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山东平度旧店金矿床成矿物质来源和成矿模式

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摘要: 山东平度旧店金矿床地处胶东金成矿省西北部、招平断裂带南段, 是该地区石英脉型金矿床的典型代表。虽然该矿床开采历史悠久, 但由于规模较小, 矿床成因研究较为薄弱, 制约了对区域成矿规律的认识。本次研究在详细的野外地质调查基础之上, 开展了含矿石英脉中石英氢氧同位素和黄铁矿硫同位素分析, 以期揭示成矿流体和金属来源, 厘定矿床成矿模式。研究表明, 旧店金矿床成矿过程可分为早成矿阶段(I)、主成矿阶段(II)和晚成矿阶段(III)3个阶段, 其中主成矿阶段是金矿化主要发育阶段, 包括中粗粒黄铁矿阶段(II₁)、细粒–微细粒黄铁矿阶段(II₂)和多金属硫化物阶段(II₃)3个亚阶段。不同成矿阶段的石英氢氧同位素特征显示, 早成矿阶段成矿流体 $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ 值为 2.22‰~4.64‰, $\delta\text{D}_{\text{V-SMOW}}$ 值为 -90.2‰~-72.2‰, 主成矿阶段成矿流体 $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ 值为 1.17‰~4.36‰, $\delta\text{D}_{\text{V-SMOW}}$ 值为 -84.3‰~-77.1‰, 晚成矿阶段成矿流体 $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ 值为 -0.01‰~4.28‰, $\delta\text{D}_{\text{V-SMOW}}$ 值为 -97.7‰~-87.1‰。旧店金矿床早阶段成矿流体以岩浆水为主, 随着流体演化, 大气降水逐步混入且所占比例不断增加。不同成矿阶段的黄铁矿硫同位素无明显差异, $\delta^{34}\text{S}$ 值大多为 7.65‰~10.10‰(平均值为 8.87‰), 这一范围与前寒武纪变质基底和中生代花岗岩特征十分类似。综合前人研究成果, 建立了旧店金矿成矿模式: 晚侏罗世, 旧店地区大量含金的古老变质岩基底发生重熔, 形成玲珑型花岗岩; 在早白垩晚期受华北克拉通东部岩石圈大规模伸展–减薄和软流圈上涌影响, 花岗岩和变质岩残留体中的金发生活化富集, 并在岩浆水和大气降水组成的混合流体作用下, 在招平断裂带下盘次级断裂构造中富集成矿。

关键词: H-O-S 同位素; 成矿物质来源; 成矿模式; 旧店金矿床; 胶东

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Sources of Ore-forming Materials and Ore-forming Model of the Jiudian Gold Deposit in Pingdu, Shandong Province

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Abstract: The Jiudian gold deposit in Pingdu, Shandong Province, is located in the northwestern part of the Jiaodong gold metallogenic province and the southern segment of the Zhaoping Fault Zone. It serves as a typical representative of quartz-vein type gold deposits in this region. Despite its long mining history, genetic studies of this relatively small-scale deposit have remained inadequate, constraining the understanding of regional metallogenic regularity. Based on detailed field geological investigations, this study conducted hydrogen-oxygen isotopic analyses of quartz from ore-bearing veins and sulfur isotopic analyses of pyrite from ores, aiming to elucidate ore-forming fluid and metal sources and establish a metallogenic model for the deposit. Research indicates that the mineralization process of the Jiudian gold deposit can be divided into three stages: early mineralization stage (I), main mineralization stage (II), and late mineralization stage (III). The main mineralization stage represents the primary period of gold mineralization, further subdivided into three sub-stages: medium-coarse grained pyrite stage (II₁), fine-micron fine grained pyrite stage (II₂), and polymetallic sulfide stage (II₃). Isotopic characteristics of quartz from different mineralization stages show: $\delta^{18}\text{O}_{\text{H}_2\text{O}}=2.22\text{‰}\sim 4.64\text{‰}$ and $\delta\text{D}_{\text{V-SMOW}}=-90.2\text{‰}\sim -72.2\text{‰}$ for the early stage; $\delta^{18}\text{O}_{\text{H}_2\text{O}}=1.17\text{‰}\sim 4.36\text{‰}$ and $\delta\text{D}_{\text{V-SMOW}}=-84.3\text{‰}\sim -77.1\text{‰}$ for the main stage; $\delta^{18}\text{O}_{\text{H}_2\text{O}}=-0.01\text{‰}\sim 4.28\text{‰}$ and $\delta\text{D}_{\text{V-SMOW}}=-97.7\text{‰}\sim -87.1\text{‰}$ for the late stage. The early-stage ore-forming fluids were predominantly magmatic water, with progressive mixing and increasing proportion of meteoric water during fluid evolution. Sulfur isotopic compositions of pyrite from different mineralization stages show no significant variation, with $\delta^{34}\text{S}$ values ranging from 7.65‰ to 10.10‰ (average 8.87‰), closely resembling the characteristics of Precambrian metamorphic basements and Mesozoic granites. Integrating previous research findings, a metallogenic model for the Jiudian gold deposit is proposed: During the Late Jurassic, remelting of gold-bearing ancient metamorphic basement rocks in the Jiudian area formed Linglong-type granites. In the Late Early Cretaceous, under the influence of large-scale lithospheric extension-thinning in the eastern North China Craton and asthenosphere upwelling, gold from granites and metamorphic rock relics underwent mobilization and enrichment. Driven by mixed fluids composed of magmatic water and meteoric water, gold precipitation occurred in secondary fracture structures within the footwall of the Zhaoping Fault Zone, forming the ore deposit.

Keywords: H-O-S Isotopes; source of ore-forming materials; ore-forming model; Jiudian gold deposit; Jiaodong

中国胶东地区资源丰富, 已发现金矿床超过 240 个(宋英昕等, 2017), 累计探明黄金资源储量近 5 800 t (Deng et al., 2023; 宋明春等, 2023), 占国内黄金总资源量的近 1/3, 是仅次于南非兰德和乌兹别克斯坦穆龙套的世界第三大金矿集中区(杨立强等, 2022)。区内招平断裂带是胶东地区 6 条主要构造-金矿带之一, 带内目前已探明含玲珑、大尹格庄、夏甸金矿等大型-

超大型金矿床在内的 40 余处金矿床(宋明春等, 2022a, 2022b)。金矿床主要位于招平断裂中段和北段, 累计探明黄金资源储量超过 1 500 t (邓军等, 2023; 杨立强等, 2024)。目前, 前人对这些位于招平断裂中段和北段的大型-超大型金矿床已经开展了大量的研究, 揭示了它们的成岩成矿时代、成矿流体、金属来源以及流体活动规律(Song et al., 2020; Xiong et al., 2020)。

