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东昆仑夏日哈木矿区岩体含矿性特点与形成机理探讨

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摘要:东昆仑造山带有众多基性-超基性岩体,但目前仅在夏日哈木 I 号岩体发现有超大型镍矿。对其含矿性特点与形成机理的研究有助于东昆仑地区的铜镍矿勘查。笔者梳理了目前诸多学者对夏日哈木铜镍矿的研究工作,并结合野外实际,发现铁质系列的镁铁-超镁铁质岩体才可能形成有经济价值的岩浆铜镍矿体,镁质系列的镁铁-超镁铁质岩体不可能成就大规模岩浆铜镍硫化物矿体;大断裂为地幔岩浆上侵提供了有利条件;分异良好的超基性岩体更具有成铜镍矿的潜力;结晶分异促进了硫化物饱和,地壳硫混染是成矿的关键;橄榄石和单斜辉石具有低的 CaO 和 FeO 含量的岩体特征,更有利于成铜镍矿。

关键词:镁铁-超镁铁质岩体;铜镍矿;含矿特点;成矿机理;夏日哈木

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Mineralization Characteristics and Formation Mechanism of the Intrusions in Xiarihamu Magmatic Ni-Cu Sulfide Deposit, East Kunlun Orogenic Belt, Northwest China

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Abstract: Some mafic-ultramafic intrusions have been found in the East Kunlun Orogenic Belt, but only the Xiarihamu I intrusions produces the giant Ni deposit. The study on ore-bearing characteristics and formation mechanism of the Xiarihamu I intrusion contributes to the exploration of copper-nickel deposits in the east Kunlun area. The research work of many scholars on the Xiari-

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hamu nickel ore deposit has been summarized in this paper. Based on the field investigation, it's found that the iron series mafic-ultramafic intrusions can form the magmatic copper-nickel ore body with economic value, but magnesia series mafic-ultramafic intrusions are unlikely to achieve copper-nickel sulfide ore bodies. Large faults provide favorable conditions for the intrusion of mantle magma. The well-differentiated mafic-ultramafic intrusions have the better potential to form a copper-nickel ore. Crystallization differentiation promotes sulfide saturation, and the crystal sulfur contamination is the key to mineralization. The intrusions with olivine and clinopyroxene have low CaO and FeO content, which is more conducive to copper-nickel ore.

Keywords: mafic-ultramafic intrusion; magmatic Ni-Cu sulfide deposit; mineralization characteristics; metallogenetic mechanism; Xiarihamu

岩浆铜镍硫化物矿床大都赋存于镁铁-超镁铁质岩体内,其岩浆属于铁质系列的基性-超基性岩浆。该类矿床一般多发育于稳定陆块边缘裂谷系统,或与大火成岩省的岩浆活动密切相关,是深部地幔上涌硫化物不混溶作用发生的具体表现(ZHANG et al., 2018)。在中国,岩浆铜镍硫化物矿床还发育在造山带内,如新疆东天山的黄山岩带,产出了黄山、黄山东、图拉尔根等大型岩浆铜镍硫化物矿床和东北吉林红旗岭矿床(郗爱华等,2006;李文渊,2015,2018)。尤其是近几年在东昆仑造山带新发现了夏日哈木岩浆铜镍硫化物矿床,120万t镍金属储量已达到超大型规模(张照伟等,2015;LI et al., 2015;ZHANG et al., 2017),这也是继1996年加拿大沃尔斯贝(VOISEY'S BAY)岩浆硫化物矿床发现以来近20年全球镍矿最重要的发现(张照伟等,2016)。这无疑会引起矿床学家研究的兴趣和热情。青海夏日哈木镍矿床在大地构造位置上处于东昆仑造山带的昆中带古生代岛弧带内(姜春发等,1992;李荣社等,2008),造山带内也能产出超大型岩浆铜镍硫化物矿床,其铁质系列的基性-超基性岩浆是如何形成和演化的呢?不仅如此,在夏日哈木矿区发育有5个镁铁-超镁铁质岩体,但仅有I号岩体含矿,是一个低铜贫铂族元素的岩浆硫化镍矿床(丰成友等,2016),其余II、III、IV、V号岩体均不含有经济价值的工业矿体,具体是一种什么样的构造背景和成矿机理形成的呢?如对该问题能够客观回答,一方面可以提升造山带背景岩浆成矿理论的认识水平,另一方面可极大推进东昆仑造山带岩浆铜镍硫化物矿床的找矿实践与新突破。

