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走滑断层研究进展及启示

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**Abstract:** Since the emergence of the strike-slip fault concept, the importance of strike-slip fault in geological science research has gradually been realized. Understandings of strike-slip faults in terms of geometry, kinematics, dynamics and tectonic significance have accelerated the study on strike-slip fault. However, there are still some limitations in its classification and mechanism analysis. Based on the literature review of strike-slip fault, the principle, concept and related terms are summarized, then the displacement characteristics, identification marks, mechanical mechanism, strike-slip derived or associated structures, characteristics of strike-slip basins, classification and examples of strike-slip faults are analyzed. Combined with the mechanical mechanism of strike-slip fault, a new classification method of strike-slip fault is proposed, and a brief analysis of typical faults is made using the new classification method, such as the San Andreas fault on the west coast of the United States, the Alpine fault in New Zealand, the Tanlu fault and the Altyn Tagh fault in China.

**Key words:** strike-slip fault; classification; dynamics; transform fault; compound transfer fault

**摘 要:** 自走滑断层概念提出之后, 走滑断层在地质科学研究上的重要性逐渐体现出来, 并在几何学、运动学、动力学及其构造意义等方面取得了重要的认识, 使得走滑断层的研究得到快速的发展, 但其分类及其成因机制分析还存在一定的局限性。在走滑断层相关文献调研的基础上, 文章对走滑断层原理、概念和相关术语发展历程进行了归纳总结, 同时也对走滑断层的位移特征、识别标志、力学性质、走滑派生或伴生构造、走滑盆地特征、走滑断层分类及走滑断层实例等研究成果进行了系统性研究分析。在此基础上, 结合走滑断层的力学机制, 提出走滑断层新分类方式, 并运用新的分类方式进一步对美国西海岸圣安德列斯断层、新西兰的阿尔卑斯断层和中国著名的郯庐断裂带以及阿尔金断裂带等典型断层进行简要分析。

**关键词:** 走滑断层; 分类; 动力学; 转换断层; 复式变换断层

**中图分类号:** P542      **文献标识码:** A

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## 0 引言

断层是地壳岩石（地质体）发生破裂并顺破裂面发生位移的一种构造，其广泛发育，是地壳中最重要的构造类型（徐开礼等，1989；Twiss et al., 1992）。走滑断层作为断层的三大类之一，自19世纪末被提出以来，其在地质研究上的重要性逐渐体现出来，吸引着无数学者对走滑断层进行研究，使得走滑断层的研究得到了快速地发展。经过一百多年的研究积累，已经取得了一系列的认识。

文章主要对走滑断层发展历史以及研究成果进行总结分析，其中包括：①走滑断层概念（Mckay, 1890, 1892；Suess, 1892；Noble, 1926, 1927；Freund, 1971；Sengör et al., 1985；Bates and Jackson, 1987；Sylvester, 1988；徐开礼和朱志澄，1989；徐嘉炜，1995；陆克政等，2001；漆家福等，2006）及其相关术语（Anderson, 1905；Reid et al., 1913；Gill, 1935；Kennedy, 1946；Moody and Hill, 1956, 1958；Wilson, 1965；Dahlstrom, 1969；Wilcox et al., 1973；Gibbs, 1983；Christie-Blick et al., 1985；Woodcock, 1986；Sylvester, 1988；Morley, 1990；Twiss and Moores, 1992；Mueller and Talling, 1997；刘和甫，1999, 2004；Escalona and Mann, 2006；漆家福等，2006）的提出和发展；②走滑断层位移特征的识别（Gill, 1935；Billings and Rabbitt, 1947）；③走滑断层识别标志（Crowell, 1962；Christie-Blick and Biddle, 1985；Harding et al., 1985；Woodcock and Fischer, 1986；Zolnai, 1991；漆家福和陈发景，1995）；④走滑派生或伴生构造（走滑应变椭圆）（Cloos, 1928；Riedel, 1929；Pakiser, 1960；Skempton, 1966；Clayton, 1966；Morgenstern and Tchalenko, 1967；Ramsay, 1967；Tchalenko and Ambraseys, 1970；Tchalenko et al., 1970；Harland, 1971；Wilcox et al., 1973；Crowell, 1974；Harding, 1974；Harding and Lowell, 1979；Aydin and Page, 1984；Christie-Blick and Biddle, 1985；Tuminas, 1988；Sylvester, 1988；Allen and Allen, 2005；刘和甫等，1999）；⑤走滑作用形成的盆地特征（Burchfiel and Stewart, 1966；Crowell, 1973, 1982；Aydin and Nur, 1982；Hempton and Dunne, 1984；Sylvester, 1988；Miall, 1990；刘和甫，1993）；⑥走滑断层的力学

机制（Anderson, 1905；Riedel, 1929；McKinstry, 1953；Moody and Hill, 1956；Morgenstern and Tchalenko, 1967；Tchalenko, 1970；Tchalenko and Ambraseys, 1970；Freund, 1974；Aydin and Page, 1984；Sylvester, 1988；徐嘉炜，1995；刘和甫等，1999）；⑦走滑断层的分类方法（Hill, 1959；Woodcock, 1986；Sylvester, 1988；Miall, 1990；刘和甫等，1999, 2004；夏义平等，2007）。

