段与孔隙演化特征

Fig. 8 Diagenetic evolution stages and pore evolution characteristics of lithologically different sandstones from the Benxi Formation, East Yanchuan Block

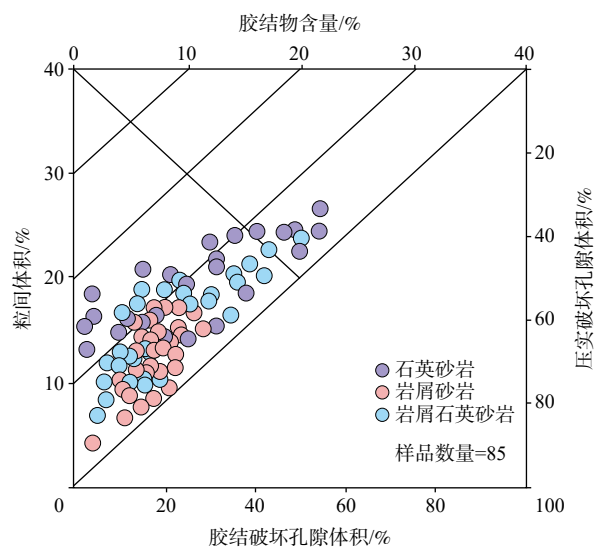


图 9 延川东地区本溪组储层孔隙度压实作用及胶结作用 (图版引自 Houseknecht, 1987)

Fig. 9 Relationship between compaction, cementation, and porosity of Benxi Formation reservoirs in the East Yanchuan Block (modified after Houseknecht, 1987)

压实与胶结作用相对较弱的成岩环境中, 颗粒排列

较为疏松, 以点-线式接触为主, 原生粒间孔相对完整, 孔喉结构连通性较好。此外, 储层尚未经历强烈成岩改造, 保留了较多原始孔隙, 在本溪组早期成岩阶段具备较好的储集潜力。

5.2 中等压实中等胶结相

中等压实中等胶结相主要发育于潮坪-潮上泥坪过渡区或边缘三角洲前缘斜坡带, 沉积水动力处于中等水平。常见于岩屑石英砂岩及部分胶结较弱的岩屑砂岩(表 2, 表 3, 图 10), 其主要形成于压实与胶结强度均处于中等水平的成岩环境中, 颗粒以线接触与局部凹凸接触并存为主, 孔隙结构受成岩改造程度适中, 原生粒间孔部分保留, 局部可见晶间孔或细小溶孔, 储层连通性适中。

5.3 胶结致密相

胶结致密相主要形成于潮上泥坪与三角洲前缘等低能沉积区。主要见于岩屑砂岩和杂基含量较高的岩屑石英砂岩, 少量石英砂岩也可在深埋与强胶结下演化为此类成岩相(表 2, 表 3, 图 10)。其早成岩阶段压实显著、中成岩阶段胶结强烈, 石英

表 2 延川东地区本溪组砂岩成岩相特征及划分标准

成岩相类型	弱压实弱胶结相	中等压实中等胶结相	胶结溶蚀复合相	胶结致密相
命名依据	机械压实与胶结作用均弱	机械压实与胶结作用适中	胶结基础上发育局部或强烈溶蚀	胶结强度高, 封闭孔喉
显微特征	石英颗粒点接触或线接触, 胶结物零散	石英线、凹凸接触, 胶结物含量低	胶结物大面积溶蚀, 见蜂窝状溶蚀构造	胶结物充填孔隙, 颗粒凹凸接触
储层质量	好	中等, 非均质性强	中等, 非均质性强	差
常见岩性	石英砂岩为主, 少量岩屑石英砂岩	岩屑石英砂岩为主, 少量岩屑砂岩	三者均有	三者均有