然而,作为胶东金矿集区的重要组成部分,招平断裂带南段由于金矿床规模普遍较小,研究一直较为薄弱,对成矿物质来源等关键问题还未形成统一认识,成矿机制和成矿模型也存在争议(岳石等,1988;李逸凡等,2015; Tian et al., 2022; 汪浩等, 2024)。这严重制约了对区域金矿成矿规律的认识。

旧店金矿床位于山东省平度市旧店镇,是招平断裂带南段发育的典型石英脉型金矿床,矿床规模已达中型(牛警微等, 2023a; 李经纬等, 2023)。本次研究在详细的野外地质调查基础之上,基于矿石矿物组合、脉体穿插包裹关系和结构构造特征等,识别出3个成矿阶段,并对来自不同成矿阶段的石英和黄铁矿分别开展了石英 H-O 同位素和 S 同位素分析。研究结果可以深刻揭示旧店金矿成矿流体和金属来源,厘定矿床成矿模式,从而有助于深入认识招平断裂带金矿成矿规律,并为该地区下一步找矿勘查提供理论指导。

1 区域地质背景

旧店金矿床所处的胶东地区,地处华北克拉通东部(图 1a),以郯庐断裂为界和鲁西相邻,包括胶北地体和苏鲁地体,二者以五莲-烟台断裂为界(Deng et al.,

2020)。

胶北地体由胶北隆起和胶莱盆地构成(图 1b),胶北隆起中变质基底发育,可见新太古代胶东群斜长角闪岩、浅粒岩、黑云变粒岩(王惠初等, 2015; 万渝生等, 2017), 古元古代荆山群和粉山子群石英岩、片岩、大理岩、变粒岩、透辉岩(陈燕平等, 2024), 新元古代蓬莱群灰岩、板岩、大理岩、石英岩(旷红伟等, 2023)。胶莱盆地主体为早白垩世沉积岩和火山岩,自下而上可见莱阳组,青山组,王氏组,岩性以砾岩、砂砾岩、粉砂岩、泥岩、山岩及火山碎屑岩为主(李守军, 1998; 任凤楼等, 2008; 庞崇进, 2015); 中生代花岗岩广布,以晚侏罗世玲珑花岗岩为主,其次为早白垩世早期郭家岭花岗闪长岩和艾山花岗岩; 断裂构造发育,以 NE 向-NNE 向三山岛断裂、焦家断裂、招平断裂、栖霞断裂、桃村断裂、牟平-即墨断裂规模最大,是胶东地区重要的金矿控矿构造,各断裂带内探明金矿床超过 160 处,累计探明黄金资源储量超过 5 000 t(杨立强等, 2024)。苏鲁地体由超高压变质带和中生代侵入岩组成(徐扬等, 2015; Liu et al., 2022), 超高压变质带以新元古代花岗质片麻岩为主(刘利双等, 2018); 中生代侵入岩可见晚侏罗世昆嵛山花岗岩(Goss et al., 2010; Yang et al., 2017)、早白垩世晚期伟德山花岗岩

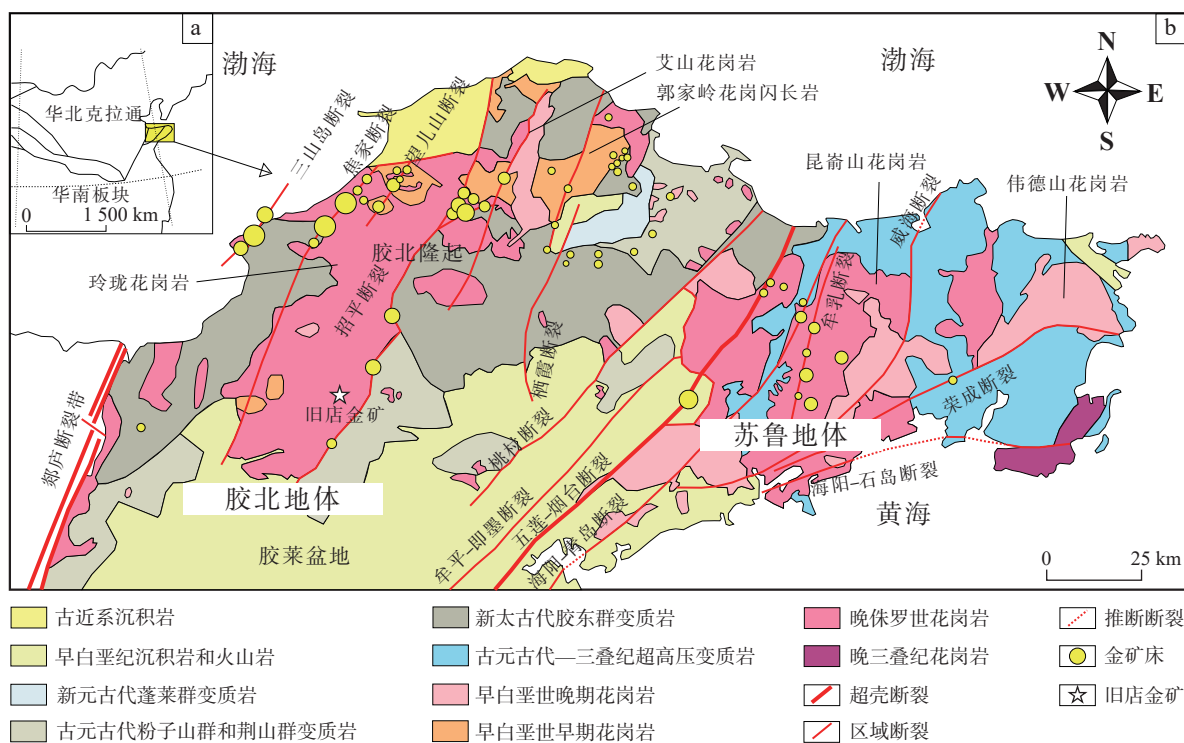


图1 旧店金矿床大地构造位置示意图(据邓军等, 2023 修改)

Fig. 1 Schematic diagram of the geotectonic location of the Jiudian gold deposit

(邓军等, 2023); 断裂构造以海阳-青岛断裂、牟乳断裂、威海断裂、荣成断裂和海阳-石岛断裂为代表, 断裂带内矿体赋存和发育程度不及胶北地体, 累计明金矿床超过 90 处, 累计探明黄金资源储量超过 450 t(杨立强等, 2024)。

2 矿床地质特征

旧店金矿床地处胶北地体招平断裂带南段下盘。区内地层出露有古元古代荆山群和新生代第四系(图2), 荆山群呈孤岛残留体零星分布于中生代玲珑超单元云山单元花岗岩体中, 岩性可见野头组蛇纹大理岩、长石石英砂岩、透辉变粒岩、斜长透辉岩和黑云变粒

