

帕米尔弧形构造带的构造过程与地貌特征

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摘 要: 帕米尔弧形构造带是印度-欧亚板块碰撞变形最强烈的地区之一, 是研究构造过程、地貌演化以及气候变化及其相互作用的理想场所。本文基于前人的研究成果, 对帕米尔弧形构造带新生代构造单元、地貌特征和动力学演化模型进行了总结归纳, 包括: 主要构造单元的活动起止时间、活动量及活动速率; 帕米尔弧形构造带现今的地貌特征(水系和冰川的分布); 帕米尔弧形构造带 6 种主要的地球动力学演化模型的主要样式、优点及限制。论文提出了帕米尔弧形构造带晚新生代构造研究的三个重要的科学问题: 精细厘定构造带内部的不同断裂带运动学特征和相互关系; 深部地质过程与浅部响应相结合, 探讨构造带形成的深部地质过程控制; 将构造过程、气候特征与地貌演化作为一个耦合系统开展研究。

关键词: 帕米尔弧形构造带; 新生代; 构造过程; 地貌特征; 地球动力学模型

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The Tectonic Processes and Geomorphic Characteristics of Pamir Salient

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Abstract: Pamir salient is one of the intensest deformation zones caused by India-Eurasia collision, and is an ideal region to study tectonic processes, geomorphic characteristics, climate changes and their interactions. On the basis of previous research results, the Cenozoic tectonic units, geomorphic characteristics and geodynamic evolution models of the Pamir salient are synthesized in this paper, which include the onset and stop timing, motion magnitude and rates of the main tectonic units; the present geomorphic characteristics of the Pamirs salient (distribution of river systems and glaciers); and the main configuration, advantages and limitations of six geodynamic evolution models of the Pamir salient. We propose three key scientific issues in the study of late Cenozoic tectonic processes in Pamir salient, i.e., precise determination of kinematic characteristics and relationships of different faults; combining deep geological processes with surface responses so as to explore the control of deep geological processes on the formation of Pamir salient; and taking the tectonic processes, climatic characteristics and geomorphic evolution as a coupling system.

Key words: Pamir salient; Cenozoic; tectonic processes; geomorphic characteristics; geodynamic models

帕米尔弧形构造带(Pamir salient)位于喜马拉雅—青藏高原的西北部, 主要由北帕米尔、中帕米尔和南帕米尔 3 个地体在晚古生代—中生代拼贴而成(图 1)(Angiolini et al., 2013; Robinson, 2015), 由于印度-欧亚板块的持续汇聚, 其在晚新生代整体向

北弧形突进了约 280 km(Burtman and Molnar, 1993; Cowgill, 2010; Sobel et al., 2011, 2013)。帕米尔强烈的向北扩展, 导致地壳厚度由前新生代的 25~30 km(Burtman and Molnar, 1993)增加至现今的 55 km 以上, 局部地区甚至达到 70 km(Belousov et

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al., 1980; Mechie et al., 2012; Schneider et al., 2013), 并使得地壳物质被大规模垂向挤出(Robinson et al., 2004, 2007; Stübner et al., 2013a)、剥蚀(Lukens et al., 2012; Carrapa et al., 2014; Fuchs et al., 2014)和俯冲断离(Negredo et al., 2007; Schneider et al., 2013; Sippl et al., 2013b)。帕米尔的扩展也使得塔里木盆地西南缘被强烈改造, 使这一地区成为新生代构造变形最强烈的地区之一(陈汉林等, 2010, 2014; 程晓敢等, 2011, 2012; 廖林等, 2012; Wang et al., 2013, 2016a; 李康, 2014; Cheng et al., 2016, 2017; Wang et al., 2016b)。

随着新都库什和喀喇昆仑山脉的隆升, 印度夏季风在晚第四纪几乎不能为帕米尔提供水气, 而西风所带来的水气大部分降落在帕米尔西缘, 从而造成帕米尔中—东部地区的半干旱—干旱化(Fuchs et al., 2013, 2014), 并使得公格尔山—慕士塔格峰和塔什库尔干冰川消退(Seong et al., 2009a, b; Owen et al., 2012)。而构造活动和气候影响造就了帕米尔现今的地貌特征, 导致帕米尔高原西缘发育深切河谷, 在高原的中-东部发育公格尔山—塔什库尔干谷地、卡拉库勒谷地等, 并接受新生代沉积

(Brookfield, 2008)。同时, 地貌演化也影响了构造活动, 帕米尔东北缘公格尔山穹窿在 3–1 Ma 的快速剥露是对作用于穹窿的冰川作用和盖孜河的侵蚀作用的响应, 而河流的侵蚀和冰川作用则是对伸展构造的反馈(Cao et al., 2013a)。

因此, 帕米尔弧形构造带是研究构造过程、地貌演化以及气候变化及其相互作用的理想场所。

1 帕米尔弧形构造带的构造过程

印度板块与欧亚板块~50 Ma 以来的持续碰撞, 造成了中亚地区强烈的陆内变形和帕米尔高原的隆升, 同时形成了帕米尔弧形构造带(Molnar and Tapponnier, 1975; Rowley et al., 1996)。帕米尔弧形构造带周缘以一系列逆冲、走滑和走滑逆冲断层带为界, 内部则以发育了一系列与伸展相关的构造为特征(图 1)。南缘为科依斯坦—拉达克弧及向北俯冲的印度板块; 北缘为向北逆冲的主帕米尔逆冲断层(Main Pamir Thrust, MPT)和帕米尔前缘逆冲断层(Pamir Frontal Thrust, PFT); 东缘为喀喇昆仑断层(Karakorum Fault, KKF)、喀拉喀什断层(Karakax Fault, KXF)和喀什—叶城走滑系统

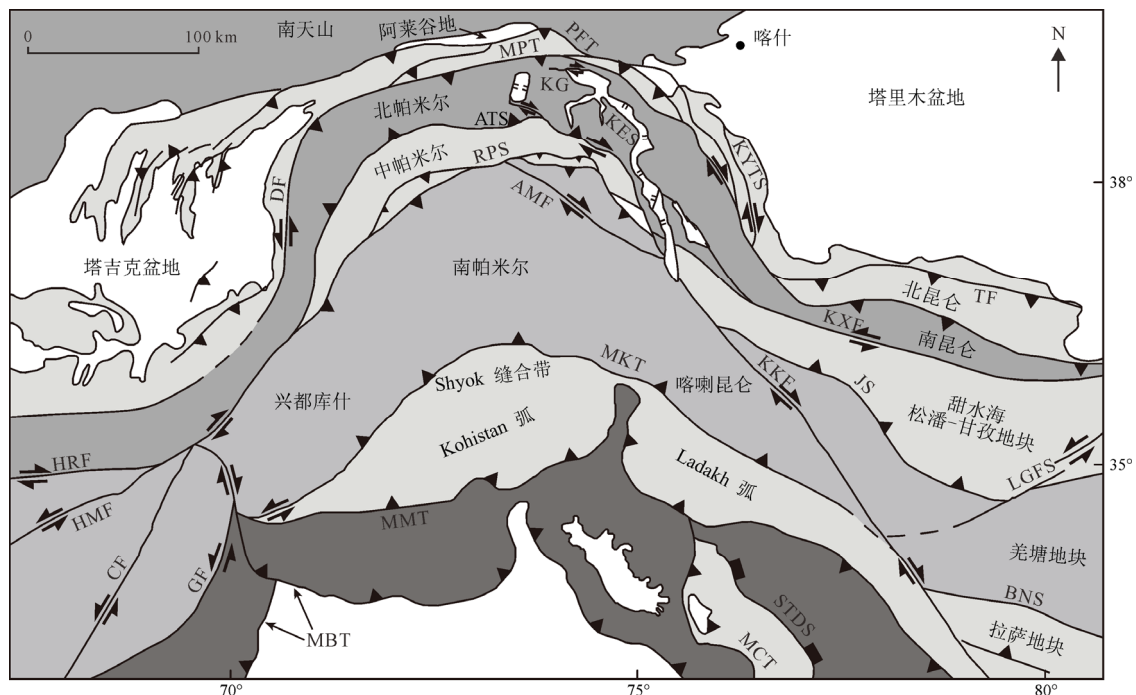


图 1 帕米尔弧形构造带构造纲要图(据 Robinson et al., 2004; Cowgill, 2010)

Fig. 1 Tectonic outline of the Pamir Arcuate Tectonic Belt (modified after Robinson et al., 2004; Cowgill, 2010)

