

*HYDROCHEMISTRY OF KARST SPRINGS FROM TWO  
CARBONATIC UNITS IN ZAGROSIDES OF IRAN*

*E. Raeissi\**

*F. Moore\**

*\*Department of Geology, College of Sciences,  
Shiraz University. Shiraz, Iran.*

**ABSTRACT**

119 Karst Springs, all with electric conductivity of less than 500  $\mu\text{m}/\text{cm}$ , discharging two carbonatic aquifers in Zagrosides of Iran are chosen and studied. The results indicate that the lithology of each Unit is well reflected in the chemistry of springs and the overall mineralogical composition of the aquifer can successfully be reconstructed from the hydrochemistry of the springs using a mass balance approach. The comparison of the chemistry of springs reveals that the significant difference are due to the dolomitic nature of one of the studied aquifers.

**INTRODUCTION**

The hydrochemistry of ground water is mainly a function of the lithology and mineral composition of the aquifer that it passes through. In fact differences in the chemistry of aquifers is the main reason behind the observed differences in the ground water hydrochemistry (Flint, 1971; Palmer & Cherry 1985, Rogers 1989). In carbonatic units, as infiltrating rain and meltwater passes through pore spaces and secondary structures of various size, reaction takes place between the  $\text{CO}_2$  bearing ground water and the carbonatic host. As a result the chemistry of ground water continuously changes as a function of time and the distance travelled by the water. Subsequently considerable dissolution of carbonatic and dolomitic units will also takes place. There is no doubt that if other soluble phases such as evaporitic minerals are also present as interlayers or some other form, these too will react and modify the major-ion chemistry of the

ground water. This is why the chemistry of ground water is traditionally taken by the hydrogeologists as a good representation of the lithology of the aquifer through which the water flows. This approach is preferred to direct analysis of the geological units for the following reasons.

- a) The lithology of sedimentary units is usually so variable that a large number of samples should be analysed in order to obtain an average composition.
- b) The exposed and available sections of the formations are generally weathered and thus analysis would not reflect the true chemistry of the fresh units.
- c) Sampling from depth is expensive and uneconomical.

In this paper the hydrochemistry of springs discharging from two carbonatic units in Zagrosides of Iran, namely Sarvak and Asmari-Jahrum Formations is studied and compared.

### ***GEOLOGICAL SETTING***

The stratigraphy of Zagros sedimentary sequence has been described in detail by James & Wynd (1965) and Falcon (1974). The stratigraphy of the study area is presented in fig 1. The formations of interest in a decreasing order of age consist of Kazhdomi Formation (Bituminous shale of Aptian to Cenomanian), Sarvak limestone (Albian to Turonian), Gurpi shales (senonian to Maestrechtian), and Asmari. Jahrum limestone (Paleocene to Early Miocene).

Sarvak Formation with a thickness of about 800 m is a neritic, massive feature forming limestone containing in some parts siliceous nodules and thin interlayers of marl. The Undifferentiated Asmari-Jahrum Formation with a thickness of 700 to 900 m in the study area is also a massive feature forming dolomitic limestone with shaly interlayers.

Considerable karstification occurs both in Sarvak and Asmari-Jahrum Formations and many karst springs can be found at the local base levels of erosion, especially at the contact between the carbonatic units and the

impermeable marly and shaly formations that overlie and underlay these formations.

## DATA AND METHODS

72 karst springs from Asmari-Jahrum Formation and 47 karst springs from Sarvak formation were sampled. Fig. 1. represents the location of the sampled springs. The electric conductivity of all sampled springs is below 500  $\mu\text{m}/\text{cm}$  and care was taken to select those springs that have no apparent connection with the adjacent evaporitic Formations. PH and electric conductivity of the samples was measured and the major ions Na, K, Mg, Ca,  $\text{HCO}_3$ ,  $\text{SO}_4$  and Cl were also determined using standard wet chemical methods.

*Table 1 represents the mean, maximum and minimum of the measured ions in each formation. The contribution of each ion to the total dissolved solids is also presented in percent values.*

## HYDROCHEMISTRY OF ASMARI-JAHRUM SPRINGS

Table 1 indicates that 85.7% of the total dissolved ions consist of Ca, Mg and  $\text{HCO}_3$ , the rest being made of  $\text{SO}_4$ , Cl, Na and K in that order.

The correlation coefficient of measured cations and anions are calculated using the student t-test and the results are presented in table 2. It is clear that in Asmari-Jahrum springs a good correlation exists between  $\text{HCO}_3$  and Ca and Mg;  $\text{SO}_4$  and Ca, Mg, Na; and between Cl and Na.