岩, 第四系主要分布于黄同水库上游庙东河和陈家沟河道内, 岩性可见河床相砂砾、含砾粗砂, 洪冲积相泥、砂、砾、巨砾混杂堆积。区内构造以断裂构造为主, 可分为 NE 向、NNE 向和 NW 向 3 组, 其中以 NE 向和 NNE 向为主, 该两组断裂是 I 级招平断裂和 II 级涧里断裂的次级断裂, 是区内重要的控矿和容矿构造, 北西向断裂以碎石山断裂为代表且规模最大, 为成矿后断裂。区内岩浆岩发育, 以玲珑超单元云山单元弱片麻状细粒含石榴二长花岗岩为主, 次为崔召单元片麻状细粒二长花岗岩, 郭家店单元中粗粒二长花岗岩仅在矿区西北角小面积出露, 脉岩可见煌斑岩、闪长玢岩、石英闪长玢岩、花岗细晶岩和伟晶岩脉等。

矿区及附近共分布大小矿脉 20 多条, 矿石类

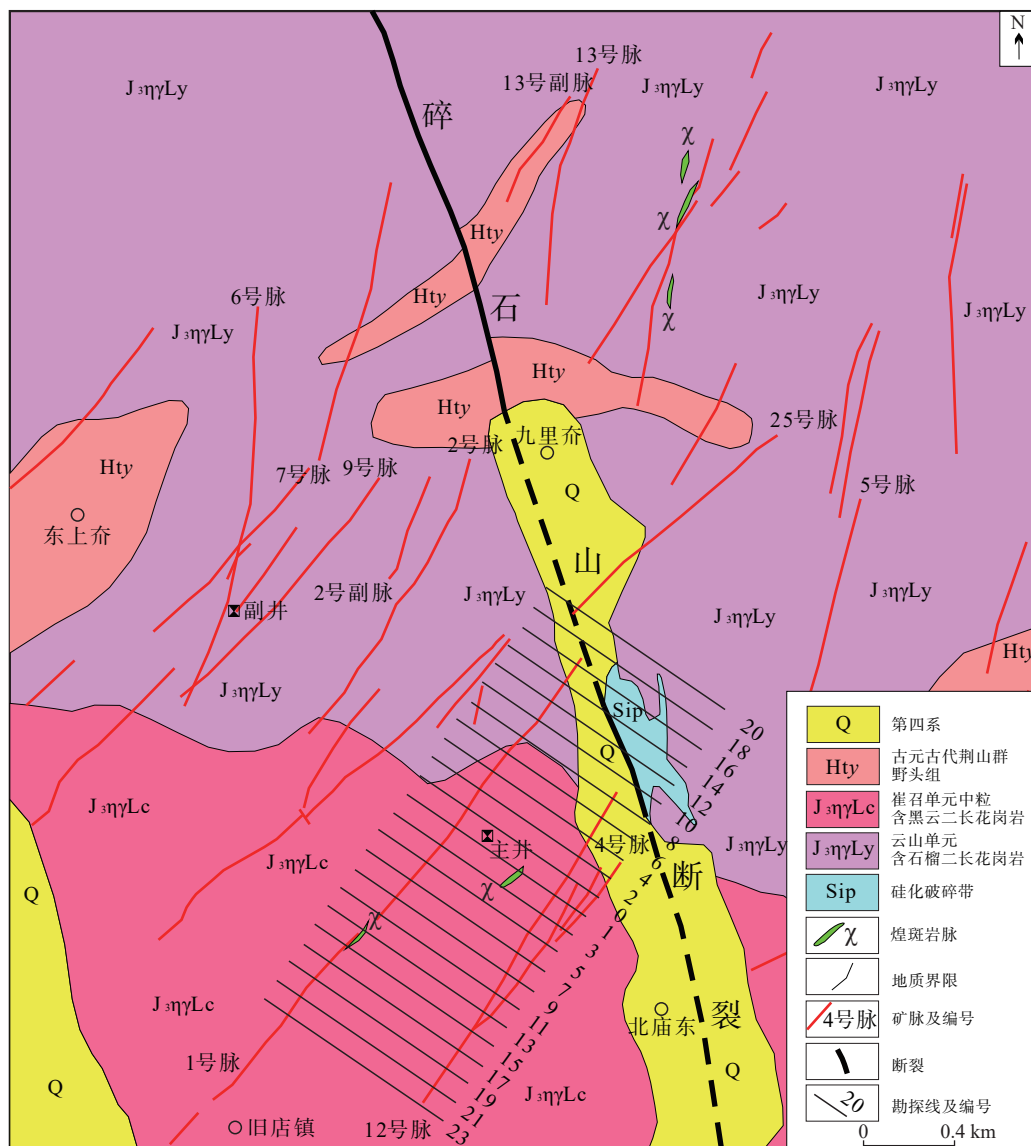


图2 旧店金矿床矿区地质图(引自矿山内部报告)

Fig. 2 Geological map of the Jiudian gold deposit mining area

型以石英脉型为主(图2),局部出现少量交代蚀变岩型,矿体以1号脉规模最大,矿体连续性好,已累计探明金资源量超过20 t(姜文峰等,2008;文志民等,2013),包括I、II两个矿段。其中,II矿段分布在旧店镇至北双山一带,为旧店金矿早期主要开采矿段,受北东向断裂构造控制,区内延伸长度约为1.5 km,走向为 $36^{\circ}\sim 44^{\circ}$,总体走向为 40° ,倾向NW,倾角为 $60^{\circ}\sim 84^{\circ}$,矿体与围岩界限清晰,石英脉两侧黄铁绢英岩化蚀变发育(图3)。

矿石类型以石英脉型、多金属硫化物石英脉型和黄铁绢英岩型为主,矿石结构可见自形-半自形粒状结构、不等粒结构和交代残余结构,矿石构造主要为块状构造、网脉状构造和浸染状构造,矿石金属矿物以黄铁矿为主,其次为黄铜矿、磁黄铁矿、方铅矿、闪锌矿,含有少量的银金矿、金银矿、自然金及自然银,金主要以银金矿形式产出。

通过野外和镜下观察,依据矿石矿物组合、相互间的穿插包裹关系、矿石结构特征等,结合前人研究成果(谢成连等,2008),初步认定成矿过程可划分为早成矿阶段(I)、主成矿阶段(II)、晚成矿阶段(III)3个阶段,早阶段主要发育不含金的石英-黄铁矿脉(图4a),主成矿阶段是金矿化主要发育阶段,从早到晚可以进一步分为3个亚阶段,包括中粗粒黄铁矿阶段(II₁)(图4b、图4d)、细粒-微细粒黄铁矿阶段(II₂)(图4c、图4g)、多金属硫化物阶段(II₃)(图4e、图4h),晚阶段以大量出现无矿石英-碳酸盐脉为特征(图4f、图4i)。