在此基础上，研究发现目前所提出的走滑断层的分类方式具有一定的局限性，主要表现为：分类界限不够明确，同一分类系统标准不统一、考虑因素不全面以及同一分类中不同类型具有包含、从属关系。从而对具体走滑断层类型的判定和识别造成困难。因此，文章以走滑断层的基本概念、原理、观点以及走滑断层的几何学、运动学、动力学以及构造意义为基础，根据走滑断层的动力学机制，对走滑断层提出新的分类方案，并运用新的分类方案对全球著名的走滑断层进行实例分析，进一步阐述了走滑断层的力学机理。

## 1 走滑断层相关术语的发展历程

关于走滑作用这一现象描述最早可以追溯到公元前347年Zechariah的著作中（Freund, 1971）。虽然书中没有明确提出走滑的概念，但Zechariah对相互错动的两盘山体进行了总体描述，因此可以将其看作是对走滑这一现象论述的起源（Sylvester, 1988）。

18世纪50年代，著名的瑞土地质学家Arnold Escher von der Linth可能是第一个发现走滑断层并对其进行初步描述的学者（Suess, 1892；Sengör et al., 1985）。

按照Freund（1971）的观点，Mckay（1890, 1892）最早将与地震相关的走滑现象记录并出版公布。随后在1906年旧金山大地震中，圣安德列斯断层突然发生了最大位移为4.7 m的右旋滑动，走滑这一现象从而在科学界得到充分印证（Hill, 1982；Sylvester, 1988）。

Noble（1926, 1927）通过对圣安德列斯断层区域的一些独特的非海相砂岩进行地层对比，研究走滑断层的走滑位移量，首次提出了走滑位移量可达数十千米，打破了以往断层位移量较小的

观点, 随后其他学者通过对不同的走滑断层研究认为走滑断层的位移量可达数百千米, 例如圣安德列斯断层 (Wallace, 1949; Hill and Dibble, 1953)、阿尔卑斯断层 (Wellman, 1955)、死海裂谷 (Quennell, 1958, 1959; Bartov et al., 1980) 以及大格林断层 (Kennedy, 1946), 进一步促进了对走滑断层的认识。

随着对走滑断层认识的不断深入, 其他相关术语也陆续提出。其中, 平移断层 (transcurrent fault) 被用来描述火成岩、变质岩以及沉积岩中近乎直立的走滑断层 (Geikie, 1905; Kennedy, 1946; Moody and Hill, 1956, 1958; Wilcox et al., 1973; Christie-Blick and Biddle, 1985), 然而许多学者认为用平移断层描述走滑断层不能详细解释走滑断层成因上的关系, 故将平移断层作为不切割整个岩石圈的走滑断层的统称 (Sylvester, 1988); 扭动断层 (wrench fault) 概念的提出 (Moody and Hill, 1956) 是用来描述区域性的、直立的、并卷入结晶基底的走滑断层 (Kennedy, 1946)。

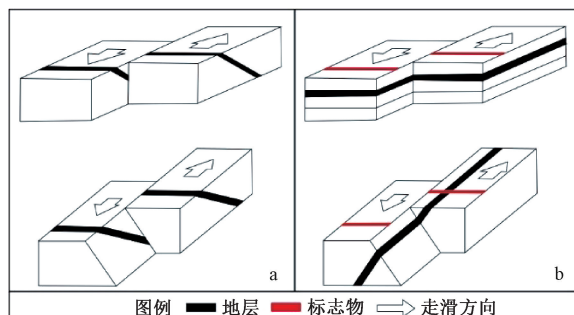
走滑断层 (strike-slip fault) 的基本定义不同学者有不一样的见解, Bates and Jackson (1987) 将走滑断层定义为断层的主运动方向与断层的走向基本平行的断层 (Sylvester, 1988; 徐开礼和朱志澄, 1989; 徐嘉炜, 1995; 漆家福等, 2006); 陆克政等 (2001) 认为走滑断层的基本含义是由扭应力或剪应力引起地壳或岩石圈沿着某些构造边界或特定的构造带发生走滑作用的断层。

对于走滑性质的断层, 还提出了其他术语, 包括: ① Wilson (1965) 提出的转换断层 (transform fault), 其作为重要的板块边界类型, 是切穿岩石圈或地壳并调节板块之间运动的走滑断层; ② Gibbs (1983) 引入伸展盆地中的变换断层 (transfer fault), 将其作为伸展变形 (或位移) 的变换带; ③ Gill (1935) 提出撕裂断层 (tear fault) 并进行定义, 也称掀断层, 随后 Christie-Blick and Biddle (1985) 将其定义为调节某一外来体内或与外来体相邻构造单元之间的差异位移, 之后也用来描述与逆冲断层突然终止相关的走滑断层 (Twiss and Moores, 1992; Mueller and Talling, 1997; Escalona and Mann, 2006), 国内主要将撕裂断层定义为在逆冲断层系统中的横向或斜向断层, 并与逆冲断层的位移过程有关的, 具有走滑位移性质的断层 (刘和甫, 1999, 2004; 漆家福等, 2006)。