AMF-Aksu-Murgab 断层; ATS-Akbaytal-Tanyamas 缝合带; BNS-班公—怒江缝合带; CF-恰曼断层; DF-达瓦孜断层; GF-Gardez 断层; HMF-Helmand 断层; HRF-Herat 断层; JS-金沙江缝合带; KES-公格尔山伸展系统; KG-卡拉库勒地堑; KKF-喀喇昆仑断层; KXF-喀拉喀什断层; KYTS-喀什—叶城转换系统; LGFS-龙木错—郭扎错断层系统; MBT-主边缘逆冲断层; MCT-主中央逆冲断层; MKT-主喀喇昆仑逆冲逆冲断层; MMT-主地幔逆冲断层; MPT-主帕米尔逆冲断层; PFT-帕米尔前缘逆冲断层; RKF-Rangkul 断层;

RPS-Rushan-Pshart 缝合带; STDS-藏南拆离系; TF-铁克里克断层

AMF-Aksu-Murgab fault; ATS-Akbaytal-Tanyamas suture; BNS-Bangong-Nujiang suture; CF-Chaman fault; DF-Darvaz fault; GF-Gardez fault; HMF-Helmand fault; HRF-Herat fault; JS-Jinsha suture; KES-Kongur Shan extensional system; KG-Karakul graben; KKF-Karakoram fault; KXF-karakax fault; KYTS-Kashgar-Yecheng transfer system; LGFS-Longmu Co-Guoza Co fault system; MBT-main boundary thrust; MCT-main central thrust; MKT-main Karakoram thrust; MMT-main mantle thrust; MPT-main Pamir thrust; PFT-Pamir frontal thrust; RKF-Rangkul fault; RPS-Rushan-Pshart suture; STDS-south Tibetan detachment system; TF-Tiklik fault

(Kashgar–Yecheng Transfer System, KYTS); 西缘为达瓦孜断层(Darvaz Fault, DF)和恰曼断层(Chaman Fault, CF); 内部则主要发育与伸展相关的公格尔伸展系统(Kongur Shan Extensional System, KES)、卡拉库尔地堑(Karakul Graben, KG)和新生代穹窿。

1.1 南缘印度板块的俯冲历史

~50–25 Ma, 印度-欧亚大陆的碰撞导致了亚洲岩石圈的逐步增厚(Klootwijk et al., 1992; Najman et al., 2010)。古地磁重建、地质平衡恢复和全球层析成像结果显示, 在印度克拉通岩石圈与欧亚板块碰撞之前, 印度克拉通北缘的被动大陆边缘(~1 350 km 宽)已经发生了俯冲, 称为 Greater India(Van der Voo et al., 1999; Guillot et al., 2003; Van Hinsbergen et al., 2011)。持续的 N–S 向汇聚作用导致了帕米尔内部的强烈缩短和地壳的增厚。

25–20 Ma, 由于印度克拉通板块浮力大、厚且稳定, 加之亏损的岩石圈地幔(Kumar et al., 2001), 可能阻止了俯冲, 导致了 Great India 板片沿着其与印度克拉通岩石圈之间的过渡带发生了断离(图 2)(DeCelles et al., 2002; Negredo et al., 2007; Stearns et al., 2013, 2015)。之后, 汇聚模式由俯冲变为下插(Kufner et al., 2016), 速度从~44 km/Ma 减少到~34 km/Ma(Molnar and Stock, 2009)。持续向北下插的印度克拉通岩石圈地幔与帕米尔地壳耦合, 增加了造山带的边界挤压应力。

~12 Ma, ~N–S 向的重力垮塌在中帕米尔终止, 重力垮塌导致重力势能减弱, 使得亚洲地壳的应力由拉张变为挤压, 变形向前陆传播, 造成了北帕米尔与南天山之间的加速缩短。~11 Ma, 深部持续向北的印度克拉通岩石圈与亚洲克拉通岩石圈发生碰撞, 导致了亚洲克拉通板片的拆沉和后撤(Kufner et al., 2016), 也形成了帕米尔之下的地震带(Sippl et al., 2013a)。

1.2 北缘正向逆冲

帕米尔北缘发育南倾的主帕米尔逆冲断层(MPT)、帕米尔前缘逆冲断层(PFT)及其两者之间的乌泊尔背驮盆地(Wupoe Piggyback Basin), 并以此与北侧的阿莱山谷和南天山分界(图 1)。

1.2.1 主帕米尔逆冲断层(MPT)

MPT 是向北后撤的俯冲亚洲板片在地表的投影(Sobel et al., 2013), 近 EW 向展布的 MPT 由多条向北逆冲的叠瓦状断层组成, 位于 MPT 南侧的 Trans Alai 山脉沿 MPT 逆冲到北侧的阿莱谷地之上(陈杰等, 2011)。

MPT 的活动时间可能起始于 20–25 Ma(Thomas et al., 1994, 1996; Sobel and Dumitru, 1997; Bershaw et al., 2012; Cao et al., 2013b, Thompson et al., 2015)

或者中始新世(Yin et al., 2002)。晚中新世后, 其活动速率明显下降(Sobel et al., 2011)。Coutand et al.(2002)利用平衡剖面法估算出阿莱谷地 25 Ma 以来的缩短速率为 0.7~0.8 mm/a。Arrowsmith and Strecker(1999)通过对位于阿莱山谷的 Syrinadjar 河中被错断阶地的分析, 计算出 MPT 中段(阿莱谷地南缘)全新世的倾滑速率为 ≥ 6 mm/a。GPS 数据显示 MPT 在阿莱盆地南缘现今仍以 10~15 mm/a 的速率吸收南北向的缩短(Zubovich et al., 2010; Ischuk et al., 2013), 但这一断层在帕米尔东北缘可能已不再活动(陈杰等, 2011)。

1.2.2 帕米尔前缘逆冲断层(PFT)

PFT 位于 MPT 以北~30 km, 是帕米尔弧形构造带最前缘最新的变形带, 其前缘逆断层又称乌帕尔断层带(陈汉林等, 2010), 且具有明显的右旋走滑分量(肖安成等, 2000)。PFT 由多个向北逆冲的次级推覆体及其间的横向撕裂断层组成(刘胜等, 2005; 陈汉林等, 2010; 陈杰等, 2011), 由西向东包括卡巴加特断层、吉勒格由特断层、托姆洛安断层、木什背斜和乌帕尔压扭断层(陈杰等, 1997; 尚新璐等, 2004; 陈汉林等, 2010)。其形成可能与乌拉根古隆起的存在有关(Wang et al., 2016a)。

6–3.5 Ma, 帕米尔前缘逆冲断层开始活跃(Arrowsmith and Strecker 1999; Fu et al., 2010; Thompson et al., 2015)。Thompson et al.(2015)通过生长地层及磁性地层学研究, 认为 PFT 于~5–6 Ma 开始活动。PFT 几乎吸收了帕米尔北缘晚第四纪以来的所有变形, ~0.35 Ma 以来的平均缩短率为 6~8 mm/a(Li et al., 2012b)。野外露头及磁性地层学数据显示 PFT 现今最小缩短量为 2.1 km, 最小缩短速率约为 6.2 mm/a(陈杰等, 2011)。Zubovich et al.(2016)通过 GPS 数据分析, 认为 PFT 现今倾滑速率至少在 (5.6 ± 0.8) mm/a, 缩短速率在 (92.2 ± 0.8) mm/a。潘家伟等(2009)根据乌帕尔断裂带造成的水系错断量, 计算出该断层的右旋走滑速率为 4.0~6.8 mm/a, 活动时间为>2.2–3.0 Ma。

1.3 东缘右旋-走滑逆冲

帕米尔东缘发育三条主要的走滑断层, 从西向东依次为喀喇昆仑断层(KKF)、喀拉喀什断层(KXF)和喀什-叶城转换系统(KYTS)(图 1), 并以此与东侧的甜水海地块、西昆仑和塔里木盆地分隔(Burtman and Molnar, 1993; Cowgill, 2010), 这三条断层都有可能参与了帕米尔相对于西昆仑向北的运动(陈杰等, 2011)。

1.3.1 喀喇昆仑断层(KKF)