3 样品采集和测试方法

挑选4号脉3中段、5中段、7中段、8中段、14中段和12号矿体7中段、8中段、14中段、16中段涵盖各成矿阶段且有代表性的含黄铁石英脉样品用于H-O-S稳定同位素分析,可以代表各成矿阶段的成矿流体特征。H、O、S同位素分析工作委托核工业北京地质研究院分析测试研究中心完成。样品测试前经严格预处理,包括粉碎,清洗,镜下黄铁矿和石英矿物挑选后进行送样测试。测试过程中,H同位素和O同位素的测试仪器为MAT 253型稳定同位素质谱仪,S同位素的测试仪器为Delta V plus型硫稳定同位素质谱仪。O同位素的测试直接分析了各成矿阶段矿石石英晶格中的O元素,测试方法采用用BrF₅法进

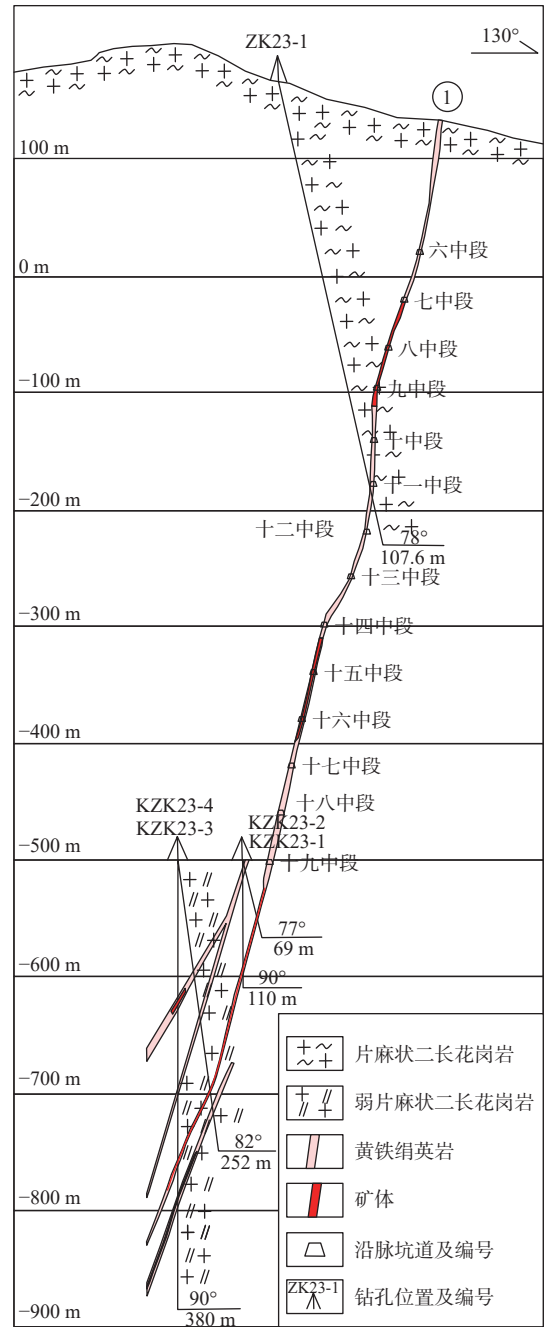
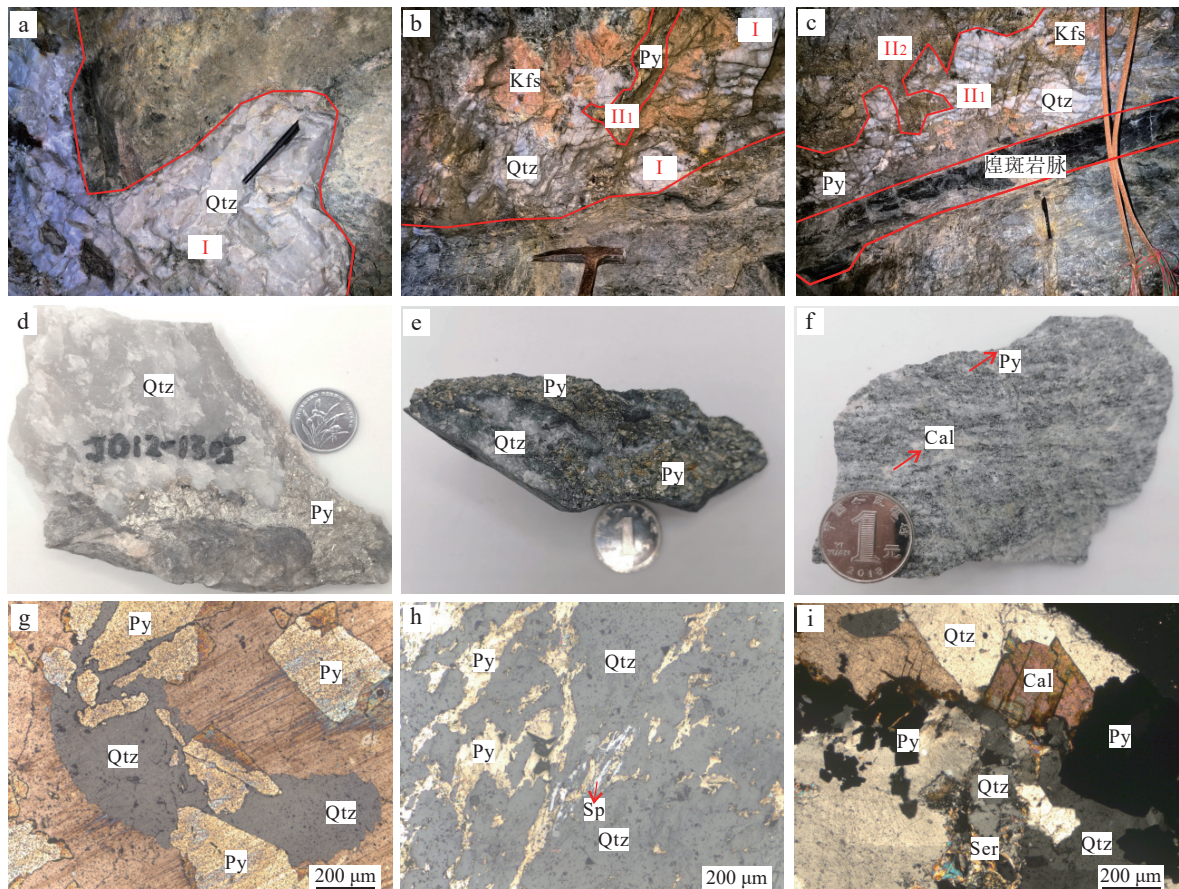


图3 旧店金矿床1号脉II矿段23号勘探线剖面图
(引自矿山内部报告)