## 2 走滑断层的研究现状

### 2.1 走滑断层的位移特征

走滑断层的位移主要通过与其走向斜交的非水平地层 (标志层) 进行识别, 也可以通过标志物的错动进行识别 (图 1); 但在其他复杂情况下会导致标志层在平面上或剖面上的视错动, 称为断层效应 (Gill, 1935; Billings and Rabbitt, 1947)。



a—地层走向与断层走向一致; b—地层走向与断层走向斜交

图 1 走滑断层与地层交互关系示意图 (据 Sylvester, 1988 修改)

Fig. 1 Schematic diagram of the interactive relationship between strike-slip fault and stratum (modified after Sylvester, 1988)

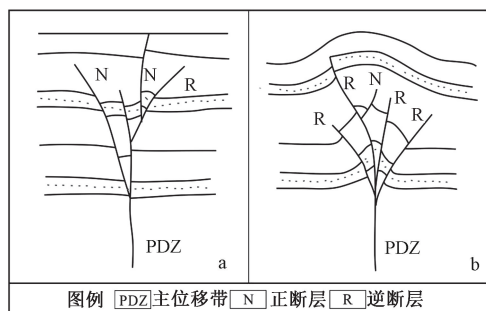
### 2.2 走滑断层的识别标志

走滑构造变形往往会形成一些特定的识别标志, 也是识别走滑构造的基础。走滑断层识别标志分为剖面标志和平面标志。其中剖面标志有: 正花状构造和负花状构造 (图 2) (Christie-Blick and Biddle, 1985); 海豚效应和丝带效应 (图 3) (Zolnai, 1991); 走滑断层内部构造和夹块 (Zolnai, 1991); 剖面上地层的不连续现象等 (漆家福和陈发景, 1995)。平面标志有: 线性延伸或带状展布 (Harding et al., 1985; Woodcock et al., 1986); 走滑带两侧地质界线的水平错开 (Crowell, 1962; Harding et al., 1973) 等。

### 2.3 走滑断层的派生或伴生构造

走滑断层作用常形成派生或伴生构造, 它们受沿走滑断裂带不同位置上若干因素的控制, 其中包括走滑断层走向的变化及其排列和组合方式、岩石的力学性质、应力的大小和方向、应变速率、先存构造和滑动矢量方向及位移量等。

走滑派生或伴生构造的规律可以用走滑应变椭圆 (图 4) 来表达, 最初来自 Cloos (1928) 和 Riedel (1929) 的简单剪切物理模拟实验。相关的



a—负花状构造；b—正花状构造图

图2 花状构造示意图 (据 Christie-Blick and Biddle, 1985 修改)

Fig. 2 Schematic diagram of flower structure (modified after Christie-Blick and Biddle, 1985)

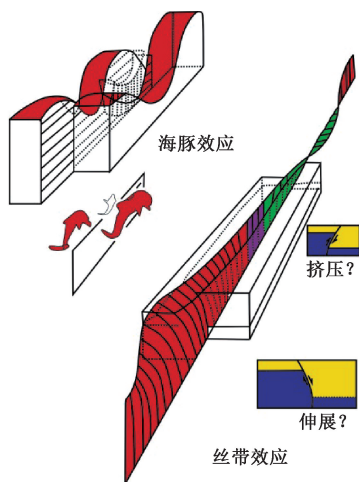


图3 海豚效应和丝带效应示意图 (据 Zolnai, 1991 修改)

Fig. 3 Schematic diagram of dolphin effect and dextral ribbon effect (modified after Zolnai, 1991)

构造 (图4) 包括: ①主位移带 (principal displacement zone, 简称 PDZ), 即与走滑构造带走向一致、连续的走滑断层位移带 (Tchalenko, 1970; Tchalenko and Ambraseys, 1970); ②里德尔 (R) 剪切 (Riedel shear), 走滑方向与主位移带方向相同两者呈小角度 (一般小于  $15^\circ$ ) 相交, 其夹角相当于岩石内摩擦角的一半 ( $\varphi/2$ ), 锐角角顶指示本盘断块的运动方向 (Tchalenko, 1970), 也称为同向走滑断层 (synthetic strike-slip fault) (Cloos, 1928); ③共轭里德尔 (R') 剪切 (conjugated Riedel shear), 走滑方向与主位移带反向相反, 两者呈大角度相交, 与 R 剪切共轭, 共轭角通常为  $60^\circ \sim 70^\circ$  (Tchalenko, 1970; Tchalenko and Ambraseys, 1970), 也称为反向走滑断层 (antithetic strike-slip fault) (Cloos, 1928); ④同向剪切破裂 (P), 走滑方向及与主位移带的

夹角同里德尔 (R) 剪切, 但锐夹角角顶指向对盘断块的位移方向, 也称为次级同向走滑断层 (secondary synthetic strike-slip fault) (Skempton, 1966; Tchalenko, 1970; Tchalenko and Ambraseys, 1970); ⑤张性 (T) 破裂 (extension fracture), 是剪切应力作用下产生的局部张性破裂, 与主位移带大角度相交的延伸不长的正断层组或张节理组, 走向与走滑应变椭圆中的局部伸展方向垂直 (Tchalenko and Ambraseys, 1970); ⑥雁行式褶皱 (en echelon folds) 和逆断层组, 一般只在构造带中局部发育, 走向与应变椭圆的局部收缩方向垂直 (Ramsay, 1967; Harding and Lowell, 1979); ⑦Y 剪切破裂与主位移带平行 (Morgenstern and Tchalenko, 1967)。在走滑应变椭圆中, PDZ、R、P 和 Y 剪切变形过程中相互作用, 形成走滑断裂带 (图4)。