KKF 总体走向 NW–SE, 长约 1 200 km, 自青藏高原东南端的冈仁波齐山脉一直延伸到中帕米

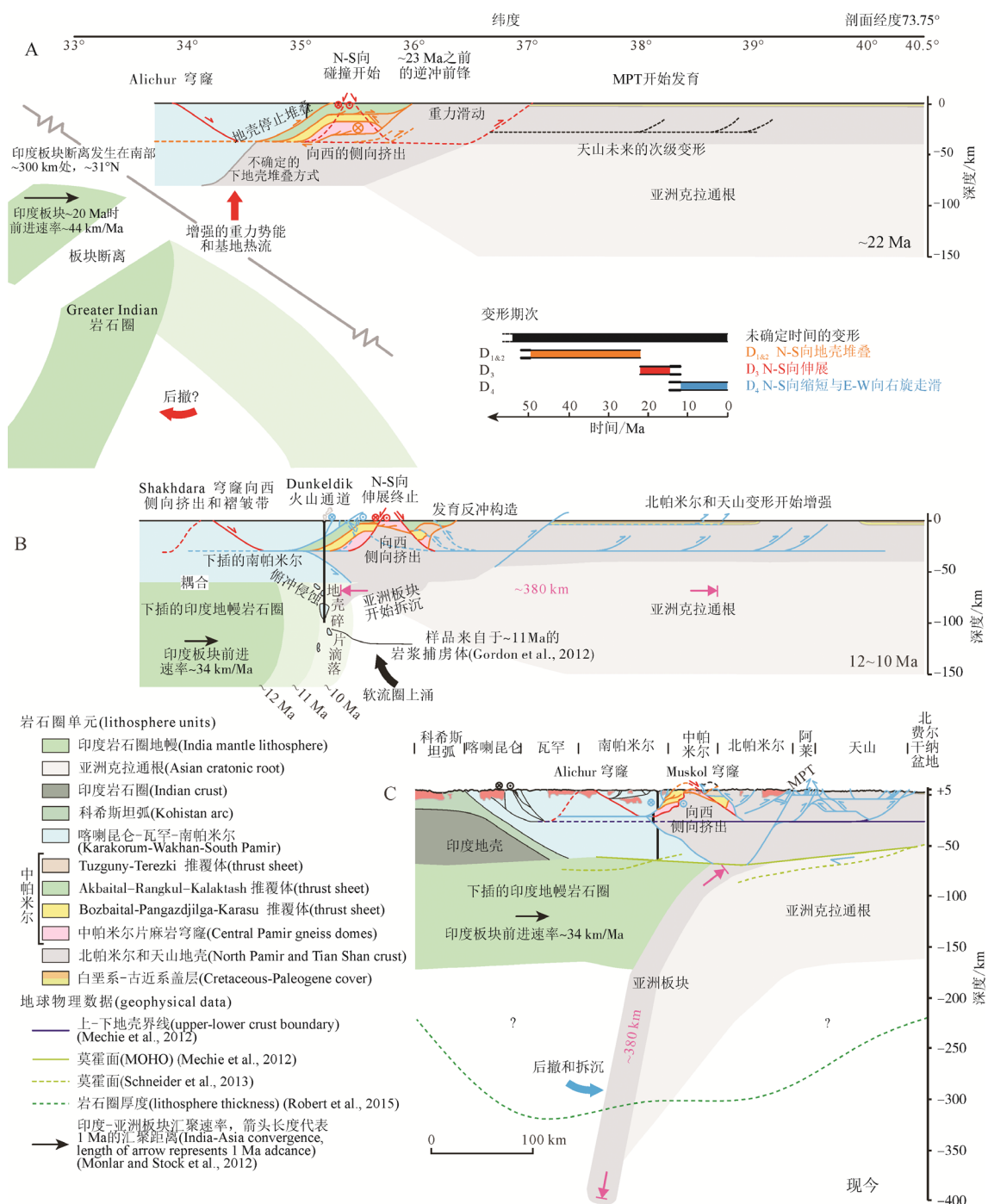


图2 帕米尔深部构造过程及其浅部响应剖面示意图(据 Rutte et al., 2017b)

Fig. 2 Schematic sections through Pamir Arcuate Tectonic Belt illustrating the spatiotemporal relationships between crustal and mantle processes (after Rutte et al., 2017b)

A- ~22 Ma: 始于~37 Ma 的逆冲推覆体褶皱堆叠使中帕米尔地壳强烈增厚, 北帕米尔地壳表现为下插。部分地壳物质向西侧向挤出。Greater India 岩石圈在~25-20 Ma 发生断离, 引发重力坍塌, 在中、南帕米尔发生了广泛的 N-S 向伸展作用, 北帕米尔与天山之间开始发生汇聚; B- ~12-10 Ma: 板块断离之后, 印度岩石圈地幔由俯冲变为下插, 且与南帕米尔地壳耦合, 近 N-S 向的伸展作用停止, 变为 N-S 向的挤压作用。南帕米尔地壳岩石由于俯冲侵蚀作用, 地壳碎片滴落至地下 100 km 深, 之后作为捕虏体被从 Dunkeldik 岩浆通道喷发至地表。~11 Ma: 印度岩石圈地幔与亚洲克拉通岩石圈发生碰撞, 导致了亚洲克拉通板片的拆沉和后撤。C- 现今: 亚洲板片的拆沉仍在继续, 地壳发生了 N-S 向的缩短和地壳物质向西的侧向挤出, 以及 E-W 向的伸展作用

A- ~22 Ma: starting prior ~37 Ma thrust sheets and fold nappes thickened the Central Pamir (and South Pamir, not shown), tripling the upper crust; North Pamir crust is underthrusting. Part of the crust is extruding laterally (westward). Greater Indian lithosphere breakoff at ~25-20 Ma triggered gravitational collapse, i.e., ~N-S extension in the South and Central Pamir. Dominant top-to- ~N sliding initiated contraction in the North Pamir and possibly the Tian Shan; B- ~12-10 Ma: after breakoff, Indian mantle lithosphere has been underthrusting Asia and coupling with the Pamir crust, ending ~N-S collapse and resuming N-S shortening (in the Central Pamir at ~12 Ma). Through a combination of subduction erosion and transient dripping, South Pamir rocks are buried to depth ≤ 100 km and erupted as xenoliths in the Dunkeldik volcanic pipes. ~11 Ma: Indian mantle lithosphere impinged on the southern edge of the Asian cratonic root (Tarim and Tajik basin lithosphere) and forced its delamination; C- recent: Delamination of the Asian slab is ongoing. Crust experiences dominant ~N-S shortening and westward extrusion through dextral wrenching and ~E-W extension (not displayed)

尔, 构成了帕米尔高原东南部的边界(Peltzer and Tapponnier, 1988)。对于 KKF 在调节印度和欧亚板块之间变形中所起的作用, 有两种不同的观点: (1)KKF 新生代调节了南帕米尔向北~300 km 的位移量, 并造成了地壳的显著增厚, 形成了~70 km 的陆壳(Tapponnier et al., 1982; Peltzer and Tapponnier, 1988; Armijo et al., 1989; Hamburger et al., 1992; Burtman and Molnar, 1993; Lacassin et al., 2004; 李海兵等, 2007); (2)KKF 作为转换系统在连接帕米尔的逆冲带和青藏高原西缘之间所起的作用有限(Burtman and Molnar, 1993; Searle, 1996; Seeber and Pecher, 1998; Murphy et al., 2000; Wright et al., 2004; Robinson et al., 2007; Searle and Phillips, 2007; Robinson, 2009a; Wang et al., 2012)。

KKF 走滑活动可能起始于 13~15 Ma(Murphy et al., 2000; Murphy et al., 2002; Phillips et al., 2004; Phillips and Searle, 2007)或者(23±1) Ma (Lacassin et al., 2004; Valli et al., 2008)。Brookfield et al.(2017)通过花岗岩中锆石结晶年龄, 认为 KKF 活动时间在~12 Ma。

Robinson(2015)和 Robinson et al.(2016)通过测试穿过 KKF 北段主干而未变形的宽谷阶地中次生碳酸盐岩帽的年龄和穿过其北段分支未被错动的冰期沉积物的年龄, 认为 KKF 北段至少~200 ka 以来已不再活动。Robinson(2009a, b)通过对遥感影像的解译, 认为 KKF 北段活动的停止可能与~3 Ma 龙木错—郭扎错左旋走滑断裂的活动有关。对于 KKF 北段的延伸, 有学者认为 KKF 走滑量向北被帕米尔内部的 RPZ 和 Aksu-Murgab 断层系所吸收并终止, 并未传到北部的 KES 和 MPT(Burtman and Molnar, 1993; Strecker et al., 1995; Robinson et al., 2007)。而 Chevalier et al.(2016)认为 KKF 北段在第四纪仍有活动, 且向北与公格尔伸展系统连接成为同一个断裂体系。

断层北段右旋走滑量为 149~167 km, 平均滑动速率为 6.9~10.8 mm/a(Robinson, 2009a)。而其南段走滑量只有~65 km(Murphy et al., 2000; Wang et al., 2012)。

1.3.2 喀拉喀什断层(KXF)

KXF 又名康西瓦断层, 位于 KKF 以东, 走向近 NWW, 延伸近 700 km, 可能为阿尔金断裂的西延部分, 为一左行走滑断层(Peltzer et al., 1989; Tapponnier and Molnar, 1997; 刘栋梁等, 2011)。Avouac and Peltzer(1993)认为左旋走滑的 ATF 在其西南端分为 SW 走向的龙木错—郭扎错左旋走滑断层和 NWW 走向的 KXF。且由于龙木错—郭扎错左旋走滑断层的位移量只有 25~30 km(Rateman et al., 2007), 所以 Cowgill et al.(2003)推测 ATF 的大部分