Fig. 3 Section diagram of the No. 23 exploration line in the No. 1 ore block of the No. 1 ore vein in the Jiudian gold deposit
H同位素则通过测定石英包裹体中水直接获得,测试方法为爆裂取水 Zn 还原法(刘天航等,2024)。S同位素的分析目标为硫化物中的S元素,测试结果为 $\delta^{34}\text{S}$ 与Cañon Diablo Troilite(CDT)的比值,分析误差为 $\pm 0.2\text{‰}$ 。同位素具体测试的工作由研究中心根据国家同位素检测标准开展,可参考Wang等(2016)的论述。



a.早成矿阶段(I); b.主成矿阶段中粗粒黄铁矿矿体(II₁)侵入早成矿阶段纯白石英脉中(I); d.主成矿阶段细粒-微细粒黄铁矿矿体(II₂)切入中粗粒黄铁矿矿体(II₁); d.中粗粒黄铁矿阶段(II₁); e、h.多金属硫化物阶段(II₃); f、i.晚成矿阶段(III); g.细粒-微细粒黄铁矿阶段(II₂); Py.黄铁矿; Sp.闪锌矿; Qtz.石英; Ser.绢云母; Kfs.钾长石; Cal.方解石

图4 旧店金矿床矿石宏观和显微特征

Fig. 4 Macroscopic and microscopic characteristics of ore in the Jiudian gold deposit

4 分析结果

4.1 石英 H-O 同位素特征

笔者对各成矿阶段的 13 个样品的石英流体包裹体进行了 H-O 同位素分析(表 1), 总体上旧店金矿床石英的 O 同位素 $\delta\text{O}_{\text{V-SMOW}}$ 值为 8.08‰~12.77‰, H 同位素 $\delta\text{D}_{\text{V-SMOW}}$ 值为 -97.7‰~-72.2‰。根据 O 同位素组成, 采用了 Clayton 等(1972)的分馏方程, 结合各成矿阶段内石英中流体包裹体的平均均一温度(牛警徽等, 2023b)限定分馏系数, 计算得出成矿流体的 O 同位素值 $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ 值为 -0.01‰~4.64‰(Sharp et al., 2016)。其中, 早成矿阶段(I)成矿流体 $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ 值为 2.22‰~4.64‰, $\delta\text{O}_{\text{V-SMOW}}$ 值为 10.33‰~12.77‰, $\delta\text{D}_{\text{V-SMOW}}$ 值为 -90.2‰~-72.2‰; 主成矿阶段(II)成矿流体 $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ 值为 1.17‰~4.36‰, $\delta\text{O}_{\text{V-SMOW}}$ 值为

9.27‰~12.48‰, $\delta\text{D}_{\text{V-SMOW}}$ 值为 -84.3‰~-77.1‰; 晚成矿阶段(III)成矿流体 $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ 值为 -0.01‰~4.28‰, $\delta\text{O}_{\text{V-SMOW}}$ 值为 8.08‰~12.48‰, $\delta\text{D}_{\text{V-SMOW}}$ 值为 -97.7‰~-87.1‰。

4.2 黄铁矿 S 同位素特征

对旧店金矿床各成矿阶段的 15 个样品的黄铁矿进行了 S 同位素分析(表 2)。结果显示, 旧店金矿床主成矿阶段黄铁矿 $\delta^{34}\text{S}$ 值大多为 7.65‰~10.10‰(平均值为 8.87‰), $\delta^{34}\text{S}$ 值标准差小, 表明硫源相对均一、稳定, S 同位素分馏较弱(马健等, 2019), 暗示深部岩浆流体是主要硫源。其中, 早成矿阶段(I) $\delta^{34}\text{S}$ 值为 8.46‰~8.67‰(平均值为 8.53‰), 主成矿阶段(II) $\delta^{34}\text{S}$ 值为 7.65‰~10.1‰(平均值为 9.11‰), 晚成矿阶段(III) $\delta^{34}\text{S}$ 值为 7.81‰~9.03‰(平均值为 8.47‰), 各成矿阶段黄铁矿 $\delta^{34}\text{S}$ 值基本一致(图 5)、成矿物质来源相同。

表 1 旧店金矿及其他金矿成矿流体 H-O 同位素组成

Tab. 1 H-O isotope compositions of the ore-forming fluids in the Jiudian gold deposit and other gold deposits

金矿床	样品数	样号	测试矿物	δD_{V-SMOW} (‰)	$\delta^{18}O_{V-SMOW}$ (‰)	温度 (℃)	$\delta^{18}O_{H_2O}$ (‰)	数据来源
旧店	13	I阶段-JD4-0301	石英	-90.2	12.77	270.0	4.64	本文
		I阶段-JD4-0501		-72.6	10.33	270.0	2.22	
		I阶段-JD4-0801		-72.2	11.18	270.0	3.07	
		II阶段-JD4-0802		-81.9	12.39	270.0	4.27	
		II阶段-JD12-0706		-80.5	12.21	270.0	4.09	
		II阶段-JD12-1 601		-80.9	12.48	270.0	4.36	
		II阶段-JD4-0302		-84.3	10.52	270.0	2.41	
		II阶段-JD12-0809		-77.1	12.31	270.0	4.19	
		II阶段-JD12-1408		-78.9	9.27	270.0	1.17	
		II阶段-JD12-1602		-80.4	12.16	270.0	4.04	
		III阶段-JD4-0806		-97.7	12.4	270.0	4.28	
		III阶段-JD4-1402		-87.1	8.08	270.0	-0.01	
		III阶段-JD12-0704		-91.1	12.39	270.0	4.27	
旧店	6	JD-TW1等	石英	-83.4~-74.2	10.7~12.8	300.0	3.81~5.91	Tian et al., 2022
旧店	7	17S32等	石英	-92~-74.2	9.1~13.4	265.2	0.84~5.14	安梦莹, 2022
夏甸	9	XD-wd-1等	石英	-73.22~-46.22	10.4~12.4	240.0	-1.6~2.89	李逸凡等, 2021
夏甸	3	XD-3等	石英	-99.8~-95.1	8.6~9.3	253.0	-0.2~0.5	杜佛光, 2019
夏甸	3	713-41等	石英	-72.1~-66.2	13~14.9	172.0~363.7	4.6~9.2	吴迪, 2016
大尹格庄	11	DYGZ-4-2等	石英	-103~-77.3	8.7~12.2	173.8~353.8	2.75~7.18	严子清等, 2024
大尹格庄	11	Y309Cc3等	石英	-95~-68	7.3~12.7	280.0~350.0	-0.3~7.8	Yang et al., 2010
山后	5	SH-CM4-TW2等	石英	-90.7~-77	10.2~11.1	230.0	0.55~1.25	王巧云等, 2023
山后	4	WH2-1等	石英	-82.1~-78.6	7.5~11.2	215.0	-3.28~0.42	安梦莹, 2022
姜家窑	2	K3等	石英	-66.5~-61.78	7.9~8.5	311.0~344.0	1.57~2.1	王金辉, 2020
曹家洼	3	K12等	石英	-74.16~-61.55	10.4~10.8	299.0~333.0	3.89~5.33	王金辉, 2020
谢家沟	15	XJG19等	石英	-106~-85	7.7~11.3	209.0~351.0	0.8~5.7	Du et al., 2023
三山岛	2	S615-1等	石英	-83.1~-79.1	13.55~13.67	257.0~335.0	7.49~7.61	王金雅, 2020
焦家	5	J450-1等	石英	-95.8~-68.3	10.4~14.58	200.0~400.0	3.94~8.38	辛洪波, 2005
新立	9	XJG19等	石英	-77.63~-68.77	11.4~14.17	165.0~310.0	-0.17~7.02	毛兴强, 2022