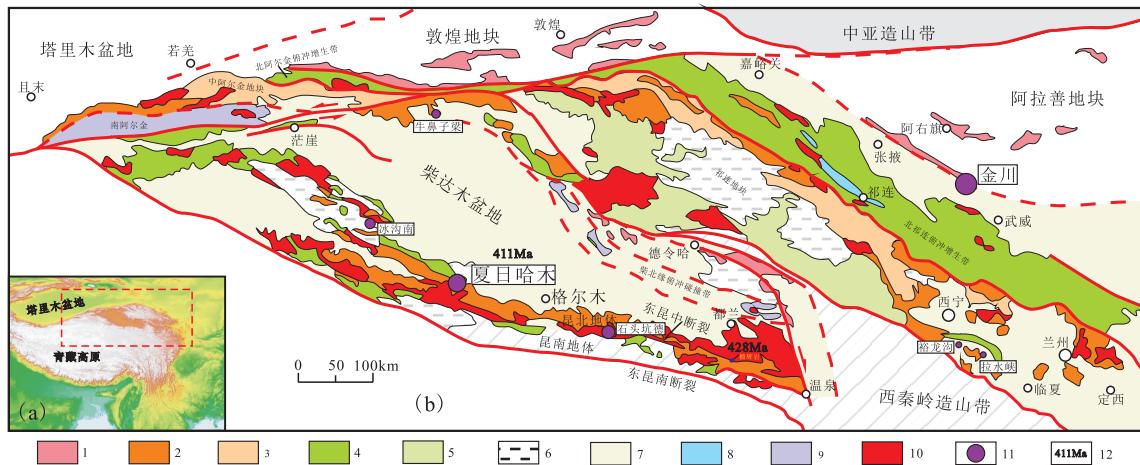
1 区域地质构造背景

东昆仑造山带整体位于青藏高原的东北部,柴

达木盆地南缘(图1a)。东昆仑造山带整体区域构造演化先后经历5个阶段:分别是太古宙—古元古代古陆核形成阶段、中新元古代古大陆裂解与超大陆汇聚阶段、南华纪—早古生代洋陆转换阶段、晚古生代—早古生代洋陆转换阶段以及中新生代陆内造山阶段(ZHANG et al., 2014;校培喜等,2014;钱兵等,2015;张玉等,2017;何书跃等,2018)。东昆仑造山带依据区域性大断裂划分了几个大的次级构造单元:昆北断裂带、黑山-那陵格勒断裂、昆中断裂带以及昆南断裂带。将东昆仑自北而南分为祁漫塔格早古生代岩浆弧、中昆仑微陆块以及昆南增生楔杂岩带3个部分(图1b)。在昆北造山带的东部,发现有榴辉岩约为428 Ma(MENG et al., 2013,2015)。其他几处蛇绿混杂岩的年龄变化为467~518 Ma,并且这些蛇绿混杂岩的玄武质岩石表现出了典型的MORB特征(BIAN et al., 2004;宋谢炎等,2009;姜常义等,2015)。由此推测,昆北造山带地体大约于428 Ma拼贴到柴达木克拉通的南部边缘。在昆北地体的北部边界产出了夏日哈木超大型岩浆铜镍硫化物矿床(图1b)(张照伟等,2015,2016,2017;LI et al., 2015;SONG et al., 2016;ZHANG et al., 2017; LIU et al., 2018)。矿区出露地层主要为古元古代白沙河岩群;岩石类型为黑云斜长片麻岩、眼球状混合片麻岩、大理岩、二云石英片岩等;原岩恢复为碎屑岩-碳酸盐岩-火山岩建造;经历了角闪岩相区域变质作用(张雪亭等,2007;耿林等,2007;李荣社等,2008;范丽琨等,2009;校培喜等,2014;张照伟等,2015,2017;LIU et al., 2018;孙小攀等,2018)。

2 岩体地质特征

夏日哈木矿区已发现的镁铁质-超镁铁质岩体共5个,呈北西向带状展布,单个岩体出露面积较



1. 太古宙—古元古代变质基底;2. 中新元古代深变质岩石;3. 新元古代浅变质岩;4. 早古生代俯冲增生杂岩;5. 早古生代沉积建造;6. 晚古生代—中生代沉积建造;7. 新生代沉积岩;8. 高压低温变质岩;9. 超高压变质岩;10. 花岗岩;11. 铜镍硫化物矿床;12. 成岩年龄

图1 东昆仑岩浆铜镍硫化物矿床区域地质构造简图(据孟繁聪等,2017;张照伟等,2018)

Fig. 1 The sketch regional geological map of magmatic Ni - Cu sulfide deposits in eastern Kunlun