走滑量可能被 KXF 和向南逆冲的甜水海逆冲带所吸收。

早期的研究认为, KXF 的左旋走滑速率为 20~30 mm/a(Peltzer et al., 1989; Ryerson et al., 1997)。而付碧宏等(2006)根据不同年代地表地貌特征的左旋错位距离, 估算出康西瓦断层晚第四纪以来的长期走滑速率为 8~12 mm/a。Li et al.(2012a)和 Gong et al.(2017)的研究认为 KXF 晚第四纪的平均走滑速率为分别为 6.5 mm/a 和(7.8±1.6) mm/a, 这与 GPS 观测结果(7±3) mm/a 相一致(Shen et al., 2001)。

1.3.3 喀什-叶城转换系统(KYTS)

KYTS 走向 NNW, 长约 350 km, 宽约 50 km, 由四条近平行的右旋走滑断层组成, 自西向东有库斯拉甫断层、库姆塔格断层、叶尔羌断层和阿尔塔什断层。

Sobel and Dumitru(1997)通过采自 5 个剖面的 AFT 数据的分析, 识别出从晚渐新世—早中新世时期存在一期与 KYTS 的活动有关的隆升剥蚀和冷却, 认为 KYTS 于 20 Ma 之前开始活动。Cowgill(2010)认为 Yin et al.(2002)文章中阿尔塔什剖面 37 Ma 的快速沉降是 KYTS 开始活动的标志, 且该转换系统是分割东帕米尔和西昆仑—塔里木块体的主要边界右旋压扭断裂带, 吸收调节了北帕米尔相对于西昆仑—塔里木块体向北的运动量; 晚新生代以来的累积右旋断错量约为 280 km, 平均滑动速率为 7~8 mm/a。Cao et al.(2013b)通过对采自克孜勒陶河河谷的样品的 AFT 年龄分析, 结合野外剖面的调查, 认为 KYTS 启动时间在~50 Ma, 但此时的断裂系比现在的要窄, 直到~10~6 Ma 才形成与现在规模相当的断裂系。

Sobel et al.(2011)通过对采自塔什库尔干—叶尔羌河的碎屑样品的热年代学分析, 认为 KYTS 中的库斯拉甫断层和库姆塔格断层至少在 5~3 Ma 以来已停止活动, 而最东缘的阿尔塔什断层在 5~3 Ma 以来的走滑速率仅为 1.7~5.3 mm/a。这一现象表明晚中新世或上新世以来, 帕米尔东缘与塔里木板块近乎作为一个整体向北运动(Sobel et al., 2011)。

1.4 西缘左旋逆冲

帕米尔西缘发育两条主要的左行走滑断层: 达瓦孜断(DF)和西南端的恰曼断层(CF)(图 1)。

1.4.1 达瓦孜断层(DF)

NNE 走向的 DF 是帕米尔与塔吉克盆地的分界, 其启动时间未被确定, 推测可能与 MPT、KYTS 一致。根据全新世和晚更新世的地貌, 尤其是早全新世的阶地和冲积扇的错断距离, 推测该断层左旋走滑速率为 10~15 mm/a(Kuchai and Trifonov, 1977;

Trifonov, 1978)。Mohadjer et al.(2010)根据 Khorog (MANM)和 Shaartuz(SHTZ)两观测站所记录的 GPS 速度数据,通过计算其在 DF 走向上的速度分量,得出 DF 的最大左旋剪切速率为 (11.4 ± 2) mm/a。Ischuk et al.(2013)通过 GPS 数据得出帕米尔西缘与塔吉克盆地之间的左旋剪切速率为 ~ 10 mm/a,但由于 GPS 测点零散的分布使得不能很好地限定其分布和位置。Tomas et al.(1994)结合已有的地震和地质资料,认为 DF 存在向西的逆冲分量。

1.4.2 恰曼断层(CF)

CF 位于帕米尔高原西南端,是印度板块与亚欧板块的分界,走向 NE,南北长约 $\sim 1\ 000$ km,东西宽约 20~30 km,南端连接麦克兰(Makran)增生汇聚带,北端与喜马拉雅汇聚带相连(Lawrence et al., 1981; Tapponnier et al., 1981; Dewey et al., 1988; 许志琴等, 2011)。

CF 开始活动时间为 $\sim 25\sim 20$ Ma,总的位移量为 (460 ± 10) km,早中新世以来的平均走滑速率为 $24\sim 35$ mm/a(Beun et al., 1979; Lawrence et al., 1992)。GPS 数据显示 Chaman-Gardiz-Konar System 北部现今的左旋剪切速率为 (18 ± 1) mm/a(Mohadjer et al., 2010)。合成孔径雷达干涉测量技术(InSAR)对北部的现今滑移速率进行分析得到了一个相对较低的滑动速率 ~ 8 mm/a(Furuya and Satyabala, 2008), Fattahi and Amelung(2016)利用该方法也得到了相似的结果 $(\sim 8.1\pm 2)$ mm/a;对 CF 的中部和南部, Barnhart(2017)通过 InSAR 分析,得出走滑速率的峰值为 $9\sim 12$ mm/a。在 CF 中南部的 Spinatizha 地区北部的 GPS 数据显示了 $(\sim 8\sim 10)$ mm/a 的走滑速率(Crupa et al., 2017)。

1.5 内部伸展作用

帕米尔弧形构造带内部,晚新生代的构造活动主要表现为片麻岩穹窿的剥蚀和伸展系统的扩展。其中,北帕米尔主要发育与东西向的伸展作用,且在其东北缘发育与之相关的公格尔伸展系统和公格尔—慕士塔格穹窿;中、南帕米尔主要发育南北向伸展作用和之相关的穹窿。

1.5.1 EW 向伸展构造

在新生代,帕米尔内部的 EW 向伸展构造主要为卡拉库勒地堑(Karakul Graben, KG)和公格尔伸展系统(Kongur Shan Extensional System, KES)以及除了 KES 以外的公格尔—慕士塔格穹窿边界的断层(图 1)。

(1)卡拉库勒地堑(KG)

在北帕米尔中部,发育了近 NNE-SSW 走向的卡拉库勒地堑(图 1),内部形成了卡拉库勒盆地。KG 北部以 NW-SE 走向的右旋走滑 Markansu 断层

带为界,南部以 NW-SE 走向的右旋走滑 Aksu-Rangkul 断层带为界(Strecker et al., 1995)。

如果拉张只沿东西两侧的边界正断层发育,KG 的 EW 向总拉张量 <3 km,考虑拉张速率为 $\sim 0.5\sim 1$ mm/a,则推算出 KG 的 EW 向扩张可能开始于 $\sim 9\sim 3$ Ma(Amidon and Hynek, 2010)。同时,沿地堑的拉张还伴随着左旋走滑分量,加之南北两侧边界的右旋走滑断层,该地区除了受到 EW 向拉张之外,还收到了 NNW-SSE 向的挤压(Strecker et al., 1995)。

(2)公格尔伸展系统(KES)

KES 走向 NW-SE,延伸约 250 km,由北向南依次为木吉断层、公格尔山正断层、塔合曼断层和塔什库尔干正断层(Robinson et al., 2007)。KES 东西向的伸展量在木吉断层为 30 km,在公格尔山附近最大,约 34 km,向南减少至慕士塔格山附近的 20 km,最南边的塔什库尔干正断层只有 <3 km(Robinson et al., 2004, 2007),总体上呈北宽南窄。

KES 北部活动时间较早,为 $8\sim 7$ Ma(Robinson et al., 2004);南部略晚,为 $6.2\sim 5$ Ma(Robinson et al., 2007; Sobel et al., 2011; Cao et al., 2013a)。Robinson et al.(2004)对采自北部 Qimugang 山谷地区和公格尔山附近 KES 下盘的样品进行了独居石 U-Pb 年龄、云母 $^{40}\text{Ar}/^{39}\text{Ar}$ 年龄和钾长石多重扩散域 MDD (Multi-diffusion Domain, MDD)模式分析,表明断层下盘 $8\sim 7$ Ma 以来发生了快速稳定隆升。而公格尔正断层在慕士塔格山附近的起始活动时间则为 $6.2\sim 5$ Ma(Robinson et al., 2007; Sobel et al., 2011; Cao et al., 2013a)。第四纪的断层陡坎和地震活动表明公格尔正断层现今仍然在活动(Seong et al., 2009b; 李文巧等, 2011)。