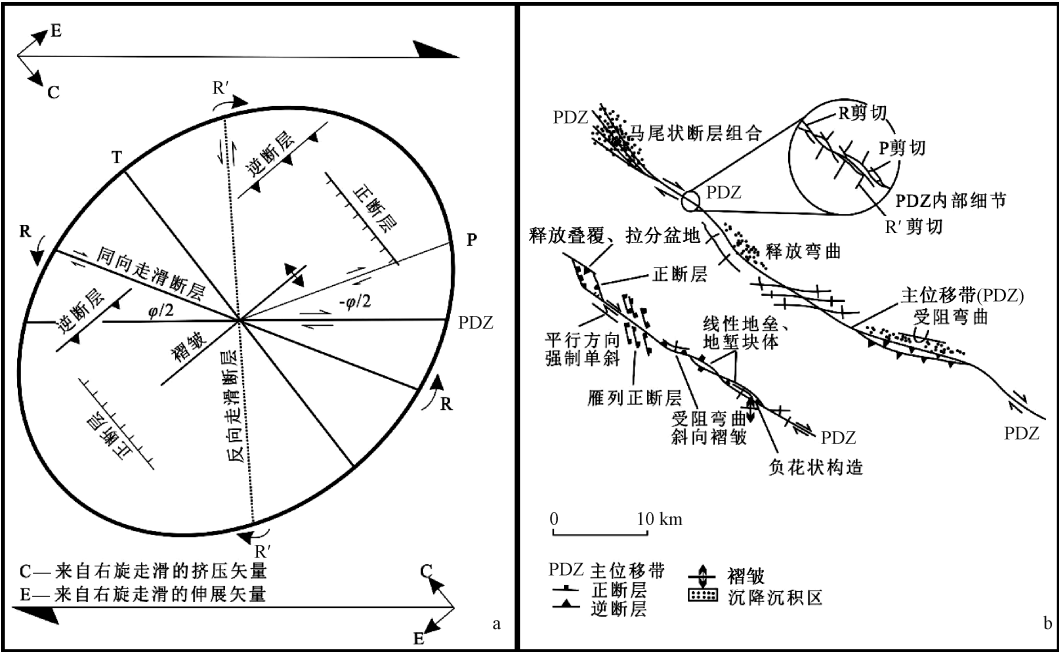
在走滑断层主位移带构造组合中, 局部构造带走向和走滑矢量方向决定断裂带局部会聚和离散 (Pakiser, 1960; Clayton, 1966; Crowell, 1973; Christie-Blick and Biddle, 1985; Allen and Allen, 2005), 走滑会聚区域会造成收敛挤压或扭压, 造成地壳缩短和隆起 (Harland, 1971; Wilcox et al., 1973; Crowell, 1974); 走滑离散区形成释放弯曲, 并伴随地壳伸展, 沉降以及拉分盆地的形成 (Harland, 1971; Crowell, 1974; Sylvester, 1988); 主位移构造带构造组合也受到 R 和 P 剪切作用的控制, 并常伴生有正断层、褶皱、线性地垒、地堑块体等构造组合 (图4) (Allen and Allen, 2005)。

雁列褶皱通常在走滑断层附近出现, 认为是剪切变形的结果 (Sylvester, 1988; 刘和甫等, 1999), 雁行褶皱可以发育在主走滑断层上方或附近一个相对狭窄而稳定性的区域内; 也可以形成于两个走滑断层之间的广阔地带 (Aydin and Page, 1984)。雁列褶皱排列的倾斜方向与剪切主应力方向相同 (图5) (Harding, 1974; Dibblee, 1977; Harding and Lowell, 1979; Harding and Tuminas, 1988), 即右旋形成右阶褶皱, 左旋形成左阶褶皱 (Sylvester, 1988)。

## 2.4 走滑断层作用与沉积盆地

走滑断层作用也可以控制盆地的形成, 当走滑断层伸展弯曲时, 形成走滑-拉分盆地 (pull-apart basin) (Burchfiel and Stewart, 1966); 当走滑断层挤压弯曲时, 发育挤压山岭, 形成走滑-挠曲盆地; 走滑断层呈雁列时, 形成盆地或山岭; 走滑