小,最大出露面积仅为 0.9km^2 ,主要呈岩盆状或岩墙状侵位于元古代金水口群变质岩系及新元古代花岗片麻岩中(图2)。其中,I号岩体为含矿岩体,其余4个II、III、IV、V号岩体矿化较弱,尽管II号岩体地表见较强的铜镍矿化,但仍未发现有经济价值的矿体。

夏日哈木矿区I号岩体据钻孔施工及勘探工作,初步确定其岩体形态,长约为1.4 km,宽约为0.6 km,长轴方向近东西向,西段略向南偏转(图3a),岩体顶界面东高西低,东段出露于地表,西段隐伏于地下,且越向西埋深越深(图3b),总体形态为向西倾伏的岩床(图3)。夏日哈木矿区I号岩体在地表有氧化蚀变带及铁帽出露,主要集中在0号勘探线东西两侧(图3a)。岩体岩性主要是辉石岩、橄榄岩、辉长岩及少量的花岗岩脉(图3a、图3b),并且橄榄岩越向西橄榄石含量逐渐增多,同时埋深加大,围岩地层厚度增厚。在钻孔横剖面图上,上述情况则更加明显直观(图3b),沿着NM线 248° 方向,岩石基性程度变深,岩体埋深增大,橄榄石含量增多,矿体增厚变富(张照伟等,2015,2016)。

II号岩体位于矿区东侧,地表出露2个露头,围岩为金水口群片麻岩。其中东侧II-1号岩体呈北东东向展布,走向约 75° ,长约为550 m,宽为50~240 m,出露面积约为 0.15km^2 。岩体主要由辉长岩组成,其次含少量辉石岩。辉长岩中可见二长花

岗岩的捕虏体(图2)。地表岩石普遍发育透闪石化、蛇纹石化和碳酸盐化蚀变,局部可见孔雀石、镍华、褐铁矿、磁黄铁矿、镍黄铁矿、黄铜矿等矿化。西侧II-2号岩体呈东西向展布,走向约为 90° ,长约为400 m,宽约为200~350 m,出露面积约为 0.1km^2 。主要由辉长岩组成,局部见有少量的辉石岩,辉石岩普遍有碳酸盐化、褐铁矿化,地表探槽中仅见有少量镍黄铁矿化,镍黄铁矿呈团块状、星点状分布。

III号岩体位于矿区北西部,岩体形态呈圆形展布,出露面积约为 0.35km^2 ,侵位于金水口群变质岩系中。岩体主要由中细粒的蛇纹岩和石榴子石斜长角闪岩组成,其次含少量辉石岩(图2)。根据地表槽探揭露和少量钻孔验证,蛇纹岩普遍具有弱镍矿化,偶见铜镍矿化的辉石岩转石。

IV号岩体位于矿区中部偏南,围岩为金水口群片麻岩。岩体呈长条状,近东西向展布,长约为0.7 km,宽约为60~150 m,面积约为 0.7km^2 。岩体主要由蛇纹岩和辉长岩组成,局部为辉石岩(图2)。在岩体内见有少量分布的金水口群白沙河组和花岗岩体,经初步槽探揭露,岩体中偶见有镍黄铁矿和黄铁矿。