表 3 延川东地区本溪组不同砂岩成岩过程及孔隙特征

岩石类型	石英砂岩	岩屑石英砂岩	岩屑砂岩
成岩演化路径	弱压实、弱胶结→石英、碳酸盐胶结→局部溶蚀	中等压实、胶结→胶结致密→强烈溶蚀	强压实、胶结→碳酸盐胶结致密→局部溶蚀
主控成岩作用	弱压实	强溶蚀	局部溶蚀
典型孔隙组合	残余粒间孔、溶孔	粒间溶孔、岩屑溶孔	岩屑溶孔、晶间孔
主要成岩相	弱压实弱胶结相	胶结溶蚀复合相	胶结致密相

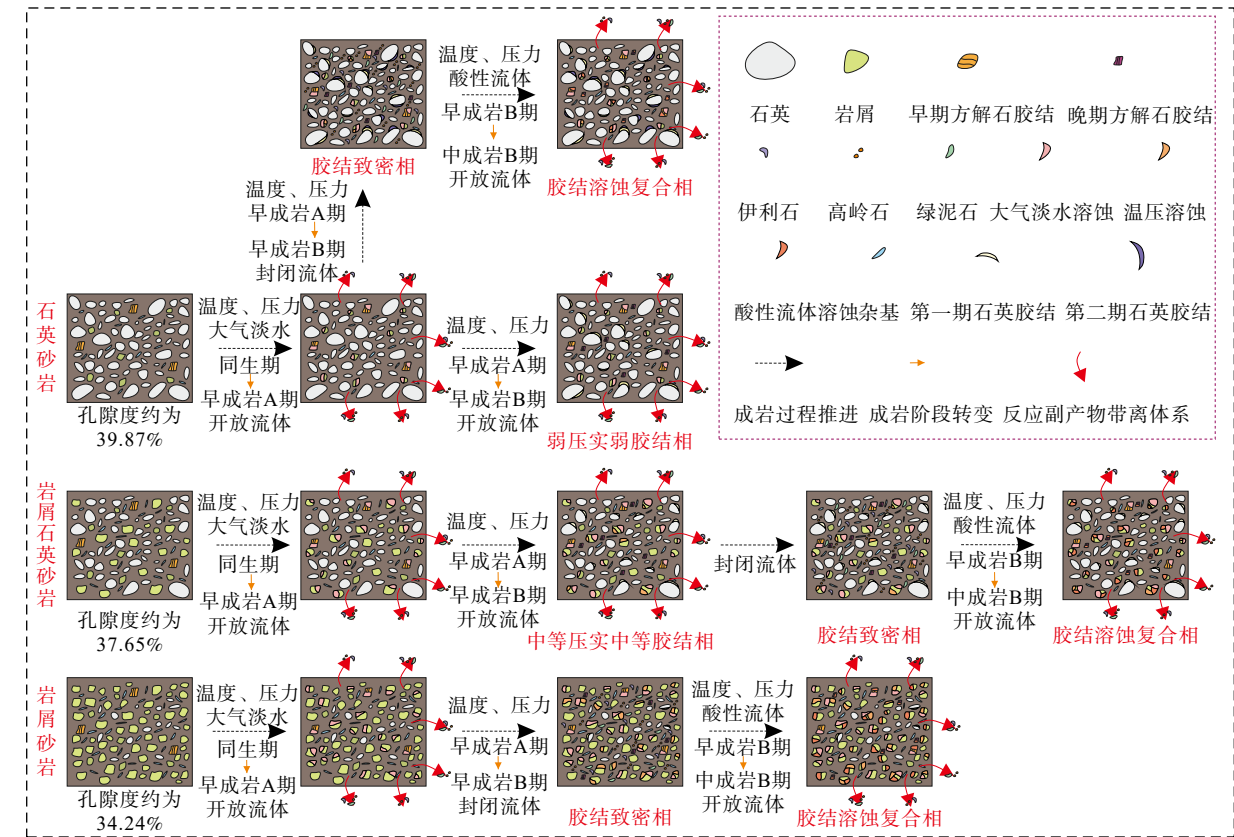


图 10 延川东地区本溪组不同矿物组成砂岩成岩演化路径与成岩相划分依据

Fig. 10 Diagenetic evolution pathways and criteria for diagenetic facies classification for Benxi Formation sandstones with different mineral compositions, East Yanchuan Block