Using a mass balance approach and assuming that all bicarbonate is used up by Mg and Ca, and that excess Mg and Ca is consumed by  $\text{SO}_4$  anion, and also considering that Na is mainly used up by Cl, the source minerals of Asmari-Jahrum aquifer were reconstructed from the data in table 1 and are presented in table 3. Hence, the mass balance approach reveals that the mineralogical composition of the aquifer should predominantly consist of calcite and dolomite with minor amounts of gypsum, halite and probably epsomite. Although this is a rather perfect reflection of what is seen in the field, but it must be noted that this is a somewhat

Table 1: Mean, Maximum, Minimum, Sum and percent of concentration of measured ions from karst springs of Sarvak and Asmari-Jahrum Formations.

	Asmari Formation (N=72)	Sarvak Formation (N=47)
CA:		
MEAN	2.196	2.233
MAXIMUM	3.800	3.700
MINIMUM	1.000	0.800
% CA	28.400	31.790
MG:		
MEAN	1.386	1.059
MAXIMUM	2.400	1.800
MINIMUM	0.200	0.035
% MG	17.928	15.074
NA:		
MEAN	0.239	0.213
MAXIMUM	1.050	1.120
MINIMUM	0.029	0.026
% NA	3.096	3.026
K:		
MEAN	0.017	0.014
MAXIMUM	0.097	0.110
MINIMUM	0.000	0.000
% K	0.219	0.203
HCO <sub>3</sub> :		
MEAN	3.046	2.668
MAXIMUM	4.000	4.000
MINIMUM	1.700	1.300
% HCO <sub>3</sub>	39.394	37.985
SO <sub>4</sub> :		
MEAN	0.582	0.531
MAXIMUM	1.900	1.824
MINIMUM	0.000	0.060
% SO <sub>4</sub>	7.530	7.560
Cl:		
MEAN	0.265	0.306
MAXIMUM	0.600	0.600
MINIMUM	0.020	0.100
% Cl	3.433	4.362
PH:		
MEAN	7.387	7.544
MAXIMUM	8.430	8.140
MINIMUM	6.400	6.800
EC:		
MEAN	377.514	347.532
MAXIMUM	494.000	486.000
MINIMUM	244.000	220.000

Concentration in m.eq/l; PH in Units; Ec is electric conductivity in  $\mu\text{mhos/cm}$ .

Table 2: Correlation Coefficient between the chemical Variables

		HCO <sub>3</sub> <sup>-</sup>	SO <sub>4</sub> <sup>2-</sup>	Cl <sup>-</sup>
As-Ja	Ca <sup>++</sup>	0.58*	0.27*	0.12
	Mg <sup>++</sup>	0.39*	0.46*	-0.08
	Na <sup>+</sup>	0.12	0.74*	0.52*
	K <sup>+</sup>	0.08	0.03	0.044
Sa	Ca <sup>++</sup>	0.67*	0.02	-0.06
	Mg <sup>++</sup>	0.06	0.37*	0.29*
	Na <sup>+</sup>	0.25	0.37*	0.39*
	K <sup>+</sup>	0.18	0.21	0.28

simplified picture of the true situation, as there are undoubtedly minor minerals of various composition that are present in insignificant amounts and are not accounted for here. The original composition of the rain water should also be considered.

The relation between electric conductivity and the concentration of various ions is presented in fig 2. It can be seen that Ca, Mg, HCO<sub>3</sub>, Na and SO<sub>4</sub> show good correlation with electric conductivity while Cl and K reveal a very poor correlation. Using a multiple regression method the relation between the dissolved ions and electric conductivity is calculated as logarithmic, semi logarithmic and non logarithmic equations. The F test and the coefficient of determination (R<sup>2</sup>) show that the non logarithmic calculated equation is the best. The significant factor of F was also calculated and it was indicated that K and Cl hold no significance and so were omitted from the equation. After all these considerations, the equation of the line defining the relation between electric conductivity and the dissolved ions may be expressed as follow.

$$Ec = 41.42 Ca + 39.25 Mg + 72.85 Na + 47.55 HCO_3 + 46.56 SO_4 + 42.1$$

#### HYDROCHEMISTRY OF SARVAK SPRINGS

As it can be seen from table 1, Ca, Mg and HCO<sub>3</sub> comprise 84.80%

of the total dissolved solids, the rest being made of  $\text{SO}_4$ , Cl, Na and K respectively. The same mass balance approach is used and the results are presented in table 4. It is obvious that the mineralogical composition of Sarvak aquifer should predominantly be made of calcite with some minor amounts of gypsum, epsomite, dolomite and halite. Here again we can see a perfect agreement between the reconstructed mineralogical composition and field observations.