V号岩体位于矿区最南端,围岩为金水口群片麻岩(图2)。岩体呈漏斗状,北宽南窄。岩体主要由蛇纹岩组成,局部偶见辉石岩,岩体中未见镍黄铁矿化和黄铁矿化。

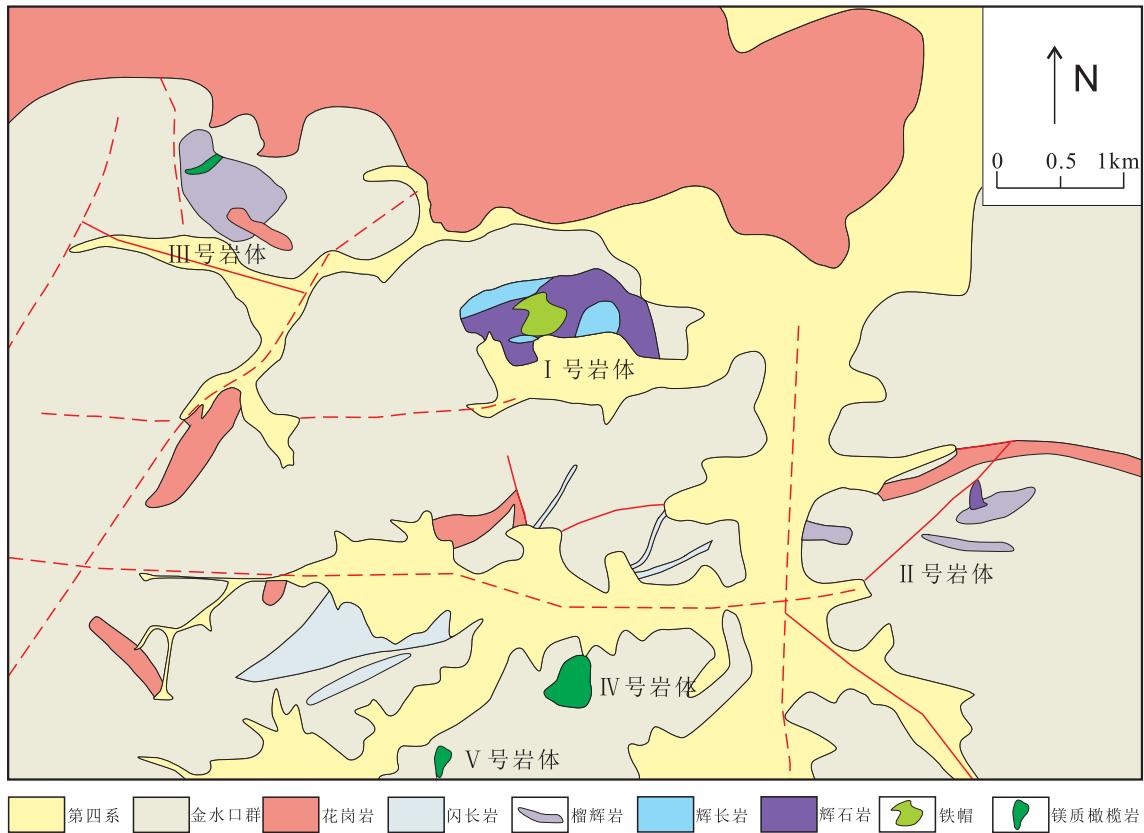


图 2 东昆仑夏日哈木岩浆铜镍硫化物矿区岩体地质分布略图(据张照伟等, 2015)

Fig. 2 Schematic geological map for mafic – ultramafic intrusions from Xiarihamu magmatic Ni – Cu deposit in Eastern Kunlun orogenic belt (After ZHANG, et al., 2015)

3 岩体含矿性特点

在东昆仑夏日哈木铜镍矿区中,仅有 I 号岩体含有岩浆铜镍硫化物矿体,其他 4 个岩体均不含矿。夏日哈木 120 万 t 镍金属全部集中在 I 号岩体内。II 号和 III 号岩体主体是榴辉岩,在 I 号岩体的深部钻孔岩心内也见有榴辉岩。该榴辉岩的年龄为 440Ma(张照伟等, 2017)。IV 号和 V 号岩体主体为蛇纹岩,或称为镁质橄榄岩,是冷侵位构造活动的产物,这类镁铁-超镁铁质岩体基本不含铜镍矿。即便是表现有微弱程度的铜镍矿化,也基本是橄榄石结晶时所包裹的 Ni 元素,无利用价值。