5 讨论

5.1 成矿流体来源

笔者收集并整理前人关于旧店金矿床及胶东地区其他主要金矿床的 H-O 同位素数据(图 6, 表 1)。H-O 同位素特征表明, 旧店金矿床的成矿流体与招平断裂带内夏甸(吴迪, 2016; 杜佛光, 2019; 李逸凡等, 2021)、大尹格庄(Yang et al., 2010; 严子清等, 2024)、山后(安梦莹, 2022; 王巧云等, 2023)等金矿床, 以及三山岛(王金雅, 2020)、焦家(辛洪波, 2005)和新立(毛兴强等, 2022)等其他金成矿带内金矿床, 其成矿流体性质具有高度相似性。除焦家、新立、夏甸和姜家窑(王金辉, 2020)等金矿床少量 H-O 同位素落点于

岩浆水区域, 大部分金矿床 H-O 同位素特征位于岩浆水附近且向大气降水线逐步靠近, 表明成矿过程中旧店金矿床等胶东地区金矿床成矿流体表现出岩浆水参与的显著特征(李逸凡等, 2021), 同时在成矿过程中大气降水的参与逐渐增多。

此外, Tian 等(2022)和 Sun 等(2022)在对旧店金矿床各成矿阶段包裹体测温结果显示, 从早期到晚期, 旧店金矿床成矿总体上呈现中温(198.5~329.2 ℃)和中低盐度(3.95~12.50 wt% NaCl_{eqv}), 温度从早期(303.1~329.2 ℃)到主成矿期(240.3~319.7 ℃)到晚期(198.5~311.7 ℃)呈现不断降低的趋势, 表明成矿流体向上运移过程中, 温度和压力逐步降低, 可能触发了流体的沸腾作用(Du et al., 2024), 导致流体相分离(倪培等, 2020)和金属沉淀(张德会, 1997), 进而

加速了成矿流体系统逐步转变为开放系统的进程、导致大气降水的混入。综上所述,胶东地区金矿床的成矿流体在整体性质上差异不大,均为岩浆水与大气降水的混合体系(Wen et al., 2016),随着成矿阶段的推进,大气降水的加入比例呈现逐渐增多的趋势,这进一步指示胶东地区金矿床成矿流体体系的一致性。

5.2 金属来源

本次研究同时收集了胶东地区部分典型金矿床和花岗岩体的 S 同位素数据结果见表 2, 并绘制了胶东部分金矿床和地质单元 S 同位素组成(图 7)。一方面, 旧店金矿床的 S 同位素特征与夏甸金矿、焦家

金矿和新城金矿等胶东地区大型金矿床 S 同位素特征基本一致,表明旧店金矿床主要载金矿物来源与招平断裂带内乃至胶东地区其他金矿床基本一致。另一方面,旧店金矿床的 S 同位素组成(7.65‰~10.10‰)主要与胶东群(6.9‰~9.4‰)(张潮等, 2014)和荆山群(9.3‰~9.8‰)(张竹如等, 1999)等老地层 S 同位素特征重叠;且与玲珑花岗岩(6.1‰~10.1‰)(毛景文等, 2005)、郭家岭花岗岩闪长岩(2.4‰~9.7‰)(宋明春等, 2013)也有一定的重合;此外,早白垩世晚期的中基性脉岩的侵入时代(116 Ma~124 Ma)是胶东地区主金成矿期(Song et al., 2020; Deng et al., 2023),其

表 2 旧店金矿床及其他地质单元 S 同位素特征

Tab. 2 Sulfur isotope characteristics of the Jiudian gold deposit and other geological units

金矿/地质单元	样品数	样号	矿物/对象	$\delta^{34}\text{S}$ 值(‰)	数据来源
旧店金矿	13	I阶段-JD4-0301	黄铁矿	8.46	本文
		I阶段-JD4-0501		8.67	
		I阶段-JD4-0801		8.46	
		II阶段-JD4-0802		10.1	
		II阶段-JD12-0706		9.51	
		II阶段-JD12-1601		7.65	
		II阶段-JD4-0302-1		9.23	
		II阶段-JD4-0302-2		9.15	
		II阶段-JD12-0809		8.29	
		II阶段-JD12-1408-1		9.75	
		II阶段-JD12-1408-2		9.63	
		II阶段-JD12-1602		8.72	
		III阶段-JD4-0806		7.81	
玲珑金矿 东风金矿 大尹格庄金矿 夏甸金矿 夏甸金矿2 谢家沟金矿 焦家金矿1 焦家金矿2 三山岛金矿1 三山岛金矿2 新城金矿 新立金矿 金青顶金矿 大柳行金矿 胶东群 荆山群 玲珑花岗岩 郭家岭花岗岩闪长岩 中基性岩脉	5 5 10 7 10 42 5 8 7 20 33 6 8 6 3 6 8 5 6	III阶段-JD4-1402	黄铁矿	8.56	Wen et al., 2015 Wen et al., 2015 严子清等, 2024 Liu et al., 2018 李逸凡等, 2021 Du et al., 2023 Qiu et al., 2023 陈阳阳, 2017 姜晓辉等, 2011 Qiu et al., 2023 张潮等, 2014 毛兴强等, 2022 李旭芬, 2011 黄鑫, 2021 张潮等, 2014 张竹如等, 1999 毛景文等, 2005 宋明春等, 2013 李俊建等, 2006
		III阶段-JD12-0704		9.03	
		JQ-Q-04等		5.9~7.4	
		10LL18等		5.8~7.0	
		DYGZ-4-2等		3.1~7.6	
		713-1等		6.95~8.1	
		XD-wd-1等		6.3~8.1	
		XJG34-2-1-Py等		5.7~9.0	
		17JJ04-1等		8.7~11.2	
		\		8.6~11.3	
		09S37等		11.5~12.4	
		17SSD-1-1等		9.4~13.0	
		XC12D001B1等		5.7~10.6	
		XL48508-21等		7.7~13.2	
		J-335-05等		6.95~8.69	
		DLH-1等		6.3~7.7	
		XC10D218B1等		6.9~9.4	
		\		9.3~9.8	
		LL-190-9等		6.1~10.1	
		\		2.4~9.7	
		\		5.4~11.2	