木吉断层位于 KES 最北段,为具有拉张分量的右旋走滑断层,木吉断层下盘 $8\sim 7$ Ma 以来自地下 10 km 深处发生了快速而稳定的冷却,其剥蚀速率为 $1.3\sim 1.4$ mm/a(Robinson et al., 2004),木吉断层的右旋走滑速率为 (4.5 ± 0.2) mm/a 到 (11.1 ± 0.9) mm/a 之间,或 ≤ 7 mm/a,但后者要求垂向移动速率为 $\sim 0.3\sim 0.5$ mm/a(Chevalier et al., 2011)。

公格尔山正断层下盘自 $8\sim 7$ Ma 以来,在 Qimugang 山谷总剥蚀厚度为 $10\sim 11$ km,在公格尔山附近剥蚀厚度为 29 km(Robinson et al., 2004);自 $7.5\sim 5$ Ma 以来,在慕士塔格山附近剥蚀量约 9 km(Robinson et al., 2007)。Robinson et al.(2010)基于热对流和断层几何特征的模拟计算表明公格尔山附近自约 7 Ma 以来,断层下盘剥蚀速率为 4.2 mm/a,倾滑速率约 6.5 mm/a。若假定断层倾角为 40° ,则垂直滑动速率约 4.2 mm/a,近 EW 向拉张

物质的剥蚀(Schwab et al., 2004; Robinson et al., 2004; Schmidt et al., 2011; Stübner et al., 2013a)。

(1)北帕米尔穹窿

在帕米尔东北缘,主要发育公格尔山穹窿(Kongur Shan dome)和慕士塔格峰穹窿(Muztaghata dome)。

公格尔山穹窿在 9 Ma 仍处于进变质作用(Robinson et al., 2004),于 7~5.5 Ma 开始剥露(Robinson et al., 2010; Thiede et al., 2013; Cao et al., 2013a),并持续至今,平均剥露速率为 4.2 mm/a(Robinson et al., 2010)或 $>2\sim3$ mm/a(Thiede et al., 2013),剥露量达到 29 km(Robinson et al., 2004),其在 3~1 Ma 可能发生了加速剥露(Arnaud et al., 1993; Cao et al., 2013a, b)。

慕士塔格峰穹窿在 30~20 Ma 发生进变质作用,原岩被埋深至 30~35 km,并于 14 Ma 发生了片岩的部分熔融,而退变质作用开始于 10~8 Ma,表明该穹窿处于第一期剥露(Robinson et al., 2007);该穹窿的第二期剥露则开始于 6.2~5 Ma(Robinson et al., 2007; Cao et al., 2013b),但剥露速率明显比第一期时小,仅为 (1 ± 0.5) mm/a(Thiede et al., 2013)。

(2)中、南帕米尔穹窿

中帕米尔的 Yazgolom、Sarez、Muskol、Shatput 穹窿在 35~20 Ma 处于进变质作用(Schmidt et al., 2011; Stearns et al., 2013; Smit et al., 2014),于 20~16 Ma 开始剥露(Rutte et al., 2013; Stearns et al., 2013, 2015),并持续至 10 Ma(Schmalholz, 2004; Cao et al., 2013b)。

南帕米尔的 Shakh-dara-Alichur 穹窿是帕米尔内部出露面积最大的晚新生代片麻岩穹窿,被 Turumtai 地堑分割成 Shakh-dara 和 Alichur 穹窿(Stübner et al., 2013a);这两个穹窿的进变质作用发生在 37~20 Ma,于 20 Ma 开始发生退变质作用,即开始剥露(Schmidt et al., 2011; Smit et al., 2014; Stearns et al., 2015),并持续至 2 Ma(Stübner et al., 2013b)。

(3)帕米尔内部东西走向穹窿的形成机制与演化

Stearns et al.(2015)则提出了一种针对南北向伸展作用的解释方案,认为俯冲的印度板片在 32~20 Ma 发生后撤和拆离,导致帕米尔地壳由挤压转为伸展,中下地壳的物质被快速剥露,形成 Yazgolom、Sarez、Muskol、Shatput 及 Shakh-dara-Alichur 穹窿。本文依据此模型,对帕米尔内部穹窿的演化史进行总结。

$>37\sim22$ Ma,由于印度板块的向北汇聚,在

中、南帕米尔形成了大范围的褶皱和逆冲推覆体,加之与造山带平行的物质侧向挤出(Ratschbacher et al., 1991),导致了中、南帕米尔广泛的 N-S 向缩短、地壳增厚和进变质作用的发生。由于中帕米尔逆冲断层和褶皱推覆体的堆叠,阻止了变形继续向北传播,变形没有传到北帕米尔(Rutte et al., 2017b)。

$\sim23\sim20$ Ma,由于地壳增厚和 Greater India 板块 ~25 Ma 的拆离,加之软流圈物质的上涌促使基底的热流值增加,导致岩石圈强度的减弱,引起了中、南帕米尔厚、热且高海拔的地壳的重力垮塌(图 2),造成了中、南帕米尔的 N-S 向的伸展和退变质作用的发生(Stearns et al., 2013, 2015),在中、南帕米尔地区形成了变质核杂岩(Rey et al., 2010; Rutte et al., 2017a)。中、南帕米尔的重力垮塌也可能造成了北帕米尔的变形,变形前锋从中帕米尔移动到了北帕米尔。由于重力垮塌、北帕米尔缩短和侵蚀作用的增强,造成了北帕米尔中部 $\sim25\sim16$ Ma 的加速剥蚀(Amidon and Hynek, 2010)。

20~12 Ma,双向(南倾和北倾)的剪切带沿中、南帕米尔穹窿的南北边界发育,且主要的伸展作用发生在穹窿北缘(Rutte et al., 2017a)。

12 Ma 至今,由于重力势能的衰减和印度克拉通的下插和向北传播,导致了亚洲克拉通板片 ~11 Ma 的拆沉和后撤(Kufner et al., 2016),中、南帕米尔 N-S 向伸展作用及其相关的穹窿剥露停止,N-S 向的挤压作用沿着双向的褶皱逆冲带发育(图 2)。Schurr et al.(2014)通过帕米尔的地震构造学研究,认为帕米尔高原西缘持续的重力垮塌和地壳物质向西部塔吉克盆地的侧向挤出,导致了一系列的构造变形,发育了 S-N 向的缩短、E-W 向与造山带平行的伸展作用和右旋走滑。

1.6 帕米尔隆升历史

基于低温热年代学研究,帕米尔弧形构造带新生代共经历了四次隆升,分别在 $\sim50\sim40$ Ma, $\sim25\sim16$ Ma, $\sim10\sim7$ Ma 和 <5 Ma (Robinson et al., 2004, 2007; Amidon and Hynek, 2010; Wang et al., 2011; Sobel et al., 2011, 2013; Cao et al., 2013a, b; Carrapa et al., 2014)。

$\sim50\sim40$ Ma,帕米尔发生了初始隆升,这一隆升是对印度与 Kohistan-Ladakh 弧在 $\sim55\sim47$ Ma 碰撞的响应(De Sigoyer et al., 2004; Leech et al., 2005; Guillot et al., 2008)。帕米尔弧形构造带东北缘奥依塔格剖面沉积物的氧同位素分析结果显示,帕米尔的初始隆升使得其在始新世—渐新世时期已达到可以阻挡大气环流的高程(Bershaw et al., 2012)。

Amidon and Hynek(2010)通过卡拉库尔地堑的磷灰石和锆石的(U/Th) - He 年龄分析, 认为该地区在~50–40 Ma 发生了与隆升有关的加速剥蚀。在此阶段, 帕米尔西北缘发生了逆时针旋转, 帕米尔东北缘和内部发生了顺时针旋转(Liu et al., 2017)。

~25–16 Ma, 帕米尔发生了第二次隆升, 与欧亚板块向南的俯冲作用有关(陈汉林等, 2014)。Sobel et al.(2011)通过总结帕米尔东缘已发表的热年代学数据, 认为 KYTS 在渐新世至早—中中新世发生过显著的构造活动。Amidon and Hynek(2010)分析了帕米尔北缘卡拉库尔地堑的磷灰石和锆石的(U/Th) - He 年龄, 认为该地区在~25–16 Ma 发生过新的一期构造隆升, 并认为此次构造隆升可能与 Greater India 板块在 25 Ma 的拆离或者 KKF 初期的活动有关。Hubbard et al.(1999)认为帕米尔西南部的 22–17 Ma 的冷却年龄, 反映了与构造隆升相关的剥蚀。这一隆升在帕米尔弧形构造带相邻的盆地内部有着很好的记录, 帕米尔北缘山前阿莱谷地和塔吉克盆地的挠曲及沉积盆地的加速发生于早中新世(Leith et al., 1985; Coutand et al., 2002); 塔西南新生代沉积物的同位素物源的研究, 表明西昆仑在此阶段发生了第一次隆升(Blayney et al., 2016); Clift et al.(2017)通过对塔克拉玛干沙漠西南部新生代沉积地层的研究, 认为 Pamir 在 24 Ma 之前已经发生了隆升, 早于北昆仑的 17 Ma。