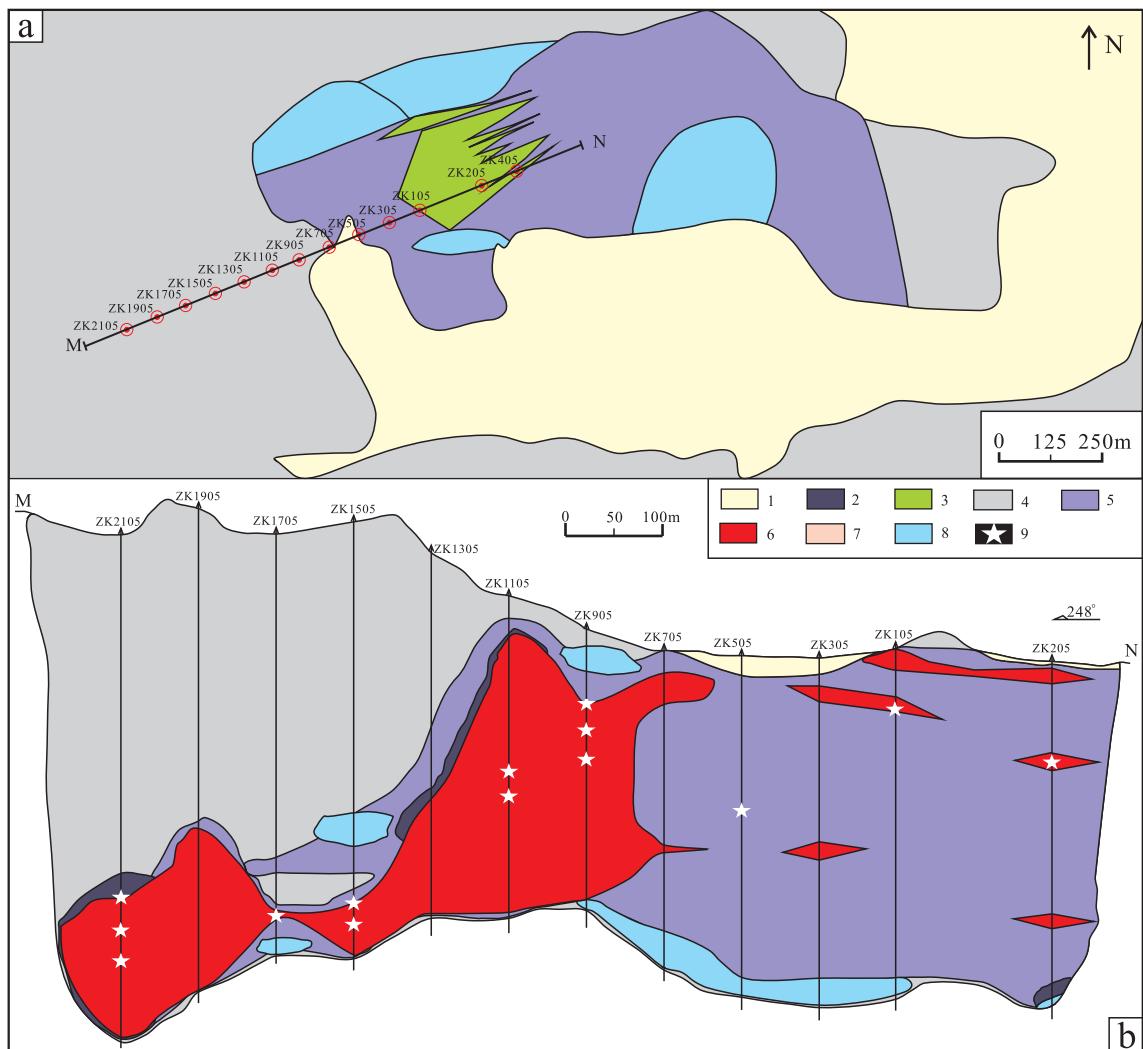
夏日哈木铜镍矿体主要赋存于 I 号岩体的辉石岩与橄榄岩中,主矿体位于 2 号勘探线以西地表以下的空间区域,在 9 号勘探线、11 号勘探线的位置区域,铜镍矿体达到了最厚(超过 300 m),随着勘探线号的变大(向西),岩体变薄,埋深增厚,橄榄石增

多,铜镍矿体变富(图 3b)(张照伟等, 2015; LI et al., 2015; ZHANG et al., 2017)。基本不含矿的辉长岩与含矿的辉石岩及橄榄岩明显不是同期的产物,岩体形成时代也说明了这一点。无矿辉长岩形成时代为 431Ma,含矿辉石岩形成时代为 411Ma(张照伟等, 2015, 2016, 2017; LI et al., 2015; SONG et al., 2016; ZHANG et al., 2017)。矿石类型以稠密浸染状(图 4a)和团块状(图 4b)为主;矿石矿物主要是镍黄铁矿、磁黄铁矿及少量的黄铜矿;结构构造中可见明显的橄榄石被辉石包裹的典型包橄结构(图 4c、图 4d)。

4 成矿机理探讨

4.1 构造环境分析

镁铁-超镁铁岩是幔源岩浆结晶分异的产物,多产在大陆边缘裂谷、克拉通裂谷、岛弧裂谷、后碰撞环境中(张照伟等, 2015; LI et al., 2015; SONG et