*Table 3: Reconstruction of Source Minerals for Asmari-Jahram Spring*

	Ca + Mg	Na	K	$\text{HCO}_3$	$\text{SO}_4$	Cl
Concentration in Spring water	2.20+1.39	0.24	0.02	3.05	0.58	0.27
$\text{CO}_3\text{Ca} + \text{CO}_3\text{Mg}$	3.05			-3.05		
	0.54	0.24	0.02	0.00	0.58	0.27
$\text{SO}_4\text{Mg} + \text{SO}_4\text{Ca}$	-0.54				-0.54	
	0.00	0.24	0.02	0.00	0.04	0.27
ClNa	0.00	0.24				0.24
	0.00	0.0	0.02	0.00	0.04	0.03

Concentration are in meq/L

*Table 4: Reconstruction of Source Minerals for Sarvak Springs*

	Ca + Mg	Na	K	$\text{HCO}_3$	$\text{SO}_4$	Cl
Concentration of Spring Water	2.23+1.06	0.21	0.01	2.66	0.53	0.31
$\text{CO}_3\text{Ca} + \text{CO}_3\text{Mg}$	-2.66			-2.66		
	0.66	0.21	0.01	0.00	0.53	0.31
$\text{SO}_4\text{Mg} + \text{SO}_4\text{Ca}$	0.53				-0.53	
	0.13	0.21	0.01	0.00	0.00	0.31
ClNa		0.21				0.21
	0.13	0.00	0.01	0.00	0.00	0.11

Concentration are in meq/L

The relation between electric conductivity and the concentration of different ions is presented in fig 3. It can be seen that electric conductivity shows maximum correlation with  $\text{HCO}_3$  while Ca, Na,  $\text{SO}_4$ , K and Mg all show moderate correlation with Ec. Similar statistical treatment of the data as that of Asmari-Jahrum indicates that the non logarithmic equation of the line best defining the relation between Ec and dissolved ions is as follow.

$$Ec = 42.97 Ca + 45.60 Mg + 68.22 Na + 45.18 \text{HCO}_3 + 45.15 \text{SO}_4 + 44.26$$

### COMPARISON OF CHEMICAL CHARACTERISTICS OF ASMARI-JAHURM & SARVAK SPRINGS

The measured ions, PH and electric conductivity of sampled karst springs from both formations were compared using non parametric Kruskal wallis and Mann-Whitney methods (Walpole & Myers). The results are presented in table 5. It can be seen that the differences in electric conductivity, Mg  $\text{HCO}_3$  of karst springs of these two formations is significant while the difference between other determined ions and PH does not bear any important significance. The higher Mg,  $\text{HCO}_3$  and electric conductivity values in Asmari-Jahrum formation seems to arise from its dolomitic lithology.

Table 5: Nonparametric comprison of Chemical Characteristic of Karst Spring from Asmari-Jahrum and Sarvak formations.

	Z	Z	H'	Result of Comparison
Ca	+0.23	0.23	0.052	$\text{Ca}_{\text{Sa}} = \text{Ca}_{\text{As-Ja}}$
Mg	-3.49	3.49*	12.205*	$\text{Mg}_{\text{Sa}} < \text{Mg}_{\text{As-Ja}}$
Na	-1.92	1.92	3.694	$\text{Na}_{\text{Sa}} = \text{Na}_{\text{As-Ja}}$
K	-1.89	1.89	3.560	$\text{K}_{\text{Sa}} = \text{K}_{\text{As-Ja}}$
$\text{HCO}_3$	-3.28	3.28*	10.743*	$\text{HCO}_3_{\text{Sa}} < \text{HCO}_3_{\text{As-Ja}}$
$\text{SO}_4$	-1.20	1.20	1.431	$\text{SO}_4_{\text{Sa}} = \text{SO}_4_{\text{As-Ja}}$
Cl	+1.91	1.91	3.633	$\text{Cl}_{\text{Sa}} = \text{Cl}_{\text{As-Ja}}$
PH	+1.59	1.59	2.532	$\text{PH}_{\text{Sa}} = \text{PH}_{\text{As-Ja}}$
Ec	-2.05	2.05*	4.200*	$\text{Ec}_{\text{Sa}} < \text{Ec}_{\text{As-Ja}}$

\* Indicates that the compared parameters in the two formations are Unequal.