~10–7 Ma, 帕米尔发生了第三次隆升。Carrapa et al.(2014)通过对采自塔吉克斯坦盆地和帕米尔河流的沉积物样品进行了磷灰石裂变径迹测试, 发现帕米尔在中新世—上新世发生了一期造山带尺度的强烈剥蚀。公格尔山伸展系统下盘的热年代学数据显示, 公格尔山和慕士塔格山分别在~8–7 Ma 和~9–7.5 Ma 发生了快速剥蚀(Robinson et al., 2004, 2007)。帕米尔东缘~10–6 Ma 的磷灰石裂变径迹年龄, 结合 KYTS 的野外调查表明帕米尔在此时期发生了一次强烈的逆冲和与之相关的剥蚀冷却(Cao et al., 2013b)。

<5 Ma, 帕米尔发生了第四次隆升。塔西南的西域砾岩和生长地层记录了此次隆升事件(Chen et al., 2002; Heermance et al., 2007)。帕米尔东缘山前在~5–4 Ma 开始发育褶皱冲断带(程晓敢等, 2012; 李康, 2014), 乌帕尔背驮盆地的物源及古水流的变化(Robinson et al., 2004; Thiede et al., 2013; Cao et al., 2013b; Sobel et al., 2013), 都记录了帕米尔在上新世至今的隆升和剥蚀。这次事件导致了帕米尔东北缘的山前变形带由 MPT 向北迁移到 PFT(李康, 2014; Thompson et al., 2015)。

2 帕米尔弧形构造带现今地貌特征

帕米尔高原平均高程约 4 000 m, 东北部公格尔山和慕士塔格峰的高程分别达到 7 719 m 和 7 546 m(崔之久, 1960; Brunel et al., 1994)。西帕米尔主要为起伏的山脉、深切峡谷和大型山谷冰川, 帕米尔西南缘的山峰平均海拔为~5 000~6 000 m, 而谷地的平均海拔为~2 000~3 000 m(Stübner et al., 2017)。而东帕米尔为整体地势较高但更为平缓的山脉, 且被大量堆积物覆盖, 仅有少量山脉被冰川覆盖(Abramowski et al., 2006)。帕米尔整体地形的起伏变化较小, 反映了宽阔的谷地和有限的河流下切作用(Fuchs et al., 2013)。

2.1 水系分布

帕米尔内部被萨雷阔勒岭分隔为东、西两大流域系统, 其中的东部流域系统主要由盖孜河和塔什库尔干河—叶尔羌河组成, 西部流域系统则由喷赤河及其支流组成(图 3)(Brookfield, 2008; Komatsu, 2016)。水系的补给皆主要来源于冰川融水和上游山区降水(Kayumov and Rajabov, 2010)。

盖孜河大致流向为 NE 向, 上游有 NW–SE 向展布的支流木吉河和康西瓦河, 于布伦口水库汇流, 之后折向 NE 穿过山谷, 流入喀什平原西部地区, 最终消失于东北部的沙漠里(图 3)。盖孜河上游的北西支为木吉河, 其源头为萨雷阔勒岭, 其两侧有沿着东西向展布的冰川, 但分布并不集中; 南东支为康西瓦河, 源头冰川十分发育, 接受慕士塔格峰和公格尔冰川的融水补给, 发源于慕士塔格的峡谷之中, 出谷后补给喀拉库勒湖, 而后流至布伦口水库与木吉河汇流(蒲健辰等, 2003; 再努尔, 2010)。喀拉库勒湖是帕米尔高原中国境内面积最大的天然湖泊, 形成于早全新世, 属于冰碛堰塞湖, 面积约 4.59 km², 呈三角形, 平均水深约 15 m, 湖面海拔 3 656 m(Seong et al., 2007)。

盖孜河流域总面积为 11 029 km², 山区流域面积占总面积的 98.2%, 为 10 827 km², 平原流域面积占总面积的 1.8%, 为 202 km²。盖孜河全长 347 km, 其中, 山区河长 184 km, 平原区河长 163 km(李燕等, 2003; 再努尔, 2010), 盖孜河多年平均径流量为 9.78 亿 m³(Su and Wang, 1998)。

叶尔羌河是东帕米尔最长的河流, 发源于喀喇昆仑山北坡的喀喇昆仑山口, 地理坐标 74°28′–80°54′E, 34°50′–40°31′N, 叶尔羌河全长 1 097 km, 多年平均径流量 66.3 亿 m³(凌红波等, 2012)。流域总面积 9.89 万 km², 山区面积 6.08 万 km², 占流域总面积 61.5%, 平原区面积 3.81 万 km², 占 38.5%(孙本国等, 2006)。叶尔羌河主要支流有上游

的克勒青河,中游的塔什库尔干河和山口以外汇入的提兹那甫河等。

塔什库尔干河是叶尔羌河最大的支流,位于 $74^{\circ}28'-76^{\circ}13'E$, $37^{\circ}11'-38^{\circ}13'N$,由明铁盖河和红其拉甫河汇流而成,发源于喀喇昆仑山脉萨雷阔勒岭(周聿超,1999)。红其拉甫河发源于红其拉甫附近海拔5 852 m的喀喇昆仑山塔木太开山达板,流向SW-NE,明铁盖河发源于海拔5 844 m的喀喇昆仑山中国与阿富汗交界的瓦赫吉尔山口,流向近E-W。塔什库尔干河流向NNW,在塔合曼村与发源于慕士塔格峰南侧的塔合曼河汇流后向东流,之后在阿克陶县境内的盖里克处汇入叶尔羌河。塔什库尔干河干流全长298 km,流域面积9 980 km²,常年流量>30 m³/s,丰水期流量>30 m³/s,洪水期流量200~500 m³/s(巴合提瓦尔等,2008;全晓霞等,2017)。

喷赤河发源于东南帕米尔,流经帕米尔南缘和西缘,最终留向西部的塔吉克盆地。喷赤河上游及其支流为东向西展布向西流动,经阿富汗Eshkashem地区折向北,与达瓦孜断层平行,且穿过了Shakhdara和Yazgulom穹窿(Fuchs et al., 2013);之后在帕米尔西北缘转向西南流入塔吉克盆地。喷赤河全场2 540 km,流域面积达46.5万 km²,河口地区多年平均流量达1 330 m³/s,年径流量430亿 m³(鲜丽菊,2017)。

水系的发育受到了帕米尔晚新生代构造活动的控制,盖孜河和塔什库尔干河的上游沿公格尔山伸展系统发育而呈南北向展布,喷赤河的上游及其支流则沿Yazgolom、Sarez、Muskol、Shatput和Shakhdara-Alichur穹窿的边缘呈东西向展布(图3)(Komatsu, 2016)。河流的下切速率反映了流经地区的构造隆升速率,如喷赤河在过去26 ka的平均下切速率为~5.6 mm/a,流经Shakhdara穹窿处的下切速率约为7.3 mm/a,在Yazgolom穹窿的下游为~4~5 mm/a,在Shakhdara和Yazgolom穹窿的南部边界仅为2~3 mm/a,在达瓦孜断层上游则为~6 mm/a(Fuchs et al., 2014)。塔什库尔干河在慕士塔格峰穹窿南缘的全新世下切速率为 (6.3 ± 1.2) mm/a,可能与慕士塔格峰的快速隆升或第四纪冰川作用有关(刘进峰等,2011)。公格尔山—慕士塔格峰现今的地貌特征则与公格尔—慕士塔格峰穹窿的剥露,该地区的岩性、气候特征和第四纪冰川作用等因素有关(Schoenbohm et al., 2014)。

2.2 冰川分布

整个帕米尔的冰雪覆盖面积在12 078 km²(Dolgushin and Osipova, 1989),冰川主要分布在其

东北缘和西北缘。

(1)帕米尔东北缘

帕米尔东北缘山峰海拔高度大,雪线在4 800~5 200 m之间。海拔超过6 000 m的山峰共有85座,在慕士塔格山主峰周围有32座,公格尔山主峰周围有53座(苏珍等,1989)。巨大的山顶和高耸的山峰,是冰川发育的良好场所,加之高大的山体阻挡了西风带,使西风所携带的水汽得以补给该区(苏珍等,1989)。因此,帕米尔东北缘地区得以发育巨大规模的冰川。

帕米尔东北缘冰川的主要类型是山谷冰川,其形态是从山谷向外呈辐射状延伸(上官冬辉等,2005)。据中国科学院兰州冰川冻土研究所的相关数据,东帕米尔共有现代冰川466条,冰川面积898.08 km²,冰川储量95.662 7 km³。其中,大小公格尔山共有现代冰川327条,冰川面积640.15 km²,冰川储量69.667 5 km³;号称“冰川之父”的慕士塔格山有现代冰川139条,冰川面积257.93 km²,冰川储量25.995 2 km³(Su and Wang, 1989)。