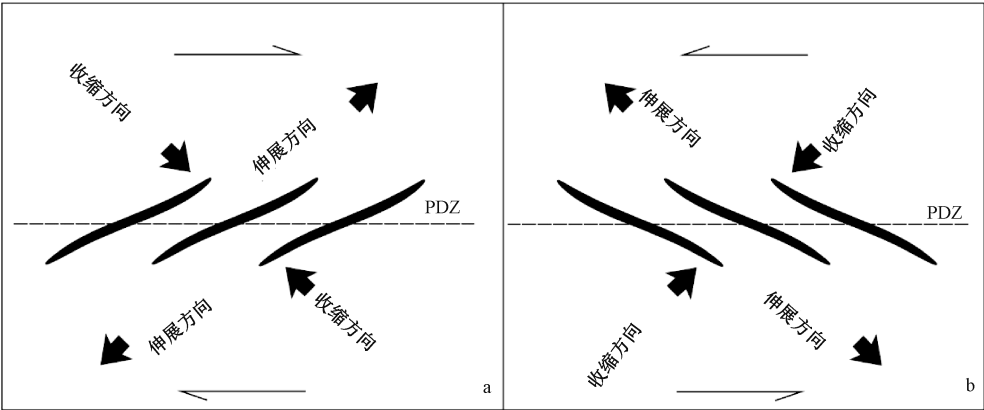




a—右行力偶产生的走滑应变椭圆（据 Harding, 1974 修改）；b—右行走滑断层主位移带（PDZ）构造组合（据 Allen and Allen, 2005 修改）

图 4 右行走滑断层特征

Fig. 4 The features of the right-lateral (dextral) strike-slip fault (modified after Harding, 1974; Allen and Allen, 2005)



a—右旋简单剪切下雁型褶皱；b—左旋简单剪切下雁型褶皱

图 5 简单剪切下雁型褶皱图（据 Sylvester, 1988 修改）

Fig. 5 Geometry of en echelon folds of a simple shear (modified after Sylvester, 1988)

断层侧向分支时，发育盆地或山脊（Crowell, 1974; Aydin et al., 1982; Sylvester, 1988; 刘和甫, 1993）。同时也有不同学者依据走滑断层边界断层的几何学和运动学对走滑盆地进行分类（Nilsen and Sylvester, 1999）。

在沉积盆地中，走滑盆地有个显著的特征，沉积中心的迁移方向与盆地的走滑运动方向一致，其中沉积物沉积特征为“威尼斯百叶窗”或“地层叠瓦状”（图 6）（Crowell, 1974, 1982; Hempton and Dunne, 1984）。

### 2.5 走滑断层作用的力学机制

目前，对走滑断层的力学机制分析主要运用纯剪切（pure shear）和简单剪切（simple shear）这两种模式解释（Aydin and Page, 1984; Sylvester, 1988; 徐嘉炜, 1995; 刘和甫等, 1999）。

纯剪切是 Anderson（1905）为了解释均匀介质中断层形成与三个主应力轴关系的模型而提出的，也被称作 Coulomb-Anderson 模型。纯剪切模型中，一旦发生剪切破裂（ $N36^{\circ}W$ ），就一定会形成与之共轭的剪切破裂（ $N36^{\circ}E$ ）（图 7a），其中共轭

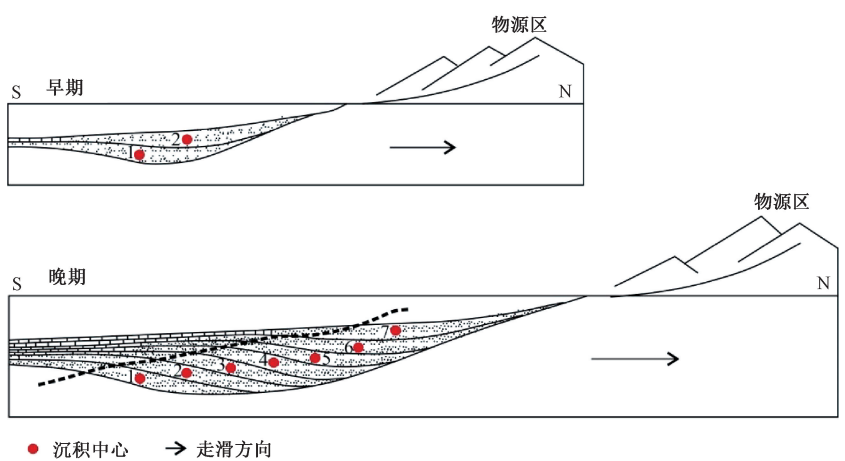


图 6 右旋走滑盆地沉积中心迁移示意图 (据 Crowell, 1982 修改)  
Fig. 6 Schematic diagram of sedimentary center migration in the right-lateral strike-slip basin (modified after Crowell, 1982)