## ACKNOWLEDGMENT

The authors would like to thank the research Council of Shiraz University for financing this research. Thanks are also due to Mr. M. Bostani for his help, in computer data processing.

## REFERENCES

- Falcon, N.L. 1974. Southern Iran: Zagros Mountains in Mesozoic-Cenozoic orogenic belts. Geol. Soc. London, Spec. Pub. 4, p. 199-211.
- Flint, R.F. 1971. Glacial and Quaternary Geology. John Wiley, New York. 892 pp.
- James, G.A. and Wynd, J.G. 1965. Stratigraphic nomenclature of Iranian Oil Consortium agreement area. Bulletin of AAPG, V. 49, No. 12.
- Palmer, C.D. and Cherry, J.A. 1985. Geochemical evolution of ground water in sequences of sedimentary rocks. J. Hydrol. V. 75, pp. 27-65.
- Rogers, J.R. 1989. Geochemical Comparison of ground water in areas of New England, New York and Pennsylvania. Ground water. V. 27, No. 5, pp. 690-712.
- Walpole, R.E. and Myers R.H. 1972. Probability and statistics for engineers and scientists. Macmillan, New York 506 pp.

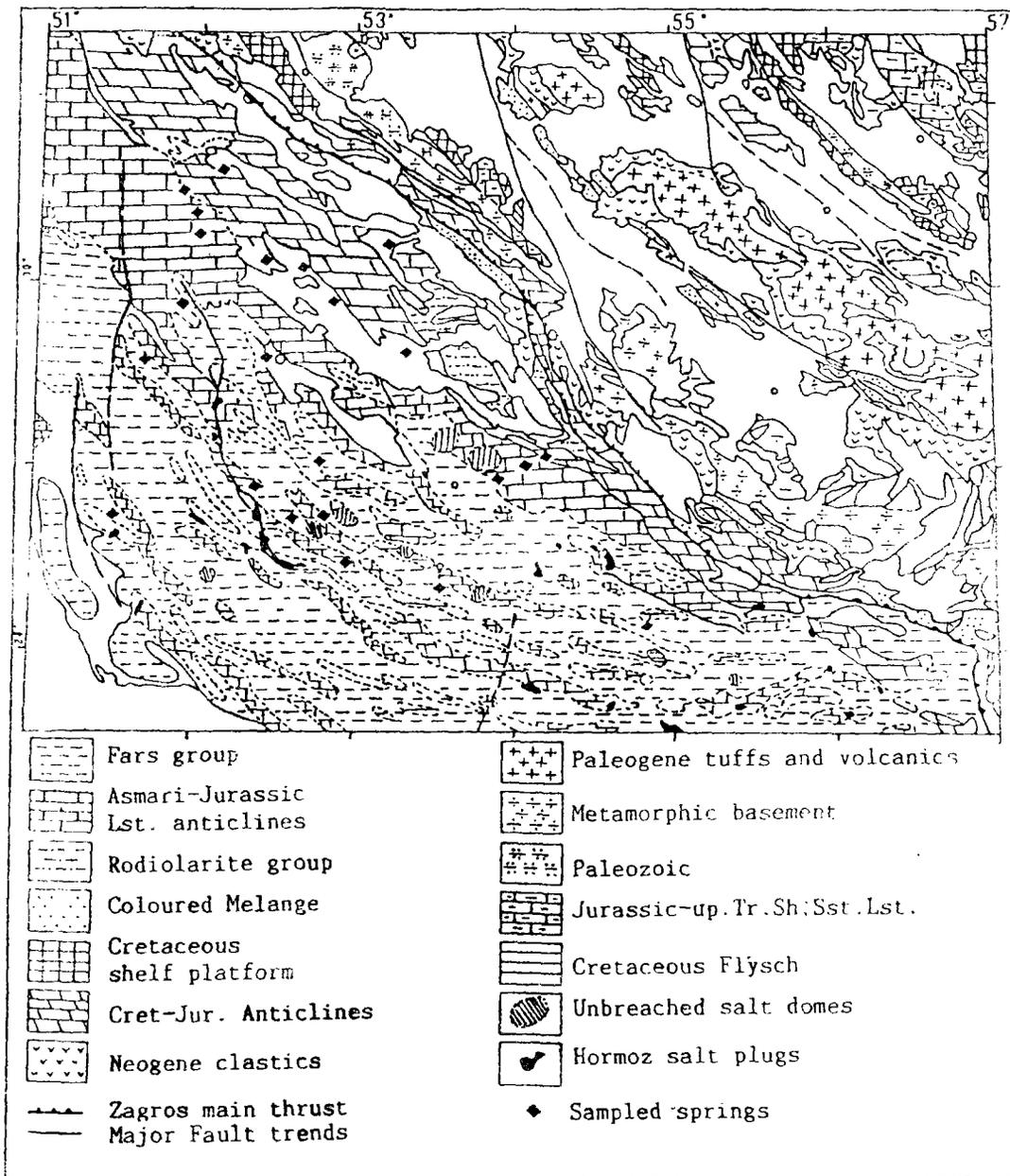


Fig 1: General geological map of the study area showing the location of the sampled springs

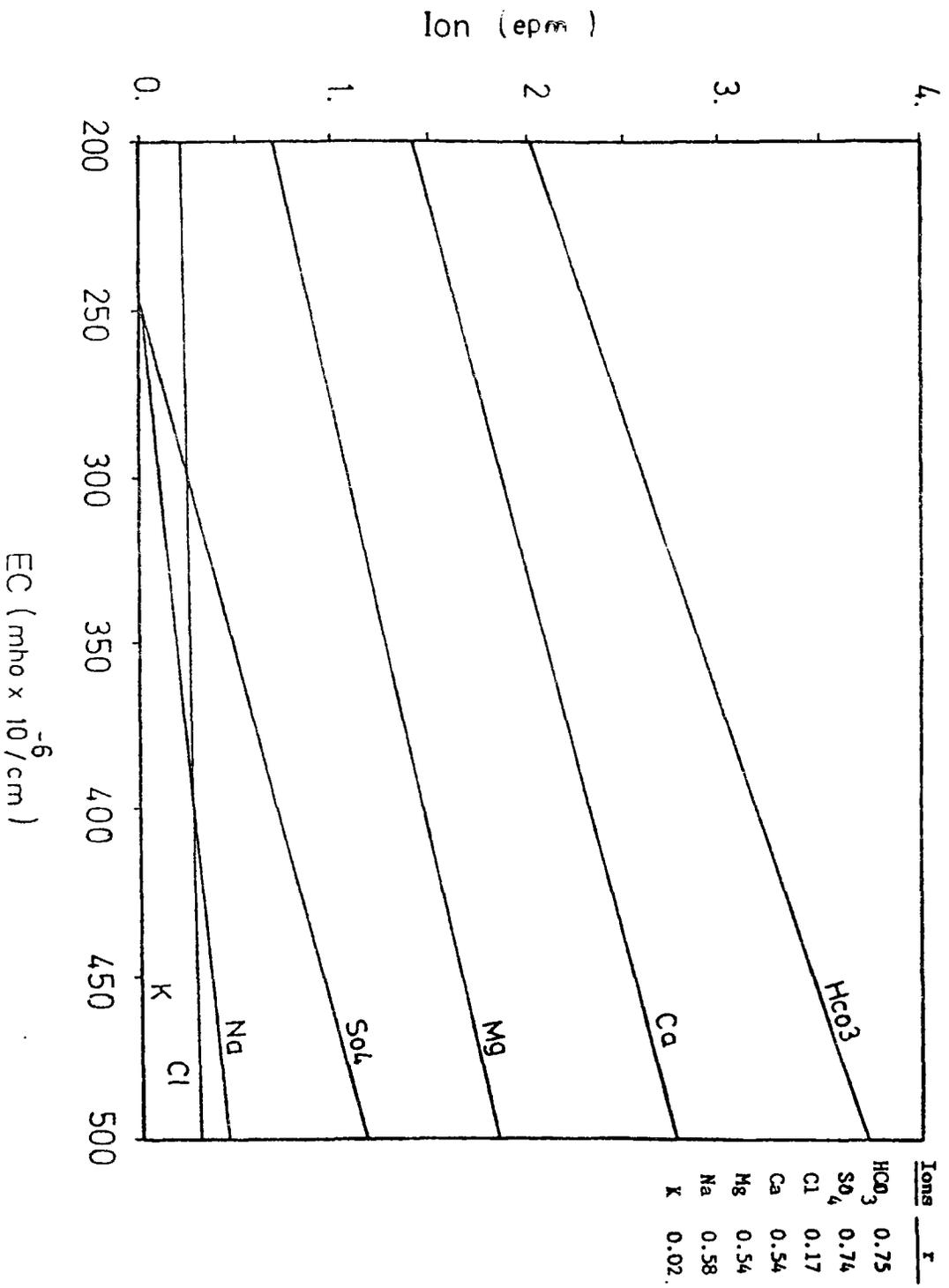


Fig. 2: Relation between electric conductivity and the concentration of various ions in Asmari-Jahrum Korst Springs.