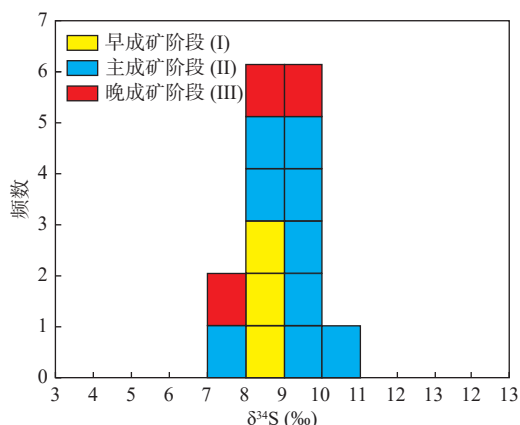


图5 旧店金矿床黄铁矿硫同位素直方图

Fig. 5 Histogram of sulfur isotopes of pyrite in the Jiudian gold deposit

在空间分布上多与矿体控矿断裂产状基本一致,且其S同位素组成与金矿床矿石中硫化物相似(图7)(李俊建等, 2006; 冯李强, 2022),表明胶东地区中基性脉岩成岩物质与金矿床成矿物质也可能具有相同的物质源区(Li et al., 2016; Yuan et al., 2019)。尽管上述研究表明这些地质体与成矿作用的直接关系较弱,但从太古代胶东岩群到中生代玲珑花岗岩体,再到旧店金

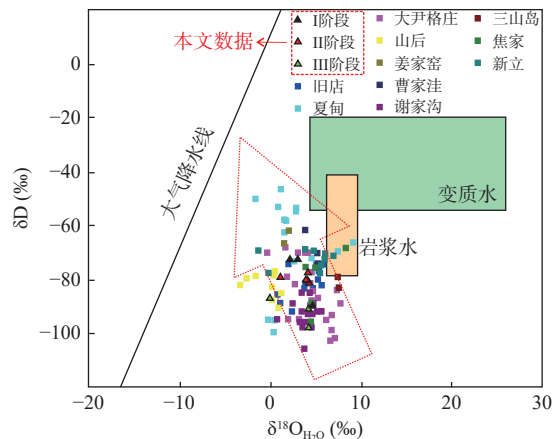
图6 部分金矿床成矿流体 $\delta D-\delta^{18}O_{H_2O}$ 同位素组成

Fig. 6 Isotope compositions of $\delta D-\delta^{18}O_{H_2O}$ of the ore-forming fluids in some gold deposits

矿床的S同位素组成变化来看,存在逐渐增大的趋势(表2, 图7)。上述证据表明,旧店金矿床乃至胶东地区金矿床的S同位素组成不仅是对胶东群、荆山群、玲珑花岗岩以及郭家岭花岗闪长岩等地层和侵入体S同位素特征的继承,而且在这一过程中有逐步富集的趋势。近年多,部分学者通过胶东地区六条金成矿带载金黄铁矿中相关均一的 $\delta^{199}Hg$ 组成(平均值为

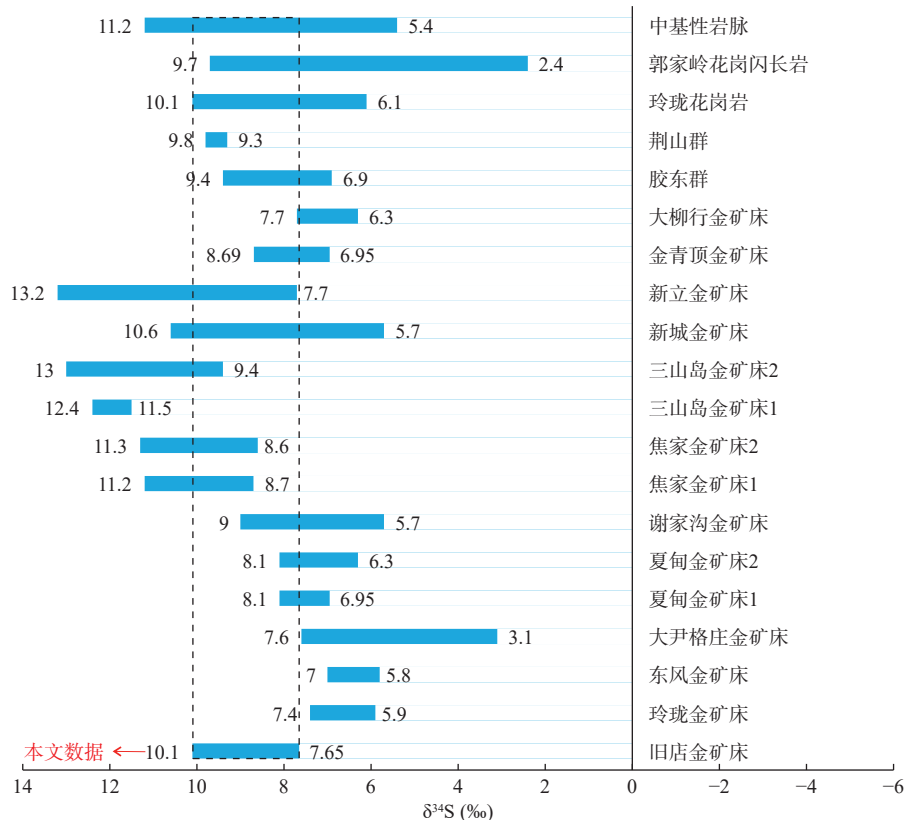


图7 胶东部分金矿床和地质单元S同位素组成

Fig. 7 Sulfur isotope compositions of some gold deposits and geological units in Jiaodong region

0.012‰)(Wang et al., 2024; Zhang et al., 2024)以及胶东地区典型金矿中黄铁矿多硫同位素中相对恒定的 $\delta^{33}\text{S}$ 组成(平均值为 -0.04‰)(Qiu et al., 2023)分析认为胶东地区含金流体主要来源于深部源区,同时交代岩石圈成矿物质。综合成矿流体特征认为,旧店金矿床成矿物质以幔源为主,上涌过程中活化并交代荆山群变质岩和玲珑花岗岩的部分成矿物质,进而达到S同位素的均衡。

5.3 成矿模式

目前研究成果显示旧店金矿床内至少发育有新元古代(709.1 Ma)花岗质片麻岩、晚侏罗世(146.5 Ma~164.1 Ma)玲珑花岗岩和早白垩世(110.6 Ma~123.0 Ma)脉岩3期岩浆(Bi et al., 2017; 牛警徽等, 2023)活动,胶东地区金成矿时代与早白垩世中基性脉岩成岩时代高度吻合(Yang et al., 2015; Wang et al., 2022),且岩浆岩锆石 Hf 同位素特征显示玲珑花岗岩成岩物质主要