公格尔—慕士塔格冰川经历了至少三次大冰期和至少10次小的冰进(Seong et al., 2007)。Yao et al.(2012)年对青藏高原及其周边冰川状态的研究,指出过去30年中青藏高原及其周边的冰川一直处于急剧的退缩状态。上官冬辉等(2005)基于ASTER影像和已有的冰川编目数据,解译了冰川从1962/1966年到2001年的变化特征,整体上是西北侧呈前进趋势,而其他大部分都有不同程度的退缩;其中,东南坡退缩比例最大,东北坡退缩量最大,而西南坡的退缩比例和退缩量均最小(曾磊等,2013)。略微增加的降雨量无法补偿冰川的消融量(Khromova et al., 2006)。东帕米尔、天山和兴都库什连续的冰川范围的减小对应了中亚日益增长的干旱(Fuchs et al., 2014)。

(2)帕米尔西北缘

中帕米尔是塔吉克斯坦境内最大的冰川区,冰川面积为2 473 km²(Schetinnikov, 1998)。Fedchenko冰川是帕米尔高原上最大的冰川,也是北半球中纬度地区最大的高山树枝状冰川之一($38^{\circ}15'N$, $72^{\circ}15'E$)(Aizen et al., 2009)。Fedchenko冰川由一些堆积区的小支流盆地和一个山谷冰川舌组成,总面积649 km²,长77 km,高程在2 900~300 m,雪线在~4 700~4 800 m(Dolgushin and Osipova, 1989),冰川总体积在124 km³左右(Schetinnikov, 1998)。1958年的地震和重力调查显示,冰川厚度在冰川舌处为~200~250 m,在海拔~4 750~4 850 m的冰川中部为800 m,在海拔5 000 m处厚度则达到1 000 m

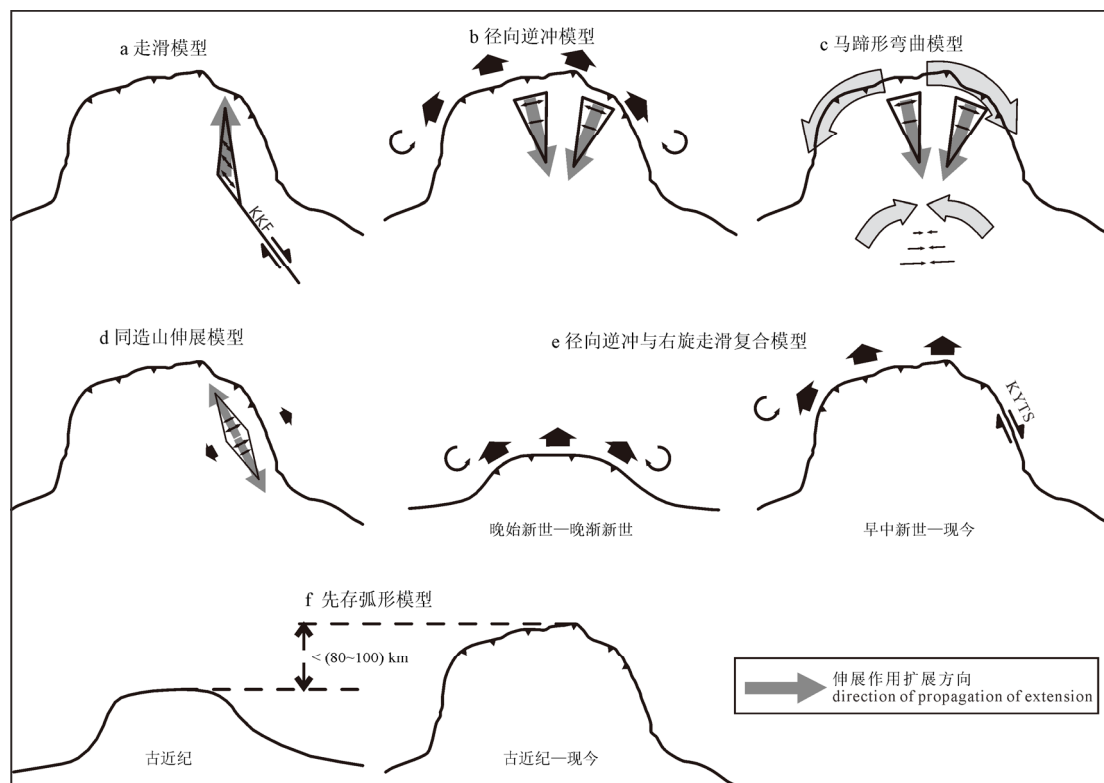


图 4 帕米尔弧形构造带演化动力学模型

(据 Robinson et al., 2004; Cowgill, 2010; Bosboom et al., 2014; Chen et al., 2018)

Fig. 4 Models for the evolution of the Pamir Arcuate Structural Belt

(modified after Robinson et al., 2004; Cowgill, 2010; Bosboom et al., 2014; Chen et al., 2018)

(Berzon et al., 1962)。两次冰川历史调查显示 Fedchenko 冰川在现今仅发生了小面积的变化, 冰川主体并未发生大的变化(Regensburger, 1963)。

3 帕米尔弧形构造带的动力学模型

关于帕米尔弧形构造带形成的动力学模型有走滑模型、径向逆冲模型、马蹄形弯曲模型、径向逆冲与右旋走滑复合模型、同造山模型和先存弧形模型 6 种(图 4)。其中, 走滑模型、径向逆冲模型和马蹄形弯曲模型可解释帕米尔弧形构造带的形成和东西向的伸展作用, 径向逆冲与右旋走滑复合模型很好地解释帕米尔弧形构造带东西段形成的差异性, 而同造山模型可解释帕米尔内部东西向和南北向的伸展作用。前五种模型都是基于帕米尔在新近纪以来的逐渐向北突进, 而先存弧形模型则提出构造带的弧形形态在古近纪之前已经形成, 之后由于印度-欧亚板块的碰撞再次向北突进。

3.1 走滑模型

走滑模型认为帕米尔弧形构造带的形成可能受控于西缘的达瓦孜左行走滑断层(DF)、东缘的喀什—叶城右行走滑系统(KYTS)和喀喇昆仑右行走滑断层(KKF)(Peltzer et al., 1988; Murphy et al., 2000; Bershaw et al., 2012), 并认为喀喇昆仑断层(KKF)在新生代的持续向北扩展导致了帕米尔内部

发生东西向的伸展(Strecker et al., 1995; Murphy et al., 2000; Chevalier et al., 2011)。但是该模型要求伸展作用最先开始于公格尔伸展系统(KES)的南端, 之后逐渐向北扩展, 且最大伸展量在南部(Robinson et al., 2004), 事实上公格尔伸展系统(KES)总体上北宽南窄(Robinson et al., 2004, 2007), 而且这一模型也很难解释帕米尔西缘发生的逆时针旋转。

3.2 径向逆冲模型

径向逆冲模型是基于帕米尔西缘塔吉克盆地在新生代发生了逆时针旋转(Bazhenov et al., 1994; Thomas et al., 1994)而被提出, 这一模型认为帕米尔弧形构造带沿主帕米尔逆冲断层(MPT)呈放射状逆冲, 并使帕米尔内部发生东西向的伸展作用(Strecker et al., 1995)。该模型要求伸展作用最先开始于公格尔伸展系统(KES)的北部, 之后逐渐向南扩展, 最大伸展量在北部(Robinson et al., 2004)。虽然这与公格尔伸展系统(KES)总体上北宽南窄的状况(Robinson et al., 2004, 2007)相符, 但是无法解释帕米尔弧形构造带东缘的塔里木盆地西南部中新世以来未发生了大角度的顺时针旋转以及喀什—叶城走滑系统(KYTS)的大规模右旋走滑。

3.3 马蹄形弯曲模型

马蹄形弯曲模型是用于解释同时发生在南迦

帕尔巴特(Nanga Parbat)的挤压缩短作用和帕米尔内部东西向的伸展作用,并认为左旋走滑的恰曼断层控制着帕米尔的弧形造山过程(Yin et al., 2001)。该模型的伸展变形机制与径向逆冲模型不同,但二者所揭示的帕米尔内部东西向伸展作用的运动学模式是相同的。