1. 新生代盖层;2. 橄榄岩;3. 铁帽;4. 变质岩基底;5. 辉石岩;6. 铜镍硫化物矿化;7. 花岗岩;8. 辉长岩;9. 采样位置

图3 (a)东昆仑夏日哈木矿区I号岩体平面及(b)剖面地质略图

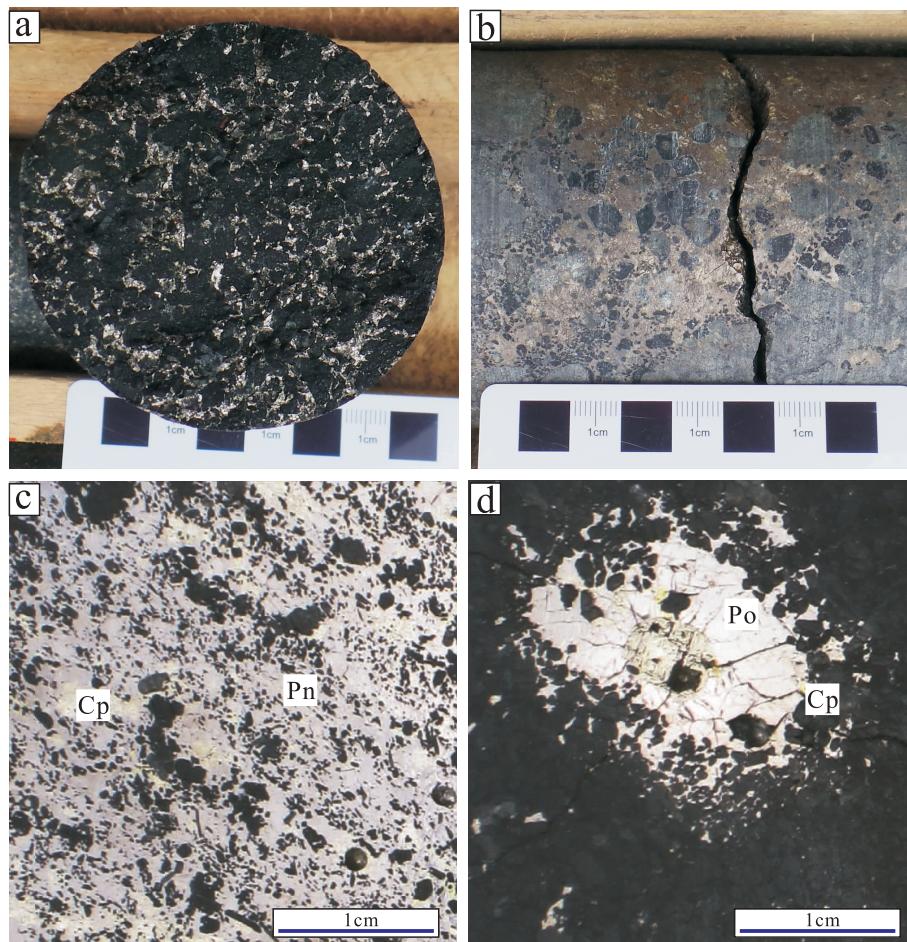
Fig. 3 (a) Schematic geological map for plan and (b) profile from Xiarihamu intrusion I in Eastern Kunlun orogenic belt

al., 2016; ZHANG et al., 2017; LIU et al., 2016, 2018)。控制岩浆硫化物矿床成矿作用的首要因素是具有连通地幔的断裂或通道,需要有超壳深大断裂作为导矿构造。东昆仑夏日哈木岩体位于黑山-那陵格勒断裂和昆中断裂附近。昆中断裂为岩石圈破裂带,呈北西—南东向,大致与昆北断裂平行展布,其东端起于格尔木以东,向西经过小灶火、那陵格勒河,穿越库木库里盆地,最终汇于呈北东向展布的白干湖断裂。昆中断裂呈北西—南东向,东段近西—东向展布,是昆中带和昆南带的分界断裂(图1)。其西端起于博卡雷克塔格北坡,向东穿过大干沟至鄂拉山,东段则被温泉-哇洪山断裂所截,西延部分可能与新疆的奥依塔格-库地断裂相接。区

域地球物理资料显示,沿该断裂带重力梯度明显,地震测深显示该断裂下延切割莫霍面。莫霍面深度和岩石圈深度均显示南高北低,落差分别为8 km和20 km(SONG et al., 2006; 祁生胜等, 2014),且这种地幔顶界面的落差可能在志留纪就已经存在(姜常义等, 2015)。所以昆中断裂是一条规模巨大的超岩石圈断裂,为深部岩浆上侵并形成铜镍矿体提供了必要的条件。

4.2 硫化物不混溶作用

夏日哈木I号岩体存在纯橄榄岩→辉石橄榄岩→辉石岩的分异过程,且有大量的斜方辉石。结晶分异过程有利于铜镍矿的硫化物饱和(LI et al., 2013; LIU et al., 2017, 2018)。硫化物饱和是幔源



a. 稠密浸染状矿石;b. 团块状矿石;c,d. 矿石显微照片;Cp. 黄铜矿;Po. 磁黄铁矿;Pn. 镍黄铁矿

图 4 夏日哈木铜镍矿区矿石及显微照片

Fig. 4 Microphotographs of rocks and ore minerals in Xiarhamu magmatic Ni-Cu sulfide deposit