来自前寒武纪变质基底岩石的部分熔融(Cho et al., 2021; 李经纬等, 2023; Chen et al., 2024),与玲珑花岗岩和荆山群等变质岩S同位素 $\delta^{34}\text{S}$ 值有一定重合的特征一致,结合区内成矿流体特征和成矿物质来源,初步建立旧店金矿床成矿模式(图8)。

第一阶段为~709 Ma(图8a),在旧店矿区,花岗质片麻岩侵入到古元古代荆山群冶头组黑云变粒岩中,标志着旧店金矿床岩浆活动的开始。第二阶段为164~146 Ma(图8b),晚侏罗世,扬子板块向华北板块持续俯冲(刘晓阳等, 2020; Chen et al., 2024),造成地壳增厚发生熔融,在胶东地区生成大量的地壳重熔型玲珑花岗岩,岩性主要包括黑云母二长花岗岩、含石榴子石二长花岗岩、花岗伟晶岩脉和花岗细晶岩脉等。第三阶段为123 Ma~110 Ma,是旧店金矿床形成阶段,早白垩世晚期,受晚中生代时期太平洋板块向华北板块俯冲影响(Zhang et al., 2020),旧店金矿床所处的胶

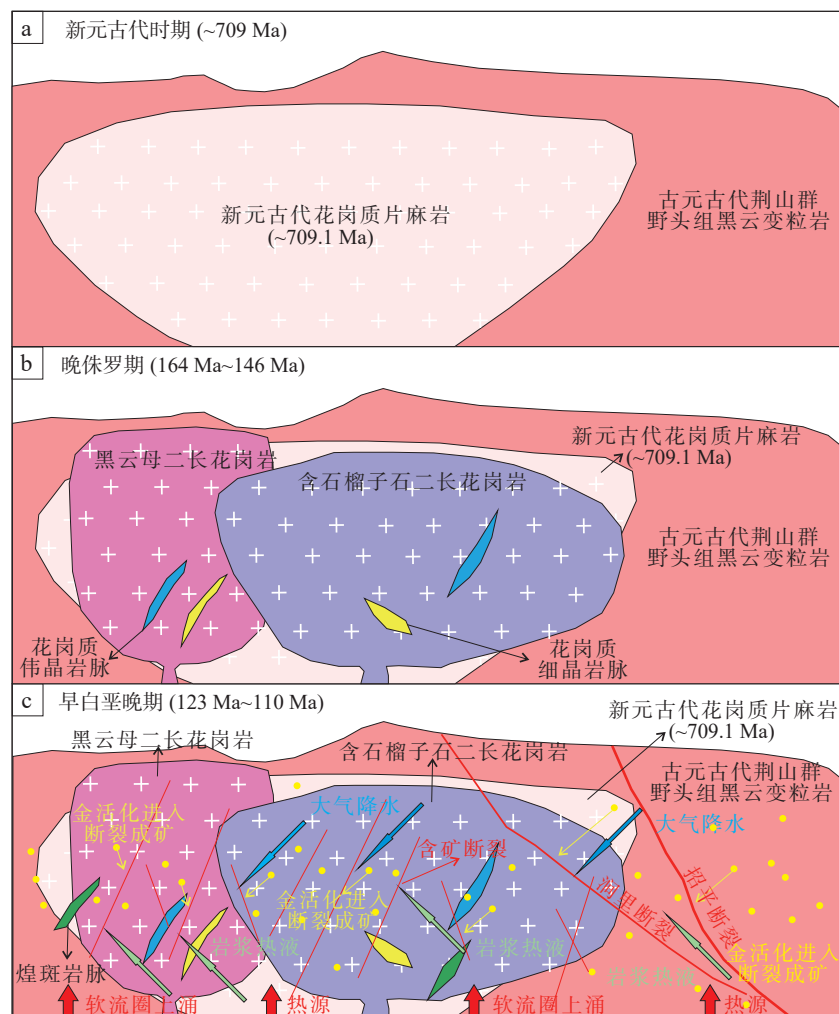


图8 旧店金矿床成矿模式示意图

Fig. 8 Schematic diagram of the mineralization model of the Jiudian gold deposit

东地区发生华北克拉通东部岩石圈大规模伸展-减薄和软流圈上涌(Yang et al., 2016; Ma et al., 2017)(图 8c),在地壳产生大量的断裂(宋明春等, 2022a),大范围的幔源岩浆活化了古元古代荆山群变质岩残留体、玲珑花岗岩体以及中基性脉岩中的 Au 物质,在岩浆水为主、大气降水为辅的混合流体作用下在招平断裂带下盘次级断裂构造中富集成矿。

6 结论

(1)旧店金矿床矿体主要赋存于 NE 向和 NNE 向断裂构造中,成矿过程可分为早成矿阶段(I)、主成矿阶段(II)、晚成矿阶段(III)3 个阶段。其中,主成矿阶段包括中粗粒黄铁矿阶段(II₁)、细粒-微细粒黄铁矿阶段(II₂)和多金属硫化物阶段(II₃)3 个亚阶段。

(2)不同成矿阶段的石英 H-O 同位素分析显示,早成矿阶段成矿流体 $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ 值为 2.22‰~−4.64‰, $\delta\text{D}_{\text{V-SMOW}}$ 值为−90.2‰~−72.2‰,主成矿阶段成矿流体 $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ 值为 1.17‰~−4.36‰, $\delta\text{D}_{\text{V-SMOW}}$ 值为−84.3‰~−77.1‰,晚成矿阶段成矿流体 $\delta^{18}\text{O}_{\text{H}_2\text{O}}$ 值为−0.01‰~4.28‰, $\delta\text{D}_{\text{V-SMOW}}$ 值为−97.7‰~−87.1‰。以上特征表明,早阶段成矿流体以岩浆水为主,随着流体演化,大气降水逐步混入且所占比例不断增加。

(3)不同成矿阶段的黄铁矿 S 同位素无明显差异, $\delta^{34}\text{S}$ 值大多为 7.65‰~10.10‰(平均值为 8.87‰),这一范围与前寒武纪变质基底和中生代花岗岩特征十分类似,表明旧店金矿床成矿物质来源于前寒武纪变质基底和中生代花岗岩。

(4)综合前人研究成果,建立了旧店金矿成矿模式:晚侏罗世,旧店地区大量含金的老变质岩基底发生重熔,形成玲珑型花岗岩;在早白垩晚期受华北克拉通东部岩石圈大规模伸展-减薄和软流圈上涌影响,花岗岩和古老变质基底中的金发生活化富集,并在岩浆水和大气降水组成的混合流体作用下,在招平断裂带下盘次级断裂构造中富集成矿。

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