3.4 同造山伸展模型

同造山伸展模型认为主帕米尔逆冲断层(MPT)持续逆冲,使得上盘的公格尔山—慕士塔格峰穹窿不断隆升,并在穹窿的西侧形成公格尔伸展系统(KES)(Arnaud et al., 1993; Brunel et al., 1994)。Thiede et al.(2013)和 Sobel et al.(2013)对公格尔伸展系统(KES)的形成提出了新的动力学机制解释方案,认为沿着阿莱俯冲板片的东边界存在一条俯冲边界转换断层,大量的热沿着该断层上涌,导致帕米尔东缘公格尔伸展系统(KES)的发生。该模型要求伸展作用最先开始于公格尔伸展系统(KES)的中部,再逐渐向南北两端扩展,且最大伸展量在中部,这与公格尔伸展系统(KES)总体上北宽南窄的现象(Robinson et al., 2004, 2007)不符。前文中提到的 Stearns et al.(2015)针对南北向伸展作用的解释方案,认为俯冲的印度板片在 23~20 Ma 发生后撤和拆离,导致帕米尔地壳由南北向挤压转为南北向伸展,中下地壳的物质被快速剥露,形成 Yazgolom、Sarez、Muskol、Shatput 及 Shakh dara-Alichur 穹窿。

3.5 径向逆冲与右旋走滑复合模型

径向逆冲与右旋走滑复合模型是为了更好地匹配帕米尔东、西缘构造活动的差异性,认为帕米尔弧形构造带的形成过程可能分为 2 个阶段,即在始新世—渐新世,以径向逆冲为主;而从早中新世开始,帕米尔东缘以右旋走滑为主,西缘则一直为径向逆冲(Cowgill, 2010; Bosboom et al., 2014)。但该模型未对造成帕米尔东、西缘不同演化形式、东缘从径向逆冲转为右旋走滑和内部发生东西向伸展的机制作出解释。

3.6 先存弧形模型

该模型提出弧形构造带的弧形形态不完全是在新近纪以来形成,弧形形态在古近纪之前已经形成,在新近纪由于印度-欧亚板块的碰撞再次向北突进(Chen et al., 2018)。

Chen et al.(2018)总结了帕米尔东北缘前陆地区 12 个沉积剖面的地层学特征,发现了三套由隆升所造成的古近纪冲积砾岩层: (1)早古新世,新特提斯洋闭合时亚洲板块南缘阿尔卑斯型大陆边缘的弧后变形; (2)晚古新世—早始新世,对印度-帕米尔碰撞的即时响应; (3)中始新世—渐新世,印度-亚洲板块的持续汇聚。这些古近纪的冲积砾岩最大粒径

在 5~10 cm,这种粒径的沉积物从扇根最多只能被搬运 60 km 远(Shukla et al., 2001),而结构分析发现这些砾岩在渐新世—中新世向北位移了 20~50 km(Coutand et al., 2002; Cheng et al., 2016; Chapman et al., 2017)。所以认为,在古近纪,帕米尔位于现今位置以南 80~110 km,即在帕米尔新生代向北位移之前有一个先存的向北突出的弧形形态。这个先存的弧形形态可能是古生代—中生代印度-亚洲板块碰撞之前,欧亚板块南缘一系列的块体拼合所造成的。

以上这些模型基本上是从某一个侧重面提出来的端元性模型,但是帕米尔弧形构造带在新生代时期经历了复杂的构造演化过程,并遭受了强烈的剥蚀作用,构造带内部的很多记录被强烈改造或消失。因此,需要将造山带与盆地相结合、深部过程与浅部响应相结合及其空间变化与时间演化相结合开展系统研究,提出更为合理的动力学模型。

4 帕米尔弧形构造带构造过程与地貌演变研究的关键科学问题

综合前述研究现状分析,前人对帕米尔弧形构造带的晚新生代的构造过程开展了系统的研究,并提出不同的扩展模型。而这些模型的提出很重要地依赖于帕米尔弧形构造带内部不同性质的断裂带的运动学特征分析,而这些断裂带的发育特征很可能是受控于深部的地质过程,它们是深部地质过程在浅部的响应。同时,构造过程决定了地貌的演化和气候的变化,帕米尔弧形构造带特殊的构造过程造就了帕米尔独特的地貌特征和水系的变迁;而地貌的演化和气候的变化也制约了构造的形成,通过对帕米尔水系和地貌演化过程的分析,可以获知某些断裂带或某区域的构造活动特征,从而进一步限定帕米尔弧形构造带晚新生代的演化模型。

综上分析帕米尔弧形构造带晚新生代构造研究,有以下 3 个方面的科学问题值得进一步关注。

(1)问题一:帕米尔弧形构造带内部不同性质的断裂带在弧形构造带的扩展过程中扮演了十分重要的角色,要建立合理的扩展过程与动力学模型,如何系统厘定构造带内部主要断裂带的运动学特征及其相互关系是关键。

帕米尔弧形构造带内部发育大量不同性质的断裂带(图 1),前人对帕米尔弧形构造带周缘的主帕米尔逆冲断层(MPT)、帕米尔前缘逆冲断层(PFT)、喀什—叶城右行走滑系统(KYTS)、达瓦孜左行走滑断层(DF),以及对应于片麻岩穹窿剥露的公格尔山断层、南帕米尔剪切带等拆离断层得到了一定程度的研究,这些断层的启动时间、运动量、

平均运动速率等得到揭示。但是, 这些数据多来自断层的一个部位, 不能简单地将其代表整条断层的活动特征, 如主帕米尔逆冲断层(MPT)在阿莱盆地的南缘现今仍以 10~15 mm/a 的速率吸收南北向的缩短作用(Zubovich et al., 2010; Ischuk et al., 2013), 但这一断层在帕米尔东北缘已不活动(陈杰等, 2011)。对于帕米尔弧形构造带的形成过程, 特别是构造带的东部, 现有的单一模型都无法完美地解释, 不同时期的演化模型可能不同(Bosboom et al., 2014), 但不同阶段演化模型的转换时间、转换机制等问题的解决需要对构造带内部的不同断裂带运动学特征和相互关系的精细厘定。

(2)问题二: 帕米尔弧形构造带的形成不仅受水平方向上差异性的控制, 而且还受控于垂向上的变化, 也就是深部地质过程对构造带形成的控制, 因此将深部结构与浅部构造相结合、深部地质过程与浅部响应相结合开展研究, 探讨构造带形成的深部地质过程控制, 也是揭示帕米尔弧形构造带形成机制的关键。

Stearns et al.(2015)认为帕米尔弧形构造带中——南部的南北向伸展作用可能与印度俯冲板片的后撤和拆离有关, 但其与俯冲板片后撤和拆离的具体关系仍有待研究。Sobel et al.(2013)和 Thiede et al.(2013)提出帕米尔东部公格尔山伸展系统(KSE)的形成则可能与一条存在于帕米尔东缘深部的俯冲边界转换断层有关, 但是还缺少足够的地球物理证据来支撑这一观点。另外, 帕米尔东南部塔什库尔干杂岩体和 Dunkeldik 火山岩的形成机制也仍存争议, 二者可能与印度俯冲板片的拆离(Jiang et al., 2012; Stearns et al., 2015)或俯冲边界转换断层的存在(Sobel et al., 2013)有关。同样, Sippl et al.(2013b)认为欧亚板块向南俯冲于帕米尔之下, 且有中——上地壳的物质滞留在 80~100 km 深处, 这些物质的滞留或折返与片麻岩穹窿剥露的关系仍有待研究。因此, 要揭示帕米尔弧形构造带形成机制, 必须将深部结构与浅部构造相结合、深部地质过程与浅部响应相结合开展研究, 探讨构造带形成的深部地质过程控制。

(3)问题三: 作为一个现今还在活动的弧形构造带, 构造过程对地貌演化和气候变化起到重要的控制作用, 要揭示帕米尔弧形构造带地貌演化及其对气候变化的影响, 必须将构造过程、气候特征与地貌演化作为一个耦合系统开展研究。

作为一个现今还在活动的弧形构造带, 目前对其地貌特征的研究相对比较缺乏, 讨论构造过程、气候特征与地貌演化之间的关系则更少。Schoenbohm et al.(2014)认为帕米尔东北缘公格尔

山——慕士塔格峰的地貌主要受控于公格尔——慕士塔格峰穹窿的剥露和第四纪冰川作用; 塔什库尔干谷地西侧南萨雷阔勒岭的地貌受控于塔什库尔干断层的构造活动, 其地貌特征反映出该断层至少在晚第四纪具有中段构造活动强度大于南、北两段的特征。但是帕米尔弧形构造带还有大量的地貌演化的问题, 例如叶尔羌河主河道的形态是否反映了帕米尔东缘的右旋走滑? 盖孜河在公格尔山段是先存河流还是其溯源侵蚀后形成的(Brookfield, 2008)? 这些问题的解决必须依赖于公格尔——慕士塔格峰穹窿的剥露速率、公格尔山断层的活动特征和/或第四纪冰川作用等的研究成果。

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