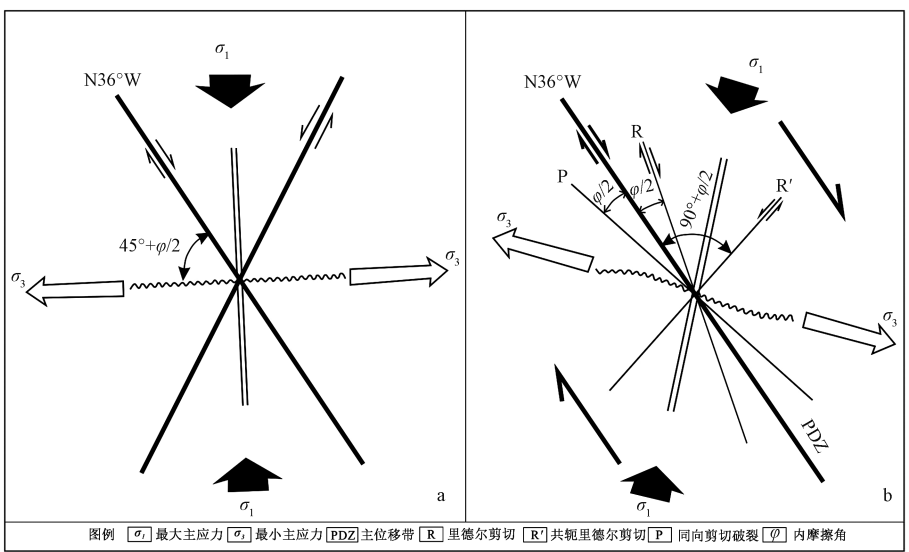
破裂面的交线与应力场  $\sigma_2$  一致，共轭破裂面的锐角平分线与应力场  $\sigma_1$  的方向一致，共轭破裂面钝角平分线与应力场  $\sigma_3$  的方向一致，纯剪切作用下应变主轴在变形过程中始终不旋转 (图 7a) (Anderson, 1905; McKinstry, 1953; Moody and Hill, 1956; Aydin and Page, 1984; Sylvester, 1988)。

简单剪切模型是 Riedel (1929) 提出的一种走滑断层力学模型 (图 7b)。简单剪切模型中，主走滑位移带方向与剪切应力的方向一致，同时也会诱导出次级应力场，在局部挤压应力 ( $\sigma_1$ ) 和局部伸展应力 ( $\sigma_3$ ) 下形成次级破裂 (R、R'、P、Y、T) 及其他伴生构造 (褶皱、正断层、逆断

层)，简单剪切变形过程中应变主轴一直旋转 (图 7b) (Riedel, 1929; Skempton, 1966; Morgenstern and Tchalenko, 1967; Tchalenko, 1970; Tchalenko and Ambraseys, 1970; Wilcox et al., 1973; Freund, 1974; Sylvester, 1988)。

2.6 走滑断层分类方案

随着对走滑断层认识的不断深入，走滑断层的分类方式应运而生，走滑断层的分类方式很多，目前学术界主流的分类方式有，①Hill (1959) 根据断层两盘的相对运动方向对走滑断层进行的分类：左旋走滑断层和右旋走滑断层；②Woodcock (1986) 根据在板块构造中的位置对走滑断层进行



a—纯剪切；b—简单剪切  
图 7 走向 N36°W 直立走滑断层相关构造平面图 (据 Aydin et al., 1984; Sylvester, 1988 修改)  
Fig. 7 Plan view of geometric relations among structures according to a vertical strike-slip fault which strikes N36°W (modified after Aydin et al., 1984; Sylvester, 1988)

的分类: 洋中脊转换断层、与海沟相连的走滑断层、与楔入相关的走滑断层、边界转换断层; ③Miall (1990) 根据在板块构造边界的不同构造部位对走滑断层进行的划分, 共分为四大类: 板块边界转换断层、离散边缘转换断层、聚敛边缘平移断层、缝合带平移断层; ④Sylvester (1988) 根据板块构造上的位置及卷入深度对走滑断层分为两大类: 板间转换断层和板内平移断层 (仅限于地壳); ⑤刘和甫等 (1999) 根据断层构造位置和卷入的深度对走滑断层的分类: 转换断层、平移断层、变换断层、撕裂断层; ⑥刘和甫等 (2004) 结合油气勘探中盆地—区带—圈闭三个层次对走滑断层进行的分类: 转换断层、变换断层、撕裂断层; ⑦夏义平等 (2007) 按照断层规模和对石油地质控制作用将走滑断层分为五类: 板块级走滑断层、盆地级走滑断层、区带级走滑断层、圈闭级走滑断层、显微级走滑断层。

总的来说, 通过对不同学者走滑断层分类方式总结分析, 认为目前存在的主要问题是: 走滑断层分类界限不够明确、同一分类体系中分类标准不一致、考虑因素不系统、在分类结果中具有

包含、从属关系, 从而使得对具体的走滑断层类型的判别造成困难, 例如在 Woodcock (1986) 和 Miall (1990) 的分类中只考虑了板块边界上的走滑断层; 在 Sylvester (1988) 和刘和甫等 (1999, 2004) 的分类中撕裂断层属于变换断层, 同一级别分类体系中具有包含关系; 在夏义平等 (2007) 的分类中考虑因素太少 (只有规模这个要素), 过于笼统; 从而易对走滑断层类型的准确识别造成困难。

因此, 在已有的分类基础上, 文章提出走滑断层新的分类方案, 首先根据动力学机制把走滑断层划分分为两大类: 主动走滑断层 (active strike-slip fault) 和被动走滑断层 (passive strike-slip fault)。主动走滑断层是指在走滑应力体制下形成的走滑断层 (纯剪切或简单剪切变形); 被动走滑断层是指在非走滑应力体制 (伸展应力体制和挤压应力体制) 下形成的走滑断层 (简单剪切变形)。被动走滑断层可分为转换断层和变换断层, 其中变换断层可进一步划分为单式变换断层和复式变换断层 (童亨茂等, 2015), 详细的分类方案另文报道。