岩浆形成铜镍硫化物矿床的关键机制。岩浆铜镍硫化物矿床是幔源岩浆经历了复杂的演化最终达到硫化物饱和而发生硫化物熔离的产物(NALDRETT, 2004, 2009)。岩浆中硫化物的熔离需要岩浆体系中的硫含量达到一定含量,这个临界值称为硫化物饱和时的 S 含量[sulfur content at sulfide saturation (SCSS)]。在结晶分异进行过程中,SCSS 在逐渐降低,该值降低有利于硫化物的出现,进而促进了硫化物的饱和(LIU et al., 2019),这也解释了为什么分异良好的基性-超基性岩体具有好的成矿潜力。

地壳混染过程主要从 2 个方面影响硫化物饱和:①混染过程导致岩浆中的主量元素含量变化,如无硫化物大理岩的混染对新疆坡北的硫化物饱和起到了抑制作用,主要原因在于无硫化物大理岩的混染使岩浆体系的 CaO 含量升高, CaO 升高会抑制

硫化物饱和(LIU et al., 2017);而 SiO₂ 的混染可以促使硫化物饱和(IRVINE, 1975; 张招崇等, 2003)。②含硫化物围岩的混染使岩浆体系的 S 含量增高进而有利于硫化物饱和,如 Noril'sk 矿床硫同位素数据表明,沉积围岩中的膏盐层对成矿具有一定的贡献(LI et al., 2009)。

LIU et al. (2018) 发现夏日哈木围岩片麻状花岗岩中星点状硫化物的原位 S 同位素为 $\delta^{34}\text{S}_{\text{V}-\text{CDT}} = 11.2\text{\textperthousand}$, 而夏日哈木的岩体中不同岩相硫化物 $\delta^{34}\text{S} = 2.4\text{\textperthousand} \sim 7.7\text{\textperthousand}$, 均值为 $4.5\text{\textperthousand}$, 并通过同位素两端元混染模型发现, 地壳硫对夏日哈木总 S 的贡献程度为 40%~60%。可见地壳硫混染是成矿的关键因素。

4.3 矿物学分析

LIU et al. (2019) 总结了东昆仑不含矿基性-超

基性岩体与夏日哈木I号岩体(成超大型镍矿)的橄榄石与单斜辉石的成分差异,发现:①不含矿岩体相对于夏日哈木的橄榄岩有低的MgO、Fo和SiO₂的含量,但是有高的CaO和FeO含量。②不含矿岩体的单斜辉石相对夏日哈木矿区的单斜辉石具有大致相同的MgO和SiO₂含量,但具有相对高的FeO和CaO含量。根据橄榄石和单斜辉石的成分结果可以推测,不含矿岩体相对于成矿的夏日哈木I号岩体,是从一

个相对高的FeO和CaO含量与相对低的MgO和SiO₂含量的母岩浆中结晶的,而相对高的FeO和CaO且相对低的SiO₂含量的母岩浆更难达到硫化物饱和,这可能是夏日哈木I号岩体成矿的一个原因(LIU et al., 2019)。其根本原因在于SCSS与压力、FeO、TiO₂和CaO的含量成正比与SiO₂成反比(WENDLANDT, 1982; MAVROGENES et al., 1999; LIU et al., 2007; LI et al., 2009; LIU et al., 2017)(图5)。