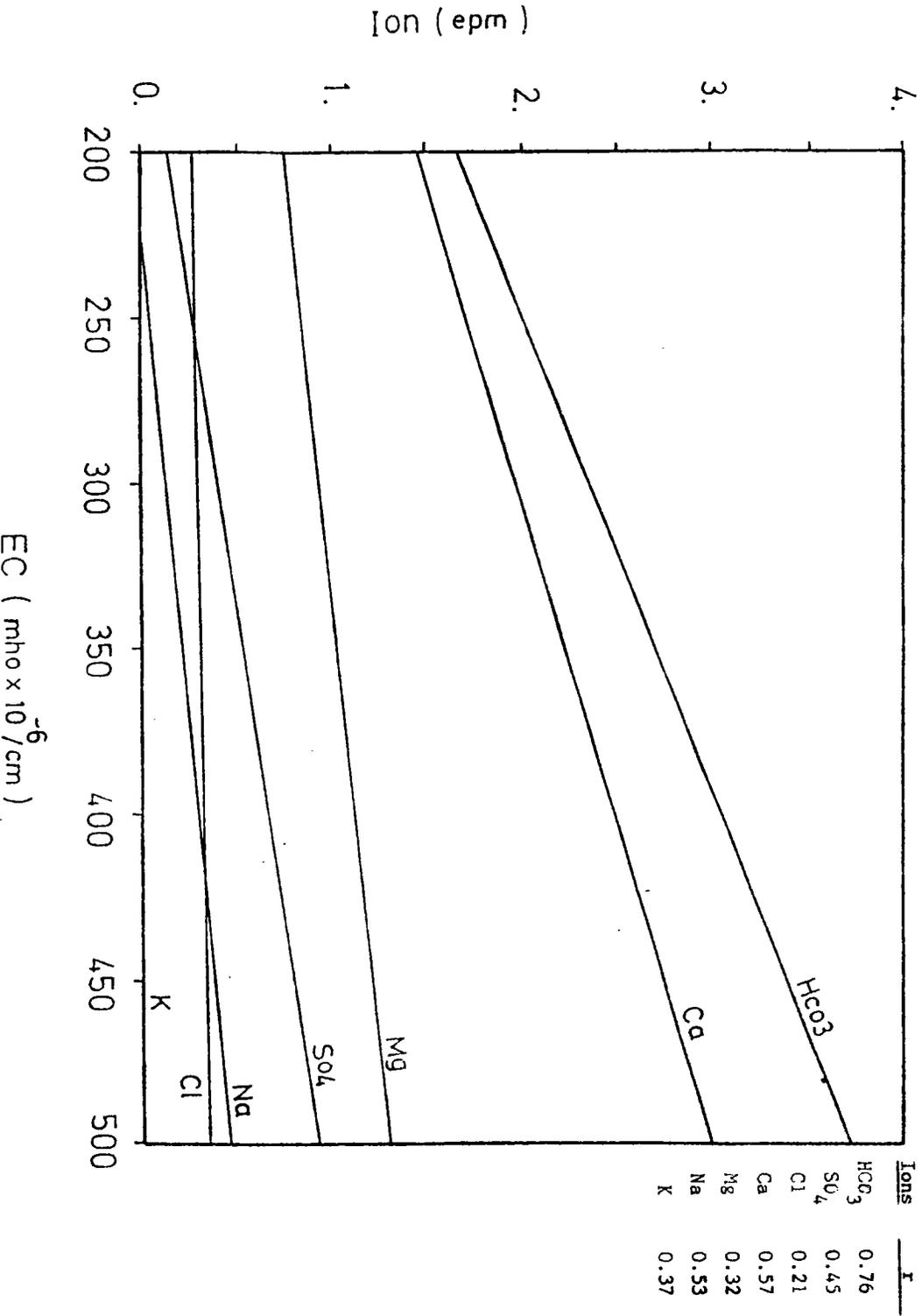


Fig. 3: Relation between electric conductivity and the concentration of various ions in Sorvak Karst Springs.