表 1 走滑断层新的分类方案

Table 1 A new classification method of strike-slip faults

| I 级类型  | II 级类型 | III 级类型 | 实例分析                             |
|--------|--------|---------|----------------------------------|
| 主动走滑断层 |        |         | 塔里木盆地塔北隆起共轭走滑断裂系 (吴涛等, 2007)     |
|        | 转换断层   |         | 圣安德列斯断层 (Crowell, 1962)          |
| 被动走滑断层 |        | 单式变换断层  | Morley et al. (1990) 变换带分类方案中的类型 |
|        | 变换断层   | 复式变换断层  | 琼东南盆地宝岛变换带 (童亨茂等, 2015)          |

### 3 典型走滑断层实例及其类型分析

走滑断层的研究离不开理论与实践的结合, 不同学者依据理论对走滑断层实例进行研究分析, 同时也通过实例的研究进一步得到了新的认识。走滑断层作为目前研究的热点, 实例有很多, 如美国西海岸的圣安德列斯断层, 中国东部的郯庐断裂、西部的阿尔金断裂等, 这一系列走滑断层实例对走滑断层认识的提高和理论验证提供了更多的数据依据。

美国加利福尼亚海岸圣安德列斯断层是世界上最著名的走滑断层之一, 是太平洋板块边界和北美板块边界之间的右旋转换断层系统 (Crowell, 1952, 1962; Allen, 1957; Allen et al., 1960;

Crowall and Sylvester, 1979; Zoback and Zoback, 1980; Zoback et al., 1987; Wakabayashi, 1999), 目前对圣安德列斯断层的研究已经从几何学 (Dickinson et al., 1979a, 1979b; Andrews, 1980; Mandelbrot, 1983; King, 1983; Aviles et al., 1987; Okubo and Aki, 1987)、运动学 (Hill and Dibble, 1953; Hamilton and Myers, 1966; Freund, 1970; Zoback and Zoback, 1980; Zoback et al., 1987; Mount and Suppe, 1987; Wesnousky, 1988; Aydin and Schultz, 1990; Khoshmanesh and Shirzaei, 2018) 以及动力学 (Moody and Hill, 1956; Hamilton, 1969; Atwater, 1970; Molnar and Atwater, 1973; Freund, 1974; Garfunkel, 1974; Coney and Reynolds, 1977; Bohannon and Howell, 1982; Argus and Gordon, 1991, 2001; Henstock et al., 1997; Wakabayashi,

1999) 方面分析取得了重要的进展;同时也从古地磁等方面进行过相关研究 (Teissere and Beck, 1973; Beck, 1976; Simpson and Cox, 1977), 该断层是世界上研究程度最高的走滑断层。通过对圣安德列斯断层在中新世到上新世时期走滑位移量的研究, 基本认为在该时期的右旋走滑位移量将近 300 km, 性质为转换断层 (Wallace, 1949; Hill and Dibblee, 1953; Allen, 1957; Allen et al., 1960; Fletcher, 1967; Ehlig et al., 1975; Crowell, 1980, 1989), 目前, 对该断层的基本认识没有太多的争议。

郯庐断裂带作为中国东部一条巨大的北北东走向的断裂带, 自中生代以来经历过长期、复杂的演化 (朱光等, 2001)。目前多数学者认为该断裂带起源于华北和华南板块的碰撞, 但对于其具体的形成机制还存在不同的观点 (Zhang et al., 1984; Hsü et al., 1987; Xu et al., 1987; Lin and Fullex, 1990; Okay et al., 1992; Yin and Nie, 1993; Li, 1994; Chung, 1999; Gilder et al., 1999; Wang et al., 2003; Zhu et al., 2009)。目前对郯庐断裂带走滑方式的研究认识有很多, 具有相对比较统一的观点为: 中生代为左旋走滑 (徐嘉炜, 1985; 徐嘉炜和马国锋, 1992; 万天丰等, 1996; 陈宣华等, 2000; 王小凤, 2000; 朱光等, 2001, 2003; 张岳桥和董树文, 2008), 新生代为右旋走滑。但对于郯庐断裂带新生代右旋走滑位移量的大小及其成因机制具有很大的争议 (Okay et al., 1992; Yin and Nie, 1993; Li, 1994; Gilder et al., 1999; 陈宣华等, 2000; 王小凤, 2000; 施炜等, 2003; 朱光等, 2006; 孙晓猛等, 2010)。文章研究发现, 郯庐断裂带在新生代时期为变换断层带 (被动走滑断层带, 断层的活动由渤海湾盆地的伸展作用引起), 其中的每一条断层均为复式变换断层 (童亨茂等, 2015), 起到传递或调节伸展位移的作用。由于复式变换断层不同段落的位移量本身就不一致 (童亨茂等, 2015), 这样, 郯庐断裂带不同段落位移量不同的争议就可以得到解决。另外, 郯庐断裂带在新生代早期 (古新世和始新世时期) 主要表现为伸展变形, 没有明显的走滑变形, 把郯庐断裂简单地认为是一条走滑断裂带是值得商榷的。