加大及伊利石、高岭石和碳酸盐胶结作用较强。颗粒多呈凹凸接触, 孔隙度与渗透率显著降低, 孔喉趋于封闭, 常为无效储层, 是成岩作用破坏最强的成岩相。

5.4 胶结-溶蚀复合相

胶结-溶蚀复合相通常发育于潮坪与三角洲前缘过渡带, 属于中等至弱能量过渡环境。主要分布于岩屑石英砂岩及少量杂基较少的岩屑砂岩(表2,

表3,图10)。其主要反映储层经历了先胶结、后溶蚀的复合演化过程。早期胶结作用使粒间孔隙封闭,而中成岩阶段酸性流体沿颗粒边缘或原始孔隙系统侵入,对胶结物或不稳定颗粒产生规模溶蚀,孔隙中胶结物与次生溶孔共存,连通性差异大,储层非均质性较强。

5.5 成岩控制下的储层演化机制

延川东地区本溪组致密砂岩储层质量受压实、胶结与溶蚀等多阶段成岩作用叠加改造控制,不同成岩演化路径决定了储层有效性与空间分布的差异性。综合成岩作用类型、时序关系与岩性特征,归纳出以下4类演化机制。

5.5.1 压实作用决定原始孔隙结构框架,为成岩响应提供前提条件

在早成岩阶段,岩石颗粒组成与结构差异导致储层的压实强度不一。石英砂岩刚性强,保孔能力优于岩屑类储层,因此,岩屑砂岩早期即遭受剧烈压实,孔隙度下降迅速。此外,压实不仅控制原生粒间孔的保存程度,还影响了中成岩阶段胶结与溶蚀过程的空间条件,是储层成岩演化中的初始控制因子。

5.5.2 胶结作用重塑孔喉结构,是储层致密化的核心机制

本溪组中成岩阶段广泛发育硅质、高岭石、伊利石、绿泥石与碳酸盐类胶结,形成致密骨架。胶结方式与时序差异控制着孔隙的封闭程度和连通性。此外,发育于弱压实背景下的胶结作用可局部保孔,反映出胶结作用在储层致密与保孔之间的双重调节效应。

5.5.3 溶蚀作用作为改善储层物性的逆向过程,具有选择性与差异性

中成岩期,在具有良好连通性的区域内酸性流体对胶结物、不稳定颗粒等发生选择性溶解。其中岩屑石英砂岩响应最强,形成多类型复合溶孔。但在结构封闭区域,反应产物迁移受限,次生孔隙常被继承性沉淀重新封闭,显示出空间选择性—连通性—再沉淀控制体系。

5.5.4 成岩相类型归结多阶段成岩作用组合的最终产物,反映储层非均质性空间格局

多阶段成岩作用的叠加演化最终导致在不同矿物组成砂岩中形成了差异化的成岩相组合,这是致密砂岩储层非均质性的直接体现。因此,成岩相类型的空间分布不仅反映了成岩过程的强弱,也为

后期储层预测与开发部署提供了地质依据。

6 结论

(1)延川东地区本溪组砂岩以低孔-特低孔、超低渗储层为主,储层非均质性较强,成岩演化路径复杂。主要发育石英砂岩、岩屑石英砂岩和岩屑砂岩3类致密砂岩储集体,碎屑组分成熟度较高,孔隙类型以残余粒间孔、粒内孔和溶蚀孔为主,局部发育晶间孔与微裂缝,孔隙结构多样。

(2)石英砂岩以强抗压、中成岩阶段致密化为特征,局部可保留早期粒间孔;岩屑石英砂岩成分复杂,经历了复合胶结-溶蚀过程,是溶蚀作用最强、孔隙结构改善潜力最大的岩石类型;岩屑砂岩压实敏感性高,胶结作用强烈,储层致密化程度最高,储集潜力有限。

(3)在识别典型成岩作用类型与演化路径的基础上,建立了由弱压实弱胶结相、中等压实中等胶结相、胶结致密相和胶结溶蚀复合相构成的成岩相划分体系。各类成岩相分别对应不同成岩过程及其孔隙演化特征,揭示了储层非均质性形成的微观机制。

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