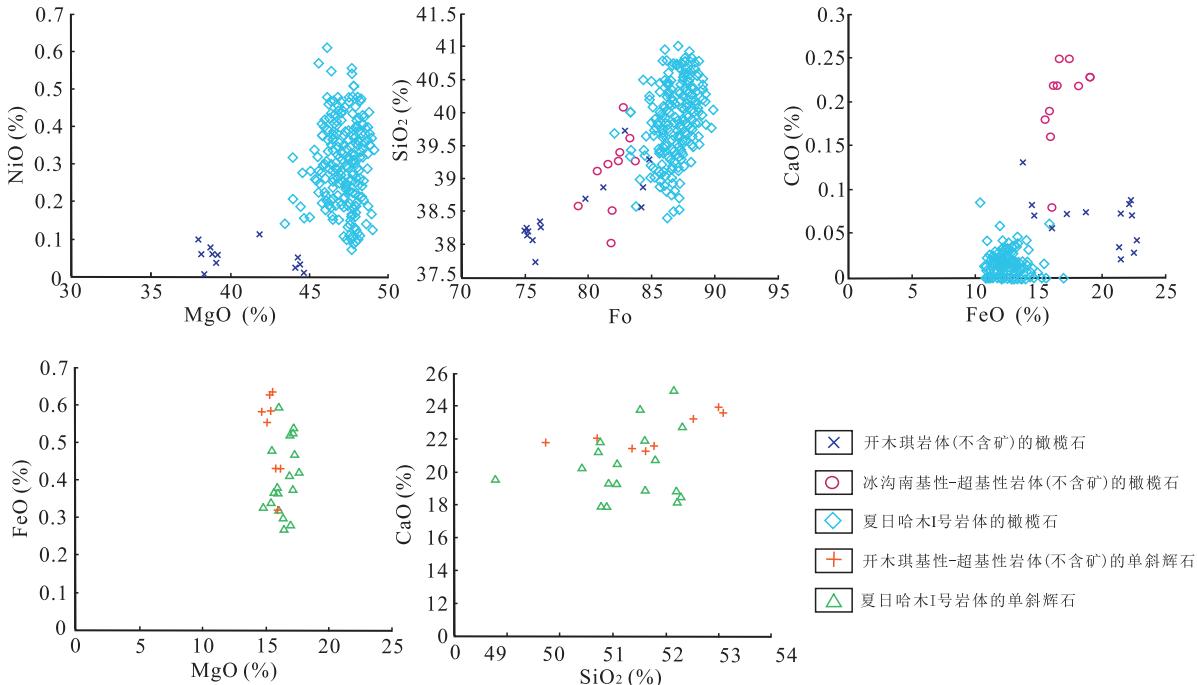


图5 成矿夏日哈木I号岩体与不含矿的冰沟南、开木琪基性-超基性岩的橄榄石和单斜辉石成分对比(LIU et al., 2019)(开木琪数据来自 LIU et al., 2019; 夏日哈木橄榄石数据来自 LI et al., 2015; 凌锦兰, 2014; 王冠, 2014; 冰沟南橄榄石数据来自王冠, 2014; 夏日哈木单斜辉石数据来自王冠, 2014)

Fig. 5 Comparison of the olivine and clinopyroxene compositions from the Xiarihamu and Kaimuqi and Binggouan mafic-ultramafic intrusions.

4.4 成矿机理认识

硫化物不混溶作用的发生是岩浆铜镍矿床形成的关键环节。从硅酸盐岩浆中熔离出来且富含亲铜元素的硫化物液滴发生聚集、就位,固结形成岩浆铜镍硫化物矿体(MAIER et al., 2010; SONG et al., 2016; ZHANG et al., 2018; LIU et al., 2018)。东昆仑造山带夏日哈木矿区尽管发育有5个镁铁-超镁铁质岩体,但仅I号岩体含有超大型规模的岩浆铜镍矿体,其余4个岩体并不含矿。可能存在以下几个方面的因素:①这5个岩体本身就不是一期的产物,岩浆源区性质存在区别。②不同构造背景的

产物,岩浆属性本身就不同,只有铁质系列的镁铁-超镁铁质岩浆才可能形成铜镍矿体,而镁质系列的镁铁-超镁铁质岩浆不可能形成有经济价值的铜镍矿体,尽管能见到铜镍矿化。③岩浆上侵就位的方式和机制明显不同,硫化物不混溶作用发生的空间位置各异,如果硫化物不混溶作用发生在岩浆深部,即便形成铜镍矿体也因深埋地下而无法利用。在东昆仑造山带夏日哈木矿区的5个岩体中,仅I号和II号岩体属铁质系列的镁铁-超镁铁质岩体,而III号、IV号和V号岩体是镁质系列的镁铁-超镁铁质岩体,本身不能形成岩浆铜镍矿体。在I号和II号岩

体中,只有Ⅰ号岩体的岩浆在上侵过程中发生了硫化物不混溶作用,是多次岩浆脉动上侵成矿的结果,地壳混染作用在成矿过程中起有重要作用(ZHANG et al., 2017)。而Ⅱ号岩体的岩浆由于规模太小,并且与榴辉岩接触,地壳混染作用非常有限,没有促进大规模硫化物饱和,因而在Ⅱ号岩体内只能见到弱的铜镍矿化(段建华等,2017;张照伟等,2018)。

5 结论

(1)东昆仑造山带夏日哈木铜镍矿区仅Ⅰ号岩体含岩浆铜镍矿体,属铁质系列的镁铁-超镁铁质岩浆多次脉冲叠加硫化物不混溶作用的结果。

(2)如果铁质系列的镁铁-超镁铁质岩体的橄榄石和单斜辉石具有低 CaO 和 FeO 含量,就更有利于形成富矿体。

(3)在夏日哈木矿区 5 个镁铁-超镁铁质岩体中,仅Ⅰ号赋存超大规模铜镍矿体,Ⅱ号岩体可见弱的铜镍矿化。除此之外,其余 3 个岩体属于榴辉岩和镁质橄榄岩,不具备成铜镍矿的岩浆条件。

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