阿尔金断裂带是中国西部一条规模巨大的、总体呈北东东向、现在仍在活动的走滑断裂带

(Molnar and Tapponnier, 1975; Tapponnier and Molnar, 1977)。阿尔金断裂带目前研究的焦点为断裂带活动方式及起始时间 (Vincent and Allen, 1999; Yue and Liou, 1999; 刘永江等, 2001; Wang et al., 2005; 李海兵等, 2006; Liu et al., 2007; 吕宝凤等, 2019)、新生代以来走滑活动历史以及走滑位移量大小 (Wang, 1997; Meyer et al., 1998; Yue et al., 2001, 2005; 陈正乐, 2001, 2006; Yin et al., 2002; Sun et al., 2005)、第四纪以来的滑移速率 (England and Molnar, 1997; Bendick et al., 2000; Zhang et al., 2004, 2007; Elliott et al., 2008) 等。与郯庐断裂带类似, 文章研究认为阿尔金断裂在新生代时期也应该是变换断层带 (被动走滑断层带, 断层的活动是由于印度大陆和欧亚大陆的碰撞作用引起), 调节或传递昆仑山以北、阿尔金山以东地区的挤压变形或逆冲位移。由于柴达木盆地在新生代发生了明显的挤压构造变形, 不断地吸收挤压构造变形量, 导致阿尔金断裂的左旋走滑位移量自西南向东北方向不断减小, 这样位移量大小的争议自然也可以得到解决。另外, 阿尔金断裂的左旋走滑活动导致塔里木盆地整体的顺时针旋转的认识自然也是值得商榷的。

此外, 世界上著名的走滑断层还有新西兰阿尔卑斯断层 (Alpine fault) (Laubscher, 1971; Ratschbacher, 1986); 苏格兰大格林断层 (Great Glen fault) (Kennedy, 1946); 死海断层 (Dead Sea fault) (Bartov et al., 1980); 阿拉斯加迪那里断层 (Denali fault) (Amand, 1957); 阿富汗恰曼断层 (Chaman fault) (Wellman, 1966); 红河断层 (Tapponnier et al., 1982) 等。

## 4 讨论

在对目前走滑断层研究成果调研的基础上, 文章进一步对走滑断层的发展历程、研究成果进展进行梳理, 方便读者对于走滑断层相关内容研究能够追根溯源;但由于篇幅有限, 文章只对走滑断层研究进展的主要成果进行了论述。通过对走滑断层分类方案的调研, 研究发现目前提出的走滑断层的分类方式虽多, 但是都具有一定的局限性, 因此, 笔者从走滑断层的形成机理出发, 依据走滑断层的动力学机制对走滑断层进行分类,



并提出新的分类模式,旨在从走滑断层形成的根源出发对走滑断层进行解释,希望能进一步提高对走滑断层不同类型的认识。

走滑断层位移量的研究是目前走滑断层研究的重点和难点,通过对相关文献调研发现,目前对走滑断层位移量的研究主要是通过对断裂带中某一段地质参考点(构造带和岩相带等)、岩石学特征分析以及同位素年代学研究等方法进行测量的(许志琴等,1999;Ritts and Biffi,2000;Meng et al.,2001;Zhang et al.,2001;Cowgill et al.,2003;任收麦等,2003;Liu et al.,2007;李海兵等,2007;房璐等,2017)。一方面,这些方法本身具有一定程度的误差;另一方面,更主要的是已有的走滑断层位移量分布的基础理论还存在局限,这是目前对同一条断层走滑位移量存在较大争议的主要原因。

经研究表明,大位移量的走滑断层均是被动走滑断层(包括转换断层和变换断层)。通常转换断层主要传递板块间扩张(大陆裂谷或洋中脊)或汇聚(大洋板块俯冲或大陆碰撞)产生的位移,由于转换断层位移的传递对象只有两个(目前还未发现连接两个以上板块边界的转换断层),同一条转换断层(如圣安德列斯断层)的走滑位移量在不同段落是一致的,因此,走滑位移量的认识也很少存在争议(如圣安德列斯断层)。而变换断层大多传递多个构造的变形(或位移),即大位移量的变换断层一般都是复式变换断层,同一断层(或断层带)的走滑位移量在不同段落本身是不一致的(如阿尔金断裂、郯庐断裂),如果按走滑位移量一致的观点去研究自然会产生争议。随着复式变换断层的概念被学术界广泛接受后,有关走滑断层位移量的争议可能会逐渐减少乃至消失。

## 5 结论

(1) 走滑断层概念的发展经历了比较漫长的历史,走滑断层在几何学、运动学、动力学及其构造意义等方面取得了重要认识和进展,但其分类及其成因机制分析还存在一定的局限性。

(2) 文章从走滑断层的形成机理出发,依据动力学机制对走滑断层进行分类,提出了新的动力学分类方案,将走滑断层分为两大类:主动走滑断层和被动走滑断层。被动走滑断层划分为转

换断层和变换断层两个亚类,其中变换断层可进一步划分为单式变换断层和复式变换断层。转换断层不同段落的走滑位移量理论上是一致的,而复式变换断层传递多个构造的变形(或位移),同一断层(或断层带)的走滑位移量在不同段落本身是不一致的。

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